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3 **Comparison of a deterministic and statistical approach for the**
4 **prediction of thermal indices in regulated and unregulated river**
5 **reaches: case study of the Fourchue River (Québec, Canada).**
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34

35 **Abstract**

36 Water temperature is an important factor modifying fish distribution patterns and community
37 abundance in streams and this is especially true for salmonids. Knowing that dams often modify
38 the thermal regime of rivers, understanding these changes is of crucial importance for fish
39 habitat management. This study aims to improve knowledge about the impact of dams on the
40 thermal regime of rivers during the summer season and to assess the relative efficiency of two
41 modelling tools used to predict water temperatures downstream of dams. A deterministic model
42 (SNTEMP) and a statistical model based on a canonical correlation analysis were calibrated on
43 the Fourchue River (St-Alexandre-de-Kamouraska, Québec, Canada) upstream and
44 downstream of a reservoir. SNTEMP was used to simulate mean water temperatures time series
45 using meteorological inputs and discharge. The statistical model was used to directly estimate
46 thermal indices (descriptive statistics of the thermal regime). The two models were compared
47 based on their efficiency to estimate thermal indices such as mean and maximum monthly water
48 temperatures and other parameters of importance in the understanding of the distribution and
49 growth of ichthyofauna. Water temperature was monitored at 18 locations in the Fourchue River
50 during the summers of 2011 and 12 locations in 2012 to describe the thermal regime and
51 calibrate the models. The statistical model achieved better results than SNTEMP in estimating
52 most of the thermal indices, especially the mean and maximum daily ranges with RMSEs of
53 4.1 °C and 4.9 °C respectively for SNTEMP as compared to 0.5 °C and 1.1 °C for the leave-one-
54 out validation and 0.6 °C and 1.4 °C for the split-sample mode for the statistical model. The
55 better performance of the statistical model for metrics related to thermally stressful events for
56 fish make it more appealing as a management tool for water resources and fisheries managers.
57 However, SNTEMP should be considered when the objective is to investigate the impact of
58 climate change, reservoir operations or other anthropogenic impacts.

59

60 **KEYWORDS:** modeling, temperature, river, SNTEMP, multivariate, geostatistics

61 **Introduction**

62 The thermal regime of rivers is of interest for fisheries management because most of the
63 physical, chemical and biological properties of fish habitat are temperature-dependent
64 (Magnuson et al., 1979; Cassie, 2006). Because fish are ectotherms, they are highly dependent

65 on water temperature to maintain important physiological and life history processes (Becker and
66 Genoway, 1979; Wood and McDonald, 1997; Beitinger *et al.*, 2000). Their suitable thermal
67 habitats are constrained by both maximum and minimum thermal tolerances (Mohseni *et al.*,
68 2003). Laboratory studies have been conducted for decades to define optimum temperatures for
69 maximum fish growth (e.g. Jobling, 1981). For instance, the optimal growth temperature of brook
70 trout (*Salvelinus fontinalis*) is 14.2 °C and the mortality rates increase when temperature
71 exceeds 24.9°C, which is the upper limit of their thermal tolerance (Hasnain *et al.*, 2010). This
72 was further ascertained by Hasnain *et al.* (2013) who reviewed thermal metrics for numerous
73 fish species, including salmonids, in North America.

74 Anthropogenic regulation of rivers also alters thermal conditions. The effects of dams on the
75 thermal regime of rivers have been widely investigated and include changes in the temperature
76 mean and variance at several temporal scales (Petts 1984; Preece and Jones 2002; Steel and
77 Lange 2007; Olden and Naiman, 2010, Maheu *et al.*, 2016). Thermal regimes downstream of
78 impoundments depend on the dam operating mode and the depth of water intake. A significant
79 number of large dams release cold hypolimnetic water establishing highly desirable habitats for
80 trout and salmon. On the other hand, smaller dams and diversions can increase water
81 temperature by releasing warm water directly from the reservoir surface (Maheu *et al.*, 2016).
82 These dam-induced modifications to the thermal conditions can have both direct and indirect
83 consequences on fish by altering the quality of their habitat or their prey's habitat (Ward, 1985;
84 Angilletta *et al.*, 2008; Olden and Naiman, 2010).

85 On regulated rivers, adequate fisheries management can be achieved by mitigating the thermal
86 stressful events via cold water releases below dams. One way to assess the impact of stream
87 regulation on a river is to compute thermal indices at an impacted site and to compare them with
88 those calculated from similar unregulated control rivers or river reaches. These indices are
89 descriptive statistics of hourly or daily mean temperatures that characterize the thermal regime
90 in terms of amplitude (mean and extremes), variability, duration and timing of events (cold or
91 warm spells). Examples of amplitude indices include the monthly means of
92 the maximum daily temperature (Arismendi *et al.*, 2013). Some jurisdictions use thermal indices
93 to manage fisheries. For instance, on the Miramichi River (Canada) angling for Atlantic salmon
94 (*Salmo salar*) is not allowed when maximum daily summer temperature exceeds 23 °C and
95 minimum temperature is greater than 20 °C (Caissie, Thistle, and Benyahya 2017). In western
96 Canada and north-western U.S., the highest average of maximum daily temperatures over any
97 7-day period (maximum weekly maximum temperature, MWMT) and the highest average of

98 mean daily temperatures over any 7-day period (maximum weekly average temperature MWAT)
99 are used as thermal metrics for fisheries management (Welsh et al. 2001).

100 Unfortunately, temperature gauging stations that could be used to calculate these thermal
101 indices are relatively scarce in Canada. To overcome the lack of data, many different simulation
102 tools are used to characterize the thermal conditions in rivers. These tools can be classified in
103 two main categories: deterministic models (Theurer *et al.*, 1984; St-Hilaire *et al.*, 2003; Cassie *et*
104 *al.*, 2007 Ouellet *et al.*, 2013) and empirical or statistical models (Bélanger *et al.*, 2005;
105 Benyahya *et al.*, 2007; Chenard and Caissie, 2008; Guillemette *et al.*, 2009). Deterministic
106 models typically calculate a heat budget at one or many points in the river using meteorological
107 inputs and information on stream geomorphology and hydraulics. However, these variables are
108 not always readily available and the gathering of these data can be a long and expensive
109 process. Statistical approaches can be an interesting alternative because they generally require
110 fewer input variables. These latter models are based on statistical relationships between water
111 temperature and correlated independent variables such as air temperature (Benyahya *et al.*,
112 2007). While most statistical models use only meteorological inputs (mostly air temperature),
113 some approaches allow for the inclusion of physiographic information. One such model was
114 adapted to water temperature modelling by Guillemette *et al.* (2009). It combines multivariate
115 methods and geostatistics. The main perceived advantage, compared to traditional deterministic
116 models, is that the simulation of temperature time series can be bypassed and thermal indices
117 can be modelled directly. This can be an attractive alternative for managers who may prefer a
118 more direct, less cumbersome approach than deterministic modelling. However, the
119 performance of this alternative needs to be equivalent to that of the more classic models. In the
120 context of impounded rivers, the performance of the two models can be compared both
121 upstream and downstream of dams, as reservoirs are often an important impediment to thermal
122 connectivity.

123 There are very few studies that compare statistical and deterministic river temperature models
124 using the same data sets. Massé and Armengol (2008) used a deterministic model and
125 compared it to a hybrid approach (deterministic hydrological model combined with a linear
126 regression between air and water temperature) on Mediterranean streams. They concluded that
127 including empirical or hybrid formulations that use air temperature as a predictor is not optimal
128 (compared to a deterministic model) when local meteorological data are available and should
129 only be preferred when meteorological stations are far from the river reaches under study. Our

130 study may be the first Canadian comparison between the two types of models on an impounded
131 river.

132 The present study therefore aims to evaluate the efficiency of the multivariate geostatistical
133 model used by Guillemette *et al.* (2009) by comparing it to a well-established deterministic model
134 called Stream Network Temperature (SNTEMP) (Theurer *et al.*, 1984). The comparison is
135 performed on two river reaches, upstream and downstream of a dam reservoir.

136 The statistical model is based on the identification of appropriate physiographical variables as
137 predictors of water temperature indices at the stream segment scale. Thermal indices are
138 obtained by interpolation in an orthogonal space constructed using a multivariate approach
139 called canonical correlation analysis (Chokmani and Ouarda, 2004). The interpolation is made
140 by using a multiple linear regressions in canonical space.

141 SNTEMP is a mechanistic, one dimensional heat transport model used to simulate daily mean
142 and maximum water temperatures. SNTEMP was selected in this study because of its extensive
143 use for regulated and unregulated rivers (Horne *et al.*, 2004; Norton and Bradford, 2009; Voss *et al.*
144 *et al.*, 2008; Shepard *et al.*, 2009).

145 The general objective of this study is to compare the two different modeling approaches in order
146 to determine which one is the most suitable for water resources managers in estimating selected
147 thermal indices.

148 **Methodology**

149 **Study site and data collection**

150 The Fourchue River is a regulated river with a drainage basin of 261 km² and a tributary of the
151 Du Loup River, located in eastern Quebec, Canada (Figure 1). The Morin dam was built to
152 regulate flows in the Du Loup River. The reservoir occupies an area of 6.8 km² at top water level
153 and has a storage capacity of 38 880 000 m³. The water level into the reservoir is kept between
154 188 m and 195 m above sea level during summer. In order to maintain these levels, the flows
155 evacuated are usually kept between 0.06 m³/s and 4 m³/s. Details on the dam, spillway and
156 draw-offs, together with a description of the operation mode, are provided by the Centre
157 d'Expertise Hydrique du Québec (CEHQ, 2008).

158 Water temperature time series were obtained for summer 2011 (July to September) and 2012
159 (June to September) with Hobo Pro V2 thermographs (± 0.2 °C) recording water temperature at
160 15 minutes intervals at approximately 15 cm from the stream bed. The loggers were deployed
161 into two reaches of the Fourchue River considered relatively similar in topography, land use and
162 climate. One reach is located directly downstream of the Morin dam and the other, which served
163 as a control reach, is located 10 km upstream of the reservoir, in the unregulated portion of the
164 river. A total of 18 loggers were deployed in 2011, seven upstream of the reservoir in a 9 km
165 reach and eleven downstream in a 5 km reach. For 2012, the downstream reach was extended
166 to include the only major tributary of the Fourchue River, the Carrier stream, for a total of 12
167 loggers deployed over 8 km. Low water levels in the upstream reach in 2012 resulted in many
168 thermographs being exposed to air and thus the 2012 upstream data could not be used.
169 Hydrological and stream geometry data were also obtained from field measurements as well as
170 the meteorological conditions for the study area.

171

172 **Meteorological inputs**

173 To calculate the energy budget equations, SNTemp requires the following meteorological
174 inputs: air temperature, relative humidity, wind speed, solar radiation and cloud cover. Daily air
175 temperature (± 0.1 °C) and relative humidity (± 0.8 %) were measured with a Rotronic HygroClip2
176 relative humidity and temperature probe (HC2-S3-L). Wind speed was measured with a RM
177 Young wind monitor (05103-10, ± 0.3 m/s) and solar radiation data were measured with a Kipp
178 and Zonen pyranometer (SP-LITE-L, ± 10 $\mu\text{V W}^{-1} \text{m}^2$). The meteorological data were averaged
179 hourly at a station located 100m north-east of the reservoir.

180 The solar radiation was used to estimate the percent possible sun (a surrogate for cloud cover)
181 using a cloud cover correction algorithm from Reifsnnyder and Lull (1965):

$$182 \quad \frac{E_c}{E_m} = 10^{-0.99C_{okt}} \quad (2)$$

183 Where:

184 E_c = Irradiance under cloudy condition

185 E_m = Irradiance under clear sky condition

186 C_{okt} = Cloud oktas

187

188 **Hydrology**

189 Rating curves were developed for the two reaches and the tributary to establish the relationship
190 between discharge and water level. The daily water levels were obtained with Hobo U20 water
191 level data loggers. Several spot measurements of discharges were taken between 1.2 and 3.8
192 m^3s^{-1} in the downstream reach, 0.1 and 2.5 m^3s^{-1} in the upstream reach and between 0 and 0.5
193 m^3s^{-1} in the tributary. The discharge data were collected using the velocity-area method with a
194 Marsh McBirney Flo-Mate 2000 flow velocimeter.

195 **Stream geometry**

196 The sites elevations were obtained with a Novalynx barometer altimeter (230-M202) with 3 m
197 accuracy. It was calibrated using the elevation of the CEHQ hydrometric station located 100 m
198 downstream of the dam.

199 A pebble count was performed to characterize the composition of the streambed. In every
200 stream segment, 100 particles were measured in the normal low flow channel. The cumulative
201 frequency curve generated from pebble counts led to the estimation of the median particle
202 diameter (D_{50}). Manning's roughness coefficient, n , was calculated from the following equation
203 (Robert, 2003):

$$204 \quad n = 0.048 D_{50}^{1/6} \quad (3)$$

205 In order to account for the riparian shade, an SNTMP component estimates an attenuation
206 factor using information on the streamside vegetation and the topography, on the average tree
207 height, the crown diameter, and the distance from the water's edge. These variables were
208 estimated from field observations. The topographic horizon angles on both sides of the river
209 were measured with a clinometer. These angles are used by the model to calculate the local
210 times of sunrise and sunset. Stream widths as a function of flow were also obtained from field
211 measurements.

212 **Thermal indices**

213 Thermal indices are used to describe the magnitude, variability, frequency and duration of
214 thermal events across space and time (Arismendi *et al.*, 2013). The thermal indices calculated
215 from the water temperature time series are monthly means and maxima of daily temperatures,
216 the mean and maximum daily ranges, cumulative degree-days, the monthly standard deviation
217 and the number of days over 24.9 °C, which is the upper incipient lethal temperature (UILT) for

218 brook trout, one of the fish species found throughout the study area (Hasnain *et al.*, 2010). Mean
219 temperatures were first selected as one of the amplitude metrics that represent the thermal
220 “climate” of a river. Daily ranges and standard deviation are important because it has been
221 shown that adequate range and variability that include low temperature at nights can allow fish
222 to recuperate from (high) stressful temperature events (e.g. Brodeur *et al.*, 2015). Temperature
223 maxima exhibited by streams during summer can affect fish species limited by low survival
224 threshold temperatures. The UILT is defined as the upper boundary to the “zone of thermal
225 tolerance” within which there is no mortality from temperature (Fry *et al.*, 1946). A metric like the
226 UILT can be used to identify affected species. The indices were first used to compare and
227 contrast the thermal regimes in the unregulated and regulated reaches. The models were also
228 compared on their ability to predict these thermal indices.

229

230 **Deterministic approach**

231 The Stream Network Temperature Model (SNTEMP) was created by Theurer *et al.* (1984).
232 SNTEMP is a steady state, one-dimensional heat-transport model used to predict daily mean
233 and maximum water temperatures. The model is composed of six components, starting with the
234 heat flux model that predicts the energy balance between the water and its environment. It is
235 defined as the arithmetic sum of the solar, atmospheric and vegetative radiations, evaporation
236 loss, heat conduction and convection, conduction and water back radiation. To predict the
237 average mean daily and diurnal water temperatures as a function of stream distance, the heat
238 transport component uses a dynamic temperature, steady flow equation. The solar component
239 predicts the amount of solar radiation penetrating the stream water as a function of the time of
240 year by calculating the radiation amount reaching the earth. The latitude is used to determine
241 the day length and the meteorological conditions are used to estimate the attenuation of the
242 radiation due to its travel through the atmosphere. Because the solar radiation reaching the
243 stream can be reduced by the local environment and the riparian vegetation, the shade
244 component estimates the attenuation using information on the streamside vegetation and the
245 topography. Finally, to consider the adiabatic process, the meteorological component corrects
246 for variations in elevation within the watershed that cause changes in atmospheric pressure, air
247 temperature and relative humidity.

248 The first step of the SNTEMP modeling process is to represent the river as homogeneous
249 segments with similar attributes like flows, width and streamside vegetation. The study area was
250 partitioned into segments based on field observations, for a total of seven segments upstream

251 and nine downstream in 2011 and twelve downstream in 2012. These homogeneous segments
 252 are called nodes. There are 14 different nodes available in the model to represent the network
 253 (presence of a tributary, structure, etc.). The use of these nodes will depend on the size of the
 254 study reach, the complexity of the system and the data availability. In the case of the Fourchue
 255 River, six nodes were required to represent the study area (figure 2). The description of the node
 256 types are presented in Table 1.

257

258 **Table 1: Description of the node types used for the composition of the network of Fourchue River in SNTEMP**

Node type	Abbreviation	Description
Source	H	The upstream boundary usually located at a gage or a zero flow headwater.
Structure	S	A point (reservoir) that may have discontinuity in discharge and will have a released temperature defined by the user.
Change	C	The upstream end of a reach with new stream shading or hydraulic properties
Validation	V	Node where the temperature is known and can be compared to predicted temperature
Point load	P	Node where a point load discharges into the river at a known temperature
End	E	The network end point (most downstream point)

259

260

261 **Model calibration and validation**

262 The deterministic model was calibrated in the downstream reach using a split-sample approach.
 263 The first two weeks of June and August 2012 were used as calibration periods in order to
 264 include the whole water temperature range in the calibration set. The calibration consists in
 265 adjusting the model parameters for a better representation of the river's environment (Table 2).
 266 For instance, the air temperature above the stream is usually lower than the temperature
 267 measured at the meteorological station. A correction factor of -0.5 °C was applied. Similarly, the
 268 relative humidity values were corrected and increased by 10 % over recorded values to account
 269 for humidity above the river. Finally, because wind speed was measured in an open area while
 270 wind above the water surface is impacted by canopy, the wind speed was reduced by 15% to

271 represent the wind speed conditions in the sheltered river channel (Nieto et al., 2019). This
272 percentage was determined by trial and error.

273

274 **Table 2: SNTMP's global calibration factors and the corrections applied for a better representation of the**
275 **Fourchue River conditions.**

SNTMP global calibration factors	Corrections applied
Air temperature calibration constant	↓0.5 °C
Air temperature calibration coefficient	-
Wind speed calibration constant	-
Wind speed calibration coefficient	↓15%
Humidity calibration constant	-
Humidity calibration coefficient	↑10%
Sunshine calibration constant	-
Sunshine calibration coefficient	-
Solar calibration constant	-
Solar calibration coefficient	-

276

277 The model temperature estimations were compared to the continuous temperature
278 measurements into two segments, referred to as verification nodes, in the upstream reach, and
279 to three segments in the downstream reach. The model was validated in the downstream reach
280 over July 2012. Finally, the thermal indices were calculated using the mean and maximum daily
281 water temperatures simulated by SNTMP. The performance of the model was assessed by
282 considering two specific performance evaluation criteria: the BIAS and the root mean square
283 error (RMSE) (See Laanaya et al. 2017 for detailed equations). Given that thermographs
284 precision is of the order of 0.5 °C, a RMSE value of the order of 1 °C can be considered as a low
285 error for a water temperature model. Bias should, of course, also be minimized, especially as it
286 relates to high temperatures.

287

288 SNTMP does not have the ability to model temperatures within impoundments so the sections
289 upstream and downstream of the reservoir were modelled separately for August 2011.

290 **Statistical approach**

291 The statistical model is based on an interpolation technique that estimates the thermal indices in
292 a mathematical multivariate space rather than a geographical space, as proposed by

293 Guillemette *et al.* (2009). The approach relies on the construction of an orthogonal space
 294 defined by the canonical correlation analysis (CCA) of the physiographical and water
 295 temperature characteristics of the stream segments. CCA is a multivariate approach that
 296 produces linear combinations of two sets of observations in order to maximize the associations
 297 (measured by the correlations) between the two data sets, while ensuring orthogonality of the
 298 canonical variates within the same group. Here, those two data sets are the matrix X of the
 299 thermal indices and the matrix Y of the predictors, which are the physiographic variables
 300 representing the environment of the river. In this case, only four metrics, strongly correlated with
 301 water temperature, were necessary to characterize the stream segment; these were the distance
 302 from the dam (positive downstream and negative upstream), the elevation, the Stralher order
 303 and the vegetation density. CCA produces the orthogonal linear combinations U of variables in
 304 matrix X , known as canonical variates that maximally correlate with the linear combinations V of
 305 variables in matrix Y . The coefficient vectors a and b are respectively associated with the
 306 thermal indices (X) and the physiographical variables (Y):

307 $a) U = aX$

308 $b) V = bY$ (4)

309 Pairs of vectors (U_i, V_i) are identified as the i^{th} canonical variate pair. There are p possible
 310 canonical covariate pairs, where p is the smallest vector length of X or Y . The vectors are found
 311 by a joint covariance analysis of the variables (Härdle and Simar, 2003). This allows to
 312 maximize the canonical correlation between (U_i, V_i), calculated as:

313 $\rho_i = \frac{cov(U_i, V_i)}{\sqrt{var(U_i)var(V_i)}}$ (5)

314

315 A multiple linear regression (MLR) was performed in the orthogonal plane composed of the first
 316 two dimensions of the canonical variates V , which constitute the axes of the physiographic
 317 space. For a given water temperature index, values at monitoring stations were projected in the
 318 V space and interpolation at ungauged sites was achieved by fitting a linear equation that best
 319 approximate all individual data points in the least square sense. It was also possible to find the V
 320 coordinates of an ungauged site by using equation 6(b). Figure 3 summarizes the main steps of
 321 the statistical model.

322 In order to assess the performance of the statistical approach, two validation techniques were
323 used: a cross validation using a leave-one-out resampling (jackknife) and a split-sample
324 validation. In the jackknife, the value of a station is temporarily removed from the data set and
325 this value is estimated using the remaining stations. This operation is repeated for the whole
326 station set. The estimated values are then compared with the observed data. For the split-
327 sample validation, almost all stations were removed from the observed sample to serve as a
328 validation group except for the stations at the most upstream and downstream points of the two
329 stream reaches. These remaining four stations in 2011 and three stations in 2012 were used as
330 calibration group. The BIAS (equation 4) and the RMSE (equation 5) were calculated for the two
331 validation techniques (Chokmani and Ouarda, 2004). The performance of SNTMP and the
332 statistical model were compared on the basis of the two aforementioned evaluation criteria
333 (BIAS and RMSE).

334 **Results and discussion**

335 The total rain amount in the region exceeded the normal in August 2011 (106.6 mm as
336 compared to the monthly mean of 89.1 mm), resulting in a water level 2.4 m over the monthly
337 mean recorded at the CEHQ hydrometric station. On the opposite, rainfall was below normal in
338 August 2012, with only 53.2 mm of total precipitations. The mean air temperature was 2.1°C
339 above the normal conditions. This resulted in low water levels and warmer water temperatures
340 as compared to 2011. Because it captures a fair range of the possible summer hydroclimatic
341 conditions, the results of the modeling approaches will be presented for these two months.

342 The canonical space was defined for every thermal index. Figure 4 shows an example of a
343 canonical space for August 2011 mean temperature. There is a clear separation between the
344 upstream and downstream sections and the two stations located downstream of the tributary.
345 The interpolation was performed within that space.

346

347 **Thermal indices based on mean temperature for August 2011 and 2012.**

348 Both models showed very similar good performance for the estimation of the thermal indices
349 based on monthly mean water temperature (Figure 5). The performance measures indicate that
350 SNTMP is slightly more accurate for the prediction of the mean monthly (August) water
351 temperature, with a RMSE of 0.2 °C compared to 0.4 °C and 0.3 °C for the leave-one-out and

352 split-sample validation of the statistical model, respectively. BIAS was much smaller than sensor
353 precision (< 0.01 °C) for these thermal indices. No estimation was performed with SNTMP for
354 stations 17 and 18 due to the lack of flow data from the tributary of the Fourchue River, the
355 Carrier River, located in that reach, just upstream of these two stations. To evaluate thermal
356 mixing below tributaries, SNTMP requires daily discharge and temperature from the tributary,
357 which were not available for 2011. The statistical model does not use discharge as a metric so it
358 was possible to estimate temperature at these stations. The accurate estimations of these
359 downstream stations are explained by the fact that the longitudinal variability of the monthly
360 means is well represented by the Strahler order, which is a component of canonical variate V1.

361
362 The same observations can be made for the cumulative degree-days, an important metric for the
363 evaluation of the growth rate for fish (Neuheimer and Taggart, 2007). The obtained RMSE are
364 5.0 °C-days, 11.5 °C-days and 9.4 °C-days for SNTMP, the leave-one-out and the split-sample
365 validations of the statistical model, respectively. The RMSEs are considered relatively low for the
366 two approaches because the observed cumulative degree-days vary between 540 and 625 °C-
367 days. There was no significant BIAS in the estimation of this thermal index with either of the two
368 approaches. Hence, SNTMP outperformed the statistical model for this metric.

369 The monthly standard deviation was estimated with more accuracy by the statistical model with
370 a RMSE of 0.2 °C and no BIAS for both leave-one-out and split-sample, as compared to a
371 RMSE of 1.0 °C and a BIAS of 0.5 °C for SNTMP.

372
373 In August 2012, the main tributary of the Fourchue River, located 3 km downstream of the dam,
374 was included in SNTMP with the point source model configuration. This means that the water
375 temperature was not simulated in the tributary but the discharge and water temperature of the
376 tributary was included in the modeling of the main river. Both models predicted mean daily water
377 temperature with a RMSE of 0.1 °C and no significant BIAS (Figure 6). In contrast with 2011, the
378 cumulative degree-day was simulated with more accuracy with the statistical model than
379 SNTMP in 2012 (RMSE of 0.6 °C-days (jackknife validation) and 2.9 °C-days, respectively).
380 However, the statistical model could not produce good estimations given only three calibration
381 stations: RMSE associated with the split-sample validation using three calibration stations is
382 22.8 °C-days. RMSE could be lowered to 0.7 °C-days with eight out of thirteen calibration
383 stations uniformly distributed over the downstream reach.

384

385

386

387 **Thermal indices based on maximum temperature for August 2011 and 2012**

388 The statistical model surpassed SNTMP in the estimation of the thermal indices based on
389 maximum temperature (figures 7 and 8).

390

391 In 2011, the Fourchue River has not experienced temperatures exceeding the zone of thermal
392 tolerance of the brook trout, which was well predicted by the statistical model. SNTMP
393 predicted 3 days over 24.9 °C, leading to a RMSE of 1.2 days. With the warmer conditions
394 experienced in 2012, one to eleven days over the UILT was recorded in the river. The jackknife
395 and split-sample RMSEs were less than 1 day and BIASes under 0.4 day, while SNTMP gave
396 8.7 days RMSE and a BIAS of 4.8 days. The UILT can hardly be used adequately by river
397 managers using this deterministic model as it would always over estimate the number of days
398 where fish experiment temperature over their zone of thermal tolerance.

399 The calculations of the daily maxima in SNTMP are based on an empirical model. Theurer *et*
400 *al.* (1984) elaborated a method to estimate average afternoon air temperature, the main
401 component for the estimation of maximum daily water temperature. Regression coefficients were
402 determined for normal meteorological conditions, based on the arithmetic mean of historical data
403 at 16 selected weather stations around the United States, which is not representative for the
404 current study site. SNTMP does not explicitly model minimum temperatures, which are
405 estimated using the daily mean and maximum temperatures.

406 SNTMP overestimated maximum daily water temperatures, especially downstream of the dam.
407 This is due to the fact that the model extends the current reach stream geometry indefinitely
408 upstream in order to simulate the conditions through which the water must travel from solar noon
409 (considered as the mean daily water temperature) to solar sunset (considered as the maximum
410 daily water temperature) and thus, does not include the reservoir in its simulation. The water
411 released in the downstream reach from the shallow reservoir is warmer compared to the
412 upstream reach. Information about the reservoir is not considered in SNTMP when it calculates

413 maxima based on the extension of the current reach stream geometry. In 2011, SNTEMP
414 resulted in a RMSE of 2.5 °C and a BIAS of 0.1 °C. However, if it is calculated separately, the
415 RMSE for the upstream reach (1.4 °C) is lower than the RMSE for the downstream reach (3.1
416 °C). Information about the location of the dam is included in the statistical model in the metric
417 “distance from the dam”, allowing the model to estimate maximum water temperature with more
418 accuracy (leave-one-out and split-sample RMSEs of 0.7 °C and 0.8 °C, respectively) and no
419 BIAS.

420 The lack of information on the dam reservoir prevented accurate estimations of the conditions
421 through which water travels from solar noon to solar sunset, which explains the differences
422 between the models for the estimations of the thermal indices based on maximum temperatures.

423 Water temperatures show diurnal variations depending on the heat energy gained and lost by a
424 stream and the volume and source of runoff contributing to discharge (Ward, 1985; Webb,
425 1996). The presence of the dam reduces the range between temperature extremes at the
426 stations located downstream (Ward and Stanford, 1979). This reduction in daily variability is
427 represented by the metric “distance from the dam” in the statistical model, which resulted in a
428 better estimation of the mean and maximum daily ranges. The overestimation of maximum
429 temperature by SNTEMP led to an overestimation of the mean and maximum daily ranges in
430 2011 (RMSEs of 4.1 °C and 4.9 °C and BIAS of 4.7 °C and 2.4 °C for the mean maximum
431 ranges, respectively). The statistical model estimated the mean and maximum daily ranges with
432 RMSEs equal to 0.5 °C and 1.1 °C in leave-one-out mode and of 0.6 °C and 1.4 °C for the split-
433 sample mode. The BIAS of the statistical validation methods was of -0.2 °C for both indices for
434 the leave-one-out and 0.2 °C for the split-sample. Similar observations were made with the
435 simulation of the mean and maximum daily ranges in 2012.

436

437 **3.7 Discussion and Conclusion**

438 The objective of this study was to compare the relative efficiency of a deterministic and a
439 statistical model in the estimation of selected thermal indices, in order to determine which one is
440 the most suitable for the river managers. SNTEMP showed good results for the estimation of
441 monthly mean temperatures and cumulative degree-days but overall, the statistical model was
442 more efficient for the estimation of most selected thermal indices.

443 SNTEMP is limited by the fact that it does not model temperatures within impoundments, nor
444 does it explicitly model minimum temperature. These limitations impacted the performance of the
445 deterministic model in the estimation of the selected thermal indices, leading to inaccurate
446 estimations of three out of seven thermal indices. The multivariate geostatistical model showed
447 good results for the seven thermal indices for both regulated and unregulated reaches. This
448 model however requires water temperatures time series for each stream segment, while
449 SNTEMP requires mean daily temperature only at the verification nodes and for the upstream
450 and downstream headwater segments. This represents six gauging stations in 2011 and four in
451 2012. The split-sample validation technique aimed to reduce the number of gauging stations
452 required for the statistical model with minimum accuracy loss. It turned out that four water
453 temperature measurement stations in 2011 and three in 2012 were sufficient to simulate the
454 selected thermal indices adequately.

455 Although many studies have compared different statistical models (e.g. Laanaya et al., 2017),
456 very few have compared statistical vs deterministic approaches. Marceau et al., (1986)
457 compared a Box-Jenkins statistical approach to the CEQUEAU deterministic model. They
458 concluded that both had similar performances. SNTEMP, which has been used extensively in
459 other studies, has seldom been compared to other models, with the exception of Norton and
460 Bradford (2009). They compared SNTEMP to CE-QUAL-W2 and concluded that both had similar
461 performances, but that the latter showed more consistent performance across space and time.
462 Our results corroborate past studies indicating some equivalence in performances of both
463 methods for simulating the mean temperature regime. However, our results also indicate a
464 superior performance of the statistical approach for temperature extremes. Of course, model
465 selection is always dependent on river management needs. For the management of brook trout,
466 thermal indices related to high temperature and daily variability are the most important. Those
467 metrics are better estimated by the statistical approach. The lower input requirements for the
468 statistical approach and its relative good performance for indices that may be indicative of
469 thermal stress for fish (e.g. number of days above a high temperature threshold) make this
470 approach very attractive for manager. However, since the statistical model does not use explicit
471 hydraulic or climatic inputs, it is not possible to evaluate different scenarios related to climate
472 change and dam operations with this model in its current form. These kinds of scenarios could
473 however be simulated with SNTEMP. The input data requirements are lower for the statistical
474 model, resulting in lower implementation cost and less field work.

475 It can thus be seen that both models offer different advantages and should perhaps be used in
476 conjunction in future studies. Therefore, if the management objective is to forecast temperature
477 extremes in a drainage basin with little anthropogenic perturbations, the CCA-MLR model is
478 adequate. However, if anthropogenic impacts are present or anticipated, SNTEMP should be the
479 preferred choice for water resource managers.

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487

488 **REFERENCES**

489 Angilletta MJ Jr., Steel EA, Bartz AA, Kingsolver JG, Scheuerell MD, Beckman BR, Crozier LG.
490 2008. Big dams and salmon evolution: changes in thermal regimes and their potential
491 evolutionary consequences. *Evolutionary Applications* **1**: 286-299

492
493 Arismendi I, Johnson SL, Dunham JB, Haggerty R. 2013. Descriptors of natural thermal regimes
494 in streams and their responsiveness to change in the Pacific Northwest of North America.
495 *Freshwater Biology* **58**: 880-894

496
497 Bartholow, JM. 1995. The stream network temperature model (SNTEMP): A decade of results.
498 Workshop on Computer Application in Water Management. Fort Collins, CO: Water Resources
499 Research Institute, CSU. p. 57-60.

500
501 Bartholow, JM. 2002. Stream Segment Temperature Model (SSTEMP) Version 2.0. Revised
502 August 2002. Fort Collins, CO: U.S. Geological Survey

503
504 Becker, C.D., Genoway, R.G. Evaluation of the critical thermal maximum for determining thermal
505 tolerance of freshwater fish. *Environ Biol Fish* **4**, 245 (1979). doi.org/10.1007/BF00005481

506
507 Beitinger TL, Bennett WA, McCauley RW. 2000. Temperature tolerances of North American
508 freshwater fishes exposed to dynamic changes in temperature. *Environmental Biology of Fishes*
509 **58**: 237-275

510

511 Bélanger M, El-Jabi N, Caissie D, Ashkar F, Ribí JM. 2005. Water temperature prediction using
512 neural networks and multiple linear regression. *Revue des Sciences de l'Eau* **18**: 403-421
513

514 Benyahya L, St-Hilaire A, Ouarda Tbmj, Bobée B, Ahmadi-Nedushan B. 2007. Modeling of water
515 temperatures based on stochastic approaches: case study of the Deschutes River. *Journal of*
516 *Environmental Engineering and Science* **6**: 437-448
517

518 Breau C, Cunjak RA, Peake SJ. 2011. Behaviour during elevated water temperatures: can
519 physiology explain movement of juvenile Atlantic salmon to cool water? *Journal of Animal*
520 *Ecology* **80**: 844-853
521

522 Brodeur, N. N., C. Hébert, D. Caissie, C. Breau. 2015. Predicting stream temperatures under a
523 climate change scenario: impacts on critical temperatures for Atlantic salmon (*Salmo salar*). *Can.*
524 *Tech. Rep. Fish. Aquat. Sci.* 3118: ix + 44p.
525

526 Caissie D. 2006. The thermal regime of rivers: A review. *Journal of Freshwater Biology* **51**:
527 1389-1406
528

529 Caissie D, Satish MG, El-Jabi N. 2007. Predicting water temperatures using a deterministic
530 model: Application on Miramichi River catchments (New Brunswick, Canada). *Journal of*
531 *Hydrology* **336**: 303-315
532

533 Caissie D, El-Jabi N, St-Hilaire A. 1998. Stochastic modelling of water temperatures in a small
534 stream using air to water relations. *Canadian Journal of Civil Engineering* **25**: 250-260
535

536 Centre d'Expertise Hydrique du Québec, 2008. Aménagement Morin X0000730 et X0000731.
537 Plan de Gestion des eaux retenues. Québec.
538

539 Chenard JF, Caissie D. 2008. Stream temperature modelling using artificial neural networks:
540 application on Catamaran Brook, New-Brunswick, Canada. *Hydrological Processes* **22**: 3361-
541 3372
542

543 Chen YD, McCutcheon SC, Norton DJ, Nutter WL. 1998. Stream temperature simulation of
544 forested Riparian areas: II. Model application. *Journal of Environmental Engineering* **124**: 316-
545 328
546

547 Chokmani K, Ouarda TBMJ. 2004. Physiographical space-based kriging for regional flood
548 frequency estimation at ungauged sites. *Water Resources Research* **40**: 1-13
549

550 Coutant CC. 1977. Compilation of temperature preference data. *Journal of the Fisheries*
551 *Research Board of Canada*, **34**: 739-745
552

553 Fry FE, Hart J, Walker JS. 1946. Lethal temperature relations for a sample of young speckled
554 trout, *Salvelinus fontinalis*. Publ. Ontario Fisheries Research Lab. Univ. Toronto Biol. Ser. **66**
555 (54): 35 pp.
556

557 Guillemette N, St-Hilaire A, Ouarda TBMJ, Bergeron N, Robichaud E, Bilodeau L. 2009.
558 Feasibility study of a geostatistical modelling of monthly maximum stream temperatures in a
559 multivariate space. *Journal of Hydrology* **364**: 1-12
560

561 Härdle W, Simar L. 2003. Applied multivariate statistical analysis. Springer: Berlin.

562
563 Hasnain SS, Mins CK, Shuter BJ. 2010. Key Ecological Temperature Metrics for Canadian
564 Freshwater Fishes, Applied Research and Development Branch. Ontario Ministry of Natural
565 Resources.
566
567 Hasnain, Sarah & Shuter, Brian & Minns, Charles. (2013). Phylogeny influences the
568 relationships linking key ecological thermal metrics for North American freshwater fish species.
569 *Canadian Journal of Fisheries and Aquatic Sciences*. 70. 10.1139/cjfas-2012-0217.
570
571 Hrachowitz M, Soulsby C, Imholt M, Malcolm A, Tetzlaff D. 2010. Thermal regimes in a large
572 upland salmon river: a simple model to identify the influence of landscape controls and climate
573 change on maximum temperatures. *Hydrological Processes* **24**: 3374-3391
574
575 Jobling, M. 1981. Temperature tolerance and the final preferendum—rapid methods for the
576 assessment of optimum growth temperature. *Journal of Fish Biology* 19(4): 439-455.
577
578 Laanaya F., A. St-Hilaire, E. Gloaguen. 2017. Modeling the water temperature: comparison of
579 the generalized additive model, logistic and residuals regression models. *Hydrological Sciences*
580 *Journal* 62(7): 1078-1093. DOI: dx.doi.org/10.1080/0262667.2016.1246799.
581
582 Maheu, A.*, A. St-Hilaire, D. Caissie, N. El-Jabi. 2016. Understanding the thermal regime of
583 rivers influenced by small and medium size dams in eastern Canada. *River Research and*
584 *Applications* 32 (10): 2032–2044 8 May. DOI: 10.1002/rra.3046
585
586 Magnuson, J.J., L.B. Crowder, P.A. Medvick. 1979. Temperature as an Ecological
587 Resource. *American Zoologist* 19(1): 331–343.
588
589 Maheu, A., A. St-Hilaire, D. Caissie, N. El-Jabi. 2016. Understanding the thermal regime of rivers
590 influenced by small and medium size dams in eastern Canada. Publié en ligne dans *River*
591 *Research and Applications* 32 (10): 2032–2044 8 May. DOI: 10.1002/rra.3046
592
593 Marcé, R., J. Armengol. 2008. Modelling river water temperature using deterministic,
594 empirical, and hybrid formulations in a Mediterranean stream. *Hydrological Processes* 22:3418-
595 3430.
596 Mohseni O, Stefan HG, Eaton JG. 2003. Global warming and potential changes in fish habitat in
597 U.S. streams. *Climatic Changes* **59**: 389-409
598
599 Morin G, Nzakimuena TJ, Sochanski W. 1994. Predicting river water temperature using a
600 conceptual model: the case of the Moisie River. *Canadian Journal of Civil Engineering* **21**:63-75
601
602 Neuheimer AB, Taggart CT. 2007. The growing degree-day and fish size-at-age: the overlooked
603 metric. *Canadian Journal of Fisheries and Aquatic Sciences* **64** (2): 375-385
604
605 Nieto, H., Kustas, W.P., Alfieri, J.G. *et al.* Impact of different within-canopy wind attenuation
606 formulations on modelling sensible heat flux using TSEB. *Irrigation Science*. **37**, 315–331
607 (2019). <https://doi.org/10.1007/s00271-018-0611-y>.
608

609 Norton GE, Bradford A. 2009. Comparison of two stream temperature models and evaluation of
610 potential management alternatives for the Speed River, Southern Ontario. *Journal of*
611 *Environmental Management* **90**: 866-878
612

613 Olden JD, Naiman RJ. 2010. Incorporating thermal regimes into environmental flows
614 assessments: modifying dam operations to restore freshwater ecosystem integrity. *Freshwater*
615 *Biology* **55**: 86-107
616

617 Ouellet V, Secretan Y, St-Hilaire A, Morin J. 2013. Daily averaged 2D water temperature model
618 for the St-Lawrence River. *River Research and Applications*. DOI: 10.1002/rra.2664
619

620 Petts GE. 1984. Impounded Rivers: Perspectives for Ecological Management. Wiley and Sons:
621 New York.
622

623 Power ME. 1990. Effects of fish in river food webs. *Science* **250**: 811-814
624

625 Preece RM, Jones HA. 2002. The effect of Keepit Dam on the temperature regime of the Namoi
626 River, Australia. *River Research and Applications* **18**: 397-414
627

628 Reifsnnyder WE, Lull HW. 1965. Radiant energy in relation to forests. U.S. Department of
629 Agriculture, Forest service. Technical bulletin No. 1344. U.S. Government Printing Office:
630 Washington D.C.
631

632 Robert A. 2003. An introduction to fluvial dynamics. Arnold, Hodder Headline Group: London.
633

634 Shepard D, Taylor R, Knudson K, Hunter C. 2009. Calibration of a Water Temperature Model for
635 Predicting Summer Water Temperatures in Rush Creek below Grant Lake Reservoir. Los
636 Angeles Department of Water Power, California.
637

638 Steel EA, Lange IA. 2007. Using wavelet analysis to detect changes in water temperature
639 regimes at multiples scales: effects of multi-purpose dams in the Willamette River basin. *River*
640 *Research and Applications* **23**: 351-359
641

642 St-Hilaire A, Ouarda TBMJ, Lachance M, Bobée B, Gaudet J, Gignac C. 2003. Assessment of
643 the impact of meteorological network density on the estimation of basin precipitation and runoff:
644 a case study. *Hydrological Processes* **17**: 3561-3580
645

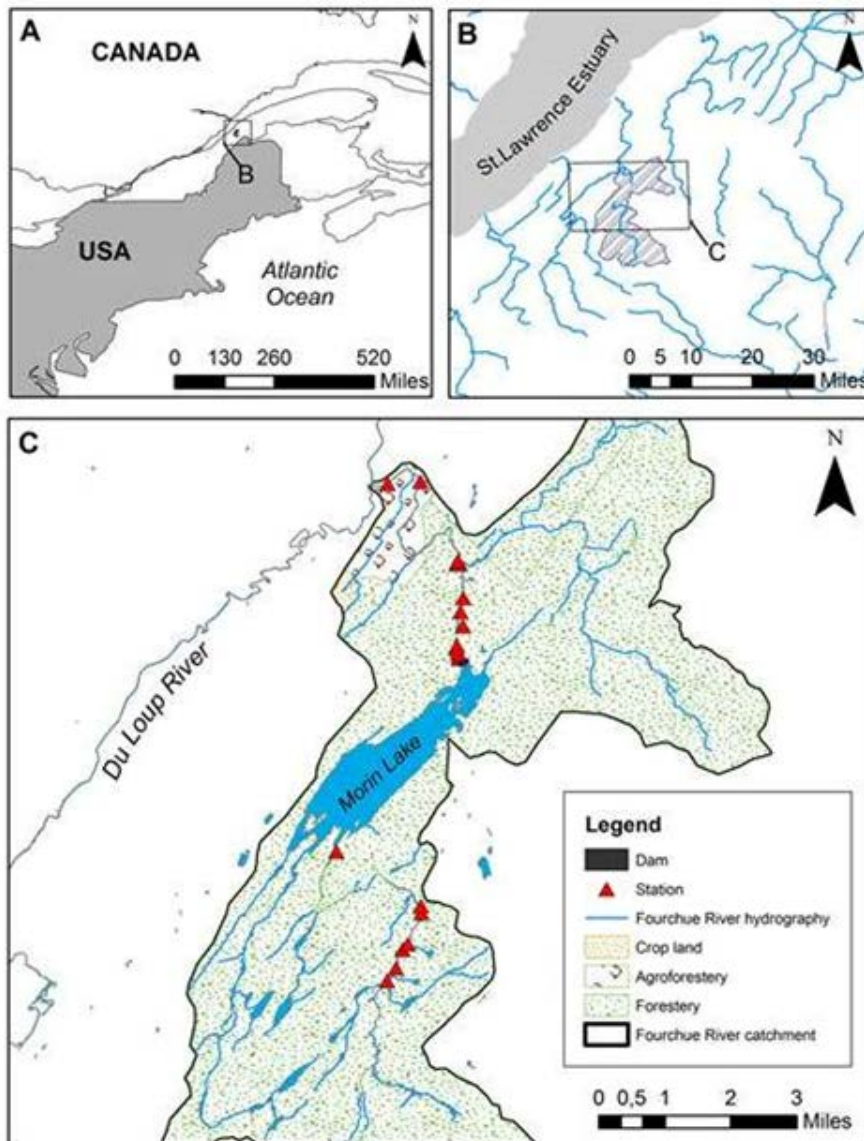
646 Stanford JA, Ward JV, Liss WJ, Frissell CA, Williams RN, Lichatowich JA, Coutant CC. 1996. A
647 general protocol for restoration of regulated rivers. *Regulated Rivers: Research & Management*
648 **12**: 391-413
649

650 Theurer F, Voos D, Kenneth A, Miller WJ. 1984. Instream Water Temperature Model. Instream
651 Flow Inf. Pap. 16 Coop. Instream Flow and Aquatic System Group, U.S. Fish & Wildlife Service.
652 Fort Collins: Colorado.
653

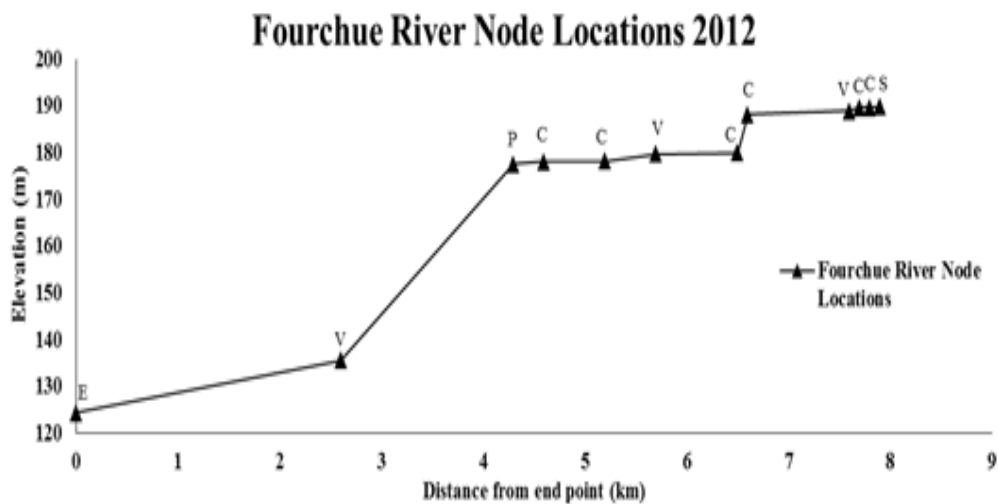
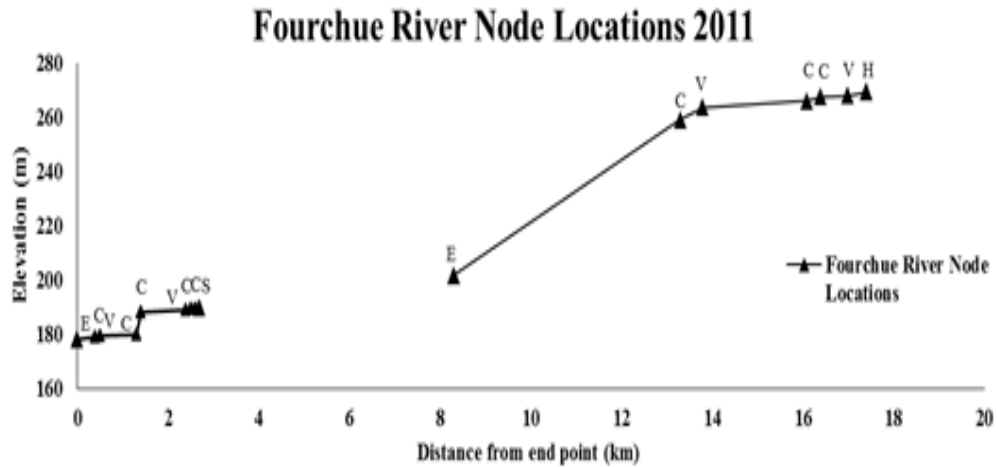
654 Voss FD, Curran CA, Mastin MC. 2008. Modeling water temperature in the Yakima River,
655 Washington, from Roza diversion dam to Prosser dam, 2005-06. U.S. Fish & Wildlife Service.
656 Fort Collins: Colorado.
657

658 Ward JV. 1985. Thermal characteristics of running waters. *Hydrobiologia* **125**: 31-46
659

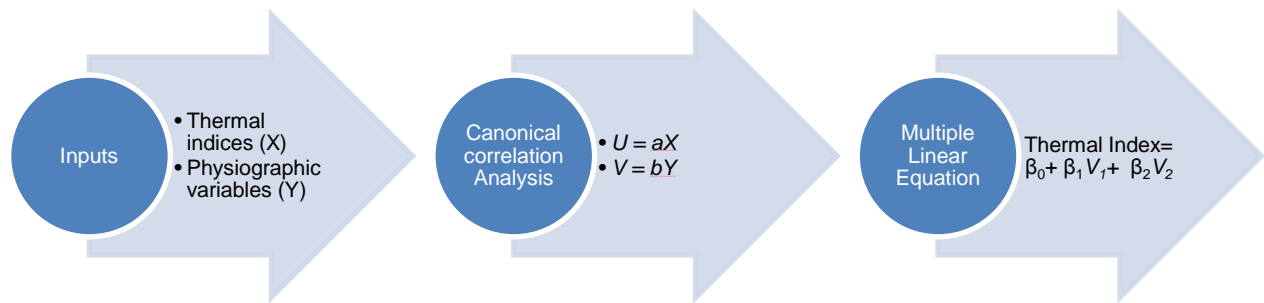
660 Ward JV, Stanford JA. 1979. Ecological factors controlling stream zoobenthos with emphasis on
661 thermal modification of regulated streams. *The Ecology of Regulated Streams*. pp. 35-55
662 Plenum Press, New York.
663
664 Webb BW. 1996. Trends in stream and river temperature. *Hydrological Processes*, **10** (2): 205–
665 226
666
667 Wood CM, McDonald DG. 1997. Global warming: implications for freshwater and marine fish.
668 Cambridge University Press: Boston.
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 678 **Figure 1. Location of the Fourchue River watershed and the water temperature monitoring**
 679 **stations.**
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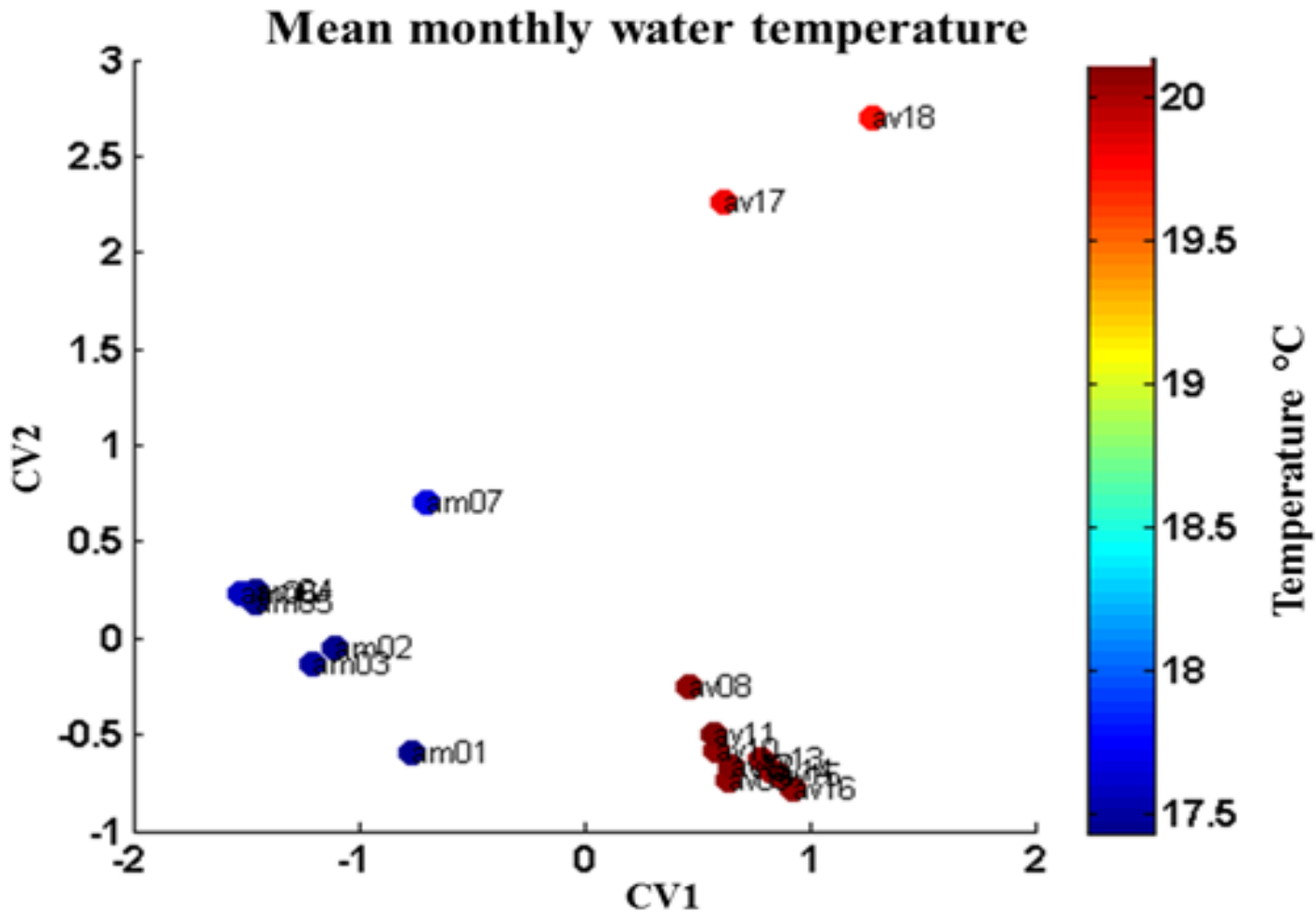


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 688 **Figure 2. Longitudinal profile of the Fourchue River illustrating the composite node**
 689 **network along the relative river gradient. Points along the diagram depict the node types**
 690 **including headwater (H), change (C), validation (V), structure (S), point load (P) and end**
 691 **(E).**
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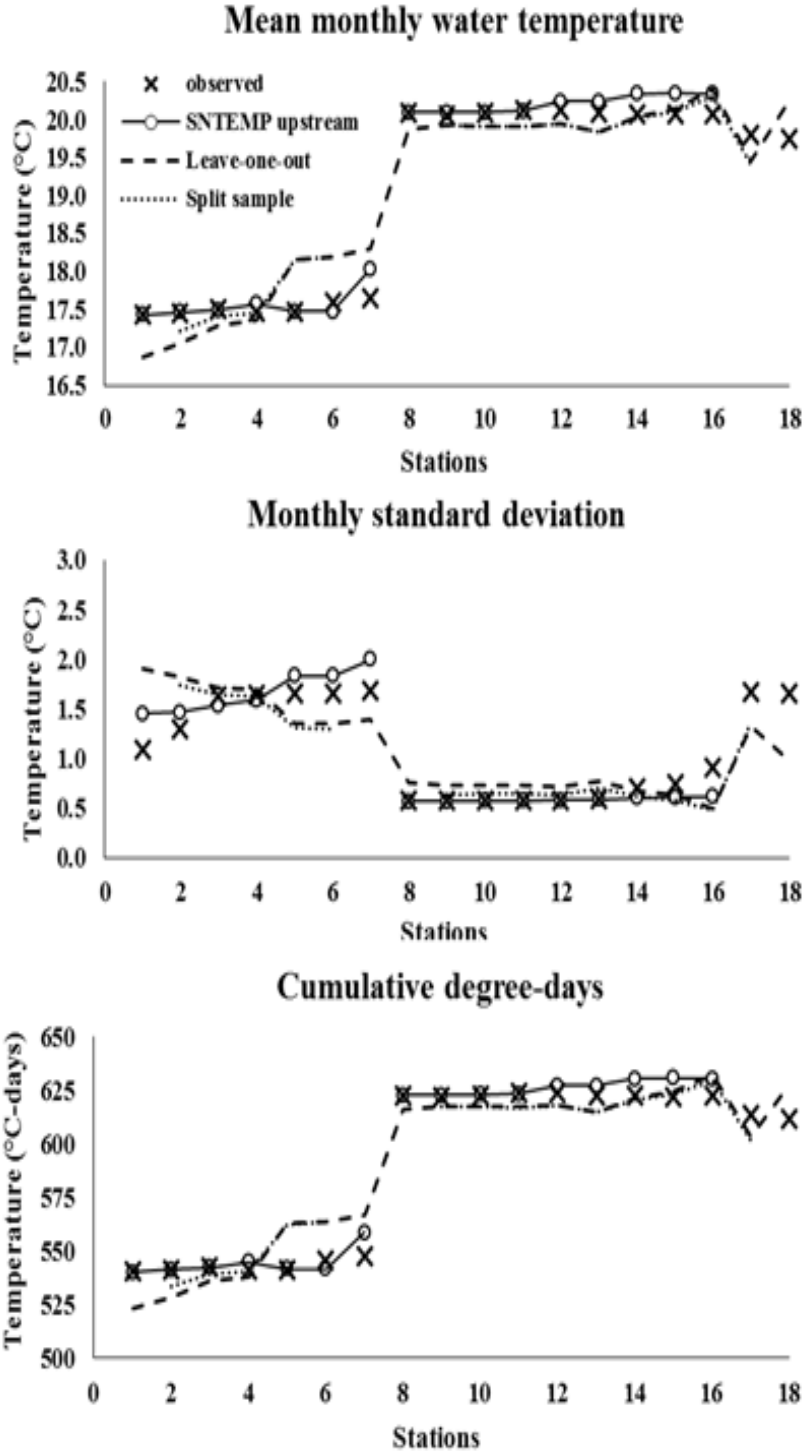
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Figure 3. Main methodological steps of the statistical model.



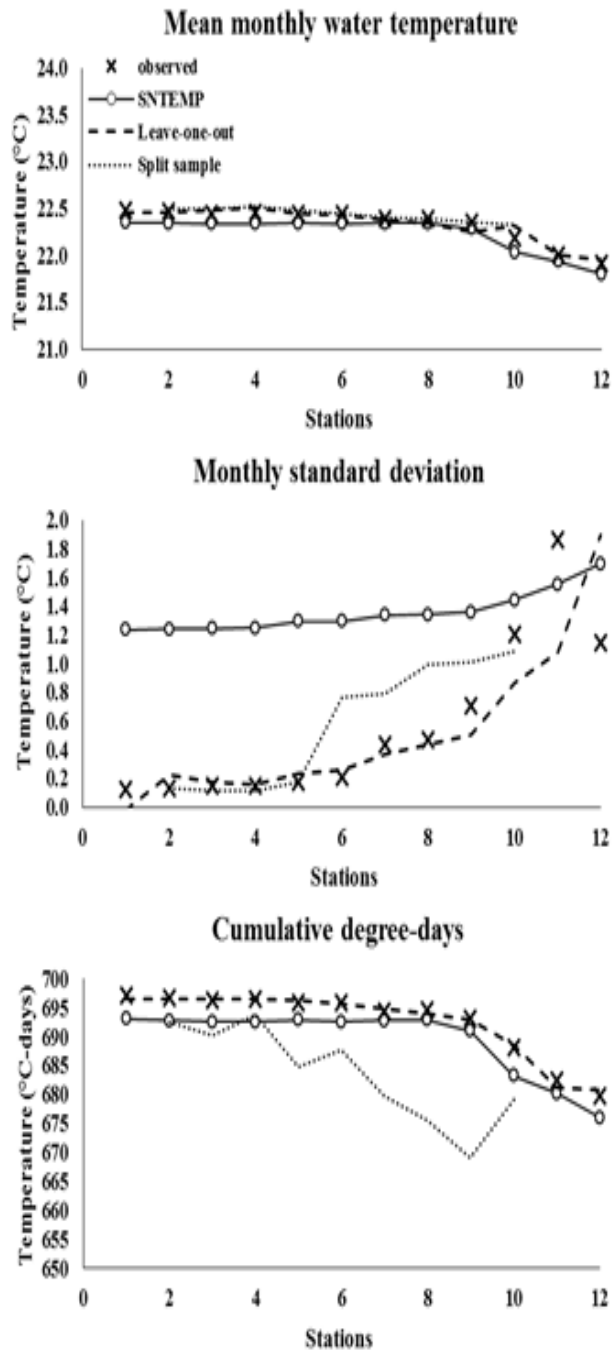
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Figure 4. Canonical space for August 2011 mean temperatures. The upstream stations are referred as am01 to am07 and the downstream stations are referred as av08 to av18.

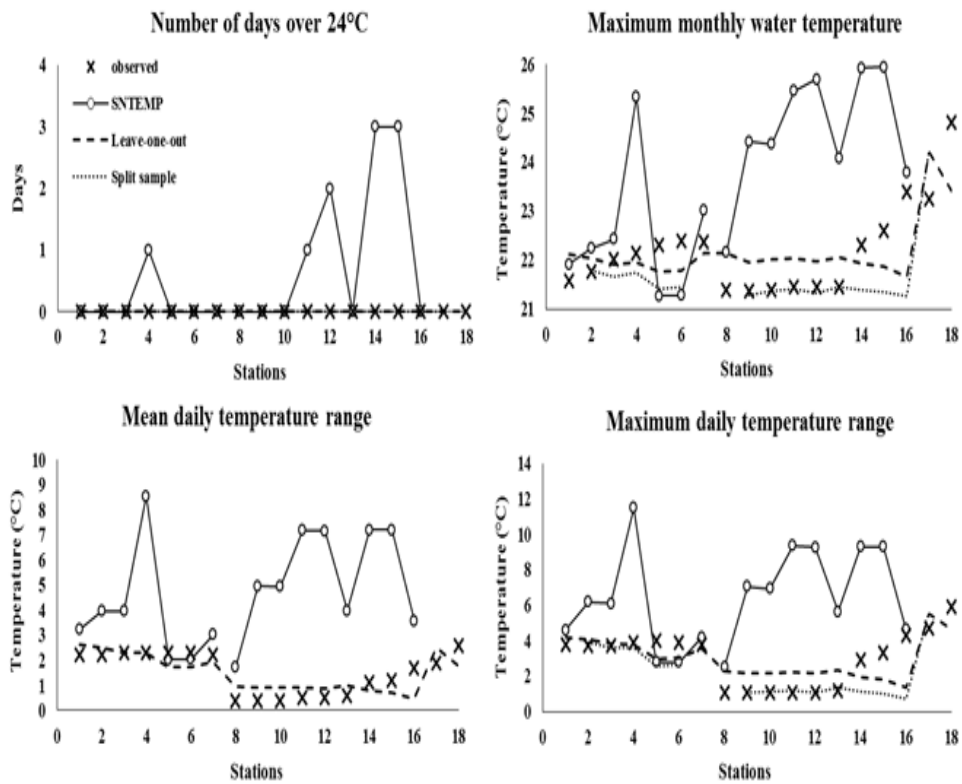


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Figure 5. Observed and simulated mean monthly water temperatures, standard deviation and cumulative degree-days for August 2011, using SNTEMP and the statistical model in leave-one-out and split sample modes. Stations 1 to 18 are from upstream to downstream.

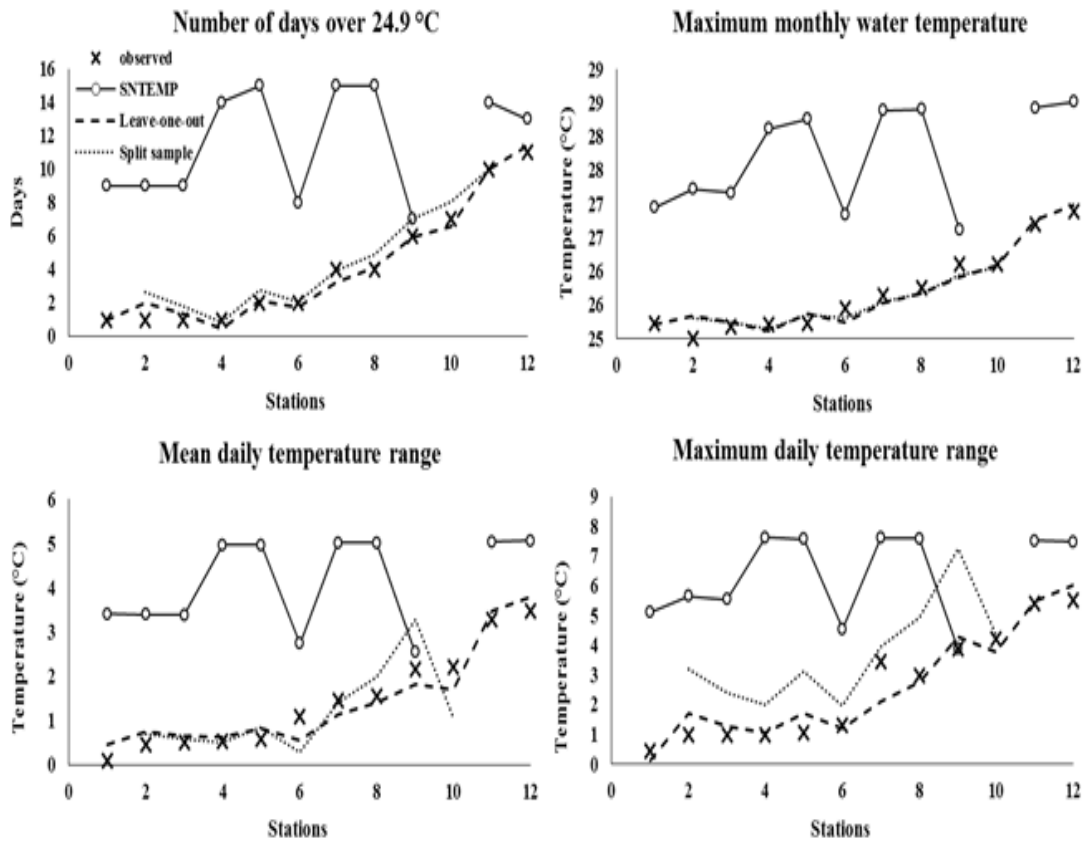


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 725 **Figure 6. Observed and simulated mean monthly water temperatures, standard deviation**
 726 **and cumulative degree-days for August 2012, using SNTemp and the statistical model in**
 727 **leave-one-out and split sample modes. Stations 1 to 12 are from upstream to downstream.**



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Figure 7. Observed and simulated mean monthly maximum temperatures, mean and maximum daily temperature ranges and the number of days over 24.9 °C for August 2011, using SNTemp and the statistical model in leave-one-out and split-sample modes. Stations 1 to 18 are from upstream to downstream.



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Figure 8. Observed and simulated mean monthly maximum temperatures, mean and maximum daily temperature ranges and the number of days over 24.9 °C for August 2012, using SNTemp and the statistical model in leave-one-out and split-sample modes. Stations 1 to 12 are from upstream to downstream.