

# Stationary and Non-stationary Temperature-Duration-Frequency Curves for Australia

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## Research Article

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1 **Stationary and Non-stationary Temperature-Duration-Frequency Curves**  
2 **for Australia**

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22        **ABSTRACT**

23        Australian summer heat events have become more frequent and severe in recent times. Temperature-  
24        duration-frequency (TDF) curves connect the severity of heat episodes of various durations to their  
25        frequencies and thus can be an effective tool for analysing the heat extremes. This study examines  
26        Australian heat events using data from 82 meteorological stations. TDF curves have been developed  
27        under stationary and non-stationary conditions. Generalised Extreme Value (GEV) distribution is  
28        considered to estimate extreme temperatures for return periods of 2, 5, 10, 25, 50 and 100 years. Three  
29        major climate drivers for Australia have been considered as potential covariates along with *Time* to  
30        develop the non-stationary TDF curves. According to the Akaike information criterion, the non-  
31        stationary framework for TDF modelling provides a better fit to the data than its stationary equivalent.  
32        The findings can be beneficial in offering new information to aid climate adaptation and mitigation at  
33        the regional level in Australia.

34        **Keywords:** Annual maximum temperature; Climate change; Climate drivers; Generalised Extreme  
35        Value; Non-stationary; Temperature-Duration-Frequency.

36

## 37 1. INTRODUCTION

38 Intergovernmental Panel on Climate Change (IPCC) estimated that the mean global temperature would  
39 exceed the higher limit of increase in temperature adopted in the Paris agreement by the 2040s (IPCC,  
40 2018). It was also reported in 2020 that the temperature in Australia has increased by  $1.44 \pm 0.24$  °C  
41 since 1910 (CSIRO and Australian Government (Bureau of Meteorology), 2020). Generally, changes  
42 in maximum and minimum temperatures impact the environment more than mean temperature (Sein,  
43 Chidthaisong and Oo, 2018) and adversely influence the events related to extreme temperature  
44 (Chowdary, John and Gnanaseelan, 2014). Extreme temperature triggers more natural hazards, i.e.  
45 droughts, bushfires, heatwaves, cyclones, and negatively affects human health, agriculture, ecosystems  
46 and infrastructure (Omer *et al.*, 2020; Suman and Maity, 2020). Therefore, predicting extreme  
47 temperatures for a particular duration and return period is important in many aspects of our lives. Some  
48 are relevant to health, energy and agriculture, urban planning, and ecology management. In this context,  
49 the design temperatures for the selected durations and return periods can be estimated following the  
50 same principle of rainfall Intensity-Duration-Frequency (IDF) curves, which have been utilised for  
51 many years for the design and management of hydrological infrastructure as well as flood risk  
52 management (Yan *et al.*, 2019). Extreme value theory has been accepted globally to construct IDF  
53 curves by fitting theoretical probability distribution functions to annual maximum rainfall time series  
54 (Cheng and Aghakouchak, 2014; Yan *et al.*, 2020).

55 Most of the works found in the literature on IDF curves are based on the temporal stationary concept  
56 (Singh and Zhang, 2007; Jakob, 2013), implying that the occurrence probability of precipitation events  
57 will not change significantly over time. However, in reality, the frequency, magnitude and duration of  
58 hydroclimatic extremes, i.e. extreme rainfall (Galiatsatou and Iliadis, 2022), floods (Berghuijs *et al.*,  
59 2019), droughts (Spinoni, Naumann and Vogt, 2017), and heatwaves (Lorenz, Stalhandske and Fischer,  
60 2019) fluctuates beyond the stationary envelope of variability. Furthermore, it is also observed that the  
61 IDF curve considering stationarity, often underestimates the return level of rainfall (Sugahara, da Rocha  
62 and Silveira, 2009; Cheng and Aghakouchak, 2014). Therefore, to increase the reliability of IDF curves,  
63 it is suggested to incorporate the non-stationarity of the distribution parameters in hydrological models  
64 (Sarhadi and Soulis, 2017), especially at their extreme levels (Ganguli and Coulibaly, 2017).

65 In the construction of IDF curves, covariates are generally introduced to apprehend the non-  
66 stationarities in time series. Any drivers that are correlated to the selected events can be considered as  
67 a candidate for such covariates (Katz, Parlange and Naveau, 2002; Hundecha *et al.*, 2008), which can  
68 be low-frequency climatological signals (Coles, 2001) as well as time (Sugahara, da Rocha and Silveira,  
69 2009; Yilmaz and Perera, 2014). In this approach, the parameters of the Generalised Extreme Value

70 (GEV) distribution are expressed as a linear or non-linear function of covariates (Kwon and Lall, 2016;  
71 Sarhadi and Soulis, 2017). This approach has been effectively employed in extreme rainfall events  
72 around the world (Ouarda, Yousef and Charron, 2019), including in Australia (Yilmaz and Perera,  
73 2014).

74 Although there have been extensive works carried out on the development of IDF curves all over the  
75 globe, very limited works are found in the literature on the construction of Temperature-Duration-  
76 Frequency (TDF) curves, which relate temperature intensities corresponding to varying durations and  
77 return periods (Wang *et al.*, 2013; Ouarda and Charron, 2018a). Moreover, no TDF curves have been  
78 constructed for Australia till date.

79 In this study, time series data of the daily maximum temperature of Australia at selected stations are  
80 analysed, and the best climate-informed covariates are selected from the most influential covariates for  
81 the annual maximum temperature to characterise the physical process related to the dynamics of the  
82 extreme temperature. For this purpose, three climate drivers, namely – El Niño Southern Oscillation  
83 (ENSO), Southern Annular Mode (SAM) and Indian Ocean Dipole (IOD), are selected along with the  
84 *Time* as potential covariates. After that, stationary TDF curves are constructed following the frequency  
85 analysis method, already adopted in the construction of IDF curves, for different durations and return  
86 periods. The present work also extends these stationary TDF curves to non-stationary conditions  
87 according to the framework outlined by Ouarda and Charron (2018b). Then, based on the Akaike  
88 Information Criterion (AIC) value, the best model for each selected station is identified, and the TDF  
89 curves are constructed under stationary and non-stationary conditions. Finally, the influence of non-  
90 stationarity on the TDF curves is assessed.

## 91 **2. STUDY AREA AND DATA**

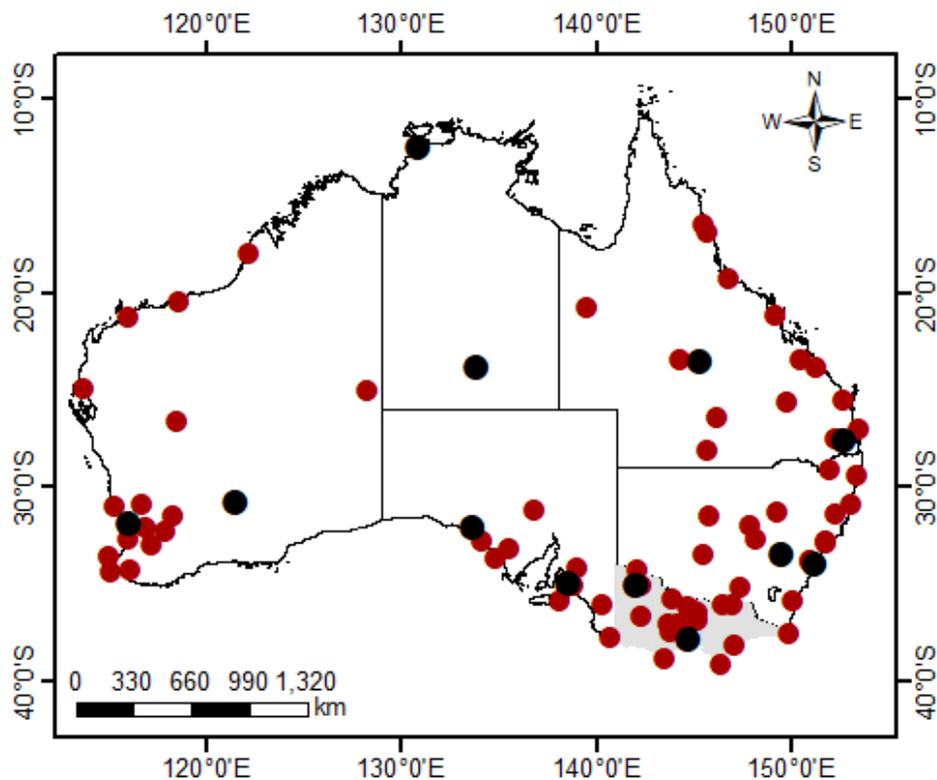
92 The whole Australia is considered as the study area except Tasmania (Figure 1). Australia experiences  
93 a variety of climates across the country because of a wide range of geographical extent consisting of  
94 tropical-influenced climate in the northern and north-eastern zones, Mediterranean-like climate in the  
95 southern coastal zones and desert climate in most of the central interior zones (Turney *et al.*, 2007).

### 96 **2.1 Temperature time series**

97 All the recorded daily maximum temperatures across Australia weather stations were examined for  
98 inconsistency, outliers and gaps. The stations, which had less than 5% missing data over the last five  
99 decades covering the time period from 1969 to 2021, were selected for analysis. According to the  
100 abovementioned criterion, 82 weather stations in Australia (out of 1415 weather stations) were selected.

101 The gaps in the daily maximum temperature time series in the selected stations were spatially  
102 interpolated with the same time series data from surrounding weather stations located within a 30 km  
103 radius. The locations of the selected temperature stations are shown in Figure 1.

104 The annual maximum daily temperatures were calculated for each selected weather station from the  
105 recorded daily maximum temperature data. Particular attention was given to the summer season of  
106 Australia, which is considered to be December to February, and the extended summer is considered to  
107 be from November to March. Therefore, the annual maximum daily temperature calculated from each  
108 calendar year would have been divided into summer or extended summer seasons into two consecutive  
109 years. Therefore, the hydrologic year in Australia, starting from April to March, instead of the calendar  
110 year, was considered in this study to calculate the annual maximum daily temperature.



111  
112 Figure 1. Study area and selected stations (stations marked with black circles are described in detail in  
113 this paper).

## 114 2.2 Covariates

115 Australia is extremely susceptible to changes in the ocean-atmosphere system. The inter-annual  
116 dynamics of Australia's seasonal weather are regulated by the climatic variability of three neighbouring  
117 oceans – the Pacific, Indian, and Southern Oceans (Cai *et al.*, 2014; Maher and Sherwood, 2014;

118 Oliveira and Ambrizzi, 2017), known as climate drivers. Three primary climate drivers strongly  
119 influence Hydroclimate in Australia – The El Niño Southern Oscillation (ENSO) (Power *et al.*, 1999;  
120 Cai and Van Rensch, 2012), Southern Annular Mode (SAM) (Thompson, Wallace and Hegerl, 2000;  
121 Risbey *et al.*, 2009) and Indian Ocean Dipole (IOD) (Ummenhofer *et al.*, 2009, 2011).

122 Therefore, in this study, three physical processes, in which two of them from the sea surface temperature  
123 – ENSO, IOD and one from air pressure – SAM, and also *Time*, were considered to be one of the  
124 potential covariates to identify the best covariate for developing non-stationarity TDF curves.

125 *i. El Niño-Southern Oscillation (ENSO) cycle*

126 The El Niño-Southern Oscillation (ENSO) is considered to be a major climate driver over many regions  
127 of the globe (Ropelewski and Halpert, 1988; Halpert and Ropelewski, 1992), including Australia  
128 (Nicholls, 1985). The SST index is considered the monthly sea surface temperature anomaly over NINO  
129 3.4 (17 °E–120 °W, 5 °S–5 °N) region (Bellenger *et al.*, 2014) and was used as a covariate representing  
130 ENSO. Nino3.4 was obtained from [https://psl.noaa.gov/gcos\\_wgsp/Timeseries/Nino34](https://psl.noaa.gov/gcos_wgsp/Timeseries/Nino34).

131 ENSO is considered to be a strong climate driver in Australia (Power *et al.*, 2006; Risbey *et al.*, 2009;  
132 Arblaster and Alexander, 2012), especially in southeastern Australia (Nicholls and Lucas, 2007) and  
133 northeast Australia (Min, Cai and Whetton, 2013; Cowan *et al.*, 2014; Perkins, Argüeso and White,  
134 2015), but weak in the far southeast of Australia (Parker, Berry and Reeder, 2013; Boschath *et al.*, 2015).  
135 In Australia, it has been observed that the strength of ENSO decreases along a north-south gradient.

136 *ii. Indian Ocean Dipole (IOD)*

137 The Indian Ocean Dipole (IOD) is represented with Dipole Mode Index (DMI) (Saji *et al.*, 1999; Meyers  
138 *et al.*, 2007; Liu *et al.*, 2014) and calculated as the difference of SST anomalies in the western equatorial  
139 Indian Ocean (50 to 70°E, 10°S to 10°N) minus those in the east (90 to 110°E, 10°S to 0°). The monthly  
140 DMI was derived from [https://psl.noaa.gov/gcos\\_wgsp/Timeseries/DMI/](https://psl.noaa.gov/gcos_wgsp/Timeseries/DMI/) and considered as a covariate  
141 that represents IOD.

142 IOD has been identified as the major driver for Eastern Australia (Meyers *et al.*, 2007; Risbey *et al.*,  
143 2009; Cai *et al.*, 2011) and has a positive correlation with maximum temperatures during winter and  
144 spring (May and November) (Min, Cai and Whetton, 2013; White *et al.*, 2013).

145 *iii. SAM*

146 The Southern Annular Mode (SAM) describes the north-south movement of the westerly wind belt that  
147 circles Antarctica and is a monthly mean sea level pressure (MSLP) anomaly time dataset. The data for  
148 this index were obtained from <http://www.nerc-bas.ac.uk/icd/gjma/sam.html>.

149 SAM considerably impacts the rainfall in Australia (Hendon, Thompson and Wheeler, 2007),  
150 particularly in southern Australia. SAM is even associated with 10-15% of the rainfall in southwest  
151 Australia. SAM also affects the temperature in Australia and is identified as the main climate driver for  
152 the southwestern land division in western Australia (Guthrie, 2021).

153 *iv. Time*

154 *Time* has been utilised as a covariate to investigate the non-stationarity in the IDF curves all over the  
155 world (Sugahara, da Rocha and Silveira, 2009; Yilmaz and Perera, 2014). Therefore, in this study, *Time*  
156 was also considered as one of the potential covariates to construct the non-stationary TDF curves. The  
157 hydrological years of the annual maximum temperature series were transformed into a series of integers  
158 from 1 to the number of years of the selected time duration and considered as a covariate representing  
159 *Time*.

160 **3. METHODOLOGY**

161 In this study, firstly, the non-stationarity of the temperature data series was assessed to justify the  
162 requirements of the incorporation of the non-stationarity in the construction of TDF curves. Then, the  
163 correlations between the selected three physical processes - ENSO, SAM and IOD were computed to  
164 identify the best covariate. In this approach, *Time* was also used as one of the covariates. Different non-  
165 stationary models were constructed considering different covariates stated earlier as well as their  
166 combination. Finally, *time* or any of the selected physical processes (ENSO, IOD and SAM) was  
167 considered a covariate in Models with one covariate.

168 Similarly, *Time* and one climate index (*Time* – ENSO, *Time* – IOD and *Time* – SAM) were adopted in  
169 models with two covariates. First, the best model for the annual maximum temperature series was  
170 chosen based on the Akaike information criterion (AIC). After identifying the best non-stationary  
171 model, the non-stationary design temperature was computed using the parameters of the best non-  
172 stationary model.

173 GEV distribution was selected to develop both stationary and non-stationary TDF curves as it is widely  
174 used to develop model climatic extremes (Ouarda and Charron, 2018a; Ouarda, Charron and St-Hilaire,

175 2020; Haddad, 2021; Devi, Gouda and Lenka, 2022). Under the condition of stationary, three  
176 parameters of the GEV distributions,  $\mu$ ,  $\sigma$ ,  $\xi$ , which are the location, scale and shape parameters,  
177 respectively, were assumed to be constant, whereas, under non-stationary conditions, the parameters  
178 were expressed as a linear or quadratic function of the covariate(s). However, the shape parameter was  
179 assumed to be constant for all the cases since the reliability of modelling the shape parameter is low  
180 (Cheng *et al.*, 2014; Ganguli and Coulibaly, 2017). The distribution parameters were calculated using  
181 the maximum composite likelihood method utilising the optimisation function *fmincon* in MATLAB.  
182 Finally, the stationary and non-stationary TDF curves were compared to assess the influence of non-  
183 stationarity in the construction of TDF curves.

### 184 **3.1 Detection of non-stationarity in Temperature time series**

185 All the historical records of annual maximum temperature at the selected stations over the selected time  
186 period were tested for non-stationary signals. An augmented Dickey-Fuller (ADF) test was applied to  
187 the annual maximum temperature data series for each of the selected stations. In this test, the null  
188 hypothesis assumes that the data series is non-stationary and thereby considered stationary when the  
189 null hypothesis is rejected at the selected significance level.

### 190 **3.2 Correlation between climate drivers and temperature time series**

191 The relations between the maximum annual temperatures for all the selected stations and climate drivers  
192 – ENSO, SAM and IOD were tested using correlation analyses considering both concurrent and time-  
193 lagged relationships.

194 The covariates representing climate drivers in this study were computed as the moving averages of  
195 corresponding climate drivers over three consecutive months starting from April and ending in March  
196 of the next year, covering the full hydrological year and denoted in this study by the first letter of the  
197 three months (AMJ, MJJ, JJA, JAS, ASO, SON, OND, NDJ, DJF, JFM). To determine the appropriate  
198 season of the climate driver acting as the predictor of the annual maximum temperature, correlations  
199 between seasonal climate drivers and annual maximum temperatures were explored. Best covariates  
200 and the suitable season were selected based on the higher correlation values between the annual  
201 maximum temperature series and climate drivers averaged over the selected season. This selection  
202 process was also validated graphically by plotting all the correlations between the annual maximum  
203 temperature time series and all the climate drivers for all seasons.

### 204 3.3 Construction of TDF curves

#### 205 i. General TDF relationship

206 The general TDF relationship can be expressed by following the formulation for IDF curves proposed  
207 by (Rossi and Villani, 1994; Koutsoyiannis, Kozonis and Manetas, 1998):

$$208 \quad t_R(d) = \frac{a(R)}{b(d)} \quad (1)$$

209 where,  $d$  is the duration and  $R$  is the return period. In this formulation, the dependency of return level,  
210  $t_R(d)$  on  $R$  and  $d$  can be modelled separately. The distribution of maximum average temperature  $T(d)$   
211 governs the function  $a(R)$  that defines the TDF curves, which remain parallel for various return periods.  
212 The function  $b(d)$  controls the shape of the TDF curves and is given as:

$$213 \quad b(d) = (d + \theta)^\eta \quad (2)$$

214 where,  $\theta$  and  $\eta$  are the shape parameters subjected to the boundary conditions of  $\theta > 0$  and  $0 < \eta < 1$ .

215 If the probability distribution of  $T(d)$  is  $F_{T(d)}(i; d)$ , it will also be the distribution of  $Y = T(d)b(d)$ ,  
216 which is the maximum average temperature scaled by  $b(d)$  (i.e.,  $F_{T(d)}(i; d) = F_Y(y_T) = 1 - \frac{1}{T}$ ).  
217 Consequently, the expression  $a(R)$  is given by:

$$218 \quad a(R) = F_y^{-1} \left( 1 - \frac{1}{R} \right) \quad (3)$$

#### 219 ii. Stationary TDF curves

220 The GEV distribution is the most widely used probability distribution that is used to model climate  
221 extremes, and hence, the GEV distribution is used here to model  $T(d)$ . The cumulative distribution  
222 function of the GEV is given by:

$$223 \quad F(x) = \exp \left\{ - \left[ 1 + \kappa \left( \frac{x - \mu}{\sigma} \right) \right]^{-1/k} \right\}, \quad (4)$$

224 where,  $\mu$ ,  $\sigma$  and  $\kappa$  are the location, scale, and shape parameters, respectively.  $F(x)$  is defined for  $1 +$   
225  $\frac{\kappa(x - \mu)}{\sigma} > 0$ , where  $\sigma > 0$ .

226 The general stationary TDF relationship based on the GEV distribution can be expressed as (Ouarda  
227 and Charron, 2018b):

228 
$$t_R(d) = \frac{a(R)}{b(d)} = \frac{\mu - \frac{\sigma}{\kappa} \left\{ 1 - \left[ \log \left( 1 - \frac{1}{R} \right)^{-\kappa} \right] \right\}}{(d + \theta)^\eta} \quad (5)$$

229 *iii. Non-stationary TDF curves*

230 To incorporate non-stationarity in Equation (5), the three parameters of the GEV distribution described  
 231 in Equation (4) are considered to be dependent on time to incorporate the non-stationarity in TDF curve  
 232 development and Equation (4) can be expressed as:

233 
$$F(x) = \exp \left\{ - \left[ 1 + \kappa(t) \left( \frac{x - \mu(t)}{\sigma(t)} \right) \right]^{-1/k} \right\}, 1 + \frac{\kappa(t)(x - \mu(t))}{\sigma(t)} > 0 \quad (6)$$

234 In this study, the location and scale parameters are considered to vary with covariates linearly or  
 235 quadratically, while all other shape parameters,  $\theta$ ,  $\eta$  and  $\kappa$  for the GEV distribution are assumed to be  
 236 constant (Katz, Parlange and Naveau, 2002; Adlouni and Ouarda, 2009).

237 The vectors of the distribution parameters  $\psi = (\mu, \sigma, \theta, \eta)$  and  $\psi = (\mu_0, \mu_1, \dots, \sigma_0, \sigma_1, \dots, \kappa, \theta, \eta)$  are  
 238 estimated using the maximum composite likelihood for stationary and non-stationary TDF curves,  
 239 respectively.

240 **4. RESULTS AND DISCUSSION**

241 The results for 12 stations (two stations from each state in Australia and shown as black dots in Figure  
 242 1) are presented in greater detail (out of the 82 selected stations) in the following sections.

243 **4.1 Detection of non-stationarity**

244 The augmented Dickey-Fuller (ADF) test was applied to each annual maximum temperature data series  
 245 separately for each selected station. Based on the ADF tests, the test statistics and p-value for 12 stations  
 246 are summarised in Table 1. Four stations showed ADF test statistics below the critical value (-3.507),  
 247 and the p-values were less than 0.05 at the 95% confidence level, as shown in boldface in Table 1. For  
 248 these stations, the null hypothesis was rejected, and the data series was considered stationary. Therefore,  
 249 the null hypothesis could not be rejected for the rest of the stations, and the non-stationarities in the data  
 250 series were detected. Although all the available stationarity tests have some drawbacks and are not  
 251 decisive alone (Cai, Cowan and Sullivan, 2009), the results of ADF tests here highlighted the presence  
 252 of the non-stationarity in the temperature time series and consequently emphasised the importance of  
 253 incorporating the non-stationarity in the construction of TDF curves.

254 Table 1. Summary of ADF test statistics and p-values for selected 12 stations in Australia.

Station ID	Station	p-value	Test statistics
066037	Sydney Airport AMO	0.465	-2.252
063005	Bathurst Agricultural Station	0.088	-3.246
014015	Darwin Airport	0.052	-3.493
015590	<b>Alice Springs Airport</b>	<b>0.008</b>	<b>-4.244</b>
009021	<b>Perth Airport</b>	<b>0.004</b>	<b>-4.579</b>
012038	Kalgoorlie-Boulder Airport	0.050	-3.506
040004	<b>Amberley AMO</b>	<b>0.006</b>	<b>-4.364</b>
036007	<b>Barcaldine Post Office</b>	<b>0.047</b>	<b>-3.532</b>
018012	Ceduna AMO	0.454	-2.275
023034	Adelaide Airport	0.060	-3.424
087031	Laverton RAAF	0.194	-2.829
076064	Walpeup Research	0.121	-3.088

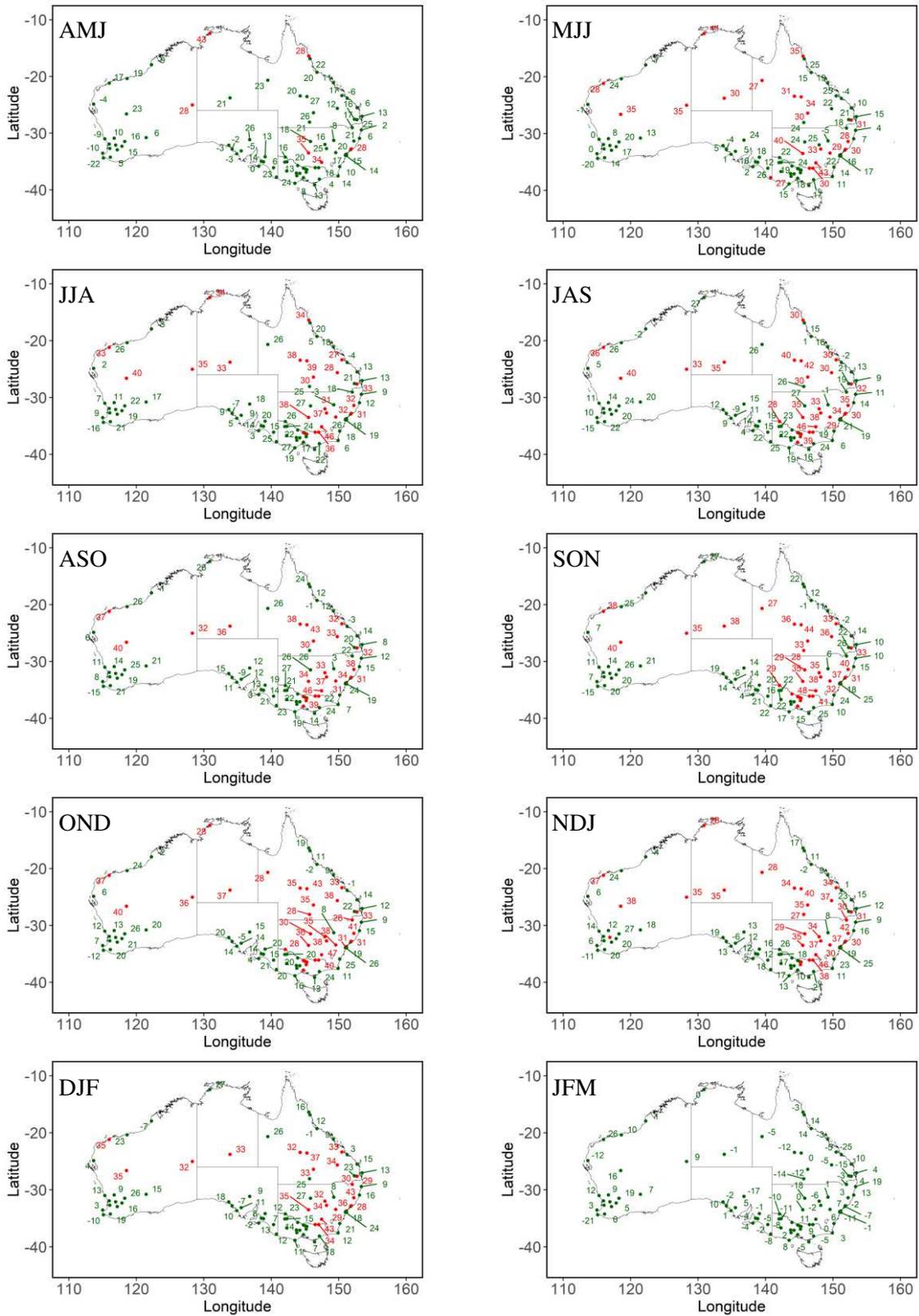
255 **4.2 Correlation analysis and covariate selection**

256 The correlations between the annual maximum temperatures and the ENSO, IOD and SAM for all the  
 257 seasons are presented spatially in Figures 2, 3 and 4, respectively.

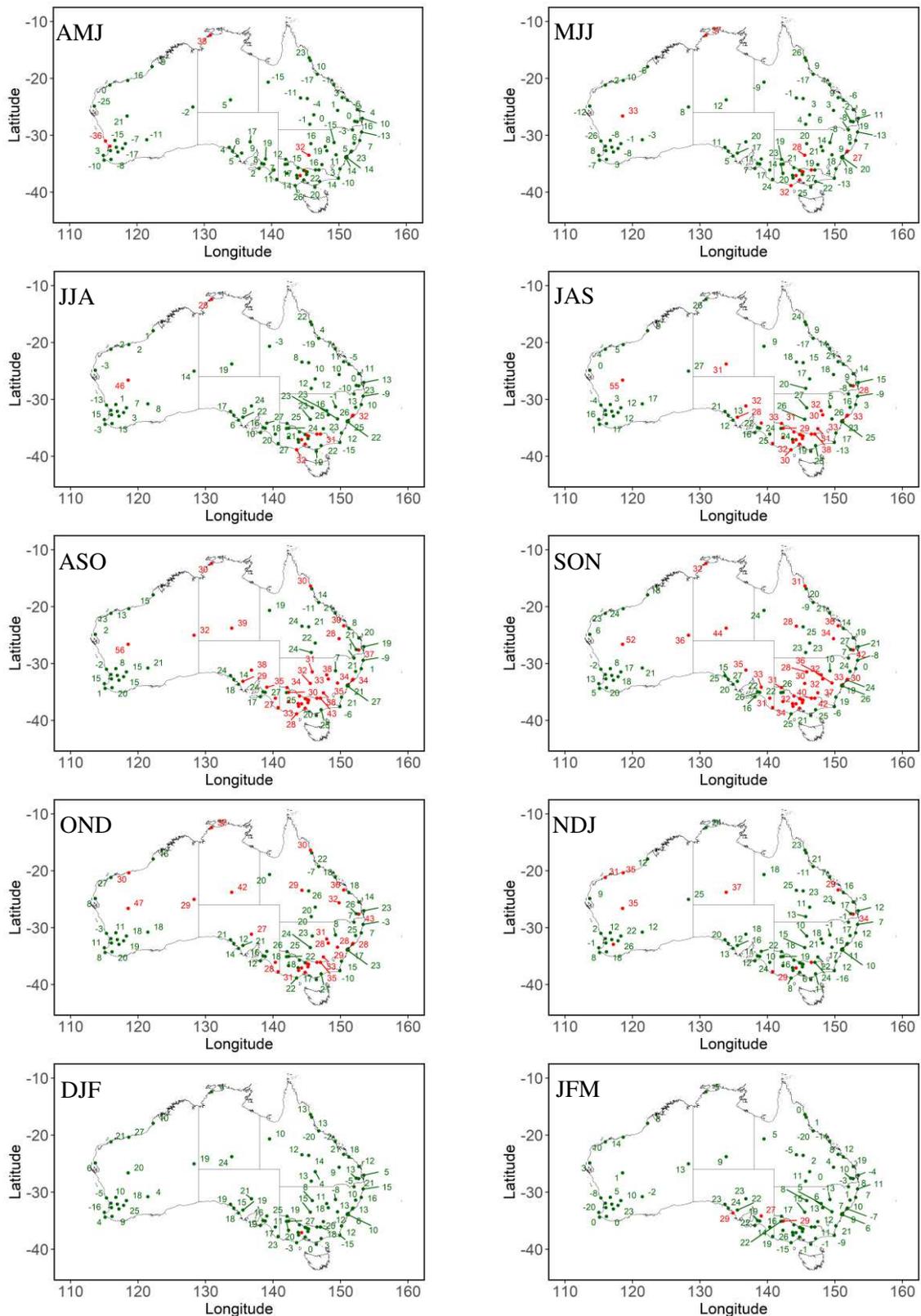
258 The correlations between ENSO and the annual maximum temperatures were the strongest and  
 259 statistically significant during the spring and summer seasons, particularly in the eastern regions of  
 260 Australia. The positive influence of the ENSO on the temperature was weaker at the beginning of the  
 261 hydrologic year and increased gradually over the year. At the end of the hydrological year, this influence  
 262 became weak, even negative for some stations.

263 Similar to the ENSO, positive influences of IOD on the temperatures were also identified all over  
 264 Australia. This influence was enhanced during the spring season and was statistically significant in the  
 265 southeastern region of Australia. However, at the end of the hydrological year, the impact of IOD also  
 266 became weaker and negative at some stations.

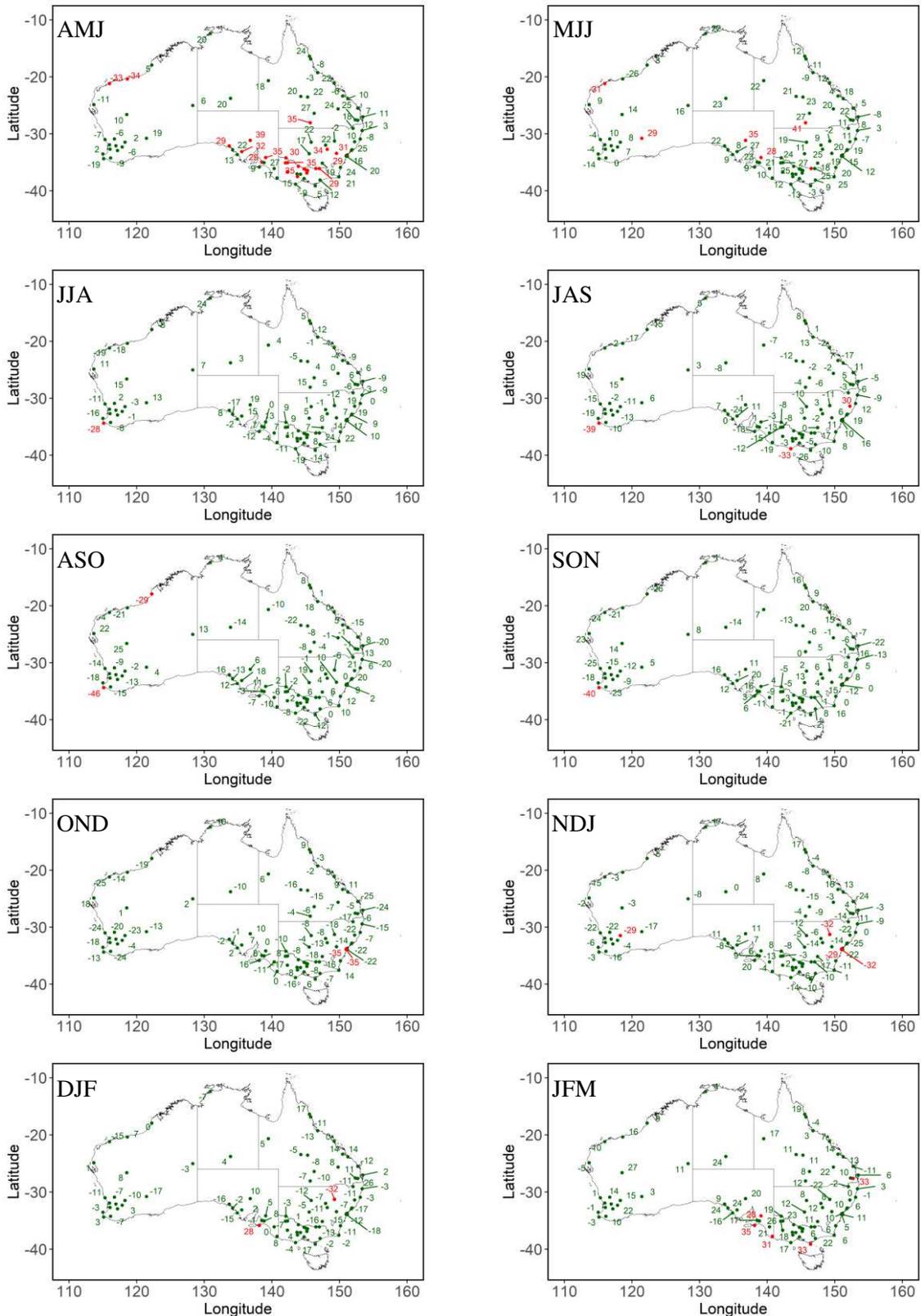
267 Unlike the other two climate drivers, SAM showed different relationships with the temperature in  
 268 Australia. SAM exhibited a statistically significant positive relationship with the temperature at inland  
 269 of the southeastern part during Autumn and negative relations with the temperature observed during  
 270 spring at the stations in the western regions of Australia.



271 Figure 2. Spatial correlation of annual maximum temperature with ENSO during 1969 – 2021.  
 272 Correlation values are multiplied by 100, and red values are statistically significant at a 95% confidence  
 273 level.



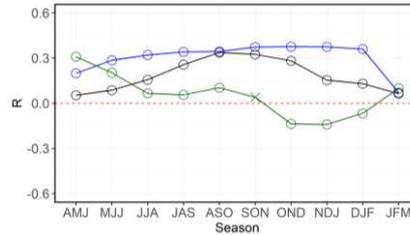
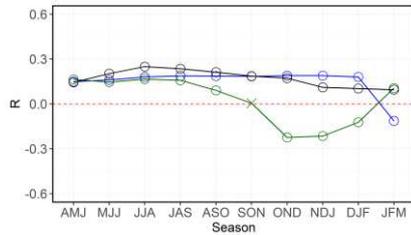
274 Figure 3. Spatial correlation of annual maximum temperature with IOD during 1969 – 2021. Correlation  
 275 values are multiplied by 100, and red values are statistically significant at a 95% confidence level.



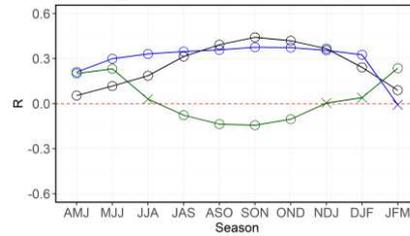
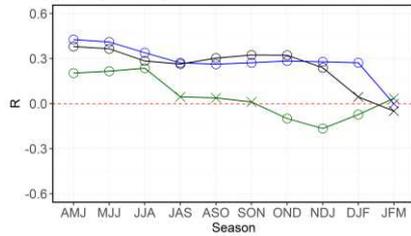
276 Figure 4. Spatial correlation of annual maximum temperature with SAM during 1969 – 2021.  
 277 Correlation values are multiplied by 100, and red values are statistically significant at a 95% confidence  
 278 level.



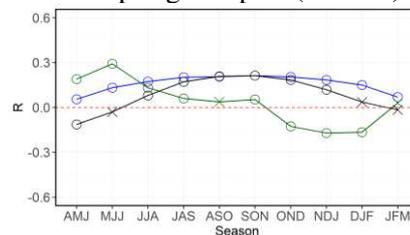
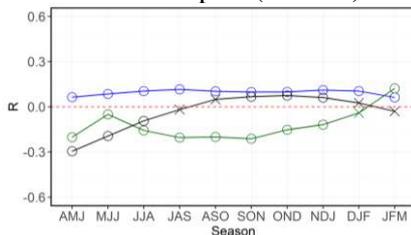
New South Wales



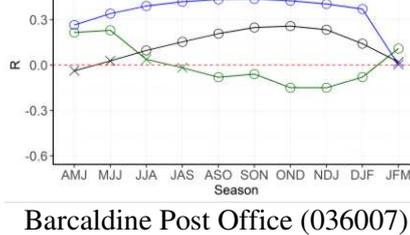
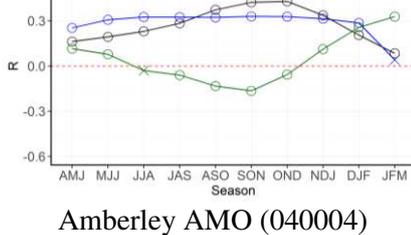
Northern Territory



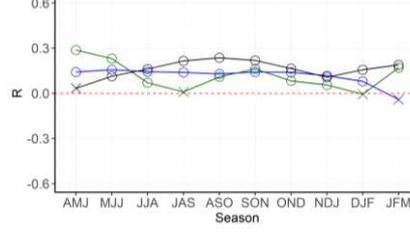
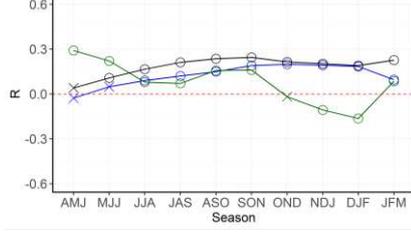
Western Australia



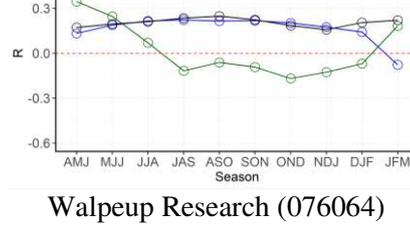
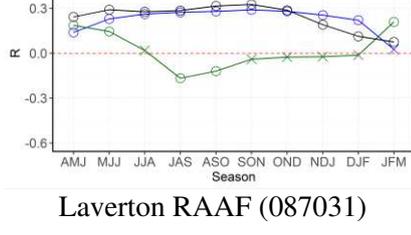
Queensland



South Australia



Victoria



300 Figure 6. Correlations between annual maximum temperature and climate drivers. The blue, black and  
 301 green lines represent correlations between AMT and ENSO, IOD and SAM, respectively.

302 Table 2. Pearson correlation coefficients between the annual maximum temperature series and the  
 303 selected seasonal climate index

<b>Station ID</b>	<b>Station Name</b>	<b>Latitude</b>	<b>Longitude</b>	<b>Climatic driver and season</b>	<b>Coefficient Value</b>
066037	Sydney Airport AMO	-33.95	151.17	IOD(JJA)	0.25
063005	Bathurst Agricultural	-33.43	149.56	ENSO(SON)	0.37
014015	Darwin Airport	-12.42	130.89	ENSO(AMJ)	0.43
015590	Alice Springs Airport	-23.80	133.89	IOD(SON)	0.38
009021	Perth Airport	-31.93	115.98	SAM(SON)	-0.21
012038	Kalgoorlie-Boulder Airport	-30.78	121.45	SAM(MJJ)	0.29
040004	Amberley AMO	-27.63	152.71	IOD(OND)	0.43
036007	Barcaldine Post Office	-23.55	145.29	ENSO(SON)	0.44
018012	Ceduna AMO	-32.13	133.70	IOD(SON)	0.25
023034	Adelaide Airport	-34.95	138.52	IOD(ASO)	0.24
087031	Laverton RAAF	-37.86	144.76	IOD(SON)	0.33
076064	Walpeup Research	-35.12	142.00	IOD(ASO)	0.25

304 The Pearson correlation coefficient values for the annual maximum temperature series from April–  
 305 May–June (AMJ) to January-February–March (JFM) of the same hydrological year recorded at 12  
 306 stations are presented in Figure 6. These illustrations validated the selection of the best climate drivers  
 307 and seasons presented in Figure 5. Selected climate drivers with the seasons for the selected 12 stations  
 308 are summarised in Table 2.

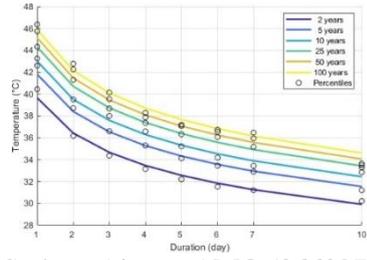
### 309 **4.3 Stationary TDF**

310 Figure 7 illustrates the stationary TDF curves based on GEV distribution as described in the  
 311 methodology section. Estimated temperatures were plotted against the selected durations (1, 2, 3, 4, 5,  
 312 6, 7 and 10 days), with each curve indicating a different return period such as 2, 5, 10, 25, 50 and 100  
 313 years at station scale. These TDF curves demonstrated a significant rise in temperature with higher  
 314 return periods and decreased with the increase in duration for all the selected stations.

315

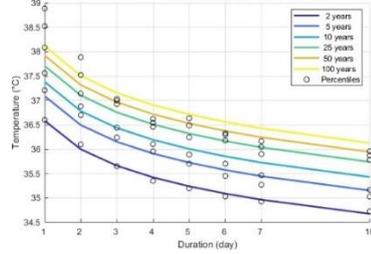
316

New South Wales



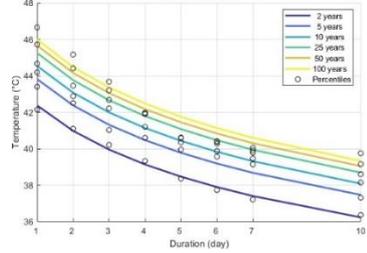
Sydney Airport AMO (066037)

Northern Territory



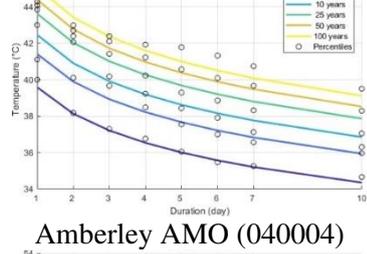
Darwin Airport (014015)

Western Australia



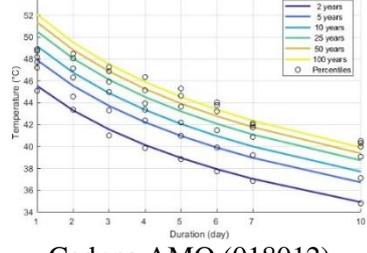
Perth Airport (009021)

Queensland



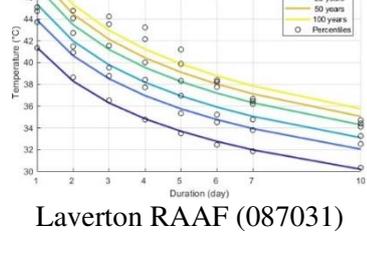
Amberley AMO (040004)

South Australia

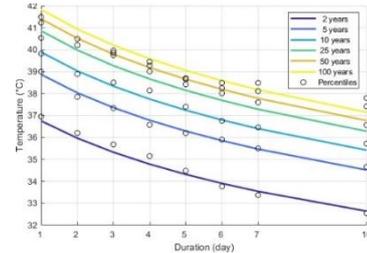


Ceduna AMO (018012)

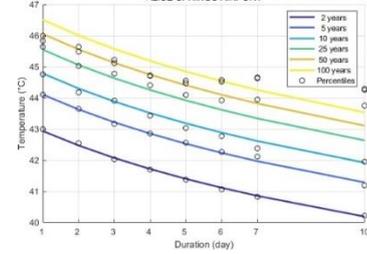
Victoria



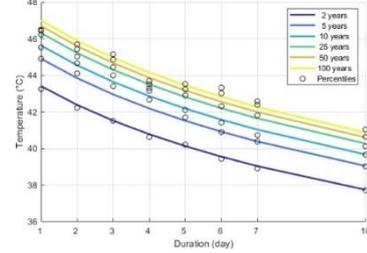
Laverton RAAF (087031)



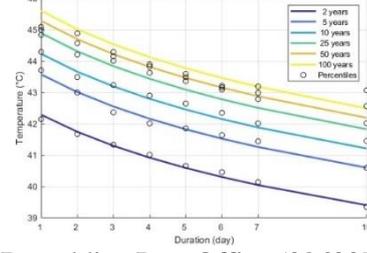
Bathurst Agricultural (063005)



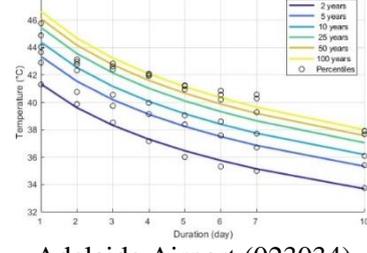
Alice Springs Airport (015590)



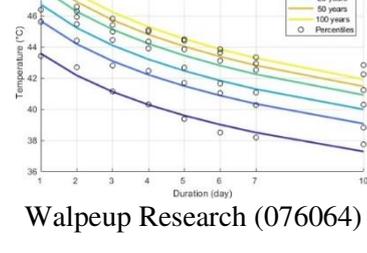
Kalgoorlie-Boulder Airport (012038)



Barcaldine Post Office (036007)



Adelaide Airport (023034)



Walpeup Research (076064)

317 Figure 7. Stationary TDF curves for 2, 5, 10, 25, 50 and 100 years return periods.

318 **4.1 Non-stationary TDF surfaces**

319 For each TDF model and station, the maximal independence log-likelihood, the CL-AIC statistic, and  
 320 the model parameters are summarised in Table 3. According to the values obtained by the AIC criterion  
 321 from each model, the table only shows the optimal parameter relationship among the constant, linear  
 322 and quadratic relationships to the covariate or the combinations of the covariates.

323 Table 3. Summary of parameters for the selected models and their error estimation.

Station	Model	$l_{ind}$	CL-AIC	Model parameters
Sydney Airport AMO	Stationary	-881.47	1791.45	$\mu, \sigma$
	<b>Time</b>	-866.07	<b>1775.35</b>	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	CD	-873.53	1787.97	$\mu_l = \mu_0 + \mu_1 I, \sigma$
	<b>Time + CD</b>	-863.46	1780.83	$\mu_l = \mu_0 + \mu_1 I + \mu_2 T, \sigma$
Bathurst Agricultural Station	Stationary	-939.29	1917.13	$\mu, \sigma$
	<b>Time</b>	-851.05	1751.47	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	CD	-895.64	1854.87	$\mu_l = \mu_0 + \mu_1 E + \mu_2 E^2, \sigma$
	<b>Time + CD</b>	-808.35	<b>1692.00</b>	$\mu_l = \mu_0 + \mu_1 E + \mu_2 E^2 + \mu_3 T, \sigma$
Darwin Airport	Stationary	-324.62	676.26	$\mu, \sigma$
	<b>Time</b>	-262.29	572.28	$\mu_l = \mu_0 + \mu_1 T + \mu_2 T^2, \sigma$
	CD	-287.11	625.63	$\mu_l = \mu_0 + \mu_1 E + \mu_2 E^2, \sigma$
	<b>Time + CD</b>	-234.22	<b>525.28</b>	$\mu_l = \mu_0 + \mu_1 E + \mu_2 T + \mu_3 T^2, \sigma$
Alice Springs Airport	Stationary	-675.70	1390.62	$\mu, \sigma$
	<b>Time</b>	-642.08	1358.22	$\mu_l = \mu_0 + \mu_1 T$ $\sigma_l = \sigma_0 + \sigma_1 T + \sigma_2 T^2$
	CD	-623.41	1304.61	$\mu_l = \mu_0 + \mu_1 I$ $\sigma_l = \sigma_0 + \sigma_1 I$
	<b>Time + CD</b>	-621.70	<b>1304.38</b>	$\mu_l = \mu_0 + \mu_1 I + \mu_2 T, \sigma$
Perth Airport	Stationary	-763.95	1557.98	$\mu, \sigma$
	<b>Time</b>	-763.95	1567.02	$\mu, \sigma_1 = \sigma_0 + \sigma_1 T$
	CD	-736.19	1526.39	$\mu_l = \mu_0 + \mu_1 S + \mu_2 S^2$ $\sigma_l = \sigma_0 + \sigma_1 S + \sigma_2 S^2$
	<b>Time + CD</b>	-732.25	<b>1524.26</b>	$\mu_l = \mu_0 + \mu_1 S + \mu_2 S^2 + \mu_3 T$ $\sigma_l = \sigma_0 + \sigma_1 S + \sigma_2 T$
Kalgoorlie- Boulder Airport	Stationary	-791.51	1616.49	$\mu, \sigma$
	<b>Time</b>	-774.88	1588.99	$\mu, \sigma_1 = \sigma_0 + \sigma_1 T$
	CD	-764.00	1582.79	$\mu_l = \mu_0 + \mu_1 S + \mu_2 S^2, \sigma$

Station	Model	$l_{ind}$	CL-AIC	Model parameters
	<b>Time + CD</b>	-741.39	<b>1554.90</b>	$\mu_l = \mu_0 + \mu_1 S + \mu_2 S^2 + \mu_3 T$ $\sigma_l = \sigma_0 + \sigma_1 S + \sigma_2 T$
Amberley AMO	Stationary	-844.36	1728.03	$\mu, \sigma$
	<i>Time</i>	-798.10	1649.35	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	CD	-806.68	1665.05	$\mu_l = \mu_0 + \mu_1 I, \sigma$
	<b>Time + CD</b>	-779.77	<b>1623.46</b>	$\mu_l = \mu_0 + \mu_1 I + \mu_2 T, \sigma$
Barcaldine Post Office	Stationary	-733.20	1505.18	$\mu, \sigma$
	<i>Time</i>	-721.59	1498.50	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	<b>CD</b>	-635.98	<b>1344.19</b>	$\mu_l = \mu_0 + \mu_1 E + \mu_2 E^2$ $\sigma_l = \sigma_0 + \sigma_1 E + \sigma_2 E^2$
	<i>Time + CD</i>	-628.05	1353.58	$\mu_l = \mu_0 + \mu_1 E + \mu_2 E^2 + \mu_3 T + \mu_4 T^2$ $\sigma_l = \sigma_0 + \sigma_1 E + \sigma_2 T$
Ceduna AMO	Stationary	-934.32	1898.72	$\mu, \sigma$
	<i>Time</i>	-924.77	1890.45	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	CD	-919.05	1888.54	$\mu_l = \mu_0 + \mu_1 I + \mu_2 I^2, \sigma$
	<b>Time + CD</b>	-892.92	<b>1861.80</b>	$\mu_l = \mu_0 + \mu_1 I + \mu_2 I^2 + \mu_3 T + \mu_4 T^2$ $\sigma_l = \sigma_0 + \sigma_1 I + \sigma_2 T$
Adelaide Airport	Stationary	-885.62	1803.53	$\mu, \sigma$
	<i>Time</i>	-867.75	1790.75	$\mu_l = \mu_0 + \mu_1 T + \mu_2 T^2, \sigma$
	CD	-884.90	1810.97	$\mu, \sigma_l = \sigma_0 + \sigma_1 I$
	<b>Time + CD</b>	-842.22	<b>1750.65</b>	$\mu_l = \mu_0 + \mu_1 I + \mu_2 T + \mu_3 T^2$ $\sigma_l = \sigma_0 + \sigma_1 I + \sigma_2 T$
Laverton RAAF	Stationary	-939.13	1908.59	$\mu, \sigma$
	<b>Time</b>	-931.24	<b>1902.91</b>	$\mu_l = \mu_0 + \mu_1 T, \sigma$
	CD	-932.31	1904.82	$\mu, \sigma_l = \sigma_0 + \sigma_1 I + \sigma_2 I^2$
	<i>Time + CD</i>	-922.50	1905.08	$\mu_l = \mu_0 + \mu_1 I + \mu_2 I^2 + \mu_3 T, \sigma$
Walpeup Research	Stationary	-915.68	1866.35	$\mu, \sigma$
	<i>Time</i>	-864.87	1788.51	$\mu_l = \mu_0 + \mu_1 T + \mu_2 T^2, \sigma$
	CD	-908.46	1865.52	$\mu_l = \mu_0 + \mu_1 I, \sigma$
	<b>Time + CD</b>	-846.70	<b>1772.17</b>	$\mu_l = \mu_0 + \mu_1 I + \mu_2 T + \mu_3 T^2$ $\sigma_l = \sigma_0 + \sigma_1 I + \sigma_2 T$

324 \*  $\kappa, \theta, \eta$  are constant for all models.

325  $T = \text{Time}, E = \text{ENSO}, I = \text{IOD}, S = \text{SAM}$

326 The CL-AIC statistic and log-likelihood suggested that the *Time* and covariate individually or  
327 combinedly increased the goodness-of-fit compared to the stationary model for all the stations. It's

328 worth noting that employing a non-stationary model or the GEV enhances the log-likelihood in every  
329 case. However, the performances of different stationary and non-stationary TDF models were compared  
330 by the CL –AIC values, which were penalised due to the inclusion of more variables and thereby  
331 provided more reliable results.

332 Most stations (9 out of 12) with a combination of *Time* and Climatic Drivers (CD) as covariates showed  
333 the best goodness-of-fit. This suggested that the combination of the two covariates considerably  
334 impacted severe temperatures. *Time* was more prominent than climate drivers and qualified as the best  
335 covariate in cases of Sydney Airport AMO and Laverton RAAF stations. On the other hand, for  
336 Barcaldine Post Office, the influence of covariate alone was stronger than *Time* or the combination of  
337 *Time* and climate driver as the covariate.

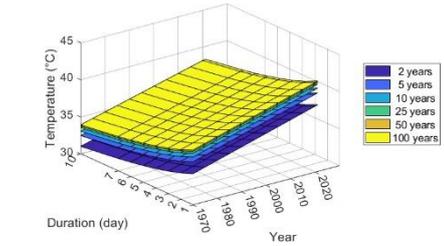
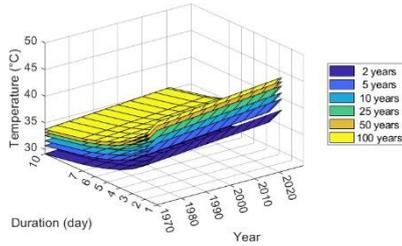
338 *i. Non-stationary TDF surfaces – One covariate (Time or Climate driver)*

339 Figures 8 and 9 present the non-stationary TDF graphs with the model considering *Time* and climate  
340 driver as a covariate for the typical stations. For all the illustrated stations in this study, non-stationary  
341 TDF models with *Time* as a covariate, either scale or location parameters were found to be varied  
342 linearly with time, except for Darwin, Alice Springs Airport, Adelaide Airport, Walpeup Research,  
343 where either of the parameters varied quadratically with time. TDF curves with *Time* as a covariate for  
344 Kalgoorlie-Boulder Airport stations linearly varied with time but were not parallel to each other for  
345 different return periods. In the case of Perth stations, the estimated temperature increased with return  
346 periods and decreased with duration but did not vary significantly with time. However, in the case of  
347 non-stationary TDF models with climate drivers as covariates, the scale or location parameters varied  
348 quadratically with selected climate drivers for most stations.

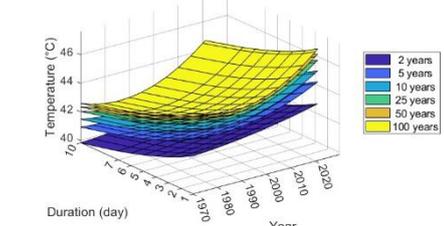
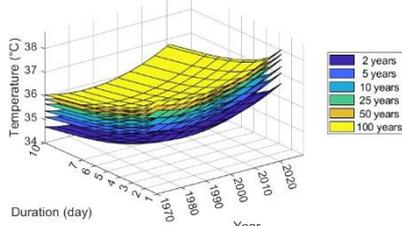
349 Consequently, the TDF curves varied parabolically with climate drivers. Exceptions were observed for  
350 Sydney Airport AMO, Alice Spring Airport, Amberly AMO, Adelaide Airport and Walpeup Research  
351 stations, where either scale or the location parameter changed linearly with climate drivers and exhibited  
352 linear TDF curves for different return periods and durations. These characteristics of non-stationary  
353 TDF models can be seen distinctively in Figures S2 – S5 (in supplementary sections).

354

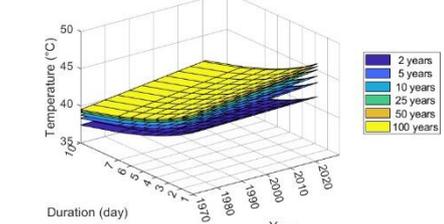
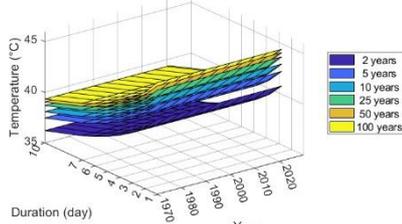
New South Wales



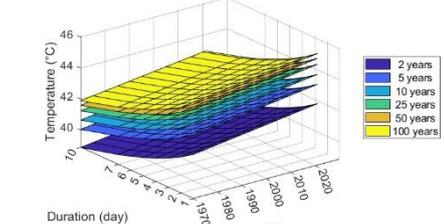
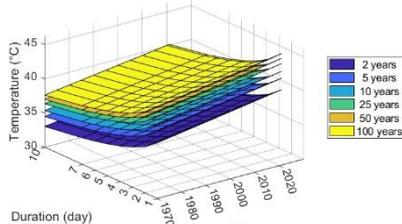
Northern Territory



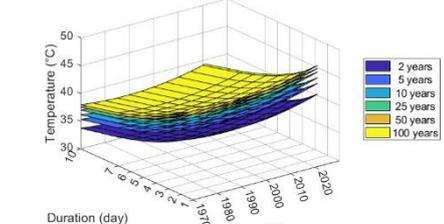
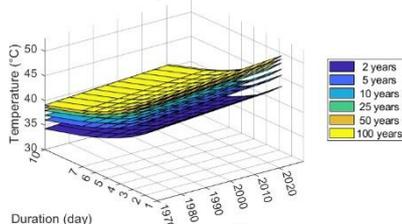
Western Australia



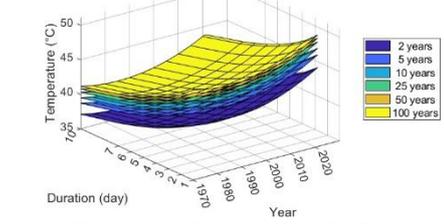
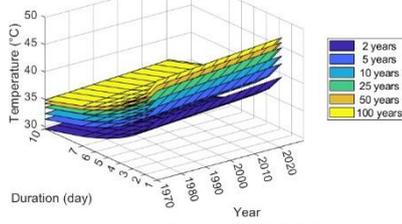
Queensland



South Australia

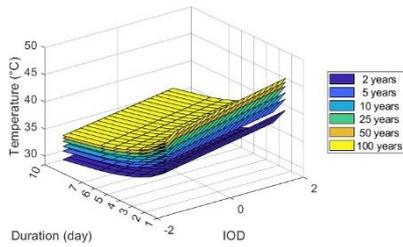


Victoria

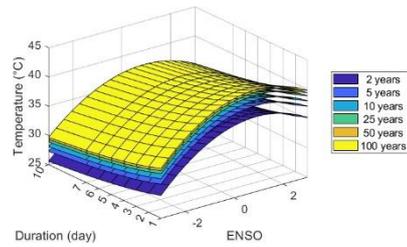


355 Figure 8. Non-stationary TDF surfaces with *Time* covariates for 2, 5, 10, 25, 50 and 100 years return  
 356 periods.

New South Wales

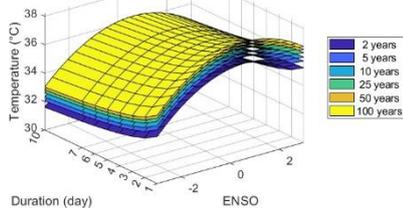


Sydney Airport AMO (066037)

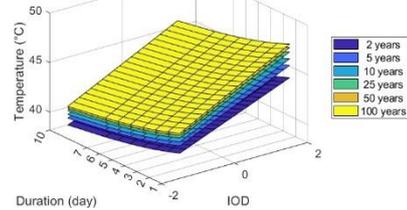


Bathurst Agricultural Station

Northern Territory

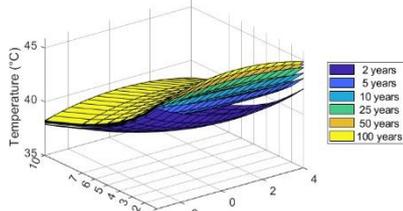


Darwin Airport (014015)

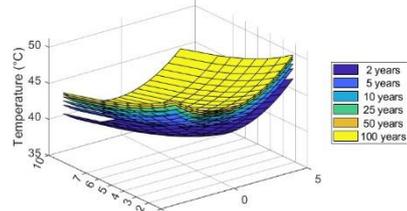


Alice Springs Airport (015590)

Western Australia

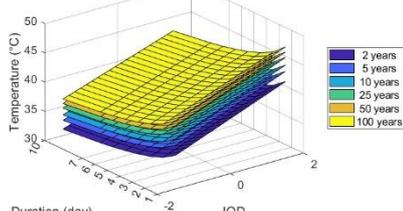


Perth Airport (009021)

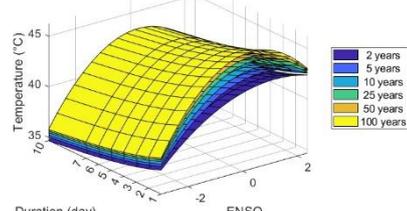


Kalgoorlie-Boulder Airport (012038)

Queensland

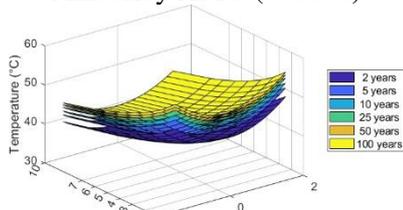


Amberley AMO (040004)

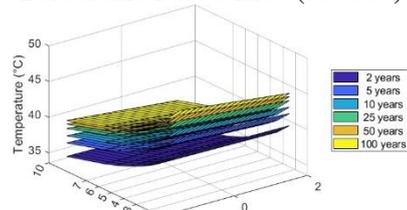


Barcaldine Post Office (036007)

South Australia

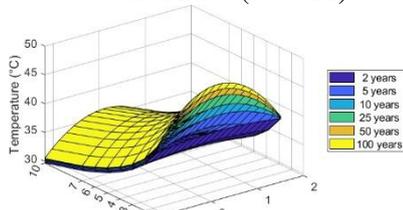


Ceduna AMO (018012)

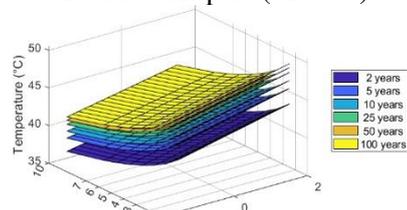


Adelaide Airport (023034)

Victoria



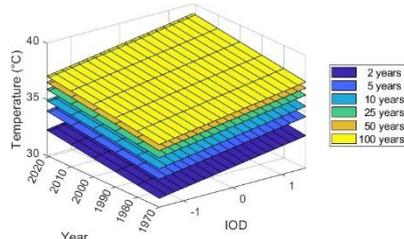
Laverton RAAF (087031)



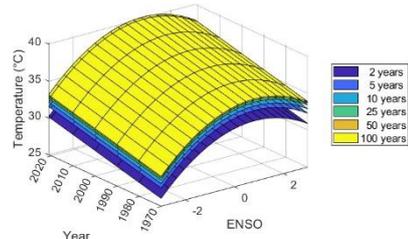
Walpeup Research (076064)

357 Figure 9. Non-stationary TDF surfaces with climate driver covariates for 2, 5, 10, 25, 50 and 100 years  
358 return periods.

New South Wales

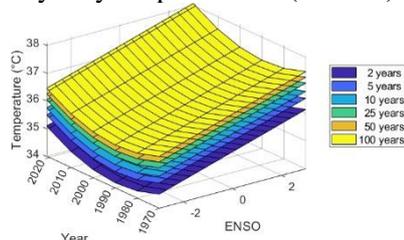


Sydney Airport AMO (066037)

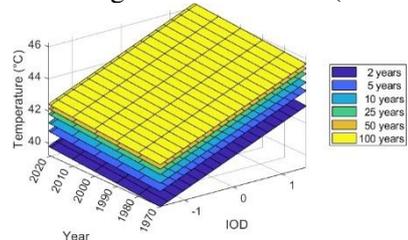


Bathurst Agricultural Station (063005)

Northern Territory

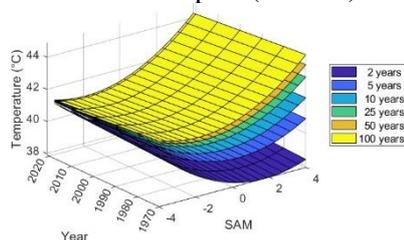


Darwin Airport (014015)

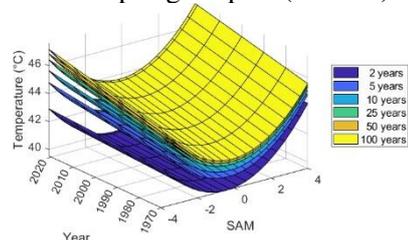


Alice Spring Airport (015590)

Western Australia

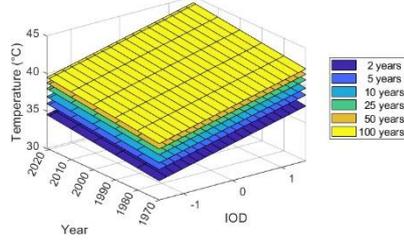


Perth Airport (009021)

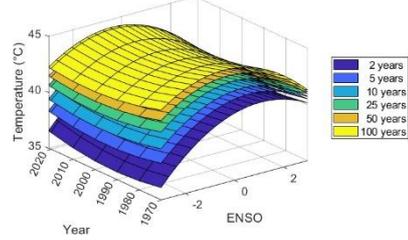


Kalgoorlie-Boulder Airport (012038)

Queensland

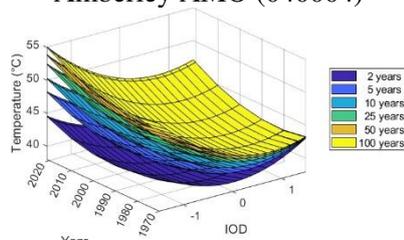


Amberley AMO (040004)

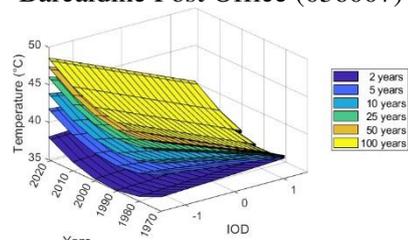


Barcaldine Post Office (036007)

South Australia

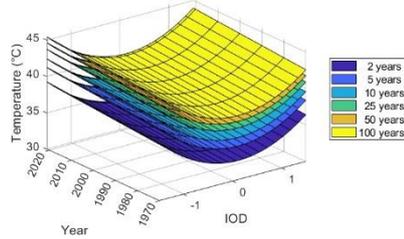


Ceduna AMO (018012)

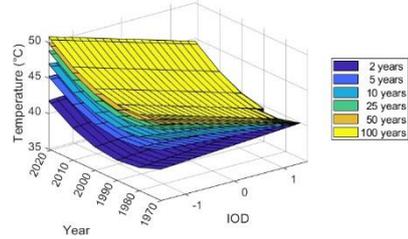


Adelaide Airport (023034)

Victoria



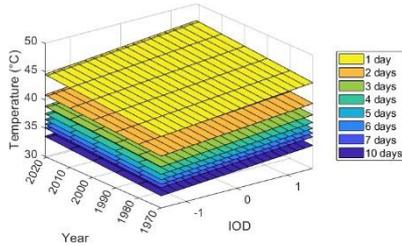
Laverton RAAF (087031)



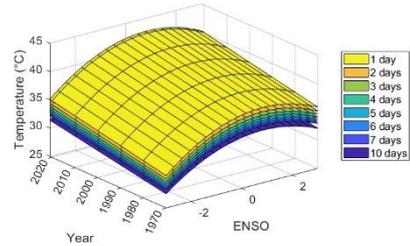
Walpeup Research (076064)

359 Figure 10. Non-stationary TDF surfaces with *Time* and climate driver as covariates of 5-days duration  
360 for 2, 5, 10, 25, 50 and 100 years return periods.

New South Wales

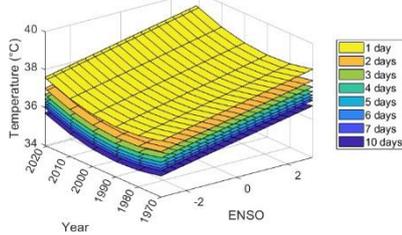


Sydney Airport AMO (066037)

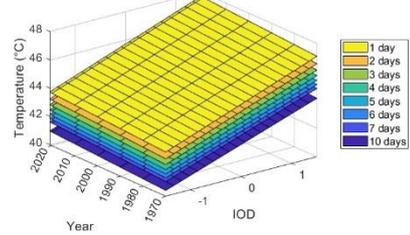


Bathurst Agricultural Station (063005)

Northern Territory

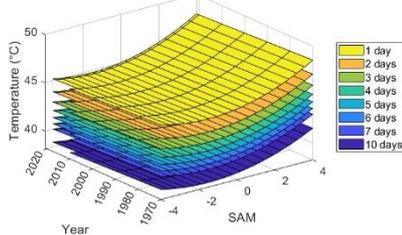


Darwin Airport (014015)

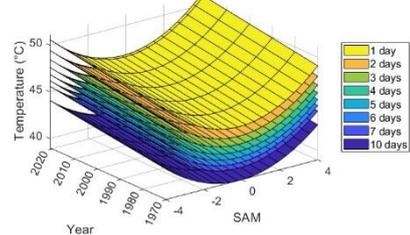


Alice Springs Airport (015590)

Western Australia

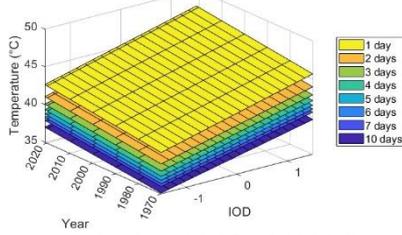


Perth Airport (009021)

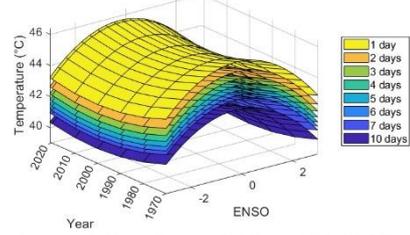


Kalgoorlie-Boulder Airport (012038)

Queensland

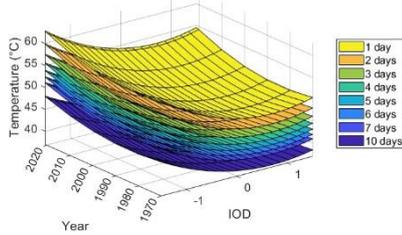


Amberley AMO (040004)

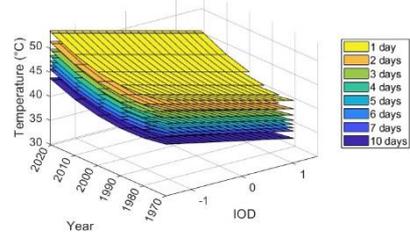


Barcardine Post Office (036007)

South Australia

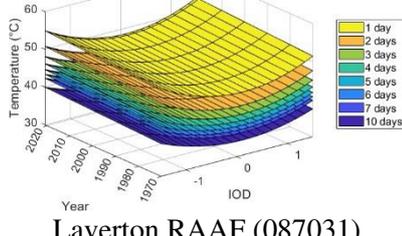


Ceduna AMO (018012)

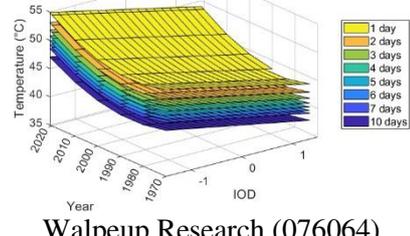


Adelaide Airport (023034)

Victoria



Laverton RAAF (087031)



Walpeup Research (076064)

361 Figure 11. Non-stationary TDF surfaces with *Time* and climate driver as covariates of 50-year return  
362 period for 1, 2, 3, 4, 5, 6, 7 and 10 days durations.

363           ii. *Non-stationary TDF surfaces – Two covariates (Time and Climate Drivers as covariates)*

364 For non-stationary TDF models incorporating two covariates, five variables are required to be present  
365 in the graph. Therefore, in this paper, non-stationary TDF surfaces are presented in two ways – either  
366 the duration is kept fixed and TDF surfaces for different return periods are presented, or return period  
367 is kept fixed and TDF surfaces for different durations are illustrated. Figure 10 shows the TDF surfaces  
368 for 5-day duration and all the return periods considered in this study, whereas Figure 11 presents the  
369 TDF surfaces for 50-year return period and all the considered durations.

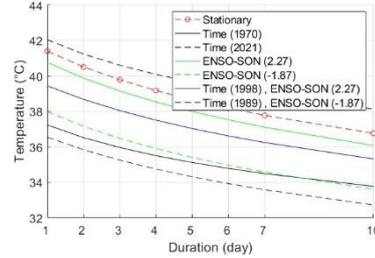
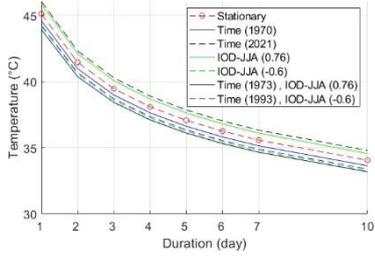
#### 370           **4.2 Impacts of non-stationarity on TDF curves**

371 The temperature quantiles calculated at a given time are significantly affected by incorporating one or  
372 more covariates. Figure 12 shows the graphs of the 50-year quantiles versus the duration for the  
373 stationary TDF model and the non-stationary TDF models considering different *Time* and covariates  
374 representing different scenarios for each station. The non-stationary quantiles for the first and last years  
375 of the study period were computed with *Time* covariate and denoted as *Time* (1970) and *Time* (2021) to  
376 display the quantiles' temporal movement. The years with the highest and lowest values of the selected  
377 seasonal climatic drivers throughout the period 1970 – 2021 were selected for "CD" and "*Time* + CD"  
378 non-stationary models to highlight the influence of the extreme conditions on the estimated quantiles.  
379 All the graphs illustrated in Figure 12 use the model that provides the greatest overall fit.

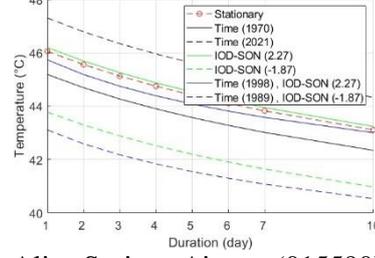
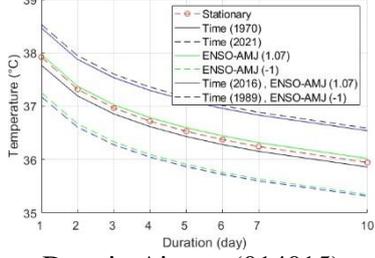
380 In all stations, it was observed that the influence of the duration on the difference in quantile estimation  
381 between the stationary and non-stationary models was negligible. For all the 12 stations, the stationary  
382 model always overestimated and underestimated the return levels in the case of the non-stationary  
383 models with the earliest (1970) and latest (2021) years of the study period as *Time* covariate,  
384 respectively. The stationary model overestimated upto 4.2 °C compare to the former cases, while  
385 underestimated upto 2.3 °C for the latter cases at all 12 stations. The only exception was for the Perth  
386 Airport station, where *Time* covariate had no significant influence. This can be explained by the fact  
387 that the Perth Airport station showed no non-stationarity (Table 1) and no temporal trend in the annual  
388 maximum temperature during the study period (Figure S1).

389 In general, for all the stations, the difference between the stationary and non-stationary models  
390 incorporating the climate drivers as covariates decreased with the return period, as shown in Figures S6  
391 – S10. Also, the stations where IOD was selected as the climate driver showed a significant difference  
392 between stationary and non-stationary models compared to the other stations where ENSO and SAM  
393 were selected as the climate drivers.

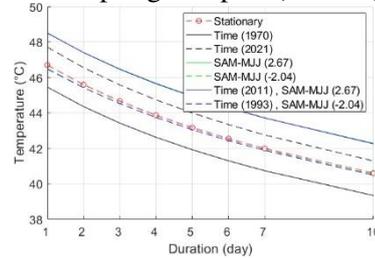
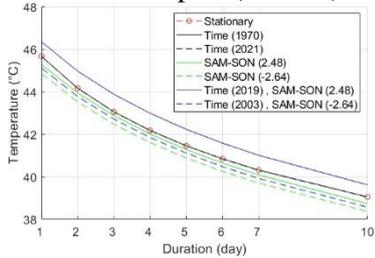
New South Wales



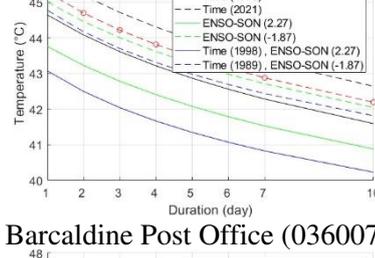
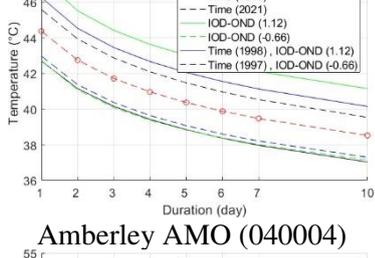
Northern Territory



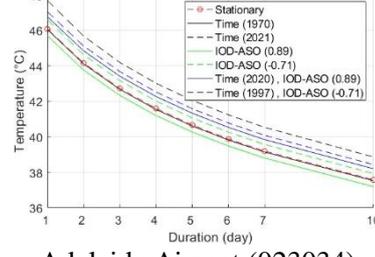
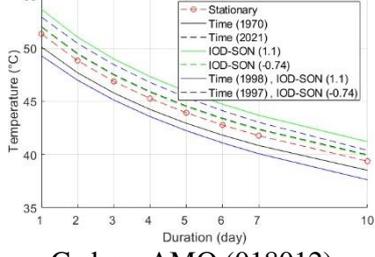
Western Australia



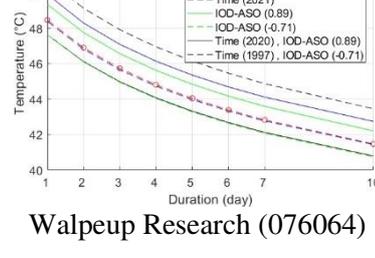
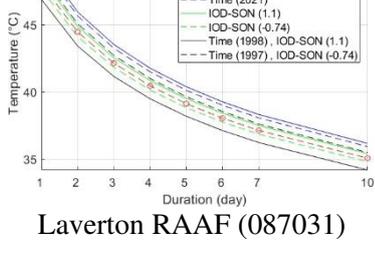
Queensland



South Australia



Victoria



394 Figure 12. Comparison between stationary and non-stationary TDF curves (50-year quantiles).

395 The average difference between the stationary model and the non-stationary models considering the  
396 climate driver alone as covariate ranged from -3.0 to 1.5 °C (+ CD) and -0.6 to 3.4 °C (– CD). In  
397 contrast, this difference became -1.9 to 2.2 °C ("*Time* + CD") and -1.6 to 4.7°C ("*Time* – CD") in the  
398 case of the combination of two covariates – *Time* and selected climate driver.

399 Although *Time* had the overall dominance as a covariate on the computed quantiles for all the return  
400 periods compared to the selected climate indices, different covariates or combinations of them yielded  
401 the maximum quantiles for different return periods and stations.

## 402 **5. CONCLUSION**

403 In this study, the formulation of stationary and non-stationary TDF curves was based on temperature  
404 time series from 82 weather stations located in Australia using the GEV distribution. Initially, the  
405 augmented Dickey-Fuller (ADF) test was conducted to identify non-stationarity in temperature time  
406 series and non-stationarity was found to be present in the data series.

407 The long-range relationships between the seasonal climate drivers ENSO, IOD, SAM and temperature  
408 data from 1969 to 2021 were investigated using the Pearson correlation coefficient to find out the best  
409 covariates in non-stationary TDF models. The magnitude of correlation coefficients of ENSO increases  
410 towards the east of Australia, and these coefficients are significant during SON and DJF seasons. The  
411 annual maximum temperature observed at stations in the southeastern region of Australia, especially in  
412 the inland and coastal region of Victoria and South Australia, showed a significant correlation with IOD  
413 during SON. SAM showed a strong correlation with the annual temperature at the stations located in  
414 the coastal regions of Western Australia

415 Stationary TDF curves showed an increase and decrease in design temperatures with higher return  
416 periods and an increase in duration, respectively. In the case of non-stationary TDF models, the location  
417 and scale parameters were modelled as being dependent on time and climate indicators for the selected  
418 stations. Inclusion of the selected covariates in non-stationary TDF models enhanced goodness-of-fit  
419 compared to the stationary TDF model for the corresponding station. Similar results were found by  
420 Ouarda and Charron (2018a) where the influence of the climate oscillation pattern was found to be more  
421 prominent than the temporal trend. Furthermore, the best goodness-of-fit of the TDF model based on  
422 the AIC values was obtained with a combination of both covariates *Time* and selected climate driver  
423 for most of the stations. These results highlighted the importance of considering the combined effect of  
424 the temporal trend caused by global warming and climate drivers in statistical models used to predict  
425 design temperature.

426 The non-stationary quantiles computed with the first and last years of the study period and with the  
427 highest and lowest values of the selected seasonal climatic drivers were compared with the stationary  
428 model to display the quantiles' temporal movement and the influence of the extreme conditions on the  
429 quantiles. In most cases, the stationary TDF model underestimated the design temperature compared to  
430 the non-stationary model, including *Time* as a covariate. This conveys a crucial message that the non-  
431 stationary framework for designing temperature facilities in Australia could be considered a stronger  
432 option than the traditional stationary approach. In addition, TDF curves developed in this study can be  
433 applied to a range of sectors such as agriculture, health care and energy production and can be a useful  
434 tool for policymakers and planners.

### 435 ***Data availability***

436 The data used in this study can be obtained by contacting the Australian Bureau of Meteorology (by  
437 paying a prescribed fee) ([Australia's official weather forecasts & weather radar - Bureau of Meteorology \(bom.gov.au\)](http://www.bom.gov.au)). Nino3.4 and DMI data used in this study can be obtained freely from NOAA  
438 ([https://psl.noaa.gov/gcos\\_wgsp/Timeseries/Nino34/](https://psl.noaa.gov/gcos_wgsp/Timeseries/Nino34/)) and ([https://psl.noaa.gov/gcos\\_wgsp/Time series /DMI/](https://psl.noaa.gov/gcos_wgsp/Time series /DMI/)); and SAM data has been obtained from NERC (<http://www.nerc-bas.ac.uk/icd/gjma/sam.html>).

## 441 **6. REFERENCE**

442 Adlouni, S. El and Ouarda, T. B. M. J. (2009) 'Joint Bayesian model selection and parameter estimation  
443 of the generalized extreme value model with covariates using birth-death Markov chain Monte Carlo',  
444 *Water Resources Research*, 45(6), pp. 1–11. doi: 10.1029/2007WR006427.

445 Arblaster, J. M. and Alexander, L. V. (2012) 'The impact of the El Nio-Southern Oscillation on  
446 maximum temperature extremes', *Geophysical Research Letters*, 39(20), pp. 2–6. doi:  
447 10.1029/2012GL053409.

448 Bellenger, H. *et al.* (2014) 'ENSO representation in climate models: From CMIP3 to CMIP5', *Climate*  
449 *Dynamics*, 42(7–8), pp. 1999–2018. doi: 10.1007/s00382-013-1783-z.

450 Berghuijs, W. R. *et al.* (2019) 'Growing Spatial Scales of Synchronous River Flooding in Europe',  
451 *Geophysical Research Letters*, 46(3), pp. 1423–1428. doi: 10.1029/2018GL081883.

452 Boschhat, G. *et al.* (2015) 'Large scale and sub-regional connections in the lead up to summer heat wave  
453 and extreme rainfall events in eastern Australia', *Climate Dynamics*, 44(7–8), pp. 1823–1840. doi:  
454 10.1007/s00382-014-2214-5.

- 455 Cai, W. *et al.* (2011) ‘Teleconnection pathways of ENSO and the IOD and the mechanisms for impacts  
456 on Australian rainfall’, *Journal of Climate*, 24(15), pp. 3910–3923. doi: 10.1175/2011JCLI4129.1.
- 457 Cai, W. *et al.* (2014) ‘Increasing frequency of extreme El Niño events due to greenhouse warming’,  
458 *Nature Climate Change*, 4(2), pp. 111–116. doi: 10.1038/nclimate2100.
- 459 Cai, W., Cowan, T. and Raupach, M. (2009) ‘Positive Indian Ocean Dipole events precondition  
460 southeast Australia bushfires’, 36(October), pp. 1–6. doi: 10.1029/2009GL039902.
- 461 Cai, W., Cowan, T. and Sullivan, A. (2009) ‘Recent unprecedented skewness towards positive Indian  
462 Ocean Dipole occurrences and its impact on Australian rainfall’, *Geophysical Research Letters*, 36(11),  
463 pp. 1–5. doi: 10.1029/2009GL037604.
- 464 Cai, W. and Van Rensch, P. (2012) ‘The 2011 southeast Queensland extreme summer rainfall: A  
465 confirmation of a negative Pacific Decadal Oscillation phase?’, *Geophysical Research Letters*, 39(8),  
466 pp. 1–7. doi: 10.1029/2011GL050820.
- 467 Cheng, L. *et al.* (2014) ‘Non-stationary extreme value analysis in a changing climate’, *Climatic Change*,  
468 127(2), pp. 353–369. doi: 10.1007/s10584-014-1254-5.
- 469 Cheng, L. and Aghakouchak, A. (2014) ‘Nonstationary precipitation intensity-duration-frequency  
470 curves for infrastructure design in a changing climate’, *Scientific Reports*, 4, pp. 1–6. doi:  
471 10.1038/srep07093.
- 472 Chowdary, J. S., John, N. and Gnanaseelan, C. (2014) ‘Interannual variability of surface air-temperature  
473 over India: Impact of ENSO and Indian Ocean Sea surface temperature’, *International Journal of*  
474 *Climatology*, 34(2), pp. 416–429. doi: 10.1002/joc.3695.
- 475 Coles, S. (2001) *An introduction to statistical modeling of extreme values*. Springer.
- 476 Cowan, T. *et al.* (2014) ‘More frequent, longer, and hotter heat waves for Australia in the Twenty-First  
477 Century’, *Journal of Climate*, 27(15), pp. 5851–5871. doi: 10.1175/JCLI-D-14-00092.1.
- 478 CSIRO and Australian Government (Bureau of Meteorology) (2020) ‘State of the Climate 2020:  
479 Australia’s changing climate’, *Medicine*, pp. 1–24. Available at: <https://apo.org.au/node/309418>.
- 480 Devi, R., Gouda, K. C. and Lenka, S. (2022) ‘Temperature-duration-frequency analysis over Delhi and  
481 Bengaluru city in India’, *Theoretical and Applied Climatology*, 147(1–2), pp. 291–305. doi:  
482 10.1007/s00704-021-03824-5.

483 Galiatsatou, P. and Iliadis, C. (2022) ‘Intensity-Duration-Frequency Curves at Ungauged Sites in a  
484 Changing Climate for Sustainable Stormwater Networks’, *Sustainability (Switzerland)*, 14(3), pp. 1–  
485 24. doi: 10.3390/su14031229.

486 Ganguli, P. and Coulibaly, P. (2017) ‘Does nonstationarity in rainfall require nonstationary intensity-  
487 duration-frequency curves?’, *Hydrology and Earth System Sciences*, 21(12), pp. 6461–6483. doi:  
488 10.5194/hess-21-6461-2017.

489 Guthrie, M. (2021) *Climate drivers of the South West Land Division*. Available at:  
490 <https://www.agric.wa.gov.au/climate-weather/climate-drivers-south-west-land-division> (Accessed: 9  
491 June 2022).

492 Haddad, K. (2021) ‘Selection of the best fit probability distributions for temperature data and the use  
493 of L-moment ratio diagram method: a case study for NSW in Australia’, *Theoretical and Applied*  
494 *Climatology*, 143(3–4), pp. 1261–1284. doi: 10.1007/s00704-020-03455-2.

495 Halpert, M. S. and Ropelewski, C. F. (1992) ‘Surface Temperature Patterns Associated with the  
496 Southern Oscillation’, *Journal of Climate*, pp. 577–593. doi: 10.1175/1520-  
497 0442(1992)005<0577:stpawt>2.0.co;2.

498 Hendon, H. H., Thompson, D. W. J. and Wheeler, M. C. (2007) ‘Australian rainfall and surface  
499 temperature variations associated with the Southern Hemisphere annular mode’, *Journal of Climate*,  
500 20(11), pp. 2452–2467. doi: 10.1175/JCLI4134.1.

501 Hundecha, Y. *et al.* (2008) ‘A nonstationary extreme value analysis for the assessment of changes in  
502 extreme annual wind speed over the gulf of St. Lawrence Canada’, *Journal of Applied Meteorology and*  
503 *Climatology*, 47(11), pp. 2745–2759. doi: 10.1175/2008JAMC1665.1.

504 IPCC (2018) *Summary for Policymakers. In: Global warming of 1.5°C. An IPCC Special Report on the*  
505 *impacts of global warming of 1.5°C above pre-industrial levels and related global greenhouse gas*  
506 *emission pathways, in the context of strengthening the global response to*, *World Meteorological*  
507 *Organization, Geneva, Switzerland*. Geneva, Switzerland. doi: 10.1017/CBO9781107415324.

508 Katz, R. W., Parlange, M. B. and Naveau, P. (2002) ‘Statistics of extremes in hydrology’, *Advances in*  
509 *Water Resources*, 25(8–12), pp. 1287–1304. doi: 10.1016/S0309-1708(02)00056-8.

510 Koutsoyiannis, D., Kozonis, D. and Manetas, A. (1998) ‘A mathematical framework for studying  
511 rainfall intensity-duration-frequency relationships’, *Journal of Hydrology*, 206(1–2), pp. 118–135. doi:

512 10.1016/S0022-1694(98)00097-3.

513 Kwon, H.-H. and Lall, U. (2016) ‘A copula-based nonstationary frequency analysis for the 2012-2015  
514 drought in California’, *Water Resources Research*, 52(7), pp. 5662–5675. doi:  
515 10.1002/2016WR018959.

516 Liu, L. *et al.* (2014) ‘Indian Ocean variability in the CMIP5 multi-model ensemble: The zonal dipole  
517 mode’, *Climate Dynamics*, 43(5–6), pp. 1715–1730. doi: 10.1007/s00382-013-2000-9.

518 Lorenz, R., Stalhandske, Z. and Fischer, E. M. (2019) ‘Detection of a Climate Change Signal in Extreme  
519 Heat, Heat Stress, and Cold in Europe From Observations’, *Geophysical Research Letters*, 46(14), pp.  
520 8363–8374. doi: 10.1029/2019GL082062.

521 Maher, P. and Sherwood, S. C. (2014) ‘Disentangling the multiple sources of large-scale variability in  
522 Australian wintertime precipitation’, *Journal of Climate*, 27(17), pp. 6377–6392. doi: 10.1175/JCLI-D-  
523 13-00659.1.

524 Meyers, G. *et al.* (2007) ‘The years of El Niño, La Niña and interactions with the tropical Indian Ocean’,  
525 *Journal of Climate*, 20(13), pp. 2872–2880. doi: 10.1175/JCLI4152.1.

526 Min, S. K., Cai, W. and Whetton, P. (2013) ‘Influence of climate variability on seasonal extremes over  
527 Australia’, *Journal of Geophysical Research Atmospheres*, 118(2), pp. 643–654. doi:  
528 10.1002/jgrd.50164.

529 Nicholls, N. (1985) ‘Towards the prediction of major Australian droughts.’, *Australian Meteorological*  
530 *Magazine*, 33, pp. 161–166.

531 Nicholls, N. and Lucas, C. (2007) ‘Interannual variations of area burnt in Tasmanian bushfires:  
532 Relationships with climate and predictability’, *International Journal of Wildland Fire*, 16(5), pp. 540–  
533 546. doi: 10.1071/WF06125.

534 Oliveira, F. N. M. and Ambrizzi, T. (2017) ‘The effects of ENSO-types and SAM on the large-scale  
535 southern blockings’, *International Journal of Climatology*, 37(7), pp. 3067–3081. doi:  
536 10.1002/joc.4899.

537 Omer, A. *et al.* (2020) ‘Natural and anthropogenic influences on the recent droughts in Yellow River  
538 Basin, China’, *Science of the Total Environment*, 704. doi: 10.1016/j.scitotenv.2019.135428.

539 Ouarda, T. B. M. J. and Charron, C. (2018a) ‘Nonstationary Temperature-Duration-Frequency curves’,

540 *Scientific Reports*, 8(1), pp. 1–8. doi: 10.1038/s41598-018-33974-y.

541 Ouarda, T. B. M. J. and Charron, C. (2018b) ‘Nonstationary Temperature-Duration-Frequency curves’,  
542 *Scientific Reports*, 8(1), pp. 1–8. doi: 10.1038/s41598-018-33974-y.

543 Ouarda, T. B. M. J., Charron, C. and St-Hilaire, A. (2020) ‘Uncertainty of stationary and nonstationary  
544 models for rainfall frequency analysis’, *International Journal of Climatology*, 40(4), pp. 2373–2392.  
545 doi: 10.1002/joc.6339.

546 Ouarda, T. B. M. J., Yousef, L. A. and Charron, C. (2019) ‘Non-stationary intensity-duration-frequency  
547 curves integrating information concerning teleconnections and climate change’, *International Journal  
548 of Climatology*, 39(4), pp. 2306–2323. doi: 10.1002/joc.5953.

549 Parker, T. J., Berry, G. J. and Reeder, M. J. (2013) ‘The influence of tropical cyclones on heat waves  
550 in Southeastern Australia’, *Geophysical Research Letters*, 40(23), pp. 6264–6270. doi:  
551 10.1002/2013GL058257.

552 Perkins, S. E., Argüeso, D. and White, C. J. (2015) ‘Relationships between climate variability, soil  
553 moisture, and Australian heatwaves’, *Journal of Geophysical Research: Atmospheres*, 120(16), pp.  
554 8144–8164. doi: 10.1002/2015JD023592.

555 Power, S. *et al.* (1999) ‘Inter-decadal modulation of the impact of ENSO on Australia’, *Climate  
556 Dynamics*, 15(5), pp. 319–324. doi: 10.1007/s003820050284.

557 Power, S. B. *et al.* (2006) ‘The Predictability of Interdecadal Changes in ENSO Activity and ENSO  
558 Teleconnections’, *Journal of Climate*, 19(19), pp. 4755–4771.

559 Risbey, J. S. *et al.* (2009) ‘On the remote drivers of rainfall variability in Australia’, *Monthly Weather  
560 Review*, 137(10), pp. 3233–3253. doi: 10.1175/2009MWR2861.1.

561 Ropelewski, C. F. and Halpert, M. S. (1988) ‘Precipitation Patterns Associated with the High Index  
562 Phase of the Southern Oscillation’, *Journal of Climate*, 2(3), pp. 268–284.

563 Rossi, F. and Villani, P. (1994) ‘A project for regional analysis of floods in Italy’, in Rossi, G.,  
564 Harmancio\uglu, N., and Yevjevich, V. (eds) *Coping with Floods*. Dordrecht: Springer Netherlands, pp.  
565 193–217. doi: 10.1007/978-94-011-1098-3\_11.

566 Saji, N. H. *et al.* (1999) ‘A dipole mode in the tropical Indian ocean’, *Nature*, 401(6751), pp. 360–363.  
567 doi: 10.1038/43854.

- 568 Sarhadi, A. and Soulis, E. D. (2017) 'Time-varying extreme rainfall intensity-duration-frequency curves  
569 in a changing climate', *Geophysical Research Letters*, 44(5), pp. 2454–2463. doi:  
570 10.1002/2016GL072201.
- 571 Sein, K. K., Chidthaisong, A. and Oo, K. L. (2018) 'Observed trends and changes in temperature and  
572 precipitation extreme indices over Myanmar', *Atmosphere*, 9(12). doi: 10.3390/atmos9120477.
- 573 Spinoni, J., Naumann, G. and Vogt, J. V. (2017) 'Pan-European seasonal trends and recent changes of  
574 drought frequency and severity', *Global and Planetary Change*, 148, pp. 113–130. doi:  
575 10.1016/j.gloplacha.2016.11.013.
- 576 Sugahara, S., da Rocha, R. P. and Silveira, R. (2009) 'Non-stationary frequency analysis of extreme  
577 daily rainfall in Sao Paulo, Brazil', *International Journal of Climatology*, 29(9), pp. 1339–1349. doi:  
578 10.1002/joc.1760.
- 579 Suman, M. and Maity, R. (2020) 'Southward shift of precipitation extremes over south Asia: Evidences  
580 from CORDEX data', *Scientific Reports*, 10(1), pp. 1–11. doi: 10.1038/s41598-020-63571-x.
- 581 Thompson, D. W. J., Wallace, J. M. and Hegerl, G. C. (2000) 'Annular Modes in the Extratropical  
582 Circulation . Part II : Trends Author ( s): David W . J . Thompson , John M . Wallace and Gabriele C .  
583 Hegerl Published by : American Meteorological Society Stable URL :  
584 <https://www.jstor.org/stable/10.2307/26244740> REF', 13(5), pp. 1018–1036.
- 585 Turney, C. S. M. *et al.* (2007) 'Quaternary climatic, environmental and archaeological change in  
586 Australasia', *Journal of Quaternary Science*, 22(5), pp. 421–422. doi: 10.1002/jqs.1139.
- 587 Ummenhofer, C. C. *et al.* (2009) 'What causes southeast Australia's worst droughts?', *Geophysical  
588 Research Letters*, 36(4), pp. 1–6. doi: 10.1029/2008GL036801.
- 589 Ummenhofer, C. C. *et al.* (2011) 'Indian and Pacific Ocean influences on southeast Australian drought  
590 and soil moisture', *Journal of Climate*, 24(5), pp. 1313–1336. doi: 10.1175/2010JCLI3475.1.
- 591 Wang, X. L. *et al.* (2013) 'Historical changes in Australian temperature extremes as inferred from  
592 extreme value distribution analysis', *Geophysical Research Letters*, 40(3), pp. 573–578. doi:  
593 10.1002/grl.50132.
- 594 White, C. J. *et al.* (2013) 'On regional dynamical downscaling for the assessment and projection of  
595 temperature and precipitation extremes across Tasmania, Australia', *Climate Dynamics*, 41(11–12), pp.

596 3145–3165. doi: 10.1007/s00382-013-1718-8.

597 Yan, H. *et al.* (2019) ‘Next-Generation Intensity–Duration–Frequency Curves to Reduce Errors in Peak  
598 Flood Design’, *Journal of Hydrologic Engineering*, 24(7), p. 04019020. doi: 10.1061/(asce)he.1943-  
599 5584.0001799.

600 Yan, H. *et al.* (2020) ‘Evaluating next-generation intensity–duration–frequency curves for design flood  
601 estimates in the snow-dominated western United States’, *Hydrological Processes*, 34(5), pp. 1255–  
602 1268. doi: 10.1002/hyp.13673.

603 Yilmaz, A. G. and Perera, B. J. C. (2014) ‘Extreme Rainfall Nonstationarity Investigation and Intensity–  
604 Frequency–Duration Relationship’, *Journal of Hydrologic Engineering*, 19(6), pp. 1160–1172. doi:  
605 10.1061/(asce)he.1943-5584.0000878.

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