

**3D hydrostratigraphical modelling of the regional aquifer system of the St. Maurice Delta Complex (St. Lawrence Lowlands, Canada)**

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| Journal:                      | <i>Canadian Water Resources Journal</i>  |
| Manuscript ID                 | TCWR-2016-0059.R2  |
| Manuscript Type:              | Original Paper   |
| Date Submitted by the Author: | 02-Apr-2017  |
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| Keywords:                     | Granular aquifers, Quaternary deposits, 3D geological modelling, High resolution seismic reflection, Champlain Sea   |
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1 **3D hydrostratigraphical modelling of the regional aquifer system of the**  
2 **St. Maurice Delta Complex (St. Lawrence Lowlands, Canada)**

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*For the use of the editors*

**Paper #:** TCWR-2016-0059.R2

**Submitted on:** October 14, 2016

**Accepted on:** April 3, 2017

**Application - Research – Commentary – Book Review:** Research (QC GW SI)

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18 **Abstract**

19 In the central part of the St. Lawrence Lowlands (Mauricie, Québec), Late  
20 Quaternary deglacial events led to the formation of a series of complex granular aquifers,  
21 such as those in the 1) Saint-Narcisse morainic complex, 2) paleodelta formed by the  
22 Saint-Maurice River, and 3) sandy littoral terraces left during marine regression. The  
23 aquifers are an important supply of potable water for most municipalities in the region,  
24 including the City of Trois-Rivières, which is a mid-size city where groundwater  
25 accounts for 46% of the water supply. The main objectives of this study were to build a  
26 3D model of the Quaternary deposits to define the main hydrogeological contexts of the  
27 Mauricie region and to characterize the regional aquifers. The compilation of existing  
28 hydrogeological data led to the selection of 5386 well logs that contained stratigraphic  
29 information of variable quality, ranging from only surficial sediment thickness to  
30 descriptions of fully cored boreholes. To supplement the existing data, fieldwork was  
31 undertaken in areas where few data were available, including 63 km of high resolution  
32 seismic reflection surveys and 34 new boreholes. The final 3D model consists of six  
33 layers, from the bedrock surface to the upper littoral and deltaic sands. The total thickness  
34 of the deposits ranges from zero, on bedrock outcrops, to 150 m beneath central Trois-  
35 Rivières. Taking into account the thickness of the saturated layer and the porosity of the  
36 sand and gravel, the upper unconfined aquifer contains an estimated 364 million m<sup>3</sup> of  
37 water. The 3D model helped refine our understanding of regional aquifers and was used  
38 to identify unexploited aquifers, notably around the Saint-Narcisse morainic complex and  
39 along the St. Cuthbert Fault. The model clarified the regional stratigraphic architecture,  
40 especially topography of the bedrock surface, the lateral extent of Late Quaternary sands  
41 and development of Holocene post-glacial sediment sequences.

42 **Keywords** : Granular aquifers, Quaternary deposits, 3D geological modelling, High reso-  
43 lution seismic reflection, Champlain Sea

## 44 **Résumé**

45 Dans la partie centrale des Basses-Terres du Saint-Laurent (Mauricie, Québec), la  
46 dernière déglaciation a mené à la formation d'aquifères granulaires complexes, tels que  
47 ceux associés à la moraine de Saint-Narcisse, au paléodelta de la rivière Saint-Maurice et  
48 aux terrasses littorales associées à la régression marine. Ces aquifères représentent la  
49 source principale d'eau potable pour la majorité des municipalités de la région, incluant  
50 Trois-Rivières, une ville de taille moyenne où les eaux souterraines contribuent à 46% de  
51 l'approvisionnement. L'objectif de cette étude était de développer un model 3D des  
52 dépôts quaternaires afin de définir les principaux contextes hydrogéologiques et de  
53 caractériser les aquifères régionaux de la Mauricie. Le recensement des données  
54 hydrogéologiques existantes a permis de compiler les résultats de 5386 puits et forages  
55 contenant des informations stratigraphiques de qualité variable. Les données existantes  
56 ont été complétées par de nouveaux relevés, incluant 63 km de sismique réflexion haute  
57 résolution et 34 forages. Le model 3D final est constitué de sept couches, de la surface du  
58 rock aux sables littoraux et deltaïques superficiels. L'épaisseur total des sédiments  
59 quaternaires varie de nulle sur les affleurements rocheux à 150 m à l'ouest de la ville de  
60 Trois-Rivières. Le volume total des sables superficiels, dans les limites de la ville de  
61 Trois-Rivières, a été estimé à 4.7 milliards de m<sup>3</sup>. En tenant compte de l'épaisseur de la  
62 zone saturée et de la porosité moyenne du sable, la nappe libre contiendrait un volume de  
63 364 millions de m<sup>3</sup> d'eau. Le model tridimensionnel a permis d'approfondir notre  
64 connaissance des aquifères régionaux et d'identifier des aquifères non exploités,  
65 notamment en marge de la moraine de Saint-Narcisse et le long de la faille de St-  
66 Cuthbert. Le model a également permis de préciser l'architecture des dépôts et la  
67 stratigraphie régionale, particulièrement en ce qui concerne la topographie de la surface  
68 du rock, l'extension latérale des dépôts pléistocènes et la séquence des sédiments  
69 postglaciaires.

## 70 **Introduction**

71 The current landscape of the St. Lawrence Valley is mainly the result of erosion  
72 and deposition during successive Late Quaternary events. Deglaciation was followed by a

73 2000-year long incursion of an arm of the Atlantic Ocean into the St. Lawrence, known  
74 as the Champlain Sea. During this marine episode, a series of coastal, deltaic and  
75 offshore depositional systems developed in the valley. Following the regression of the  
76 Champlain Sea and the drainage of its successor basin (Lake Lampsilis), fluvial incision,  
77 sedimentation and terrace formation by the early St. Lawrence River and its tributaries  
78 became the main centres for active geological processes in the valley. Elsewhere in the  
79 region, this late-glacial period was characterized by the development of peatlands,  
80 notably on the large flat surfaces of the St. Maurice paleodelta.

81         The last deglacial events led to the formation of a series of complex granular  
82 aquifers, such as the Saint-Narcisse morainic complex, the Saint-Maurice River  
83 paleodelta, and the stepped sandy terraces formed along the shores of the Champlain Sea  
84 and along the St. Lawrence River and tributaries. The characteristics and distribution of  
85 these aquifers are partially known through various hydrogeological studies conducted by  
86 municipalities, government agencies and private consulting firms. Unfortunately, these  
87 studies had limited application regionally because they were generally local in scope and  
88 scale. The first regional groundwater map of the Mauricie region was constructed by  
89 McCormack (1983), who noted the good aquifer potential of the surficial and near  
90 surface sediments of the north shore, while the aquifer potential of fractured bedrock was  
91 deemed lower. While he identified buried valleys filled with sand and gravel deposits, he  
92 had not evaluated the hydrogeological potential of these confined aquifers. In the nearby  
93 Lake Maskinongé area, Denis (1974) also mentioned that the bedrock formations of the  
94 region had a low aquifer potential, while the granular aquifers were more important.  
95 Nonetheless, and despite these preliminary attempts to map the local aquifers, there still  
96 lacks comprehensive knowledge of the different hydrogeological units and their surface  
97 or subsurface distribution. Given the crucial importance of groundwater as a source of  
98 potable water in the region, both from quality of life and economic viewpoint, we  
99 determined that a more thorough characterization and assessment of regional aquifers  
100 was needed.

101         Groundwater has been exploited in the Mauricie region since the nineteenth  
102 century to supply municipal water networks and for commercial purposes. In the town of

103 Yamachiche, for example, an aqueduct was built in 1873 to extract spring water  
104 renowned for its high quality. In the early twentieth century, a brackish water source  
105 located in Saint-Léon-Le-Grand became known for its healing properties. Nowadays, the  
106 granular aquifers of the Mauricie region are the source of potable water for most of its  
107 municipalities, including the City of Trois-Rivières, which is one of few small urban  
108 centers in Canada where groundwater constitutes a large part of the drinking water supply  
109 (46%). In addition, the use of geothermal energy for heating and cooling has expanded  
110 during the last decade. Given the relatively high groundwater temperature (10°C in Trois-  
111 Rivières) and the high aquifer potential of the region, it is likely that geothermal systems  
112 will continue to expand rapidly in the near future to capture thermal energy from the  
113 warmer groundwater.

114         Given the lack of comprehensive knowledge about the regional aquifers and the  
115 crucial role of groundwater in the Mauricie region, a regional groundwater study was  
116 initiated by Université du Québec à Trois-Rivières (UQTR) researchers in collaboration  
117 with governmental agencies and regional partners (including the regional council of  
118 elected officials, regional municipalities, the City of Trois-Rivières and watershed  
119 agencies). The study was conducted between 2009 and 2013 as part of the Groundwater  
120 Knowledge Acquisition Program (PACES), sponsored by the Quebec Ministry of the  
121 Environment (MDDELCC). One of the main objectives of this study was to build a 3D  
122 model of Quaternary deposits to better define the hydrogeological framework for the  
123 Mauricie region and characterize the aquifer properties.

#### 124 **Study area**

125         The study area (3 900 km<sup>2</sup>) is located in the southwest of the Mauricie region, between  
126 Montréal and Québec (Figure 1). The total population (2011) is 223 200 inhabitants, distributed into  
127 two main urban centers (Trois-Rivières and Shawinigan) and one regional municipality  
128 (Maskinongé). The northern part of the study area lies in the Laurentian Hills, and is mostly  
129 forested, whereas the southern part is underlain by the relative flat terrain of St. Lawrence platform  
130 and is occupied by farmlands, except for the urban area of Trois-Rivières. The shield terrain of the  
131 Laurentians is characterized by steep-sided knolls and knobs, intersected by linear through-valley

132 systems, often occupied by chains of lakes and logging roads, while the St. Lawrence platform is  
133 underlain by Paleozoic sedimentary rocks covered by thick Quaternary sediments, where  
134 agriculture is the main economic activity. The main rivers (Maskinongé, du Loup, Yamachiche and  
135 St. Maurice) flow southward from the Shield terrain to the lowlands. The hydrological regime of  
136 these catchments is characterized by spring snowmelt and low flow during summer with episodic  
137 storm events, typical of continental sub-humid, subpolar climates. Catchments receive between 850  
138 and 1200 mm of precipitation annually, evenly distributed during the year.

139 **Figure 1. Location and digital elevation model of the study area. The four main**  
140 **regional hydrogeologic contexts are shown: the Laurentian Hills, Saint-Narcisse**  
141 **morainic complex, Marine clay plain and paleodelta of the St. Maurice River.**

142 First, a compilation map of bedrock formations was realized as part of this project, based on  
143 the studies conducted by Clark et Globensky (1976), Globensky (1987) and Nadeau and Brouillette  
144 (1995). The lithological composition of the Grenville Province, a major subdivision of the Canadian  
145 Shield, consists mainly of igneous and metamorphic rocks, such as gneiss, orthogneiss, paragneiss,  
146 migmatite and marble, along with intrusive rocks such as gabbro, anorthosite, monzonite and  
147 monzogranite. The Grenville rocks are moderately folded in the allochthonous monocyclic belt and  
148 strongly folded in the polycyclic belt. The rocks of the Morin Terrane, in the south and west of the  
149 study area, were slightly deformed during the Grenville orogeny to form a large regional syncline.

150 The sedimentary rocks of the lowlands consist mainly of Ordovician sandstones,  
151 limestones and shales deposited in marine environments. Sandstones and limestones (Black River  
152 and Trenton groups) are exposed in isolated outcrops in the northern part of the lowlands, while the  
153 shales (Utica and Lorraine groups) are buried under considerable thicknesses of Quaternary  
154 sediments near the St. Lawrence River. These shales contain significant concentrations of natural  
155 gas as well as oil seeps and have been fully described in geological reports (Globensky, 1987). The  
156 fractured bedrock aquifers, both in the shield and the lowlands, only supply private wells.

157 Secondly, an updated map of surficial deposits was constructed (Figure 2) based on results  
158 of previous studies by Denis (1974), Bolduc (1999a, b) and Lamarche (2005). The  
159 Laurentian Hills are characterized by a thin till veneer with numerous rock exposures.  
160 Along the southern edge of this hilly region, valleys are floored by silty or silty-sandy

161 marine sediments that commonly overlie ice-proximal sandy gravels that are buried. The  
162 Saint-Narcisse Moraine lies along the southern margin of the Canadian Shield, but  
163 slightly beyond in the St. Maurice River valley. The moraine was emplaced by a re-  
164 advance of the Laurentide Ice Sheet during the early cold phase of the Younger Dryas. In  
165 stratigraphic sections in the moraine, a variety of sediment facies, including glacially  
166 entrained clay and till wedges, are exposed along with proximal glaciomarine deposits,  
167 melt-out till and ice-marginal outwash (Occhietti, 2007). As the Laurentide Ice Sheet  
168 retreated to the north, the Champlain Sea flooded the Saint Lawrence Valley up to almost  
169 200 m asl, an event that led to the deposition of a thick silt and clay cover in the deeper  
170 offshore basins and littoral and sublittoral sands in shallower areas. Deltas, large and  
171 small, also formed in valleys at the mouth of rivers entering the Champlain Sea.

172 **Figure 2. Quaternary geology and location of the new geophysical surveys and**  
173 **boreholes. The location of the existing boreholes and wells (5 386) used to build the**  
174 **model are not shown. See Table 2 for a detailed description of the units.**

175 The Quaternary stratigraphy of the area was described by Gadd and Karrow  
176 (1959), Denis (1974), Occhietti (1980, 2007), Ferland and Occhietti (1990) and Clet and  
177 Occhietti (1996). The most complete stratigraphic sections were observed along the St.  
178 Maurice River. The base of the Quaternary sequence is characterized by a fluvial  
179 sediment complex directly overlying bedrock. This unit is correlated with the Lotbinière  
180 Sand (Hardy and Lamothe, 1997), rather than the St. Pierre Sediments as originally  
181 thought by Gadd and Karrow (1959). The overlying unit, containing a nine meter-thick  
182 varve series, is known as the St. Maurice Rhythmites since it corresponds to a lacustrine  
183 environment younger than that of the Deschaillons Varves (Besré and Occhietti, 1990).  
184 The rhythmites are overlain in the regional stratigraphic record by the stratified Vieilles-  
185 Forges Sands, with three lithozones (lacustrine, deltaic and proglacial). This sequence is  
186 capped by the Gentilly Till, which is in turn overlain by the usual postglacial sequence  
187 containing glaciofluvial sediments, Champlain Sea sediments, Lake Lampsilis sediments,  
188 and alluvial and organic sediments (Table 2).

## 189 **Methods**

### 190 *Compilation of existing boring data*

191 Existing hydrogeological data were compiled from the sources presented in Table  
192 1. All the municipalities in the study area were visited to obtain existing hydrogeological  
193 studies. The databases from the major water supply providers were integrated, including  
194 those from the cities of Trois-Rivières and Shawinigan and from a regional water supply  
195 network (Régie d'Aqueduc de Grand-Pré). The compilation of existing hydrogeological  
196 data led to the selection of 5386 logs for the modelling that contained some stratigraphic  
197 information. Among these entries, 2185 reached bedrock and 3201 ended within  
198 Quaternary sediments. The reliability of the existing data was evaluated based on the  
199 criteria proposed by Ross et al. (2005) and Chesnaux et al. (2011), and hydrostratigraphic  
200 codes were assigned to each entry in the database. When available, the hydraulic  
201 properties were integrated to the database.

### 202 *Field work*

203 To supplement the existing data, fieldwork was undertaken in areas where data were  
204 sparse or the stratigraphic interpretations were uncertain (Figure 2). The new fieldwork  
205 included:

- 206 • 75 electrical resistivity surveys (using a TX-II transmitter and PP GRX-832  
207 receiver). The interpretations were validated with borehole logs located within  
208 500 m of the survey line;
- 209 • 50 seismic refraction surveys (using a Smartseis ST with 24 geophones);
- 210 • 63 km of high resolution seismic reflection surveys (using the Minivibe);
- 211 • 11 cone penetration tests (using a Geotech 605D);
- 212 • 23 boreholes were drilled by different techniques (rotation, percussion, pionjar  
213 and rotasonic). Monitoring wells were installed in most of the boreholes. The  
214 sediments were sampled at regular intervals or at the transition between two  
215 stratigraphic units. Particle size analysis (Analysette 22 MicroTec) was carried  
216 out on 425 samples. The hydraulic conductivity of each sample was estimated

217 from the grain size analysis. Radiometric dating (Beta Analytic Inc) was  
218 conducted on three samples of shells or organic material. Water level  
219 measurements and pumping tests were carried out in most wells.

220 Finally, the detailed Quaternary geology mapping was carried out throughout the  
221 region, but more specifically in areas along the boundary of maps created by different  
222 authors. The data collected in the field allowed us to produce a relatively uniform  
223 compilation map of surficial sediments for the study area.

#### 224 *Geostatistical modelling approach*

225 Cokriging was used to create gridded surfaces connecting the logs according to  
226 the elevation of stratigraphic contacts. This method of geostatistical estimation considers  
227 the spatial covariance and assigns a different weight to the observations based on their  
228 reliability. A digital elevation model (DEM) was used to determine the elevation of all  
229 borehole sites included in the database. The model also considered the outcrops of all  
230 deposits and the known stratigraphy shown on geological sections. Additionally, some  
231 adjustment points were manually added to the model at locations where available data  
232 were either insufficient or inadequate. For example, the range of the variogram was  
233 sometimes insufficient to properly interpolate grid values up to the boundaries of the  
234 study area and some adjustment points were added to fill in the void. The adjustment  
235 point were not included in the calculation of the error for each layer. Virtual sections  
236 were also generated using Hydro GeoAnalyst software to link, at the regional scale,  
237 similar facies where data are sparse or widely spaced (Ross et al. 2005). A final  
238 validation was performed to check whether the model-generated layers were contained in  
239 the volume between the rock surface and the ground surface. Figure 3 shows the  
240 flowchart of the GIS operations used to model the top surface of each hydrogeological  
241 unit.

242 **Figure 3. Flowchart of the GIS operations used to model the top surface of each**  
243 **hydrogeological unit.**

244 **Results**

245 *Three-dimensional mapping of Quaternary deposits*

246 The model consists of seven layers representing the topography of the bedrock  
247 surface and the main hydrogeological units encountered in the study area. Because the  
248 model is being constructed at a regional scale, some units were merged with other in  
249 order to give precedence to the more extensive and thicker aquifers and aquitards.

250 *Topography of the bedrock*

251 A total of 1584 observations from boreholes and geophysical data were used to  
252 model the surface of the bedrock. Before the addition of adjustment points, the root mean  
253 squared error between the observed and predicted values was 5.0 m. In addition, 484  
254 bedrock outcrop areas were used as control to model the bedrock surface.

255 In the lowlands, near the St. Lawrence River, the smooth bedrock surface lies as  
256 much as 100 m below present sea level. However, the bedrock surface rises to the north-  
257 east of the St. Maurice River and on the south shore of the St. Lawrence River. On a  
258 south-north axis, the bedrock surface rises slowly northwestward toward the Piedmont,  
259 except along the St. Cuthbert Fault, which is the locus of a NE-trending linear depression  
260 in bedrock. The maximum elevation of 540 masl is reached in the Laurentian Hills, north  
261 of the municipality of Saint-Alexis-des-Monts.

262 Figure 4 shows the total thickness of Quaternary deposits. This map was produced  
263 by subtracting the elevation of the DEM surface from the elevation of the bedrock  
264 surface. The thickest deposits reach 150 m in the western part of the City of Trois-  
265 Rivières where the bedrock surface lies at -80 m, an elevation which is quite similar to  
266 that underneath Yamachiche and which is amongst the lowest of the entire central St.  
267 Lawrence Lowlands (Prévôt, 1972).

268 **Figure 4. Total thickness of Quaternary deposits. The thickest deposits, reaching**  
269 **150 m, are in the western part of the City of Trois-Rivières.**

270 *Pre-LGM Quaternary sediments*

271 Pleistocene sediments deposited prior to the Last Glacial Maximum (LGM) have  
272 been observed in sections bordering the St. Maurice River and were also intersected by  
273 drilling in the southeast part of the study area. A total of 85 observations from boreholes  
274 and geophysical data were used to model the upper contact of pre-LGM sediments  
275 (Figures 5a and 6a). Before the addition of adjustment points, the root mean squared error  
276 between the observed and predicted values was 8.6 m. In addition, 3 surficial outcrops  
277 were used to control the modeled surface. Reliable drilling logs containing descriptions  
278 of the unit were only found around the City of Trois-Rivières, where the maximum  
279 thickness of the unit is 80 m. However, it is possible that other remnants of pre-LGM  
280 Quaternary sediments are preserved elsewhere in the region. Given their rare occurrence  
281 in drilling record, no attempt was made to model separately the different pre-LGM  
282 sediment subunits in the 3D hydrostratigraphical model. However, some pre-LGM  
283 subunits could be identified in high resolution seismic profiles. These deposits are  
284 discontinuous and their thickness varies depending on the topography of the bedrock  
285 surface.

286 **Figure 5. 3D model of Quaternary deposits. The model consists of six layers**  
287 **representing the main hydrogeological units observed above bedrock in the study**  
288 **area. The six layers represent the Pleistocene sediments (pre-LGM) (a), LGM and**  
289 **post-Younger Dryas tills (b), glaciofluvial sediments (c), marine silt and clay (d), St.**  
290 **Narcisse morainic complex (e) and littoral, deltaic and alluvial sediments (f).**

291

292 **Figure 6. Thickness grids (m) of Quaternary deposits.**

293 *LGM and post-Younger Dryas tills*

294 Gentilly Till is the regional surficial diamicton (till) deposited during the  
295 Wisconsinan glaciation. In the context of this paper, it includes diamictons that were  
296 deposited during and after the St-Narcisse Moraine episode. Gentilly Till is a sandy-silt

297 diamicton having a Precambrian crystalline rock content ranging from 100% in the Shield  
298 terrains to less than 10% on the Paleozoic platform. A total of 162 observations from  
299 boreholes and geophysical data were used to model the upper contact of Gentilly Till  
300 (Figures 5b and 6b). Before the addition of adjustment points, the root mean squared  
301 error between the observed and predicted values was 3.6 m. In addition, 241 surficial  
302 polygons were used to model the surface of the unit.

303 In the Laurentians, the glacial till cover is discontinuous and its thickness varies  
304 from 0.1 m to 2 m. However, the till can be 6 m thick where mapped as continuous cover.  
305 Along the St-Narcisse Moraine, the till thickness varies between 1 m and 50 m. In  
306 drilling records around the moraine, till wedges (thrust slices in most cases) deposited by  
307 the advancing ice sheet are difficult to differentiate from glaciofluvial sediment bodies  
308 emplaced subsequently and commonly reworked by littoral Champlain Sea processes.  
309 South of the moraine, in the clay plain, the till usually takes the form of a veneer varying  
310 in thickness depending on the topography of the bedrock surface.

### 311 *Glaciofluvial sediments*

312 Sediments carried by glacial meltwater were deposited in the deep valleys of the  
313 Canadian Shield and in bedrock depressions along the St. Cuthbert Fault. In the northern  
314 part of the study area, at elevations above 200 m asl, these sediments are well exposed at  
315 the surface, while in very deep valleys and on the Shield margins, they are almost  
316 continually covered with fine-grained sediments of marine or lacustrine origin. A total of  
317 341 observations from boreholes and geophysical data were used to model the upper  
318 contact of glaciofluvial sediments (Figures 5c and 6c). Before the addition of adjustment  
319 points, the root mean squared error between the observed and predicted values was 2.6 m.  
320 In addition, 117 sediment surface polygons were used to control the modeled surface.  
321 The largest thicknesses were observed under the deep valleys of the Canadian Shield,  
322 under the St-Narcisse Moraine and along the St. Cuthbert Fault.

### 323 *Marine silt and clay*

324 As the St. Lawrence valley had been isostatically depressed by the Laurentide Ice  
325 Sheet, an arm of the Atlantic Ocean, known as the Champlain Sea, inundated the valleys  
326 up to elevations equivalent to 200 m asl (Parent et Occhietti, 1988, 1999). The marine  
327 transgression lasted about 2500 years (between 13 cal ka BP to 10.5 cal ka BP), and the  
328 incursion was characterized by the deposition of thick, sparsely fossiliferous marine  
329 clays. Around 9.8 ka BP, marine waters gradually drained due to the isostatic rebound  
330 and saltwater was replaced by freshwater (Lake Lampsilis). With continued isostatic  
331 adjustment, as well as reduced meltwater influx to the basin, Lake Lampsilis was  
332 gradually drained by the Proto St. Lawrence River (7.5 cal ka BP to 3.0 cal ka BP).

333 A total of 2287 observations from boreholes and geophysical data were used to  
334 model the upper contact of marine and lacustrine sediments (Figures 5d and 6d). Before  
335 the addition of adjustment points, the root mean squared error between the observed and  
336 predicted values was 5.1 m. In addition, 118 sediment surficial outcrops were used to  
337 control the modeled surface. Marine and lacustrine sediments overlie till and glaciofluvial  
338 sediments in the major valleys of the Canadian Shield, as well as most of the St.  
339 Lawrence Lowlands. The thickest units of fine-grained sediments are located around the  
340 municipalities of Yamachiche and Louiseville, where more than 80 m of silt and clay are  
341 recorded in the drilling records. These maximum thicknesses are controlled by a  
342 depression in the bedrock around Lake St. Pierre (Prévôt, 1972). In the paleodelta of the  
343 St. Maurice River, the thickness of subsurface silty clay sediments varies from 0 to 30 m.

#### 344 *St. Narcisse morainic complex*

345 The Saint-Narcisse morainic complex was emplaced during a re-advance of the  
346 Laurentide Ice Sheet from 12.9 to 12.5 cal ka BP, along the southern margin of the  
347 Laurentians where it forms a 750 km long crest (Occhietti, 2007). From west to east, the  
348 moraine successively covers the Precambrian bedrock and then fine-grained Champlain  
349 Sea sediments in the St. Maurice valley. The moraine forms prominent ridges near  
350 Charette and Mont-Carmel. The moraine is composed of reworked clay and till, proximal  
351 glaciomarine deposits and melt-out till and ice-marginal outwash (Occhietti, 2007).

352 A total of 75 observations from boreholes and geophysical data were used to  
353 model the upper contact of the St. Narcisse morainic sediments (Figures 5e and 6e).  
354 Before the addition of adjustment points, the root mean squared error between the  
355 observed and predicted values was 7.5 m. In addition, 41 sediment outcrop polygons  
356 were used to control the modeled surface. In the Lower Mauricie region, the composition  
357 of the morainic complex varies widely, but generally it consists of gravel, coarse sand  
358 and poorly sorted medium sands deposited above a layer of marine sediments. The  
359 thickness of this layer of coarse sediments varies from 10 to 50 m.

360 *Littoral, deltaic and alluvial sediments*

361 During the Holocene, the retreat of the ice sheet was accompanied by regional  
362 isostatic uplift estimated at 9 m/100 years (Lamarche, 2005). The Champlain Sea was  
363 replaced by Lake Lampsilis, a freshwater successor basin that lasted about 1500 years.  
364 The present drainage system was gradually emplaced, thus creating successive terrace  
365 levels as base levels fell. The modeled layer that represents the surface sands includes (1)  
366 all sand units overlying fine-grained Champlain Sea sediments, including coastal and  
367 deltaic silty sands, (2) lacustrine sands and silts of Lake Lampsilis, and (3) marine and  
368 lacustrine deltaic sands of the St. Maurice River (Figures 5f and 6f).

369 The upper sands are exposed at the land surface and therefore was not modeled  
370 separately. The unit covers a total area of 1078 km<sup>2</sup>. The thickest areas of the surface  
371 sand aquifer are located in the western and eastern sectors of the City of Trois-Rivières,  
372 as well as in the localities of Sainte-Angèle-de-Prémont, Saint-Élie-de-Caxton and  
373 Charette. In the Trois-Rivières area, the total volume of the upper aquifer is estimated at  
374 4.7 billion m<sup>3</sup>, 31% of which is saturated according to our subsurface records. The unit  
375 has an estimated porosity of 25%, this means that about 364 million m<sup>3</sup> of water are  
376 contained in the unconfined aquifer. The aquifer has an average hydraulic conductivity of  
377  $6 \times 10^{-4}$  m/s. Groundwater generally flows toward southeast, its level ranging from 60 m  
378 upstream to 6 m at the outlet. Champlain Sea clays form the floor of the upper aquifer,  
379 preventing groundwater from seeping further down through the Quaternary sequence.

380 This aquifer is the most extensively exploited in the region and supplies drinking water to  
381 about half of the population of the City of Trois-Rivières.

### 382 *Regional hydrogeologic contexts*

383 Based on the 3D model, the four main regional hydrogeologic contexts of the  
384 study area were determined from the physiographic, geological and hydrogeological  
385 features (Figure 1). The hydraulic properties of the aquifers pertaining to each context are  
386 presented in Tables 3 and 4.

#### 387 *Laurentian Hills*

388 This hydrogeological context consists of fractured bedrock aquifers with low  
389 productivity (Table 4) along with granular aquifers within glaciofluvial deposits filling  
390 deep valleys. Up to an elevation of 200 masl, these deposits were covered by marine silts  
391 during the Champlain Sea marine transgression. Electrical resistivity surveys revealed  
392 over 60 m of silty clay in the Rivière-du-Loup Valley. Similar contexts were also  
393 encountered in the valleys of the Maskinongé, Yamachiche and Shawinigan rivers.

394 Elsewhere in the Laurentian Hills, deltaic sands and gravels were deposited by  
395 meltwaters at the mouths of major valleys entering the Champlain Sea. The glaciofluvial  
396 and deltaic deposits were reworked by waves and currents forming the high and low  
397 terraces along the piedmont. The aquifers found in these deposits are the main source of  
398 potable water for most municipalities in the area.

399 The geologic log for borehole FE-04-11 (Figure 7), located in the village of Saint-  
400 Élie-de-Caxton, illustrates this typical hydrostratigraphic complex. Directly overlying  
401 granitic gneiss bedrock, at depth of 65.5 m, Champlain Sea sediments from 109 m asl to  
402 140 m asl consist of ice-proximal sandy facies grading upward into a distal clay facies at  
403 134 m asl. The top layer consists of fine prodeltaic sands.

#### 404 *Saint-Narcisse morainic complex and Laurentian foothills*

405 The morainic complex frequently forms double aquifers with an unconfined  
406 aquifer within the surface sands, and a confined aquifer in the glaciofluvial sediments

407 underlying a clay or till aquitard resting on bedrock. These aquifers are relatively small  
408 and hydraulically isolated from each other. Several electric resistivity surveys indicated  
409 an overall thickness of over 60 m. Borehole FE-09-11 is representative of this  
410 hydrostratigraphic setting (Figure 7). The bedrock (garnet gneiss) was reached at a depth  
411 of 87 m. Overlying bedrock is a glaciofluvial unit consisting of sands (fine, medium and  
412 coarse) and gravel, which is overlain by the Champlain Sea clays and silts from a depth  
413 of 44 to 67 m. The overlying units include a coarsening upward sequence of silty fine sand to  
414 gravelly sand with a high hydraulic conductivity associated with the Saint-Narcisse readvance.

415 **Figure 7. Logs of selected boreholes with lithology and grain size. The locations of**  
416 **the boreholes are shown in Figure 2.**

417 The L1-2011 Minivibe line was run perpendicularly across the morainic complex  
418 (Figure 8a). The line begins over Precambrian bedrock in the north, which is buried by  
419 marine clays and thin beach sands. It passes over a narrow bedrock valley before crossing  
420 the moraine (GxT) and entering the marine clay plain. The moraine is locally covered by  
421 a thin layer of sand, and overlies heterogeneous glaciofluvial deposits. The glaciofluvial  
422 sands and gravels rest on glaciotectionized clay and till wedges emplaced by the  
423 advancing glacier.

424 Although the hydrostratigraphy of the moraine is complex, it contains an  
425 important regional aquifer, as most of the small municipalities located along the moraine  
426 and in the adjacent marine clay plain draw their drinking water from it. Moreover, the  
427 moraine acts as a recharge area for the sandy surface aquifers connected to it. New  
428 boreholes and geophysical surveys completed for this project revealed the presence of  
429 double aquifers along the moraine that have yet been unexploited for a municipal water  
430 supply.

#### 431 *Marine clay plain*

432 The fine-grained sediments of the Champlain Sea sequences represent a regional  
433 aquitard that confines the underlying aquifer unit containing groundwater with high  
434 mineral content and natural gas. The aquifer potential of the clay plain is generally low.

435 The municipalities of this area rely on a regional water supply network (Régie d'Aqueduc  
436 de Grand Pré) whose wells are located near the foothills of the Laurentian Hills.  
437 However, there are also a few private wells drawing water from the sandstone and  
438 limestone bedrock.

439 The L7-9-2010 Minivibe lines (Figure 8b) are representative of the clay plain. In  
440 the south part, surficial sediments consist mainly of silt and clay ranging in thickness  
441 from 60-70 m, gradually decreasing further north. The underlying till becomes a thin  
442 veneer over a partly buried rock ridge between kilometres 6 and 8.5. The bedrock surface  
443 then plunges abruptly to form a depression two kilometres wide and 80 metres deep. The  
444 buried depression appears to be filled with about 20 metres of gravel and sand. The  
445 depression is covered by fine-grained marine sediments. This unit could potentially hold  
446 significant quantities of groundwater. The axis of these deposits lie directly over the St.  
447 Cuthbert Fault, which is a regional structure marking the boundary between Precambrian  
448 Shield and sedimentary rocks of the Lowlands. It is possible that similar sand and gravel  
449 aquifers are present elsewhere along the fault.

450 **Figure 8. High resolution seismic survey lines L1-2011 (a) and L7-9-2010 (b). The**  
451 **profile interpretations are shown below the seismic profiles. See Figure 2 for**  
452 **location of the seismic lines and Table 2 for a detailed description of the units.**

453 *Paleodelta of the St. Maurice River*

454 The paleodelta of the St. Maurice River is the largest aquifer in the region. The  
455 delta forms a wide sand-dominated fan over a thin layer of silty marine sediments on the  
456 east side of the river, and over a thick layer of pre-LGM deposits on the west side. Over  
457 the northern part of the delta, the unconfined surficial aquifer is thin and is less  
458 productive and not exploited, except for erosion channels where the sediment is thicker  
459 and the sands better sorted. The surficial aquifer has a much greater thickness in the  
460 southern part of the delta. Borehole FE-16-12 (Figure 7) contains almost 40 m of sand of  
461 varying grain sizes. A wood sample found at 29.8 m depth in the upper delta surface was  
462 dated at  $8\,920 \pm 50$  BP (10.05 cal ka BP; Beta-318526). This date provides an age for the  
463 beginning of the construction of the St-Maurice delta in the Champlain Sea.

464 The aquifer is one source of drinking water for the greater Trois-Rivières area. A  
465 confined aquifer is also present below the marine clays, in pre-LGM sands (Vieilles-  
466 Forges). However, data characterizing the aquifer are scarce. It is seldom as a water  
467 source because the quality of groundwater is poor.

468 The L3-2010 and L3-2011 seismic lines were run parallel to the axis of the St.  
469 Lawrence River (Figure 9). These profiles illustrate the transition from the deep water  
470 depositional environment of the clay plain and the complex geology of the delta,  
471 including the Lake Lampsilis sands, pre-LGM sediments and Gentilly Till. To the west of  
472 the delta, the bedrock topography shows little variation, with a mean elevation of about -  
473 50 m below sea level. The fine-grained sediments of the Champlain Sea fill the basin to a  
474 thickness of 80 m. The recent alluvial deposits of Lake Saint-Pierre cover the sequence  
475 with 10 m of fine sands and silts. The light reflector observed in the marine sequence  
476 could correspond to the Yamachiche Diamicton (Occhietti, 1980), which is a thin,  
477 discontinuous coarse-grained deposit associated with the Saint-Narcisse Episode. The  
478 sequence below the clay unit includes the Gentilly Till, Vieilles-Forges Sand, St. Maurice  
479 Rythmites, Deschaillons Varves, and Lotbinière Sand. Only the first three units could be  
480 differentiated from the drilling logs.

481 On the L3-2010 Minivibe line (Figure 9a), the till thickness varies from 4 m at the  
482 beginning of the section, to over 20 m at kilometer 1.4, where the Lotbinière Sand begins  
483 to thin out. Champlain Sea silt and clay completely wedge out at km 2.8, where the sand  
484 of the Champlain Sea-Lake Lampsilis succession rest directly on the till which rises to  
485 shallow depths above the pre-LGM sediments. At this location, drilling indicates that pre-  
486 LGM sediments constitute the largest part of the Quaternary column (over 80 m).

487 The L3-2011 Minivibe line (Figure 9b) is the continuity of the preceding  
488 sequence, about 1 km further south. The Champlain-Lampsilis sands are no longer  
489 present and till is observed at the surface all along the 3 km long profile. The available  
490 drilling logs allowed us to differentiate the Vieilles Forges Sand and the Saint-Maurice  
491 Rythmites from the deeper and older Pleistocene sediments.

492 **Figure 9. High resolution seismic survey lines L3-2010 (a) and L3-2011 (b). The**  
493 **profile interpretations are shown below the seismic profiles. See Figure 2 for**  
494 **location of the seismic lines and Table 2 for a detailed description of the units.**

495 To the east of the delta, a series of deltaic, lacustrine and alluvial sands form a  
496 vast aquifer with a thickness ranging from 11 to 35 m. The aquifer overlies an  
497 impermeable clay (aquitard unit) with an average slope of the top surface coincident with  
498 the topography of the land surface (towards the SE). The aquifer is complex with  
499 significant vertical heterogeneity, containing interstratified medium to silty sand layers.  
500 The Quaternary deposits encountered in the borehole FE-02-11 (Figure 7) are 56.4 m  
501 thick and rest on the Ordovician sedimentary platform. Gently Till, identified at 38 m  
502 depth (-22 m below sea level), overlies a sequence of undifferentiated Pleistocene  
503 sediments that is 8 m thick. Champlain Sea clay overlies the till between 18 to 37 m  
504 depth. The uppermost unit, from 0 to 18.3 m, consist of alluvial, littoral and deltaic  
505 sediments that were reworked by the fluctuating water levels of the St. Lawrence River.  
506 The base of this unit has been dated at 7.4 ka BP (8.2 cal ka BP; Beta-318525) on the  
507 basis of a wood fragment recovered at a depth of 14.5 m. Since the alluvial unit from  
508 which the wood fragment was recovered marks the return to present-day conditions in the  
509 St. Lawrence River, Beta-31526 provides a maximum age for the end of delta  
510 construction. Taken together, the two  $^{14}\text{C}$  dates indicated that the main part of the St-  
511 Maurice delta was constructed in about 1800 years, between 10.05 and 8.2 cal ka BP.

### 512 *Hydrostratigraphic units*

513 The three-dimensional hydrostratigraphical model allowed the identification of four  
514 main hydrostratigraphic units with a regional distribution:

- 515 • Unit 1: An aquifer within fractured bedrock, divided into three sub-units: 1)  
516 aquifer in crystalline rocks (with low fracture density) of the Precambrian  
517 basement; 2) aquifer in brittle sedimentary rocks (sandstones and limestones); and  
518 3) aquifer in ductile sedimentary rocks (shales). This unit, while it has a regional  
519 extension, has a low potential for a water supply and is not being tapped by  
520 municipalities.

- 521 • Unit 2: An aquifer in glaciofluvial and glacial sediments. This unit generally lies  
522 immediately above bedrock. The aquifers in glaciofluvial sediments, despite their  
523 small size and discontinuous areal distribution, supply a number of municipal  
524 wells. Till deposits, for their part, are thin and have a low permeability. However,  
525 water yields are generally sufficient to supply domestic wells.
- 526 • Unit 3: The regional discontinuous aquitard formed by fine-grained Champlain  
527 Sea sediments. The deposits overlie almost continuously older Quaternary units in  
528 the St. Lawrence Lowlands as well as in some isolated valleys or depressions of  
529 the Laurentian Hills.
- 530 • Unit 4: An aquifer formed in the uppermost sandy sediments. The unit can be  
531 divided into two sub-units: 1) sand and gravel deposits within the Saint-Narcisse  
532 morainic complex; and 2) surficial sand deposits overlying fine-grained sediments  
533 of the Champlain Sea, including coastal and deltaic silty sands, sands and silts of  
534 Lake Lampsilis and marine and lacustrine deltaic sands emplaced by the Saint-  
535 Maurice River.

536 Integrating the 3D model, architecture of the hydrostratigraphic units and regional  
537 distribution of hydrogeological conditions lead us to define six aquifer types shown in  
538 Figure 10:

- 539 • Type 1: Unconfined fractured bedrock aquifers, which may be locally overlain by  
540 a thin layer of till;
- 541 • Type 2: Confined or semi-confined fractured bedrock aquifers overlain by marine  
542 silt and clay (aquitard);
- 543 • Type 3: Unconfined aquifers in glaciofluvial sediments or littoral sands that lie  
544 directly on the bedrock;
- 545 • Type 4: Double aquifers consisting of an upper, unconfined sandy aquifer and a  
546 confined or semi-confined aquifer in the fractured bedrock. These aquifers are  
547 separated by the regional aquitard;
- 548 • Type 5: Additional confined aquifers in glaciofluvial sediments that lie directly on  
549 the bedrock and are overlain by marine silt and clay (aquitard). These units may  
550 lie underneath other units to form aquifer complexes.

551 **Figure 10. Aquifer types of the Mauricie region based on the architecture of the**  
552 **hydrostratigraphic units.**

553 Figure 10 shows the distribution of the above aquifer types. Type 1 is the most  
554 common aquifer in the Laurentian Hills, except in the valleys where types 2-5 are found,  
555 such as the Saint-Alexis-des-Monts aquifer in the du Loup River valley. In the piedmont,  
556 the Saint-Narcisse Moraine frequently contains types 4 and 5 aquifers, all hydraulically  
557 separated from each other. The paleodelta of the St. Maurice River is an unconfined  
558 aquifer of regional extent and has attributes of types 4 and 5. Only the upper, unconfined  
559 aquifer is exploited due to the poor quality of the groundwater in the confined aquifer  
560 below, caused notably by the generalized presence of brackish water and the occasional  
561 occurrence of methane. The upper aquifer supplies several municipalities including the  
562 City of Trois-Rivières. Finally, along the north shore of Lake St-Pierre, the most common  
563 aquifer is type 2. The marine clay plain is considered an aquitard that confines aquifers  
564 (glacial till and/or the Ordovician rocks) that are heavily mineralized or contain natural  
565 gas accumulations.

566 Several local conceptual models prepared during the study are summarized in a  
567 cross-section shown in Figure 11 that illustrates the typical geologic and  
568 hydrostratigraphic units observed in southwest Mauricie. The upper panel is a geological  
569 cross-section showing the distribution of Quaternary deposits and bedrock, and is based  
570 on the 3D model. The lower panel displays a profile along the same section that identifies  
571 the major hydrostratigraphic units and groundwater flow lines from the conceptual  
572 hydrogeological model. The trace of cross-section A-A' is shown as a dashed line on  
573 Figure 10. It starts in the Rivière du Loup valley near St-Alexis-des-Monts in the  
574 northwest and runs eastward across the Saint-Narcisse Moraine near Charette, and  
575 crosses Trois-Rivières towards the southeast. This profile was selected to reflect the main  
576 hydrogeological contexts recognized in the region. While the elevation of the bedrock  
577 surface and the terrain are both accurately depicted in the cross-section, some geological  
578 and hydrogeological elements have been simplified to represent typical subsurface  
579 characteristics in the study area. In other words, while the units and conditions illustrated

580 on this cross-section are factual, they do not necessarily occur at the location of the  
581 section.

582 The northwest segment of the cross-section shows typical hydrogeological  
583 conditions in the Laurentian Hills. The surficial sediment cover is generally thin, and  
584 valleys formed along bedrock depressions are commonly floored by coarse permeable  
585 sediments that are locally overlain by fine-grained marine/estuarine sediments. Confined  
586 aquifers are locally present below the Champlain Sea sediments in valleys lying at the  
587 maximum marine limit. The Precambrian bedrock underlying the surficial sediment cover  
588 includes igneous/metamorphic rocks with no primary porosity and very low fracture  
589 porosity. It is considered as an aquifer since water flows through it, but the capacity of  
590 the unit to yield an appreciable amount of groundwater ultimately depends on the  
591 interconnections of the fracture network, which tends to increase near the fault zones. In  
592 the northern part of the region, wells in bedrock commonly yield sufficient amounts of  
593 water for domestic supply, but sometimes have to be drilled to depths exceeding 40  
594 metres.

595 **Figure 11. Cross-section illustrating the typical geologic and hydrostratigraphic**  
596 **units encountered in southwest Mauricie. The trace of cross-section A-A' is shown**  
597 **as a dashed line on Figure 10.**

598 Along the edge of the Laurentian Hills, the St-Narcisse morainic complex has  
599 hydrogeological conditions conducive for groundwater extraction. The moraine and  
600 associated glaciofluvial deposits on both its proximal and distal boundaries contain a  
601 group of aquifers that generally are poorly connected, but nonetheless connected through  
602 low-permeability glacial sediments or bedrock. The municipalities of Notre-Dame-du-  
603 Mont-Carmel and St-Paulin both extract groundwater from this unit. The Charette  
604 segment of the morainic complex has a similar hydrogeological potential, with an  
605 unconfined upper aquifer and a confined lower aquifer, both of which are not utilized. A  
606 series of depressions above the St-Cuthbert fault are filled by sand and gravel that may  
607 contain productive aquifers bounded by the marine clay aquitard. These depressions have  
608 not yet been tested or exploited for a municipal water supply. Even though the

609 hydrogeological potential of the St-Cuthbert fault area remains untested, it likely plays a  
610 key role in the regional flow system by intercepting groundwater flowing from the  
611 Precambrian Hills and redirecting groundwater into glacial sand and gravel aquifers in  
612 the adjacent Piedmont.

613 In the clay plain, the sediment cover is very thick but is almost exclusively  
614 composed of very fine-grained marine mud (silty clay) with very low permeability. The  
615 underlying till unit is more permeable than the Paleozoic bedrock by an order of  
616 magnitude, but is still not considered an aquifer. Groundwater within the sedimentary  
617 rocks below the Quaternary deposits is generally brackish. Because the rivers flowing  
618 across the plain tend to be deeply entrenched, the extent of alluvial aquifers is limited.  
619 However, the paleo-deltas of the Maskinongé and Yamachiche rivers are large enough to  
620 contain significant groundwater supplies, and are used for a potable water supply by the  
621 towns of St-Édouard-de-Maskinongé, Ste-Ursule, St-Alexis-des-Monts, St-Élie-de-  
622 Caxton, Charette and St-Mathieu-du-Parc. Municipalities that are too far from the deltas  
623 are supplied by a private aqueduct network operated by the Régie d'Aqueduc de Grand-  
624 Pré, which pumps water from wells in glacial sand and gravel aquifers near St-Boniface.

625 The cross-section ends in the St-Maurice River paleo-delta sandy aquifers (Figure  
626 11). This large deltaic system was built at the mouth of the valley during regression of the  
627 Champlain Sea and subsequent Lake Lampsilis. The delta overlies either marine clays or  
628 older Quaternary deposits. At St-Boniface and St-Étienne (upstream on the left bank), the  
629 deltaic deposits contain an unconfined aquifer which caps the stratigraphic sequence, but  
630 is too thin to supply large volumes of groundwater. In some areas though, where thicker  
631 well-sorted sands occur, commonly in narrow erosional channels cut in the underlying  
632 clay, there may be sufficient water supply for municipalities. In the Trois-Rivières  
633 (downstream right bank) and Cap-de-la-Madeleine (downstream left bank) sectors,  
634 significant thickening of the surficial sand cover makes the aquifer much more  
635 productive. The St-Maurice River is deeply entrenched into its former delta, cutting  
636 through even the upper till layer and pre-LGM Quaternary units. This creates  
637 groundwater resurgences along the river. The City of Trois-Rivières pumps water from  
638 the surficial sand deposits on both sides of the St-Maurice. The sedimentary rocks

639 underlying the Quaternary deposits are moderately permeable, but the water contained in  
640 the shale unit is mineralized and only used in a few geothermal applications.

## 641 **Discussion and conclusion**

### 642 *3D model*

643 As a result of its complex Quaternary history, the Mauricie region contains  
644 significant granular aquifers, such as the paleodelta of the St. Maurice and Yamachiche  
645 rivers, the Saint-Narcisse morainic complex and the high sandy terraces formed during  
646 Champlain Sea regression. These aquifers had been partially investigated through local  
647 hydrogeological studies conducted by municipalities. New fieldwork was undertaken to  
648 supplement existing data and to build a coherent hydrostratigraphical model presenting  
649 the stratigraphy and architecture of surficial deposits. Taking into account the reliability  
650 ratings, surfaces representing the upper and lower contacts of geological formations were  
651 modeled by cokriging. The result is a 3D model with seven layers representing the  
652 surficial deposits of the Mauricie region, from the surface of the bedrock to the upper  
653 littoral and deltaic sands. The total thickness of the sediments varies from zero, on rock  
654 outcrops, to 150 m under the City of Trois-Rivières.

655 The 3D hydrostratigraphical model has helped refine our understanding of  
656 regional aquifers, particularly the thickness, distribution and sequence of surficial  
657 deposits in the valleys, along the St. Cuthbert Fault, under the Saint-Narcisse morainic  
658 complex and in the paleodelta of the Saint-Maurice River. New, unexploited aquifers  
659 have been identified north of the City of Trois-Rivières, around the Saint-Narcisse  
660 morainic complex and along the St. Cuthbert Fault. The deposits of the St. Cuthbert Fault  
661 are confined aquifers resulting from the deposition of granular sediments into a series of  
662 bedrock depressions observed along the axis of the fault. These aquifers, as well as  
663 confined aquifers beneath the marine clay plain, are only known indirectly through the  
664 interpretation of geophysical surveys, such as seismic reflection profiling. Additional  
665 drilling is required to determine the grain size distributions of the deposits to assess  
666 aquifer potential.

667 *Regional stratigraphy*

668 The results of the 3D model is consistent with the regional stratigraphy proposed  
669 by Gadd and Karrow (1959), Denis (1972), and Occhietti (1982). However, the  
670 modelling helped to further clarify the regional hydrostratigraphic framework, notably  
671 the topography of the bedrock surface, the subsurface extent of pre-LGM Quaternary  
672 sands, and the architecture of post-Champlain Sea deposits.

673 The pre-LGM Vieilles Forges sands were identified under the marine clay layer in  
674 borehole FE-02-11. This is the first drillhole that records the presence of this unit east of  
675 the St. Maurice River, and this allows us to better assess the extent of this regional unit. It  
676 is possible that some wells of the City of Trois-Rivières draw their water from this  
677 formation, rather than from the post-Champlain upper sands. Additional drilling (with  
678 dating) would be necessary to validate the extension of pre-LGM deposits north of Trois-  
679 Rivières. The permeability of these older deposits and the quality of the water they  
680 contain remain difficult to assess at a regional scale, as well as the hydraulic links  
681 between the upper sand aquifer and the bedrock.

682 Facies encountered in drilling along the Saint-Narcisse morainic complex vary  
683 widely. In general, the upper layer consists of coarse-grained sediments, from sand to  
684 cobbles and boulders, and is associated with ice retreat from the terminal moraine. This  
685 material generally lies above groundwater table and is too coarse to form a reservoir, but  
686 it contributes to the overall recharge of the surrounding aquifers. Multiple stratigraphic  
687 wells drilled in the center of the Saint-Narcisse Moraine showed the presence of a layer  
688 of fine-grained sediments or diamicton (till) at the base, confirming that the material was  
689 emplaced in a marine environment during a re-advance of the ice sheet. Although the  
690 permeability of this unit is lower than the overlying coarse-grained material, it cannot be  
691 considered as an aquitard due to its higher proportion of fine sand. However, a sandy  
692 aquifer confined below the marine fine-grained sediments was observed in some borings.  
693 For example, an aquifer contained in 20 m of medium sand, was discovered when drilling  
694 borehole FE-09-11 in the village of Sainte-Angèle-de-Prémont. This deposit should be  
695 investigated further, given its positive aquifer characteristics.

696 The aquifer formed by the upper sands of the Trois-Rivières area is composed of a  
697 mixture of marine and lacustrine sands (littoral and deltaic), as well as sands associated  
698 with the paleodelta of the Saint-Maurice River. The unit rests on impermeable marine  
699 sediments on both sides of the river, but the thickness of the clay aquitard is much more  
700 variable on the west bank. The aquifer is highly stratified and encompasses many facies.  
701 Differentiating depositional environments based on grain size characteristics has been  
702 attempted, but did not achieve much success. In terms of thickness and grain size, the  
703 western deltaic zone shows a greater aquifer potential than the eastern zone, and our  
704 observations show that the upper sand in the northern part of the city has larger than  
705 expected thicknesses.

706 This study has highlighted yet again the major hydrogeological potential of the  
707 Champlain Sea deltas in the St-Lawrence Valley (for example, see Fagnan et al., 1998 for  
708 the Portneuf delta). Those deltaic constructions were emplaced as laterally continuous  
709 and thick units in a postglacial sea where sediment supply as well as accommodation  
710 space were elevated. They were later dissected by the post-glacial incision caused by the  
711 glacial isostatic rebound. Similar contexts can be found all along the axis of the St-  
712 Lawrence River, particularly at the mouth of its left bank tributaries which were fed by  
713 abundant meltwaters and sediments derived from the receding Laurentide Ice Sheet.

714 Also highlighted were the few spatially disconnected buried valley aquifers  
715 around the St-Narcisse Moraine, showing the highly dynamic nature of the depositional  
716 process during this event. Similar buried aquifers can most likely be found close to the  
717 moraine outside the study area.

#### 718 **Acknowledgements**

719 This work is part of the Groundwater Knowledge Acquisition Program (PACES)  
720 of the Quebec Ministry of the Environment (MDDELCC). It was conducted in  
721 collaboration with a regional council of elected officials (CRÉ Mauricie), regional  
722 municipalities (MRC Maskinongé), the city of Trois-Rivières, the Geological Survey of  
723 Canada and watershed agencies. The authors would like to acknowledge the Minivibe  
724 seismic profiling carried by André Pugin and Susan Pullan of the Geological Survey of

725 Canada. We would also like to thank the municipal officials, the local residents and the  
726 following persons for their collaboration: Myrabelle Chicoine, Isabelle Lessard, Alain  
727 Rouleau, Ali Assani, Pierre-André Bordeleau, Francis Clément, Précillia Descoteaux,  
728 Sophie Côté, Éric Dubois, Brian Bélisle, Philippe Davignon, Fannie Fortier-Fradette,  
729 Lise Lamarche, Harold Vigneault, Marc-André Carrier, Miroslav Nastev, Claude Hébert,  
730 François Hardy, René Lefebvre, Jean-Marc Ballard and Xavier Mallet.

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799 **Figure Captions**

800 Figure 1. Location and digital elevation model of the study area. The four main regional  
801 hydrogeologic contexts are shown: the Laurentian Hills, Saint-Narcisse morainic  
802 complex, Marine clay plain and paleodelta of the St. Maurice River.

803 Figure 2. Quaternary geology and location of the new geophysical surveys and boreholes.  
804 The location of the existing boreholes and wells (5 386) used to build the model are not  
805 shown. See Table 2 for a detailed description of the units.

806 Figure 3. Flowchart of the GIS operations used to model the top surface of each  
807 hydrogeological unit.

808 Figure 4. Total thickness of Quaternary deposits. The thickest deposits, reaching 150 m,  
809 are in the western part of the City of Trois-Rivières.

810 Figure 5. 3D model of Quaternary deposits. The model consists of six layers representing  
811 the main hydrogeological units observed above bedrock in the study area. The six layers  
812 represent the Pleistocene sediments (pre-LGM) (a), LGM and post-Younger Dryas tills  
813 (b), glaciofluvial sediments (c), marine silt and clay (d), St. Narcisse morainic complex  
814 (e) and littoral, deltaic and alluvial sediments (f).

815 Figure 6. Thickness grids (m) of Quaternary deposits.

816 Figure 7. Logs of selected boreholes with lithology and grain size. The locations of the  
817 boreholes are shown in Figure 2.

818 Figure 8. High resolution seismic survey lines L1-2011 (a) and L7-9-2010 (b). The  
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823 seismic lines and Table 2 for a detailed description of the units.

824 Figure 10. Aquifer types of the Mauricie region based on the architecture of the  
825 hydrostratigraphic units.

826 Figure 11. Cross-section illustrating the typical geologic and hydrostratigraphic units  
827 encountered in southwest Mauricie. The trace of cross-section A-A' is shown as a dashed  
828 line on Figure 10.

829

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Table 1. Sources of hydrogeological data.

| Sources   | Number of wells/boreholes |
|---|---------------------------|
| Hydrogeological Information System (Québec)                           | 2311                      |
| Ministry of Transportation (Québec)                                   | 2034                      |
| Hydrogeological studies by consultants                                | 1112                      |
| Municipal databases   | 312                       |
| Hydrogeological studies from the Ministry of the Environment (Québec) | 245                       |
| Private databases   | 243                       |
| Système d'Information Géoscientifique, Pétrolier et Gazier (SIGPEG)   | 164                       |

Table 2. Quaternary stratigraphy of the study area.

| Age<br>Ka | Code       | Formation  | Lithology  |
|-----------|------------|--|--|
| 11.7      | <b>O</b>   | Organic deposits   | Peat mainly  |
|           | <b>E</b>   | Aeolian sediments  | Fine sand  |
|           | <b>At</b>  | Terraced alluvial sediments  | Sand, sandy silt and gravelly sand   |
|           | <b>Ap</b>  | Alluvial sediments - Floodplain  | Sand, sandy silt, gravelly sand and gravel   |
|           | <b>Ld</b>  | Lacustrine sediments - Deltaic facies                                  | Sand, gravelly sand and well-sorted gravel, deposited at the mouth of rivers flowing into Lake Lampsilis   |
|           | <b>Lb</b>  | Lacustrine sediments - Littoral and nearshore facies of Lake Lampsilis | Sand, silt and gravelly sand and gravel deposited along or near Lake Lampsilis shorelines  |
| 75        | <b>Cg</b>  | Mass-wasting deposits - Landslides                                     | Mainly silt, clay and sand reworked by landslides in Champlain Sea sediments   |
|           | <b>Md</b>  | Marine sediments - Deltaic facies                                      | Sand, gravelly sand and gravel, stratified and well sorted, deposited at the mouth of rivers flowing into the Champlain Sea, locally including prodeltaic silty sands  |
|           | <b>Mb</b>  | Marine sediments – Littoral facies                                     | Sand, sandy silt, gravelly sand and gravel deposited along or near Champlain Sea shorelines  |
|           | <b>Ma</b>  | Marine sediments – Offshore facies                                     | Clayey silt and silty clay deposited in Champlain Sea basins   |
|           | <b>GxT</b> | Glaciofluvial and glacial sediments of the St-Narcisse Moraine         | Sand and gravel, till or diamicton deposited at the ice front during the Younger Dryas readvance   |
|           | <b>Go</b>  | Glaciofluvial sediments – subaerial outwash facies                     | Sand and gravel, well stratified and sorted, forming outwash plains locally pitted by kettles  |
|           | <b>Gs</b>  | Glaciofluvial sediments – Subaqueous outwash facies                    | Mainly sand, minor silty sand or some gravel, deposited at the mouth of sub-glacial tunnels entering in the Champlain Sea  |
|           | <b>Gx</b>  | Glaciofluvial sediments - Ice-contact facies                           | Mainly sand and gravel, minor diamicton, forming eskers, kames, delta-kames, ice-contact terraces or moraine ridges  |
|           | <b>T</b>   | Glacial sediments – Gentilly Till, Yamachiche Diamicton                | Sandy silt diamicton, compact gray matrix, deposited by the ice sheet and forming continuous blanket and discontinuous veneer, locally reworked by waves and currents along Champlain Sea shorelines         |
|           | 115        | <b>Q</b>   | Undifferentiated older Quaternary deposits: sand (Lotbinière Sand, St-Pierre sand, Vieilles Forges sand), rhythmites (Deschaillons varves, St-Maurice rhythmites), diamictons (Lévrard till, Bécancour Till) |
| 128       | <b>Q</b>   | Glacial sediments – Bécancour till                                     | Diamicton characterized by a brick red color inherited from the erosion of red shales  |

Table 3. Hydraulic properties of the regional granular aquifers: hydraulic conductivity (K), transmissivity (T), saturated thickness (b), coefficient of storage (S) and specific capacity (Q/s). See location map in Figure 1.

| <b>Granular aquifers</b>                            |            | <b>K (m/s)</b> | <b>T (m<sup>2</sup>/d)</b> | <b>b (m)</b> | <b>S</b> | <b>Q/s (m<sup>3</sup>/d/m)</b> |
|---|------------|----------------|----------------------------|--------------|----------|--------------------------------|
| <b>Laurentian Hills and piedmont</b>                |            |                |                            |              |          |                                |
| Valley bottoms                                      | Unconfined | 5.11E-04       | 436.6                      | 8.0          | 2.00E-02 | 288.0                          |
|   | Confined   | 3.36E-04       | 772.8                      | 4.3          | 1.41E-03 | 287.2                          |
| Paleodelta and terraces of the Maskinongé River     | Unconfined | 5.16E-03       | 287.0                      | 23.2         | -        | 208.6                          |
|   | Confined   | 1.24E-04       | 260.2                      | 11.2         | 1.03E-03 | 62.0                           |
| Paleodelta and terraces of the Yamachiche River     | Unconfined | 5.03E-04       | 685.5                      | 8.4          | 7.70E-02 | 240.5                          |
|   | Confined   | 8.30E-05       | 172.8                      | 24.1         | 6.10E-04 | 99.4                           |
| <b>Saint-Narcisse morainic complex</b>              |            |                |                            |              |          |                                |
| Notre-Dame-du-Mont-Carmel                           | Unconfined | -              | 20.8                       | -            | -        | 166.3                          |
|   | Confined   | 2.75E-04       | 457.4                      | 19.3         | 3.33E-04 | 172.6                          |
| Charette  | Unconfined | -              | -                          | 6.7          | -        | 165.5                          |
|   | Confined   | 1.06E-04       | 30.5                       | 3.1          | -        | 34.6                           |
| Sainte-Angèle-de-Prémont and Saint-Paulin           | Unconfined | 9.68E-05       | 987.5                      | 26.3         | 3.21E-02 | 362.2                          |
|   | Confined   | 5.69E-04       | 752.0                      | 13.5         | 4.31E-04 | 186.3                          |
| Glaciofluvial deposits of the St.-Cuthbert Fault*   | Confined   | -              | -                          | -            | -        | -                              |
| Glaciofluvial deposits underlying the marine plain* | Confined   | 4.08E-06       | 0.8                        | 2.1          | -        | 1.6                            |
| Littoral deposits of Lake Lampsilis                 | Unconfined | -              | -                          | -            | -        | -                              |
| <b>Paleodelta of the Saint-Mauricie River</b>       |            |                |                            |              |          |                                |
| Saint-Boniface                                      | Unconfined | -              | -                          | -            | -        | 223.8                          |
| Shawinigan  | Unconfined | 2.85E-04       | 563.8                      | 8.4          | -        | 366.7                          |
| Saint-Étienne-des-Grès                              | Unconfined | 5.67E-04       | 267.7                      | 5.0          | 1.19E-01 | 173.4                          |
| Notre-Dame-du-Mont-Carmel                           | Unconfined | 5.56E-05       | 149.6                      | 6.8          | 1.00E-04 | 146.1                          |
| Pointe-du-Lac (Trois-Rivières) and Yamachiche       | Unconfined | 1.55E-05       | 199.5                      | 9.0          | 8.90E-02 | 203.6                          |
| Trois-Rivières and Trois-Rivières-Ouest             | Unconfined | 2.79E-04       | 874.6                      | 24.4         | 2.05E-01 | 231.7                          |
|   | Confined** | -              | 604.8                      | -            | -        | 70.8                           |
| Saint-Louis-de-France (Trois-Rivières)              | Unconfined | 7.70E-05       | 49.2                       | 15.3         | 1.37E-01 | 55.0                           |
| Cap-de-la-Madeleine (Trois-Rivières)                | Unconfined | 3.36E-04       | 352.7                      | 8.1          | 1.22E-01 | 206.4                          |
|   | Confined*  | 1.64E-04       | 86.4                       | 6.1          | -        | 29.2                           |

\* Untapped confined aquifer \*\* Sables des Vieilles Forges

Table 4. Hydraulic properties of the regional fractured rock aquifers: hydraulic conductivity (K), transmissivity (T), saturated thickness (b), coefficient of storage (S) and specific capacity (Q/s). See location map in Figure 1.

| Fractured rock aquifers                                |            | K (m/s)  | T (m <sup>2</sup> /d) | b (m) | S | Q/s (m <sup>3</sup> /d/m) |
|--|------------|----------|-----------------------|-------|---|---------------------------|
| <b>Laurentian Hills and piedmont (Canadian Shield)</b> |            |          |                       |       |   |                           |
| Mékinac-Taureau domain                                 | Unconfined | 2.15E-07 | -                     | 61.9  | - | 1.7                       |
|  | Confined   | 7.41E-07 | -                     | -     | - | -                         |
| Morin terrane  | Unconfined | 6.67E-09 | 9.7                   | 71.0  | - | 3.2                       |
|  | Confined   | 7.70E-05 | 79.2                  | 36.1  | - | 44.6                      |
| Intrusive suites                                       | Unconfined | 7.78E-06 | -                     | 20.3  | - | 6.5                       |
|  | Confined   | -        | -                     | -     | - | -                         |
| <b>St. Lawrence Platform</b>                           |            |          |                       |       |   |                           |
| Sandstone  | Confined   | -        | -                     | -     | - | -                         |
| Limestone  | Confined   | 2.14E-06 | 0.2                   | 1.5   | - | 0.9                       |
| Shale  | Confined*  | 9.14E-07 | -                     | -     | - | -                         |

\* Untapped confined aquifer

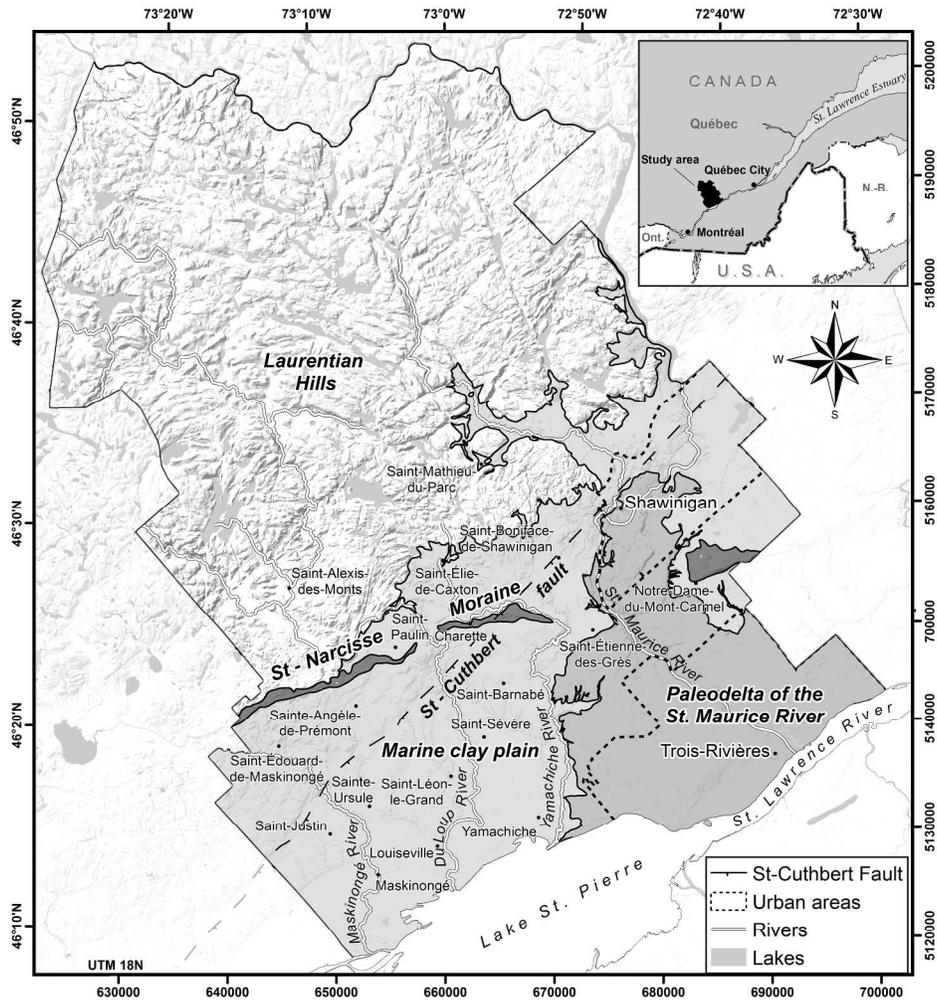


Figure 1. Location and digital elevation model of the study area. The four main regional hydrogeologic contexts are shown: the Laurentian Hills, Saint-Narcisse morainic complex, Marine clay plain and paleodelta of the St. Maurice River.

Figure 1  
228x242mm (300 x 300 DPI)

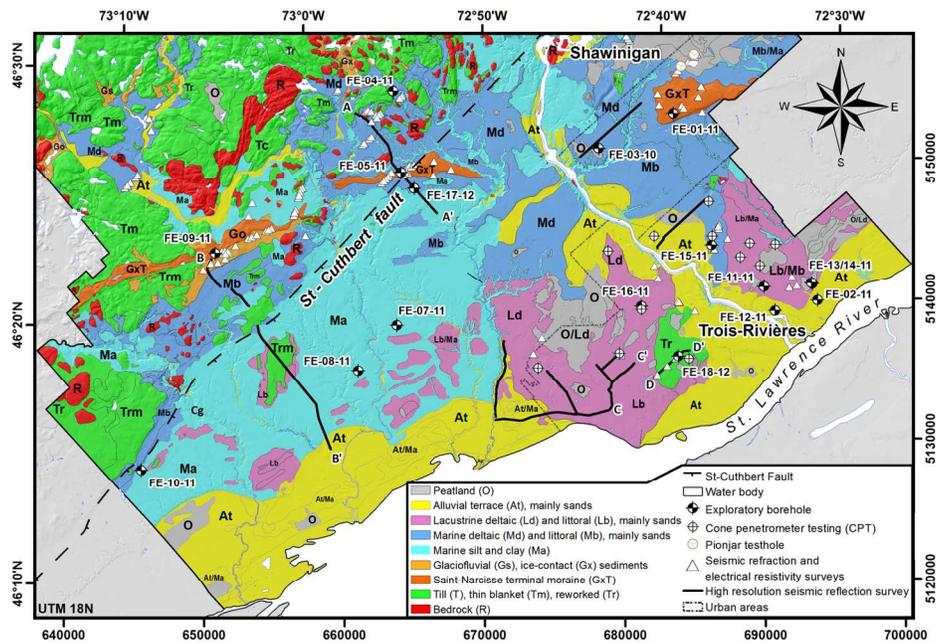


Figure 2. Quaternary geology and location of the new geophysical surveys and boreholes. The location of the existing boreholes and wells (5 386) used to build the model are not shown. See Table 2 for a detailed description of the units.

Figure 2  
152x107mm (300 x 300 DPI)

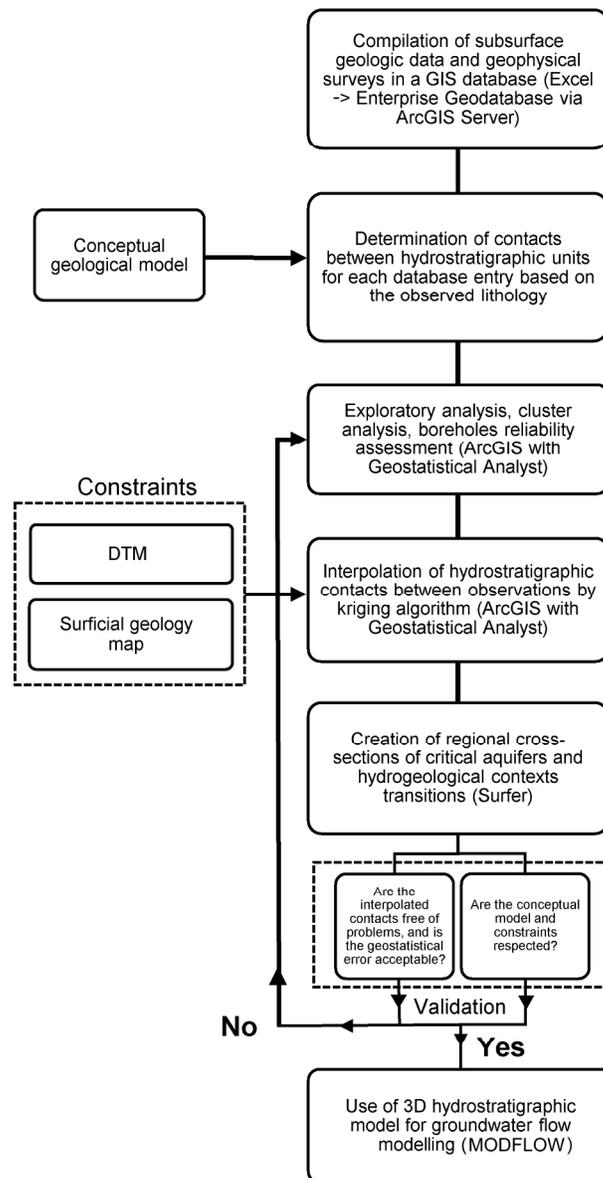


Figure 3. Flowchart of the GIS operations used to model the top surface of each hydrogeological unit.

Figure 3

160x290mm (300 x 300 DPI)

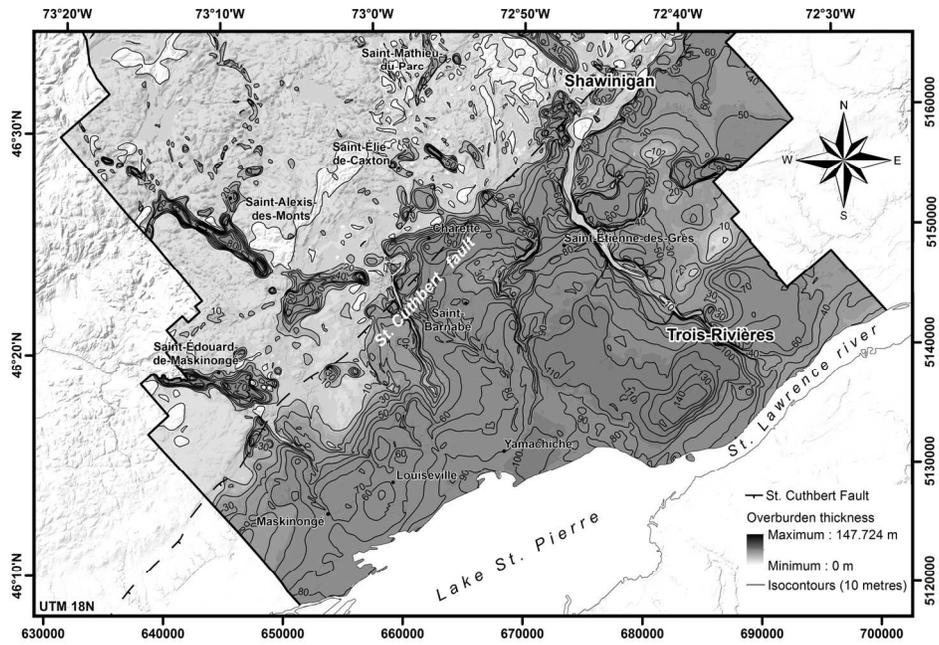


Figure 4. Total thickness of Quaternary deposits. The thickest deposits, reaching 150 m, are in the western part of the City of Trois-Rivières.

Figure 4

152x107mm (300 x 300 DPI)

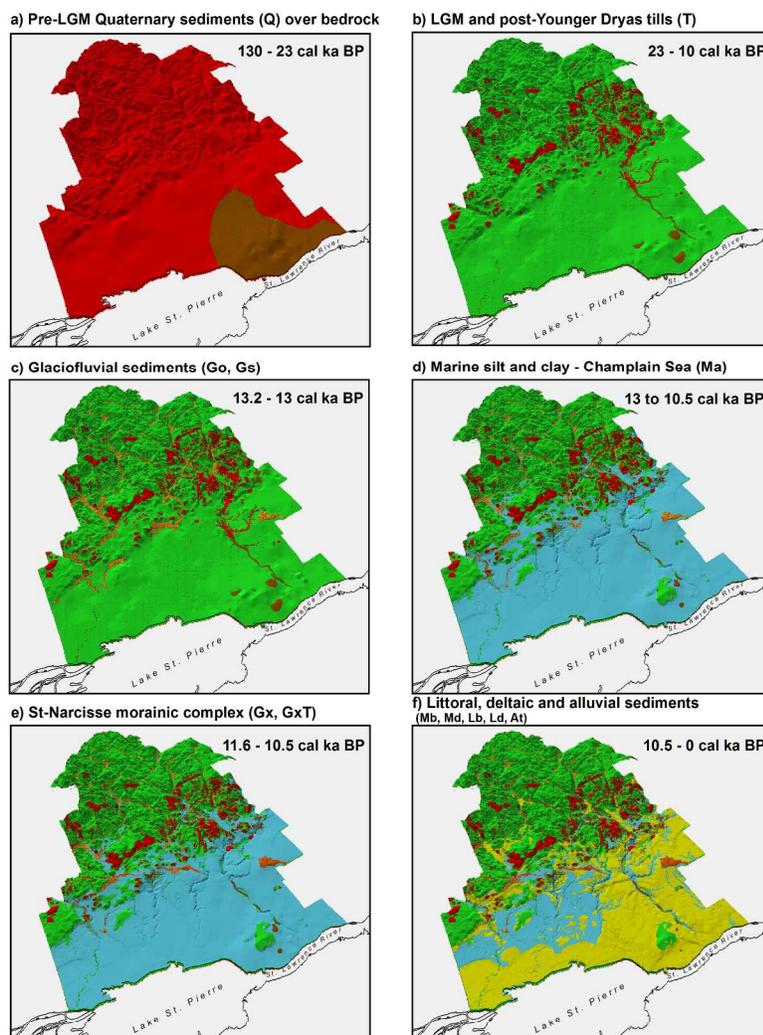


Figure 5. 3D model of Quaternary deposits. The model consists of six layers representing the main hydrogeological units observed above bedrock in the study area. The six layers represent the Pleistocene sediments (pre-LGM) (a), LGM and post-Younger Dryas tills (b), glaciofluvial sediments (c), marine silt and clay (d), St. Narcisse morainic complex (e) and littoral, deltaic and alluvial sediments (f).

Figure 5

279x361mm (300 x 300 DPI)

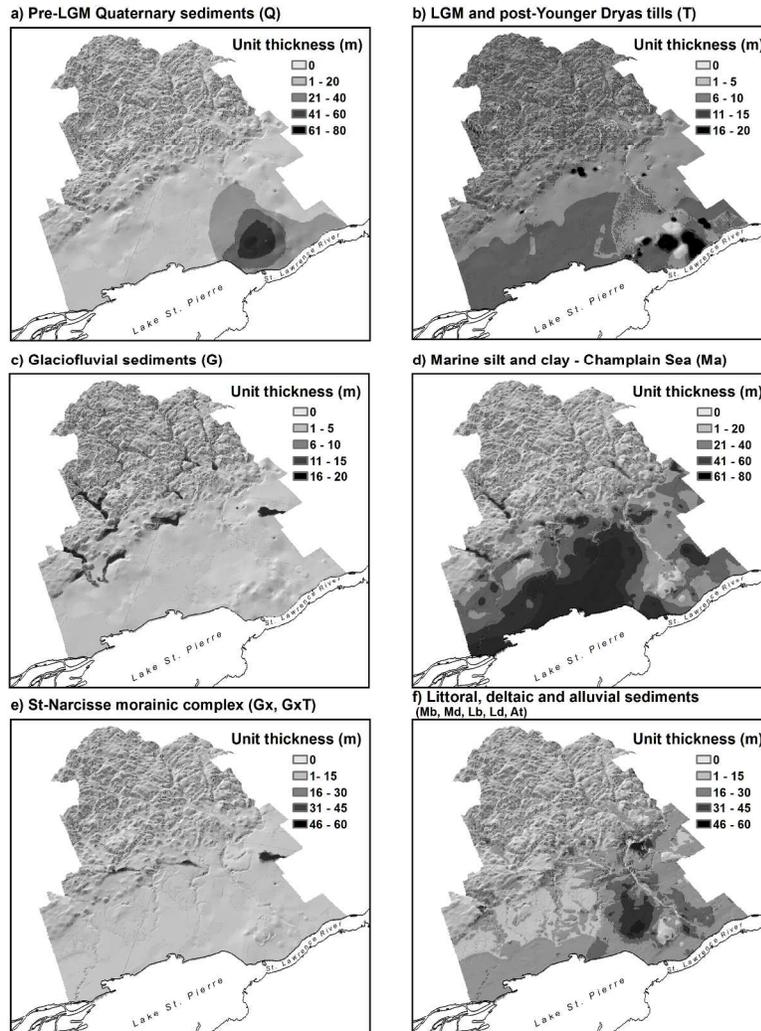


Figure 6. Thickness grids (m) of Quaternary deposits.

Figure 6

279x361mm (300 x 300 DPI)

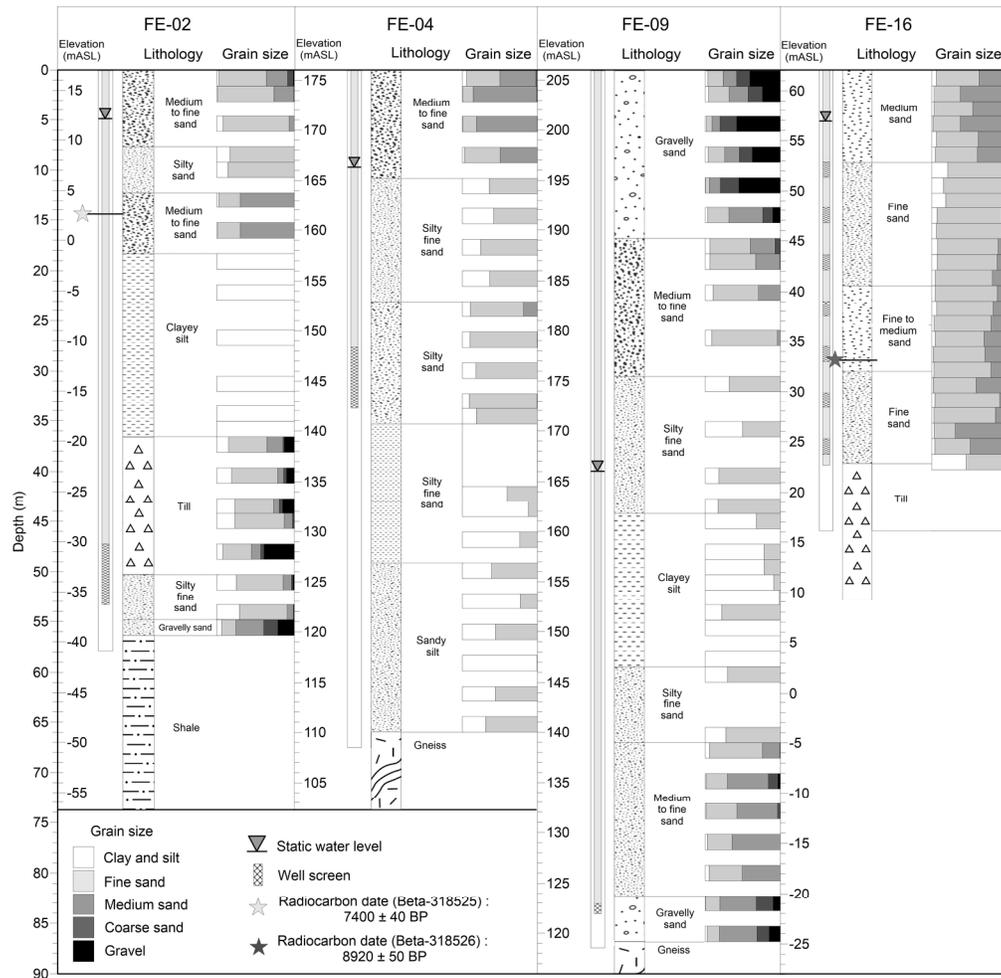


Figure 7. Logs of selected boreholes with lithology and grain size. The locations of the boreholes are shown in Figure 2.

Figure 7  
200x199mm (300 x 300 DPI)

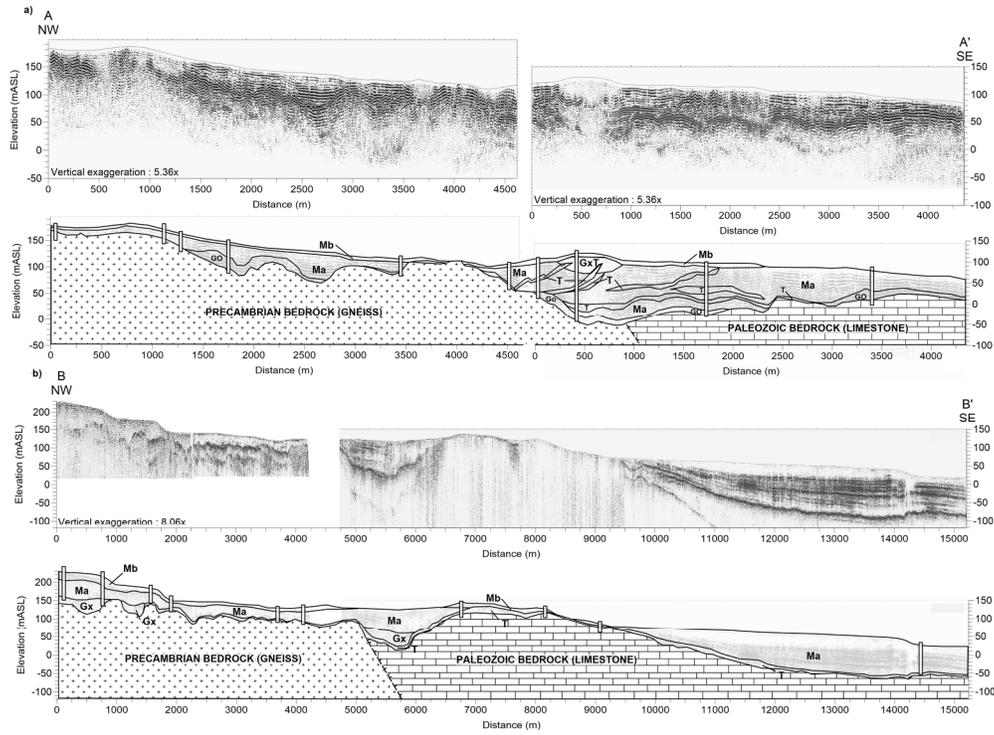


Figure 8. High resolution seismic survey lines L1-2011 (a) and L7-9-2010 (b). The profile interpretations are shown below the seismic profiles. See Figure 2 for location of the seismic lines and Table 2 for a detailed description of the units.

Figure 8  
205x150mm (300 x 300 DPI)

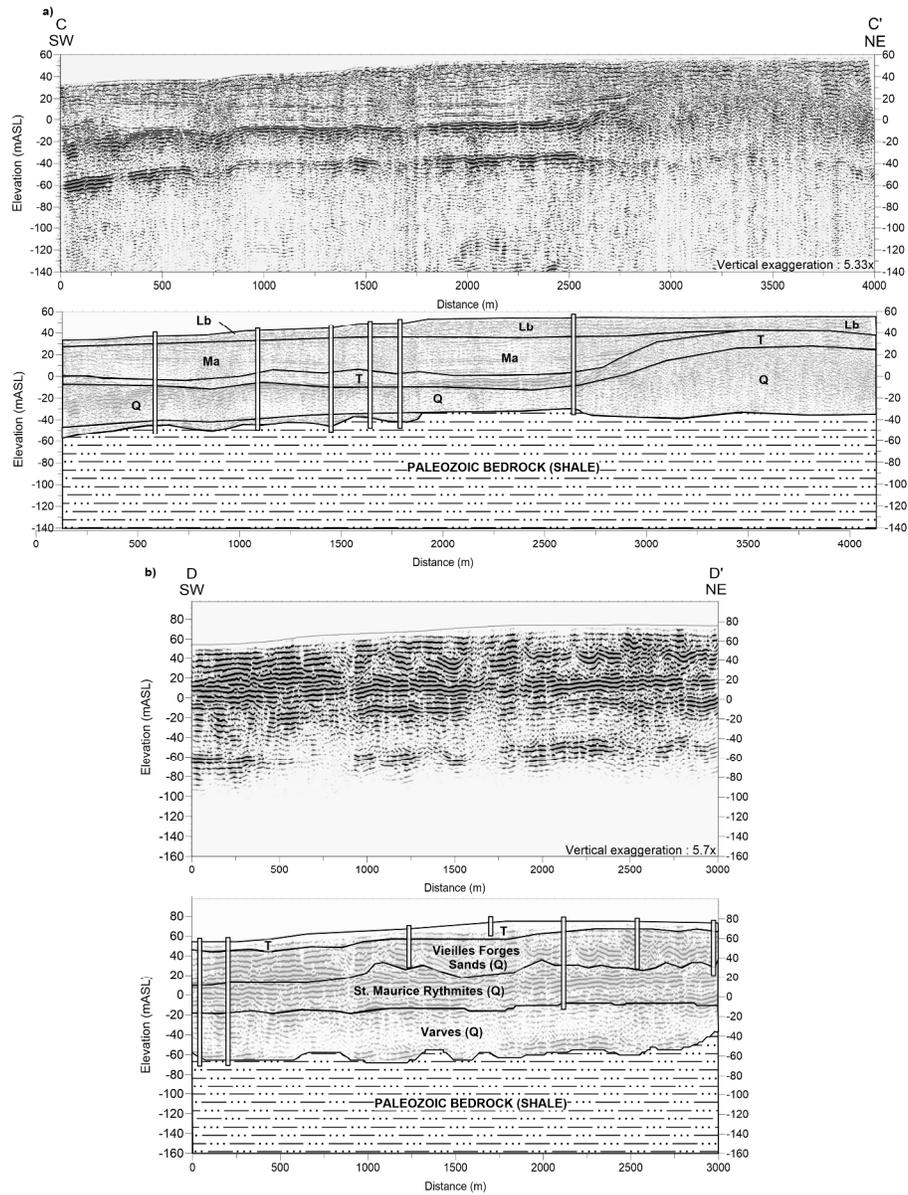


Figure 9. High resolution seismic survey lines L3-2010 (a) and L3-2011 (b). The profile interpretations are shown below the seismic profiles. See Figure 2 for location of the seismic lines and Table 2 for a detailed description of the units.

Figure 9  
271x348mm (300 x 300 DPI)

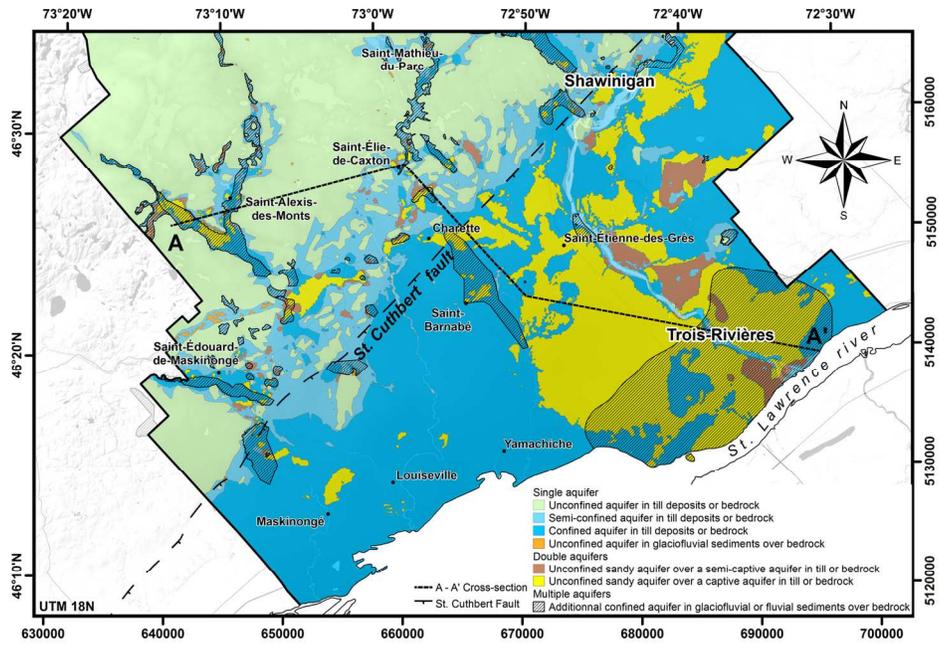


Figure 10. Aquifer types of the Mauricie region based on the architecture of the hydrostratigraphic units.

Figure 10  
152x107mm (300 x 300 DPI)

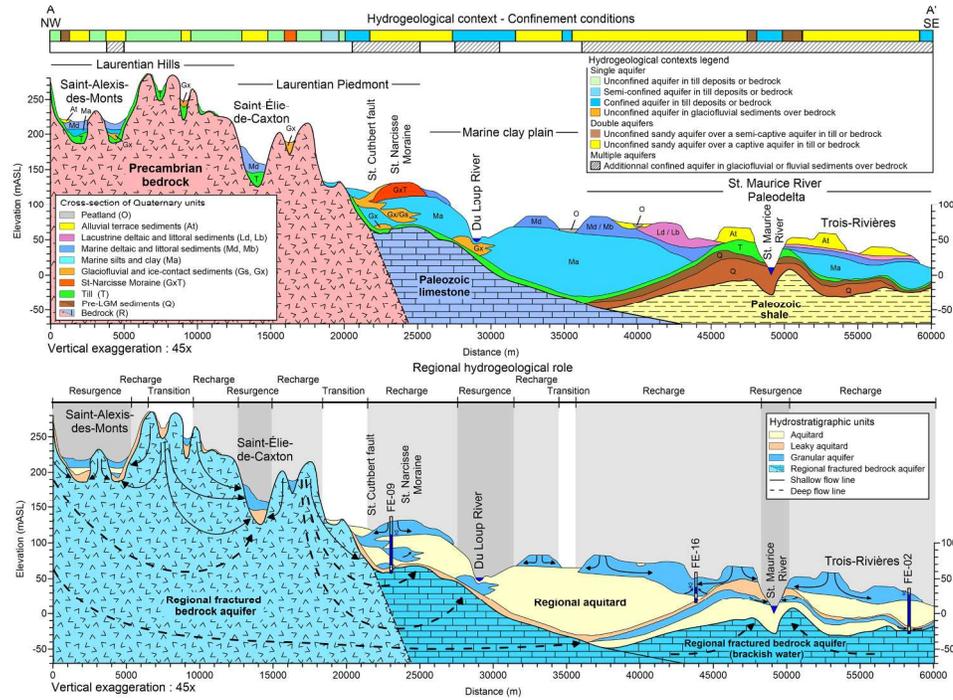


Figure 11. Cross-section illustrating the typical geologic and hydrostratigraphic units encountered in southwest Mauricie. The trace of cross-section A-A' is shown as a dashed line on Figure 10.

Figure 11  
223x176mm (300 x 300 DPI)