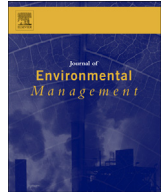




Contents lists available at ScienceDirect

Journal of Environmental Management

journal homepage: www.elsevier.com/locate/jenvman

Research article

Can isolated and riparian wetlands mitigate the impact of climate change on watershed hydrology? A case study approach

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ARTICLE INFO

Article history:

Received 21 February 2016

Received in revised form

14 August 2016

Accepted 14 September 2016

Available online xxx

Keywords:

Isolated wetlands

Riparian wetlands

HYDROTEL

Watershed-scale modelling

Climate change

Hydrological indicators

ABSTRACT

The effects of wetlands on stream flows are well established, namely mitigating flow regimes through water storage and slow water release. However, their effectiveness in reducing flood peaks and sustaining low flows is mainly driven by climate conditions and wetland type with respect to their connectivity to the hydrographic network (*i.e.* isolated or riparian wetlands). While some studies have demonstrated these hydrological functions/services, few of them have focused on the benefits to the hydrological regimes and their evolution under climate change (CC) and, thus, some gaps persist. The objective of this study was to further advance our knowledge with that respect. The PHYSITEL/HYDROTEL modelling platform was used to assess current and future states of watershed hydrology of the Becancour and Yamaska watersheds, Quebec, Canada. Simulation results showed that CC will induce similar changes on mean seasonal flows, namely larger and earlier spring flows leading to decreases in summer and fall flows. These expected changes will have different effects on 20-year and 100-year peak flows with respect to the considered watershed. Nevertheless, conservation of current wetland states should: (i) for the Becancour watershed, mitigate the potential increase in 2-year, 20-year and 100-year peak flows; and (ii) for the Yamaska watershed, accentuate the potential decrease in the aforementioned indicators. However, any loss of existing wetlands would be detrimental for 7-day 2-year and 10-year as well as 30-day 5-year low flows.

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1. Introduction

Wetlands have been recognized for their significant role on the hydrological cycle through water storage and slow release of water (Ogawa and Male, 1986; Padmanabhan and Bengtson, 2001; Liu et al., 2008; Wang et al., 2008; Wu and Johnston, 2008; Yang et al., 2010). Despite this general agreement, approximately half of their spatial extent has been lost worldwide (Zedler and Kercher, 2005). With the largest agricultural activity (Jobin et al., 2004) and the most populated region of Quebec (Li and Ducruc, 1999), Canada, the St. Lawrence Lowlands are characterized by several heavily affected ecosystems and among them, wetlands are no exception with almost 45% of affected areas (Joly et al., 2008). Their degradation or conversion in response to increase human activities (*e.g.*, agriculture, peat mining, and urbanization) is well known as one of the key disturbances of watershed hydrology and impairment to the global environment (Solomon et al., 2007). Thenceforward,

their deterioration may have severe impacts on flow regimes (*i.e.*, watershed hydrology) inducing a positive feedback under climate change conditions.

Hydrological modelling has proven to be a useful framework to assess climate change impacts on wetlands (Fu et al., 2015; Fossey and Rousseau, 2016) or watershed hydrology (Quilb  et al., 2008; Boyer et al., 2010) and to illustrate the impacts of these landscape features on flow regimes (Wang et al., 2010; Martinez-Martinez et al., 2014; Golden et al., 2015; Evenson et al., 2015; Fossey et al., 2016). However, some gaps persist regarding the evolution of wetland impacts on watershed hydrology under changing climate conditions. Over the last decade, scientific knowledge on future climate impacts in Quebec, Canada, has advanced and expected changes are well identified (Ouranos, 2015). Indeed, in the St. Lawrence Lowlands, projected annual temperatures may increase by 2–4 °C for the period 2040–2070. While the total annual precipitation may be relatively constant, seasonal modifications are expected, particularly in winter. The decrease in snowfall and the increase in rainfall during this season will affect stream flows (Appendix A). Indeed, the increase in mean temperature and the

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associated decrease in the ratio of snow to liquid precipitation should lead to an increase in the amount of water available for winter flows and a decrease in snow water equivalent reducing spring, summer and fall flows (Boyer et al., 2010).

These expected changes and resulting impacts on water availability, and thus on the water balance at the watershed scale, may affect wetlands. Indeed, the quantity and periodicity of wetland inflows could be affected. In return, wetlands could be more vulnerable and their hydrological functions and/or physical integrity be threatened (Erwin, 2009; Hebb et al., 2013; Fossey and Rousseau, 2016). Nevertheless, some questions arise such as: (i) What is the contribution of wetlands to current flow regimes? (ii) How will the role of wetlands evolve under changing climate conditions? (iii) Will their impacts on watershed hydrology be modified in the future?; and (iv) If wetlands are vulnerable, what will be the impact on the hydrological regime?

Using a distributed hydrologic modelling platform, this study assesses: (i) the current contribution of wetlands to stream flow and (ii) the expected changes at the outlet of two key watersheds of the St. Lawrence lowlands, namely those of the Becancour and Yamaska Rivers. This assessment, through analysis of some commonly used hydrologic indicators (i.e., Q_{max} , Q_7 and Q_{30}) provides a framework to distinguish climate change and wetland (i.e., isolated and riparian) effects on watershed hydrology. It also highlights the potential benefits of wetlands under changing climate conditions.

2. Methods

2.1. Study areas

Based on criteria related to land cover representativeness and wetland diversity, two watersheds were selected: the Becancour and Yamaska River watersheds. They cover 2597 km² and 4788 km² in the St. Lawrence Lowlands ecoregion (Mackey et al., 1996; McKenney, 1998) in southern Quebec, Canada, respectively (Fig. 1). Tributaries of the St. Lawrence River, the Becancour and Yamaska Rivers drain landscapes dominated by forest (54% and 35%, respectively), agriculture (23% and 41%), and wetlands (12% and 4%) (Beaulieu et al., 2012). As introduced in Table 1, in the Becancour watershed, wetlands cover a total area of 307 km² and drain 794 km² (31% of the watershed). Among them, isolated wetlands (IWs) occupy 210 km² (8%) and drain 464 km² (18%) while riparian wetlands (RWs) occupy 97 km² (4%) and drain 330 km² (13%). In the Yamaska watershed, wetlands cover a total area of 202 km² including 92 km² (2%) of IWs and 110 km² (2%) of the watershed) of RWs. From another point of view, the Yamaska wetlands drain 646 km² (13%) including 305 km² (6%) drain by IWs and 341 km² (7%) by RWs.

Located in a humid continental climate (Köppen-Geiger classification: Dfb - Peel et al., 2007), the watersheds are characterized by warm summers and severe winters with strong seasonality. Normal conditions for the 1981–2010 period (MDDELCC, 2014) are associated with an annual mean temperature of 5.1 °C with an average maximum of 19.3 °C in July and an average minimum of –11.2 °C in January. The total mean annual precipitation is 1210 mm, including 953 mm of liquid precipitation and 257 mm of solid precipitation (Appendix A). For the Becancour and Yamaska watersheds, the specific average monthly values of climate statistics are reported in Appendix B.

2.2. Hydrologic modelling platform

2.2.1. Model description

The PHYSITEL/HYDROTEL distributed hydrological modelling

platform was used to evaluate the expected changes due to climate change and to assess the effects of wetland types on watershed hydrology; quantifying their respective contribution to the flow regime (Fossey et al., 2015). More specifically, PHYSITEL is a specialized GIS and HYDROTEL is a continuous distributed hydrologic model (Turcotte et al., 2001, 2003, 2007; Fortin et al., 2001; Rousseau et al., 2011; Bouda et al., 2012, 2014; Noël et al., 2014). The model is currently used for inflow and hydrological forecasting at Hydro-Quebec, Quebec's primary power utility, and the Quebec Hydrological Expertise Centre (Centre d'Expertise Hydrique du Québec, CEHQ). This modelling platform can explicitly account for isolated (IWs) and riparian wetlands (RWs) (Fossey et al., 2015), using the Hydrologically Equivalent Wetland (HEW) concept (Liu et al., 2008; Wang et al., 2008). The basic computational units used by HYDROTEL and discretized using PHYSITEL correspond to Relatively Homogeneous Hydrological Units (RHHUs: sub-watersheds or hillslopes) and interconnected river segments (Fig. 2).

For this study, the two watersheds were discretized as follows: the Becancour and Yamaska watersheds, into 1824 and 1299 hillslopes (mean surface areas of 2.6 km² and 7.3 km², respectively), and 736 and 513 river segments (mean length of 1179 m and 1623 m), respectively. Similarly, the watersheds were discretized into 859 and 498 isolated HEWs (mean surface areas of 0.24 km² and 0.22 km², respectively) and 444 and 296 riparian HEWs (mean surface areas of 0.22 km² and 0.31 km²), respectively.

2.2.2. Model set up

HYDROTEL was run using a daily time step and basic meteorological data (i.e., precipitation, minimum and maximum temperatures obtained from CEHQ for the 1961–2010 period). Meanwhile for this study, the simulated daily stream flows at the outlet corresponded to the output of interest. Model calibration and validation were done using a manual trial-and-error strategy (Turcotte et al., 2003) over different five-year intervals of the 1969–2010 period while using 1-year spin-up period to minimize initialization errors (for a complete overview, see Fossey et al., 2015). Both temporal and spatial validations were performed, allowing a quantitative and qualitative verification of simulated flows. Statistics for calibration and validation periods are reported in Table 2.

2.2.3. Theoretical background

The addition of wetlands module has improved the model's ability to accurately reproduce the basic components of hydrograph (i.e., magnitude, frequency, timing and duration of water conditions) as reported in Fossey et al. (2015) for the Becancour watershed. These performance improvements increased the values of statistical indices from approximately 13% and 8% for the Becancour and the Yamaska (Rousseau et al., 2008, 2012) watersheds, respectively. Meanwhile, with respect to the hydrological modelling approach used for the Becancour watershed (Fossey et al., 2016), the impacts of wetlands depend on both their type (i.e., isolated or riparian) and their geographic location within a watershed (i.e., upstream/downstream gradient and stream order gradient). Moreover, for a given type and location, the results under similar hydro-climatic conditions illustrate the highly individualized and contrasting (i.e., positive or negative impacts) response that wetlands can exhibit as reported by Nilsson et al. (2013).

2.3. Climate scenarios

Daily climatic data for the 1961–2099 period were supplied by Uranos, a Consortium on Regional Climatology and Adaptation to Climate Change. Ten simulations of the Canadian Regional Climate Model (CRCM 4 ×) (Caya and Laprise, 1999; Music and Caya, 2007;

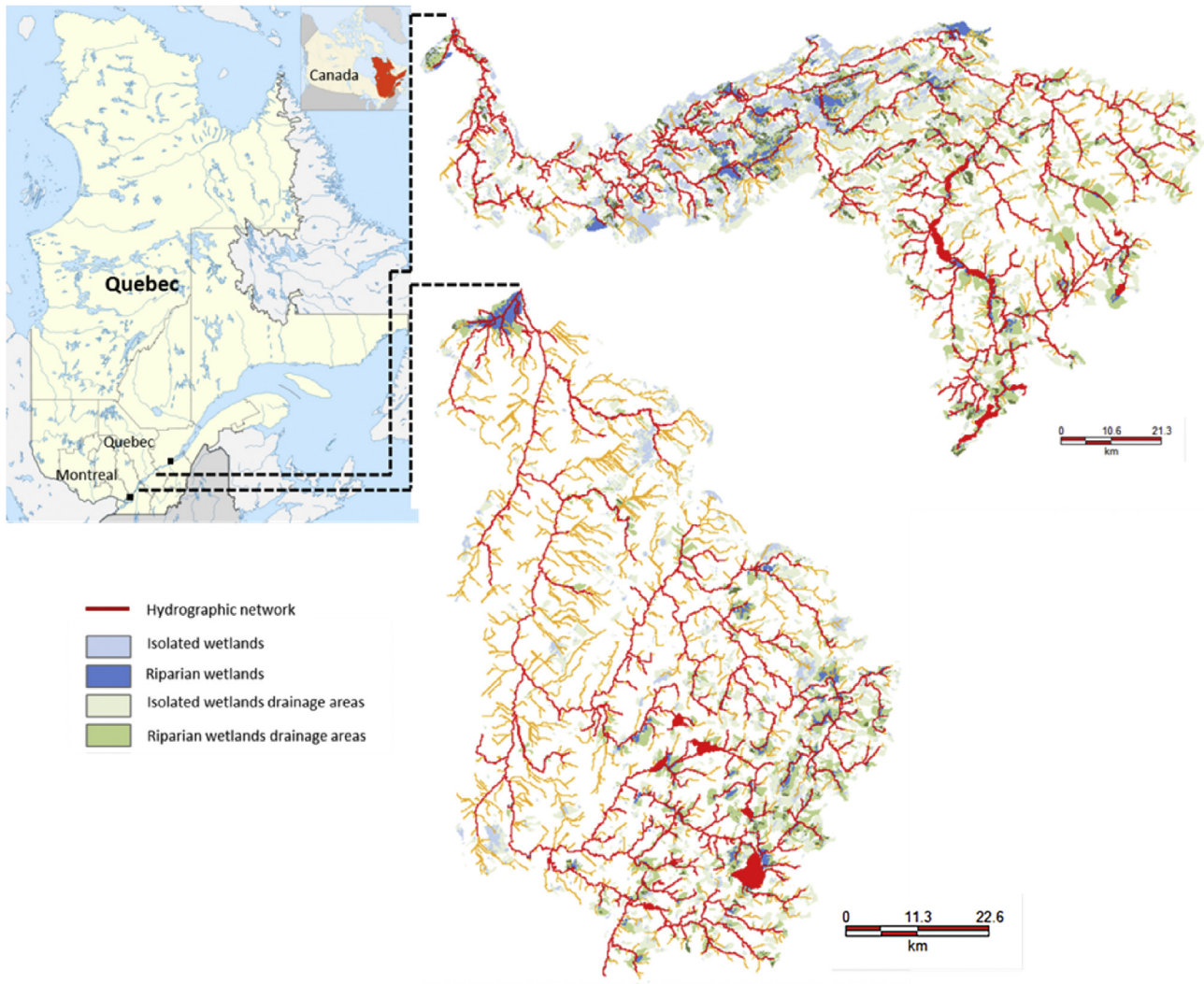


Fig. 1. Geographic locations of the Becancour River and Yamaska River watersheds, Quebec, Canada.

Table 1
Physical and wetlands statistics for the Becancour and Yamaska River watersheds.

Watershed	Surface (km ²)	Land cover (%)				Wetland area		
		Forest	Agriculture	Bush	Urban	Total km ² (%)	Isolated km ² (%)	Riparian km ² (%)
Becancour	2597	53.9	23	5.6	3.5	307 (12)	210 (8)	97 (4)
Yamaska	4788	34.6	43.1	9.3	7.3	202 (4)	92 (2)	110 (2)
		Physical characteristics				Wetland drainage area		
		Average slope (%)	Main soils	Gravelius index	Drainage density ^a	Total km ² (%)	Isolated km ² (%)	Riparian km ² (%)
Becancour	0.5	Sandy loam-loam	2.3	0.52	794 (31)	464 (18)	330 (13)	
Yamaska	0.7	Silt loam-sandy loam	1.7	0.50	646 (13)	305 (6)	341 (7)	

(%) is the area or drainage area of wetland expressed in percentage.

^a Horton concept = Σ length of river (km)/watershed area (km²).

Paquin, 2010; de Elia and Côté, 2010) using the A2 greenhouse gas and aerosol projected scenarios of the Special Report on Emission Scenarios (SREAS-A2) (Nakicenko et al., 2000) were used in this study. Note that these simulations were deemed representative of a myriad of potential climate changes using the cluster method (Hartigan and Wong, 1979) and considered equally probable. Climatic data may have a bias in the estimation of climate variables

(i.e., constant overestimation or underestimation trend) and then require post-treatment. The CRCM simulations were bias-corrected by a factor calculated from a comparison between the simulated reference period and the observed data. The bias was then removed on the entire simulated series.

These unbiased CRCM simulations were driven by two Global Climate Models (GCMs); that is two members (i.e., initial conditions

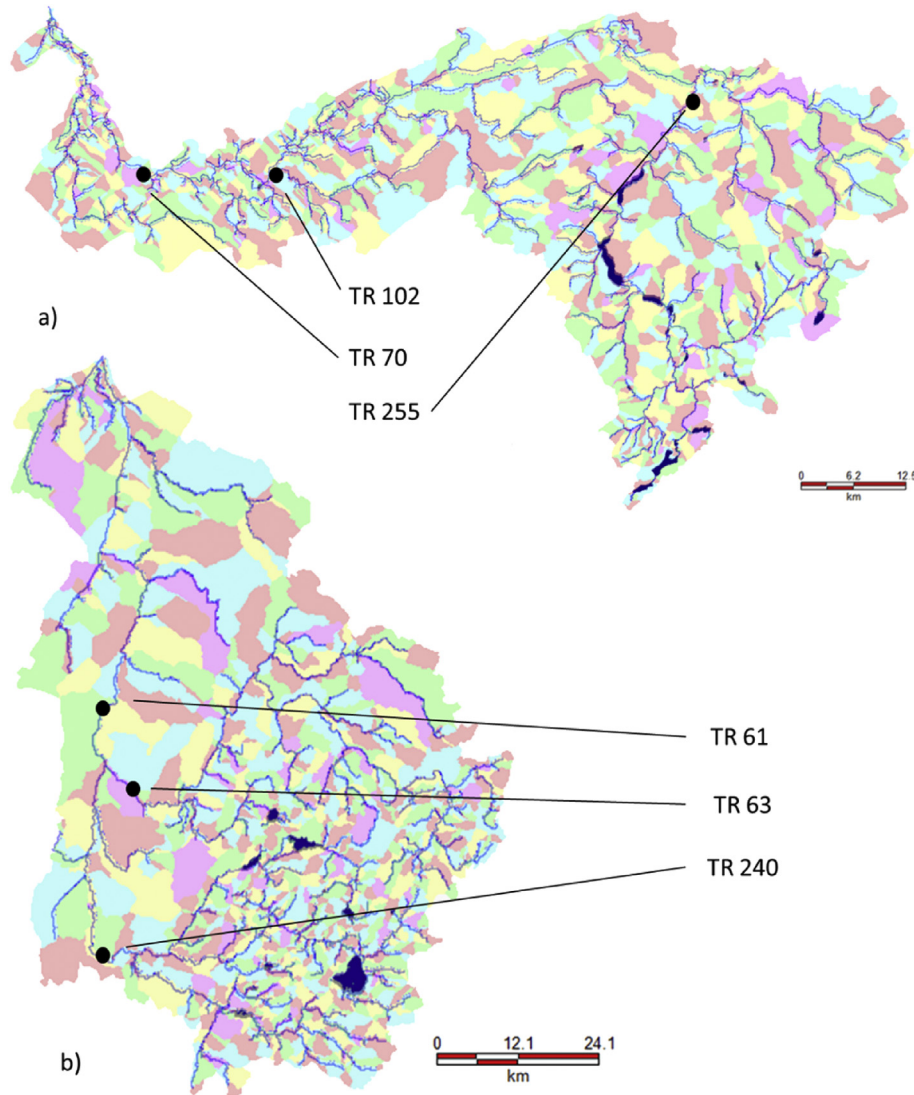


Fig. 2. Subdivision of the Becancour River (a) and the Yamaska River (b) watersheds into computational elements (hillslopes and river segments) and locations of gauge stations (TR). Note that hillslope colors do not have any specific meanings; they are solely used for illustration purposes. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Table 2

Model performances for calibration and validation periods.

Watershed	Stream flow station	Period	Statistical performance			Objective
			NSE	RMSE (m ³ /s)	Pbias (%)	
Becancour	TR_70	1988–1992	0.78	13.9	–4.4	Spatio-temporal validation
		2005–2009	0.77	28.3	–5.5	Calibration
	TR_102	2000–2004	0.80	24.3	2.6	Temporal validation
		2005–2009	0.78	13.9	–4.4	Calibration
Yamaska	TR_255	1999–2004	0.86	9.3	6.6	Temporal validation
		2005–2010	0.78	42.9	0.1	Calibration
	TR_61	2000–2005	0.78	38.3	–13	Temporal validation
		2005–2010	0.73	23.7	2.5	Calibration
Yamaska	TR_63	2000–2005	0.71	21.9	–14.2	Temporal validation
		2005–2010	0.75	16.0	–2.5	Calibration
	TR_240	2000–2005	0.77	14.4	–13.6	Temporal validation

of driver) for the third generation Canadian Coupled Global Climate Model - CGCM3 (Flato et al., 2000; Flato and Boer, 2001; Scinocca et al., 2008) and one for the German coupled Global Climate

Model - ECHAM5 (Jungclaus et al., 2006) over a domain covering North America or centred over Quebec both with a horizontal grid-size mesh of 45 km. An overview of the technical characteristics of

Table 3
CRCM versions and characteristics of simulations.

Operational name	acu	adc	adj	adl	aet	aev	afx	agr	agx	aha
CRCM version	4.1.1	4.1.1	4.2.0	4.2.0	4.2.3	4.2.3	4.2.3	4.2.3	4.2.3	4.2.3
Driving data	CGCM3	CGCM3	CGCM3	CGCM3	CGCM3	CGCM3	CGCM3	CGCM3	ECHAM5	ECHAM5
Member	4	5	4	5	4	5	4	5	1	1
Regional domain	Quebec	Quebec	Quebec	AMNO	AMNO	AMNO	Quebec	Quebec	AMNO	Quebec

these scenarios is reported in Table 3. For the 2040–2070 period, the same meteorological variables as those described for the 1981–2010 period (cf. section 2.1 Study areas), were analysed and compiled. The average monthly values of climate statistics, with respect to the Becancour and Yamaska watersheds, for this future period (2040–2070) and integrating the ten simulations, are reported in Appendix C and Appendix D respectively.

2.4. Assessment of the impact of wetlands on watershed hydrology

This assessment was defined and based on wetland typologies (i.e., IWs versus RWs) stemming from the hydrogeomorphic (HGM) approach for landscape planning (Brinson, 1993, 2009). Four wetland-based simulations were performed accounting for: (i) all wetlands (WT), (ii) isolated wetlands (IW), (iii) riparian wetlands (RW), and finally without any wetlands (WW). The latter was defined as the reference simulation. For each simulation, there was one model run from 1979 to 2010 and another from 2039 to 2070 with the first year used as a spin-up period.

The assessment was determined through the analysis of daily stream flows at the watershed outlet in terms of seasonal mean stream flow and of two sets of frequently-used stream flow parameters: (i) a first set corresponding to three high-flow indicators, the 2-year (Q_2), 20-year (Q_{20}) and 100-year (Q_{100}) daily maximum flow (Q_{max}), and (ii) the other set made up of three low-flow indicators, the 2-year (Q_{2-7}) and 10-year (Q_{10-7}) minimum flows over seven consecutive days and the 5-year (Q_{5-30}) minimum flow over 30 consecutive days. These hydrological indicators were determined from a log-normal distribution which was selected based on the BIC (Bayesian information criterion) (Schwarz, 1978).

The impact of wetlands was expressed as a variation (%) of the calculated indicator for a given wetland-based simulation (WT, IW or RW) compared to the same indicator calculated on the basis of the WW simulation. This impact can be expressed as follows:

$$WII_{x,j} = 100 \times \left[\frac{(HI_{x,j} - HI_{ww})}{HI_{ww}} \right] \quad (1)$$

where $WII_{x,j}$ (%) is the wetland impact indicator for simulation (x) for given climatic conditions j (i.e., current -c- or future -f-); $HI_{x,j}$ is the hydrological indicator calculated for the considered wetland-based simulation (x) and climatic condition j , and HI_{ww} is the hydrological indicator calculated for the reference wetland-based simulation (i.e., for the without any wetland case).

3. Results and discussion

3.1. Model performance

The Nash-Sutcliffe efficiency (NSE) (Nash and Sutcliffe, 1970) criterion, the root-mean-squared error (RMSE) (Singh et al., 2004), and the percentage bias (Pbias) (Yapo et al., 1996) were used to assess model performance. As reported in Table 2, the NSE criterion

values are greater than 0.75 for the Becancour watershed and greater than 0.70 for the Yamaska watershed either for the period of calibration or that of validation. The Pbias values are less than 10% except for the validation period of the Yamaska watershed where they range from –10% to –15%. These negative values may reflect a slight tendency of the model to underestimate the flows. The RMSE calculated for each gauge station shows a relatively stable error. According to the hydrologic model performance rating described by Moriasi et al. (2007), the model calibration may be considered very good (NSE > 0.75 and Pbias < 10%) to good (NSE > 0.65 and Pbias < 15%).

3.2. Hydrology of the study watersheds under current and future climatic periods

As shown in Fig. 3, when compared to the mean simulated hydrograph characterizing the current period (1980–2010), each mean simulated hydrograph obtained for the future period (2040–2070) presents higher flows during winter and lower flows in spring, summer and fall. These results are consistent with those of Boyer et al. (2010). Also, simulations indicate shifts and decreases in future spring floods. Summer and fall flows tend to be lower according to: (i) the shift of the spring flood period, (ii) the decrease in snowpack, reducing the period during which the water is available for flows, and (iii) the slight decrease in rainfall. Note that evapotranspiration may be time-varying due to the potential early growing period (Badeck et al., 2004) and decrease, to some extent, the water available for flows. Considering the mean seasonal specific flows (i.e., flow per unit area – $m^3/s/km^2$) for the current period, despite being twice as large as the Becancour watershed, the flows of the Yamaska watershed are of the same order (0.011–0.013) as those of the Becancour during summer and fall. Meanwhile for winter, the flow of the Yamaska is nearly twice as large (0.014) as that of the Becancour (0.008). Contrariwise, for spring, the mean flow of the Becancour watershed is larger (0.035) than that of the Yamaska (0.024). These results suggest that the Becancour watershed has a more variable flow regime when compared to that of the Yamaska which presents a more stable specific flow throughout the year apart from that of spring. Thenceforth, the expected changes due to CC vary from one watershed to another. More precisely, for the Becancour watershed, these variations range from 75% to 119% (winter), from –24% to –5% (spring), from –44% to –34% (summer) and from –42% to –18% (fall). For the Yamaska watershed, they are from 56% to 68% (winter), –8%–19% (spring), –43% to –26% (summer) and –43% to –15% (fall). The differences observed during winters and springs are the result of the dynamics of snow processes (i.e., modifications of snow accumulations, ratio of solid to liquid precipitation, maximum and minimum temperatures) mainly driven by local climatic conditions and some physical characteristics of watersheds (e.g., size, shape, drainage density, etc.) which could be factors decreasing or increasing their sensitivity to local weather conditions.

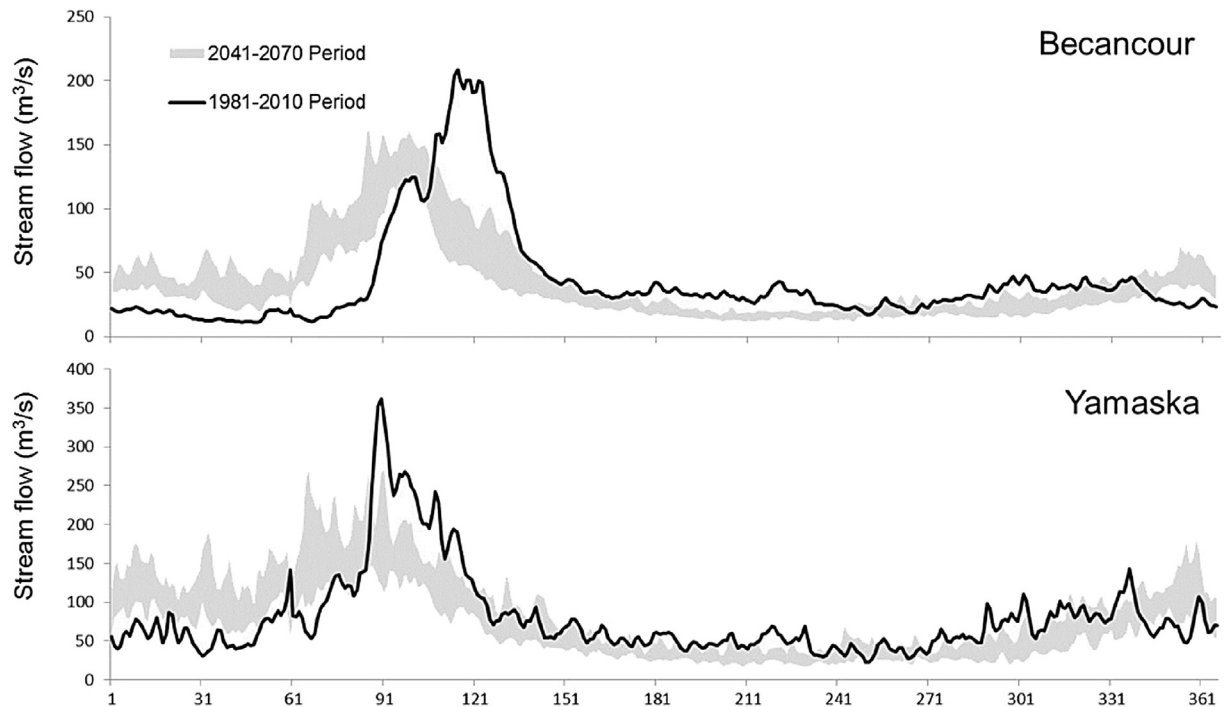


Fig. 3. Mean simulated hydrographs obtained for the reference period (1981–2010) and future climate scenarios (range of hydrographs over 2041–2070).

Table 4
Average annual high- and low-flow parameters (m^3/s) under current climate conditions (Ref – 1980–2010), and changes (%) for each parameter under future climate conditions (2040–2070).

	Flow parameter (m^3/s) and percentage of variation (%)											
	Becancour						Yamaska					
	High flows			Low flows			High flows			Low flows		
	Q_2	Q_{20}	Q_{100}	Q_{2-7}	Q_{10-7}	Q_{5-30}	Q_2	Q_{20}	Q_{100}	Q_{2-7}	Q_{10-7}	Q_{5-30}
Ref	352	557	673	7	5	6	825	1517	1954	9	7	11
acu	-9.1	-6.3	-5.1	-13.5	-20.1	3.4	-5.0	-16.3	-20.7	-28.2	-38.9	-14.4
adc	-13.4	-5.1	-1.3	-7.9	-17.0	11.3	-6.4	-5.3	-4.9	-27.6	-46.8	-21.7
adj	-14.4	2.6	10.7	-21.4	-32.5	-8	-16.2	-19.3	-20.6	-39.0	-52.2	-28.7
adl	-12.3	6.2	15.0	-3.3	-19.5	11.6	-1.8	-0.8	-0.4	-30.5	-51.3	-25.7
aet	-14.4	8.6	20.0	-19.2	-27.5	0.8	-13.0	-16.7	-18.2	-38.3	-48.2	-25.7
aev	-11.8	11.1	22.4	-7.6	-21.9	8.1	3.2	3.7	3.8	-33.3	-49.4	-22.8
afx	-11.7	-14.4	-15.4	-5.3	-11.5	13.4	-6.5	-17.7	-22.0	-24.0	-36.1	-11.6
agr	-14.5	-2.7	2.8	1.9	-4.5	23.0	-13.1	-13.1	-13.1	-22.3	-42.2	-15.7
agx	-7.4	12.8	22.5	-7.7	-18.8	5.8	-1.5	-8.2	-11.0	-24.9	-44.4	-15.0
aha	-1.0	15.6	23.4	1.3	-3.7	24.2	-0.8	3.0	4.6	-17.6	-30.4	-0.2

Also, and as reported in Table 4, these variations affect the values of the hydrological indicators both for the Becancour and Yamaska River watersheds, but with different trends. For the Becancour, the Q_2 trend is downward for all climatic scenarios indicating a decrease between 1% and 14.5%. However, for the Q_{20} and Q_{100} , the simulated trend is less clear with variations ranging between -14.4% and 15.6% and from -15.4% to 23.4%, respectively. Low-flow indicators show a downward trend for both Q_{2-7} and Q_{10-7} with mainly negative values obtained for all scenarios, ranging from -21.4% to 1.3% and from -32.5% to -3.7%, respectively. The last low-flow indicator (Q_{5-30}) displays an increasing trend up to 24.2%, with a single negative value of -8%.

While the Yamaska watershed is characterized by an overall downward trend for all indicators. The majority of high-flow

indicators show negative trends (8 scenarios at least), whereas all low-flow indicators have negative variations. The values vary from -16.2% to 3.2% for Q_2 , from -19.3% to 3.7% for Q_{20} , and from -22% to 4.6% for Q_{100} . Similarly, the values of the low-flow indicators vary from -39% to -17.6% for Q_{2-7} , from -52.2% to -30.4% for Q_{5-30} .

Similar decreasing trends for Q_2 , Q_{2-7} and Q_{10-7} were expected for both studied watersheds. Differences observed for Q_{20} , Q_{100} and Q_{5-30} may be related to several parameters (all values of hydrological indicators are reported in Tables 5 and 6). On one hand, landscape conditions (*i.e.*, slope, soil characteristics, Gravelius index and land cover – without loss of continuity not reported here) for the Becancour watershed differ from those of the Yamaska watershed, affecting spatial and temporal distributions of flows as

Table 5

Synoptic table of the impacts of all wetlands (WT), isolated wetlands (IW) and riparian wetlands (RW) on hydrological indicators for the Becancour River watersheds under current climate conditions (Ref – 1980–2010) and future climate conditions (2040–2070).

	Ref	Hydrological indicators (m ³ /s) for the Becancour watershed										
		acu	adc	adj	adl	aet	aev	afx	agr	agx	aha	
Q ₂	WT	352	320	305	301	309	301	310	311	301	326	348
	WW	–	411	401	395	414	398	413	401	398	433	452
	IW	–	335	319	315	323	316	325	326	313	343	364
	RW	–	372	360	357	369	358	370	361	355	387	411
Q ₂₀	WT	557	522	529	571	591	605	619	477	542	628	644
	WW	–	682	743	791	820	808	883	628	738	851	820
	IW	–	566	577	626	647	667	687	517	592	684	686
	RW	–	608	644	692	716	711	760	555	647	761	745
Q ₁₀₀	WT	673	639	664	745	774	808	824	569	692	824	831
	WW	–	842	959	1054	1088	1083	1210	757	954	1126	1049
	IW	–	703	738	833	863	908	936	626	770	911	892
	RW	–	745	819	911	943	945	1024	663	830	1007	954
Q ₂₋₇	WT	7	6	6	6	7	6	6	7	7	6	7
	WW	–	3	3	3	3	3	3	4	4	3	4
	IW	–	6	7	6	7	6	7	7	7	7	7
	RW	–	3	4	3	4	3	3	4	4	4	4
Q ₁₀₋₇	WT	5	4	4	3	4	4	4	4	5	4	5
	WW	–	2	2	1	2	1	2	2	2	2	2
	IW	–	4	5	4	4	4	4	5	5	4	5
	RW	–	2	2	1	2	2	2	2	2	2	2
Q ₅₋₃₀	WT	6	6	7	6	7	6	6	7	7	6	7
	WW	–	4	4	3	5	4	4	4	5	4	5
	IW	–	7	7	6	7	6	7	7	8	7	8
	RW	–	4	5	3	5	4	4	5	5	4	5

Table 6

Synoptic table of the impacts of all wetlands (WT), isolated wetlands (IW) and riparian wetlands (RW) on hydrological indicators for the Yamaska River watersheds under current climate conditions (Ref – 1980–2010) and future climate conditions (2040–2070).

	Ref	Hydrological indicators (m ³ /s) for the Yamaska watershed										
		acu	adc	adj	adl	aet	aev	afx	agr	agx	aha	
Q ₂	WT	825	784	772	692	810	717	851	771	717	813	818
	WW	–	828	798	732	855	758	899	815	758	862	863
	IW	–	776	764	687	800	711	841	763	709	805	810
	RW	–	831	818	733	860	759	899	818	761	864	866
Q ₂₀	WT	1517	1269	1436	1224	1505	1264	1573	1248	1319	1392	1563
	WW	–	1344	1464	1296	1586	1332	1659	1319	1396	1485	1656
	IW	–	1256	1418	1215	1488	1249	1551	1237	1302	1376	1545
	RW	–	1349	1532	1298	1594	1339	1659	1323	1405	1489	1662
Q ₁₀₀	WT	1954	1550	1858	1551	1945	1598	2029	1524	1698	1740	2043
	WW	–	1643	1883	1642	2049	1683	2138	1611	1799	1860	2169
	IW	–	1534	1833	1539	1924	1577	1998	1511	1674	1717	2018
	RW	–	1649	1987	1645	2058	1694	2138	1615	1811	1866	2178
Q ₂₋₇	WT	9	6	7	5	6	6	6	7	7	7	7
	WW	–	5	5	4	4	4	4	5	5	5	6
	IW	–	6	7	6	6	6	6	7	7	7	7
	RW	–	5	5	4	5	4	4	5	5	5	6
Q ₁₀₋₇	WT	7	4	4	3	3	4	4	4	4	4	5
	WW	–	3	2	2	2	2	2	3	3	3	3
	IW	–	4	4	3	4	4	4	4	4	4	5
	RW	–	3	2	2	2	2	2	3	3	3	4
Q ₅₋₃₀	WT	11	9	9	8	8	8	8	10	9	9	11
	WW	–	8	7	6	7	7	7	8	8	8	9
	IW	–	9	9	8	8	8	9	10	9	9	11
	RW	–	8	7	7	7	7	7	8	8	8	10

introduced above. On other hand, meteorological characteristics for the future period vary slightly between the two watersheds. Based on Appendix C and Appendix D, it can be noted that, under changing climate conditions, temperatures are expected to increase by 3–4 °C for the Becancour and by 2–3 °C for the Yamaska, rainfalls are expected to increase by 18%–30% and 12%–23%; while snow is expected to decrease by 31%–45% and 34%–50%, respectively. In other words, these slight variations in precipitations and

temperatures, as suggested by these first results (*i.e.*, more variable flow regime for Becancour watershed and differences with respect to snow processes) coupled with landscape conditions may partially explain some of the differences found. However, while landscape conditions are considered time-invariant in our modeling approach, their impacts on watershed hydrology are constant in time and space.

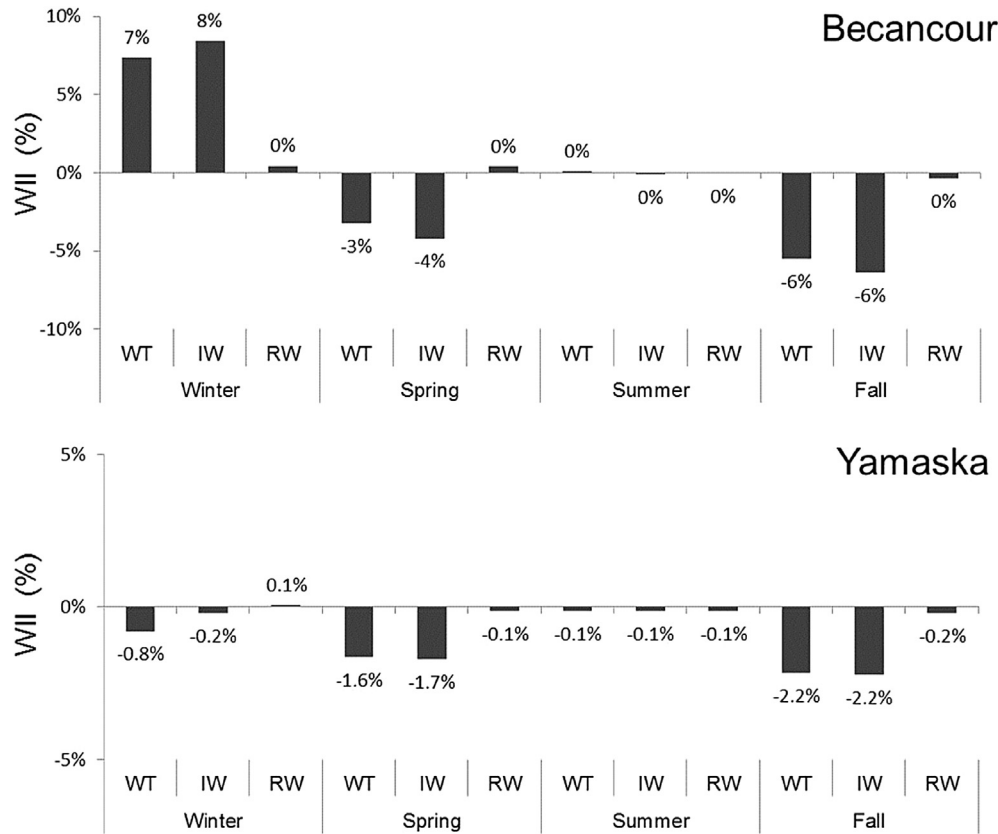


Fig. 4. Impacts of all wetlands (WT), isolated wetlands (IW) and riparian wetlands (RW) on mean seasonal specific flows in the Becancour and Yamaska River watersheds under the current period (1981–2010).

3.3. Current impacts of wetlands on watershed hydrology

As described by Fossey et al. (2016), the impact of wetlands on downgradient flows depends, to some extent, on upstream fluxes which are regulated by landscape conditions and climate. Therefore, for similar hydroclimatic conditions, the resulting impacts are governed by time-varying and space-varying fluxes. According to the hydrological modelling approach used, changes in upstream fluxes affect the hydrodynamics and hydrological functions of wetlands.

Impacts of wetlands on mean seasonal specific flows appear to be variable throughout the year (Fig. 4). Three distinct results can be highlighted: (i) during springs and falls, wetlands seem to participate to the decline of the mean seasonal flows, (ii) during summers, the impacts are almost null and (iii) during winters, while the Becancour wetlands induce a slight increase in the mean seasonal flow, for the Yamaska watershed, their presence appears to slightly decrease the mean seasonal flow. At these seasonal scales, RWs seem to have very limited effects. These findings are linked to similar IW and RW attributes (*i.e.*, wetland area and drainage area) within the Yamaska watershed (Table 1).

As shown by the changes (%) in the values of the hydrological indicators caused by the different types of wetlands (Fig. 5), at the interannual scale, characterized by more timely and more extreme hydrological events, results differ. For the Becancour watershed, the WT simulation (12% of the watershed area in wetlands) leads to a decrease in high flows of 18% for Q_2 , 16% for Q_{20} , and 14% for Q_{100} ; whereas the mere presence of IWs (8% of the watershed) leads to

decreases of 16%, 13%, and 12%, respectively. Meanwhile, RWs (4% of the watershed) are responsible for a decrease of 7%, 7% and 6% for Q_2 , Q_{20} , and Q_{100} , respectively. For low flows, WT and IWs show a similar effect on both Q_{2-7} and Q_{10-7} with an increase of 75% and 67%, respectively. For the Q_{5-30} indicator, WT leads to an increase of 50% while it is 75% for IWs. RWs seem to have not as strong an impact, although substantial with an increase of 25% for Q_{2-7} and 0% for other indicators.

Some differences are observed for the Yamaska watershed. RWs seem to have no effect (0% of variation for Q_2) or a slight negative effect (1% increase) on the Q_{20} and Q_{100} . This negative effect of RWs, described in previous work (Fossey et al., 2016), results from the modelling of RWs which accounts for their capacity to contribute to the river flow during high flow conditions, reversing the capacity of storing water and enhancing the high flow conditions. It underlines the time-varying issue of water exchange conditions between RWs, rivers, and adjacent uplands (Burt et al., 2002). Considering low flows, WT and IWs have the same impact with variation of 29%, 40% and 22% for Q_{2-7} , Q_{10-7} and Q_{5-30} , respectively. RWs have less impact with an increase of 14% for Q_{2-7} and 0% for other indicators.

Overall, at the event scale any type of wetlands plays a role in reducing peak flows (decreasing values of high-flow indicators) and sustaining low flows (increasing values of low-flow indicators). However, at the seasonal scale, wetlands generally reduce flows. These results demonstrate the damping effect of wetlands on watershed hydrology. Indeed, the presence of wetlands limits high variation between maximum and minimum flows. However, the cumulative effect of all wetlands (WT simulation) remains less than

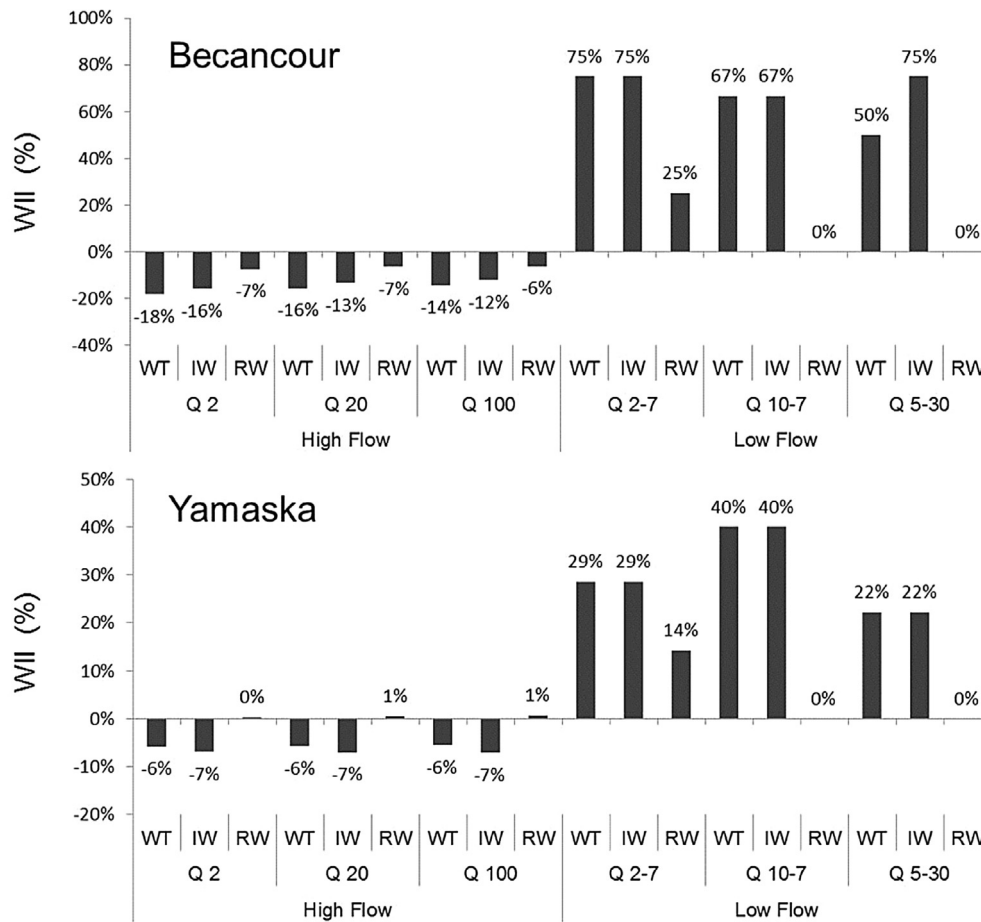


Fig. 5. Impacts of all wetlands (WT), isolated wetlands (IW) and riparian wetlands (RW) on hydrological indicators in the Becancour River and Yamaska River watersheds under the current conditions (1981–2010).

the sum of the individual effects of each typology (IWs and RWs) since the outflows of IWs are in some RHHUs routed through RWs. Moreover, for the Yamaska watershed, wetlands appear to have a constant impact regardless of the considered HIs. Whereas, for the Becancour watershed, the more the HIs are substantial (Q_2 to Q_{100}) the less wetlands are efficient flow regulators. These differences in results illustrate the complexity of the water exchanges between IWs, RWs, and surrounding landscapes as described by Fossey et al. (2016).

3.4. Evolution of the impacts of wetlands under changing climate conditions

Under CC conditions and considering the current wetland distributions, watershed hydrology will change (Fig. 3 and Table 4). At the seasonal scale, the flow variation due to wetlands appears low compared to the expected changes due to CC. So, the assessment of their impacts will be modulated by the variation of hydrological indicators (Tables 5 and 6). Not accounting for wetlands induce variable changes under CC conditions depending on both the considered watershed and the hydrological indicator of interest.

For the Becancour watershed, not accounting for IWs (8%) leads to increases of 1%–10% in the values of Q_2 . Meanwhile, for Q_{20} and Q_{100} , the conservation of existing wetlands is not enough to

maintain current flow conditions. Increases from 2.5% to 16% (Q_{20}) and from 3% to 23.5% (Q_{100}) could be expected. Thenceforth, any losses would increase flows. Moreover, the loss of all wetlands (12%) would increase the Q_{20} value to that of the Q_{100} . For low flows, results show that the conservation of existing wetlands could maintain current values or induce a decrease ranging from 17% (Q_{2-7}) to 40% (Q_{10-7}) or even lead to an increase of 17% (Q_{5-30}). Similar results are obtained with the conservation of IWs.

For the Yamaska watershed, the effects of CC related to wetlands conservation/loss are not as strong. For high flow indicators, most climatic/wetland scenarios would not produce increases and wetland conservation should be enough to maintain the current hydrological regime. In most cases, decreases are observed; ranging from 1% to 16% (Q_2), from 1% to 19% (Q_{20}) and from 0.5% to 22% (Q_{100}). It is noteworthy that in all scenarios, impacts of IWs are greater than the combined effect of IWs and RWs (i.e., WT scenario). As for the Becancour watershed, the total loss of wetlands (4%) or the loss of IWs (2%) would have similar impacts on these hydrological indicators. On the other end, the conservation of existing wetlands would not be enough to mitigate the expected decreases in low flows; ranging from 22% to 44% (Q_{2-7}), from 29% to 57% (Q_{10-7}), and from 0% to 27% (Q_{5-30}).

To summarize, wetlands would keep their current role of flow regulators. Their presence would still lead to a decrease of high

flows and an increase of low flows. However, their capacity to compensate for modifications due to climate change would be watershed-specific.

4. Conclusion

Wetlands are natural landscape features which contribute to the regulation of stream flows, reducing high flows and increasing low flows. The modifications in spatial and temporal distributions of meteorological patterns (rainfall, temperatures, snow) driven by CC conditions will affect the overall availability of water to flow through river networks. Within the Becancour River and the Yamaska River watersheds, these variations will lead to higher and earlier spring flows with shifts and decreases in future spring floods and lower summer and fall flows.

At the seasonal scale, the impacts of wetlands vary, but they contribute to a general decrease of mean seasonal flows. However, the results indicate that the impacts are relatively low when compared to the expected impacts of CC. Meanwhile, for high flow indicators, wetlands lead to a decrease in high flows between 18% and 14% and a decrease of 6%, for the Becancour River and the Yamaska River, respectively. For low flow indicators, for the Becancour the increase ranges from 50% to 75% and for the Yamaska from 22% to 40%. Under changing climate conditions, the conservation of wetlands for the Becancour watershed appears not sufficient to maintain current high flows (Q_{20} and Q_{100}) and, thus, compensate for the expected increase due to CC; for the Yamaska watershed, the results suggest that they would mitigate the decline caused by CC. On the other hand, regardless of the watershed, any loss would be harmful for low flows (Q_{2-7} , Q_{10-7} and Q_{5-30}). Based on the studied CC conditions, the current state of wetlands should provide a relevant support for low flows while maintaining a major role for flood control. In some cases, loss of wetlands would increase Q_{20} values to those of Q_{100} . However, different impacts induced by each typology of wetlands (IW and RWs) reinforce the individualized impact that wetlands have according to their local landscape conditions making them more or less sensitive to expected perturbations induced by CC.

That being said, the results presented here may be viewed as optimistic considering the rate of wetland areas loss (approximately 11% observed between 1984 and 2011) and their inherent vulnerability to climate change described in previous work (Fossey and Rousseau, 2016—40% of wetlands could be affected in their hydrological functions by climate change on the Becancour watershed). The results argue that it is only through some wetland restoration and conservation effort, increasing their total area, that wetlands could mitigate/compensate the CC effects. Finally, they suggest that it is nearly impossible to systematically transpose the impact of a given type of wetlands to another watershed. Thenceforth, if conservation/restoration programs are needed to compensate partially or totally the expected effect of CC, they should be first assessed using a watershed-scale hydrological modelling framework.

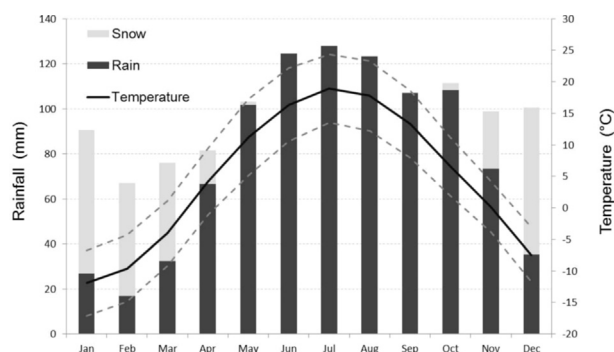
Acknowledgments

This study was done as part of a collaborative research project (*Outils d'analyses hydrologique, économique et spatiale des services écologiques procurés par les milieux humides des basses terres du Saint-Laurent: adaptations aux changements climatiques*) between Université de Sherbrooke (Richard Fournier, principal investigator, Jérôme Théau), Centre Eau Terre Environnement, INRS (Alain N. Rousseau), Université Laval (Monique Poulin), and UQAM (Jean-

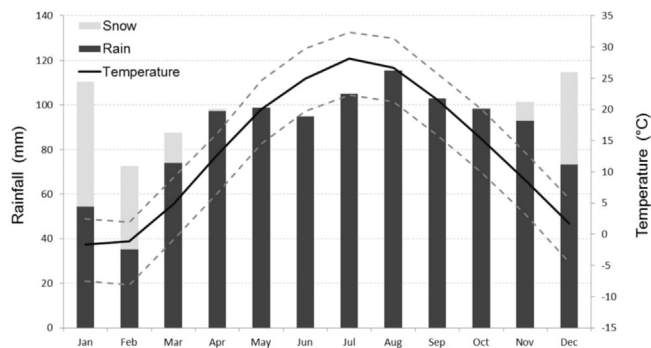
Pierre Revéret). It was conducted within the framework of the Climate Change Action Plan 26 “Biodiversity” coordinated and funded by Ouranos, a Consortium on Regional Climatology and Adaptation to Climate Change (Robert Siron, biodiversity program coordinator). Supplemental funding came from the Discovery Grant program of the Natural Sciences and Engineering Research Council (NSERC) of Canada (Alain N. Rousseau, principal investigator). We would like to thank Stéphane Savary and Alain Royer of INRS-ETE for their technical support. Finally, we would like to thank the reviewers, and editors for their meaningful comments. They significantly helped us improve the overall quality of the paper.

Appendix A

Reference climate conditions of the study area, 1981–2010
Reference climate conditions of the study area (1981–2010) (observed data).



Climate graph for the mean future conditions of the study area (2040–2070) from simulated data.



Appendix B

Monthly climate statistics for the Becancour and the Yamaska watersheds for the 1981–2010 period.

	Month	Temperatures (°C)			Precipitations		
		Max	Min	Mean	Rain (mm)	Snow (cm)	Total (mm)
Becancour River watershed	January	-6.7	-17	-11.9	26.9	63.9	89.9
	February	-4.3	-15	-9.6	16.9	50.2	68.3
	March	1.1	-9.2	-4.0	32.3	43.7	76.6
	April	9.4	-1.2	4.1	66.6	15.2	80.5
	May	17.2	5.1	11.2	101.8	1.4	103.2
	June	22.2	10.6	16.4	124.6	0	124.6
	July	24.3	13.5	18.9	127.9	0	127.9
	August	23.3	12.3	17.7	123.3	0	123.3
	September	18.6	7.9	13.3	107.2	0	107.2
	October	11.2	1.9	6.6	108.2	3.2	111.1
	November	4.2	-3.9	0.2	73.3	25.5	100.7
	December	-3.1	-12	-7.4	35.4	65.1	101.1
	Annual	9.8	-0.6	4.6	944.1	268	1214
Yamaska River watershed	January	-5.4	-15.8	-10.6	28.6	56	86
	February	-3.2	-14.2	-8.7	23.6	46.9	70.9
	March	2.3	-8.2	-2.9	35	41.9	76.6
	April	11	0.1	5.5	79.5	14.4	93.5
	May	18.4	6.1	12.2	105.7	0.5	106.2
	June	23.4	11.4	17.4	112.2	0	112.2
	July	25.5	13.8	19.6	116.2	0	116.2
	August	24.5	12.6	18.5	123.9	0	123.9
	September	19.9	8.3	14.1	103.5	0	103.5
	October	12.5	2.5	7.5	110.8	3	113.8
	November	5.4	-2.7	1.3	86.6	19.9	106.5
	December	-1.9	-10.6	-6.2	37.5	61.7	98
	Annual	11	0.3	5.6	962.8	244.2	1207

Appendix C

Monthly climate statistics (a) for the period the 1980-2010 and (b) for the 10 climate scenarios for the period 2040–2070 in the Becancour River watershed.

Period	Scenario	Climate parameter	Mean monthly values												Mean annual values
			J	F	M	A	M	J	J	A	S	O	N	D	
1980–2010 (Current Period)	Current	P (mm)	14.3	12.0	28.5	70.4	94.3	115.8	125.1	119.4	105.9	99.8	61.8	22.4	869.6
		Snow (cm)	104.3	75.9	54.8	13.7	0.9	0	0	0	0.2	3.9	35.9	98.6	388.3
		T (°C)	-12.8	-10.8	-4.7	3.2	10.8	15.7	18.3	16.8	11.8	5.6	-1.2	-9.2	3.6
2040–2070 (Climate scenarios)	acu	P (mm)	41.2	27.8	70.0	88.5	112.2	119.0	99.4	127.3	107.2	100.8	86.8	50.0	1030.2
		Snow (cm)	71.8	40.6	28.0	1.5	0.0	0.0	0.0	0.0	0.0	0.7	16.7	64.5	223.8
		T (°C)	-7.8	-6.7	-0.2	7.4	14.3	19.2	22.4	20.8	15.2	9.1	2.5	-4.5	7.7
	adc	P (mm)	51.7	29.7	74.2	99.8	104.2	101.2	122.1	138.4	109.2	114.0	75.1	67.5	1087.1
		Snow (cm)	75.1	56.5	18.2	2.2	0.0	0.0	0.0	0.0	0.0	0.5	13.3	51.1	216.9
		T (°C)	-6.8	-5.9	-1.0	6.7	14.8	19.2	21.8	20.7	15.9	9.9	2.5	-3.8	7.8
	adj	P (mm)	39.6	27.8	52.9	82.3	103.7	107.7	113.2	124.5	118.5	94.8	88.4	55.1	1008.4
		Snow (cm)	79.7	41.1	25.4	2.1	0.0	0.0	0.0	0.0	0.0	0.4	15.7	73.3	237.7
		T (°C)	-8.3	-7.0	-0.7	7.2	14.4	19.5	22.8	21.3	15.6	9.0	2.7	-4.9	7.6
	adl	P (mm)	62.5	36.4	74.5	85.3	116.7	103.6	113.3	112.9	126.2	107.4	76.6	76.7	1092.1
		Snow (cm)	69.4	48.3	23.5	3.3	0.0	0.0	0.0	0.0	0.0	0.6	15.0	44.4	204.6
		T (°C)	-6.5	-5.9	-1.2	6.7	14.7	19.0	22.4	21.1	15.8	9.7	2.5	-3.4	7.9
	aet	P (mm)	42.8	25.6	56.6	85.4	97.6	109.8	110.1	125.5	114.8	93.8	100.7	67.9	1030.5
		Snow (cm)	69.6	41.0	20.1	2.4	0.0	0.0	0.0	0.0	0.0	0.4	11.7	62.1	207.3
		T (°C)	-8.1	-7.1	-0.6	7.1	14.6	19.4	22.8	21.2	15.4	9.2	3.3	-4.2	7.8
	aev	P (mm)	56.3	34.7	76.5	98.9	114.4	97.9	105.1	127.6	121.3	111.5	89.0	79.1	1112.3
		Snow (cm)	71.6	56.0	20.0	2.4	0.0	0.0	0.0	0.0	0.0	1.0	13.4	42.2	206.6
		T (°C)	-6.3	-5.8	-0.9	6.7	14.5	19.1	22.3	21.1	16.0	9.8	3.0	-3.1	8.0
	afx	P (mm)	42.1	28.4	67.7	87.1	109.9	120.8	107.2	135.0	104.9	102.3	91.4	57.3	1054.0
		Snow (cm)	72.7	39.5	27.3	1.6	0.0	0.0	0.0	0.0	0.0	0.6	14.3	58.5	214.4
		T (°C)	-7.8	-6.8	-0.3	7.7	14.1	19.1	22.5	20.9	15.2	9.2	2.8	-4.1	7.7
	agr	P (mm)	53.5	30.7	72.2	99.6	97.9	99.0	127.9	146.0	109.3	111.8	78.1	76.0	1102.1
		Snow (cm)	72.6	56.6	18.6	1.9	0.0	0.0	0.0	0.0	0.0	0.4	10.5	44.6	205.2
		T (°C)	-6.7	-5.9	-1.1	6.9	14.5	19.2	21.9	20.7	15.9	9.9	2.9	-3.4	7.9
	agx	P (mm)	42.6	28.6	59.3	82.1	124.5	100.8	106.8	118.6	123.8	102.2	100.1	59.6	1049.0
		Snow (cm)	84.0	56.0	19.6	2.9	0.0	0.0	0.0	0.0	0.0	1.0	10.9	57.8	232.3
		T (°C)	-8.1	-8.6	-1.9	5.6	13.9	18.7	21.8	20.7	15.4	9.3	3.3	-4.8	7.1
	aha	P (mm)	39.3	25.9	56.3	82.4	106.8	93.9	111.3	152.1	135.9	102.9	78.0	63.7	1048.5
		Snow (cm)	89.9	57.8	26.4	2.7	0.2	0.0	0.0	0.0	0.0	1.1	11.7	66.7	256.5
		T (°C)	-8.3	-8.1	2.0	5.8	14.0	18.7	21.6	20.5	15.2	9.2	3.0	-4.9	7.0

Appendix D

Monthly climate statistics (a) for the period, 1980–2010 Monthly climate statistics (a) for the period 1980–2010 and (b) for the 10 climate scenarios for the period 2040–2070 in the Yamaska River watershed

Period	Scenario	Climate parameter	Mean monthly values												Mean annual values
			J	F	M	A	M	J	J	A	S	O	N	D	
1980–2010 (Current Period)	Current	P (mm)	22.1	19.3	37.2	79.1	93.1	101.2	110.1	116.1	102.7	90.8	76.1	29.4	877.1
		Snow (cm)	87.8	60.2	39.3	6.6	0.2	0	0	0	0.1	1.5	23.3	69.7	288.5
		T (°C)	-10.7	-8.7	-2.9	5.4	12.5	17.3	19.8	18.6	13.8	7.4	0.9	-6.8	5.5
2040–2070 (Climate scenarios)	acu	P (mm)	49.0	30.6	81.7	91.1	100.0	101.6	99.5	122.9	92.0	97.6	95.6	61.7	1023.4
		Snow (cm)	56.1	31.6	18.0	0.2	0.0	0.0	0.0	0.0	0.0	0.3	10.8	49.4	166.5
		T (°C)	-7.0	-5.9	0.4	8.5	14.8	19.7	23.2	21.6	16.3	10.2	3.2	-4.2	8.4
	adc	P (mm)	55.3	38.1	78.1	111.3	98.3	91.1	112.4	118.6	99.8	109.5	83.3	75.3	1071.2
		Snow (cm)	55.7	42.0	9.8	1.2	0.0	0.0	0.0	0.0	0.0	0.5	9.1	37.8	156.1
		T (°C)	-6.1	-5.4	-0.1	7.8	15.6	20.3	22.9	21.6	17.2	10.6	3.2	-2.7	8.7
	adj	P (mm)	46.1	30.7	57.8	89.4	93.2	99.9	112.5	104.0	107.0	87.5	93.6	62.7	984.3
		Snow (cm)	58.2	31.9	14.9	1.0	0.0	0.0	0.0	0.0	0.0	0.2	10.7	58.4	175.2
		T (°C)	-7.5	-6.3	0.2	8.5	14.8	19.9	23.6	22.0	16.6	10.1	3.5	-4.6	8.4
	adi	P (mm)	68.6	42.1	84.5	99.9	108.1	101.1	98.4	94.0	108.6	96.5	83.2	82.8	1067.9
		Snow (cm)	48.8	36.3	13.8	1.9	0.0	0.0	0.0	0.0	0.0	0.3	8.9	32.0	142.0
		T (°C)	-5.8	-5.3	-0.3	7.5	15.7	20.0	23.5	22.0	17.1	10.6	3.5	-2.2	8.9
	aet	P (mm)	52.0	29.4	61.8	93.6	88.4	89.4	101.4	115.5	103.2	89.5	105.3	78.2	1007.7
		Snow (cm)	50.6	29.7	12.1	1.0	0.0	0.0	0.0	0.0	0.0	0.1	7.2	42.9	143.7
		T (°C)	-7.4	-6.4	0.4	8.7	14.9	19.9	23.7	22.0	16.6	10.2	4.0	-3.4	8.6
	aev	P (mm)	63.9	44.3	90.1	108.5	99.3	90.1	93.1	105.6	109.7	101.4	93.4	85.3	1084.7
		Snow (cm)	51.3	39.7	12.2	0.8	0.0	0.0	0.0	0.0	0.0	0.4	8.6	29.6	142.6
		T (°C)	-5.6	-5.2	-0.1	7.6	15.5	20.1	23.4	21.9	17.2	10.5	3.8	-1.9	8.9
	afx	P (mm)	50.3	31.5	79.9	94.0	97.2	100.8	104.7	123.7	92.6	96.8	97.6	68.6	1037.7
		Snow (cm)	56.0	31.0	17.4	0.2	0.0	0.0	0.0	0.0	0.0	0.3	9.4	42.3	156.7
		T (°C)	-7.0	-6.0	0.2	9.1	14.5	19.7	23.2	21.6	16.3	10.2	3.5	-3.5	8.5
	agr	P (mm)	57.1	40.0	77.7	110.8	89.5	93.1	110.8	121.5	99.1	106.5	86.2	82.0	1074.3
		Snow (cm)	53.0	40.5	9.9	0.7	0.0	0.0	0.0	0.0	0.0	0.3	7.7	31.1	143.2
		T (°C)	-5.9	-5.4	-0.2	8.4	15.2	20.3	22.9	21.6	17.2	10.7	3.6	-2.1	8.9
	agx	P (mm)	56.1	34.3	62.3	90.0	112.0	93.6	107.0	112.0	100.3	99.3	103.8	66.5	1037.2
		Snow (cm)	60.8	42.7	12.6	1.5	0.0	0.0	0.0	0.0	0.0	0.3	7.2	42.4	167.5
		T (°C)	-7.0	-8.1	-0.6	6.6	14.8	19.7	22.6	21.4	16.1	9.9	4.0	-3.9	8.0
	aha	P (mm)	44.6	31.4	64.9	84.1	100.2	87.3	111.1	138.0	117.1	99.0	86.2	69.7	1033.5
		Snow (cm)	69.9	47.5	17.0	0.9	0.0	0.0	0.0	0.0	0.0	0.5	6.5	47.7	190.1
		T (°C)	-7.5	-7.8	-0.9	6.7	14.8	19.7	22.3	21.2	15.9	9.9	3.8	-4.1	7.8

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