

# **The stratigraphic position and the age of the Ordovician organic-rich intervals in the northern Hudson Bay, Hudson Strait and Foxe basins – evidence from graptolites**

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# The stratigraphic position and the age of the Ordovician organic-rich intervals in the northern Hudson Bay, Hudson Strait and Foxe basins – evidence from graptolites

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## Abstract

Graptolites recovered from the organic-rich intervals, previously named the Boas River Formation in the Upper Ordovician succession on Southampton, Akpatok and southern Baffin islands provide a reliable age assessment for the Upper Ordovician petroleum source rocks in the northern Hudson Bay, Hudson Strait and Foxe basins. They are characterised by *Anticostia lata* and *An. hudsoni* in the lower Red Head Rapids Formation on Southampton Island; *An. decipiens* and *Rectograptus socialis* in the lower Foster Bay Formation on Akpatok Island; and *Diplacanthograptus spiniferus* and *Amplexograptus praetypicalis* in the lower Amadjuak Formation on southern Baffin Island. These data suggest that the organic-rich intervals in the northern Hudson Bay and Hudson Strait basins can be correlated to the *Dicellograptus anceps* and *Paraorthograptus pacificus* zones of the upper Katian, and the horizon in the Foxe Basin to the *Diplacanthograptus spiniferus* Zone of the lower Katian. The Boas River Formation is not deemed appropriate to use as it occurs as an organic-rich interbed in different stratigraphic units in different basins; therefore, it is suggested to abandon it as a stratigraphic term.

Keywords: Ordovician organic-rich intervals; graptolites; northern Hudson Bay Basin; Foxe Basin; Hudson Strait Basin

## Introduction

The Hudson Bay, Hudson Strait and Foxe basins are three intracratonic basins in northern Canada (Pinet et al. 2013; Fig. 1). The basins are commonly filled by shallow marine Paleozoic sediments, which are represented by the Paleozoic succession exposed on Southampton, Akpatok and Baffin islands, and the Melville Peninsula. To the south, the Hudson Bay Basin succession is also exposed on Hudson Bay Lowlands in northern Ontario and northeastern Manitoba.

Although these basins have never produced hydrocarbons, recent studies have proved significant hydrocarbon source rock potential in the Upper Ordovician succession (Macauley 1986, 1987; Macauley et al. 1990; Zhang 2008, 2012; Zhang and Mate 2015; Lavoie et al. 2013, 2015; Reyes et al. 2016).

The Upper Ordovician hydrocarbon source rocks were known as the “Boas River shale” or the Boas River Formation, and stratigraphically were interpreted to occur at the same level in the three basins (Heywood and Sanford 1976; Sanford and Grant 1990; 1998, 2000) (Fig. 2). More recently, however, this stratigraphic correlation has been revised based on detailed field investigations and studies of conodont microfossils on Southampton, southern Baffin and Akpatok islands, indicating that the number, age and stratigraphic position of the organic-rich intervals vary from basin to basin (Zhang 2008, 2011, 2012, 2018) (Fig. 2). It is worth noting that the type locality of the Boas River Formation was also revisited during these field investigations, and it was found that it did not provide sufficient stratigraphic support to maintain the name Boas River Formation as a valid stratigraphic unit (Zhang 2008).

This paper reports on the graptolites from outcrops on Southampton and southern Baffin islands and from rubble on Akpatok Island, provides further support for the conclusion of Zhang (2008, 2011, 2012, 2018) regarding the stratigraphic position and relative age for the organic-

rich intervals in the three intracratonic basins in northern Canada from the graptolite point of view, and proposes abandoning the Boas River Formation as a stratigraphic unit.

### **Organic-rich intervals in the Upper Ordovician succession in the northern Hudson Bay, Hudson Strait and Foxe Basins**

The Upper Ordovician strata are dominated by various shallow marine carbonates in which thin, organic-rich argillaceous limestone or limy shale intervals are present. The stratigraphy is divided upward into the Bad Cache Rapids Group, the “Boas River shale”, the Churchill River Group and the Red Head Rapids Formation in the northern Hudson Bay Basin as exposed on Southampton Island (Heywood and Sanford 1976) (Fig. 2). A similar carbonate-dominated stratigraphic succession was partitioned into the Amadjuak, Boas River, Akpatok, and Foster Bay formations in Foxe Basin and Hudson Strait Basin cropping out on southern Baffin Island, Melville Peninsula and Akpatok Island (Sanford and Grant 1990, 1998, 2000) (Fig. 2). Generally, the Bad Cache Rapids and Churchill River groups and their equivalent Amadjuak and Akpatok formations are limestone units; the Red Head Rapids Formation and its counterpart Foster Bay Formation are dolomitic limestone and dolostone units (Heywood and Sanford 1976; Sanford and Grant 1990, 1998, 2000; Zhang 2008, 2011, 2012, 2013, 2018). The “Boas River shale” or the Boas River Formation is a thin, organic-rich unit; its stratigraphic position and age have been significantly revised (Zhang 2008, 2011, 2012, 2018), since the unit was first introduced (Heywood and Sanford 1976) and formally used (Sanford and Grant 1990, 1998, 2000).

### **Organic-rich intervals on Southampton Island, the northern Hudson Bay Basin**

The “Boas River shale” is a thin unit of organic-rich brown laminated argillaceous limestone and shale more than one meter in thickness (Zhang 2008). It was first recognised in a stream cut of Boas River in the central part of Southampton Island, and its stratigraphic position was placed between the Bad Cache Rapids and the Churchill River groups (Heywood and Sanford 1976); and then formally named the Boas River Formation (Sanford and Grant 1990). The “Boas River shale” has yielded the graptolite *Glyptograptus hudsoni* Jackson and trilobite *Pseudogygites* sp. (Jackson 1971). Several decades later, three thin organic-rich intervals, >1 m, 0.3–0.4 m and 0.35–0.5 m in upward order, were identified in a *circa* 17 m interval in unit 1 of the lower part of the Red Head Rapids Formation in a creek in the Cape Donovan area, north-eastern Southampton Island (Zhang 2008). These excellent petroleum source rocks yield maximum and average total organic carbon (TOC) of 17.3% and 9.8%, 34.1% and 22.4%, and 31% and 18.3% in the lower, middle and upper intervals, respectively (Zhang 2008). Based on the unique laminated nature of organic-rich argillaceous limestone, the trilobites and the conodonts, the lower of the three organic-rich intervals was correlated to the “Boas River shale” (Zhang 2008, 2011). Therefore, the “Boas River shale” is located within the unit 1 of the lower Red Head Rapids Formation, rather than between Bad Cache Rapids and Churchill River groups (Fig. 2). The “Boas River shale” together with the Red Head Rapids Formation was biostratigraphically correlated with the conodont *Aphelognathus divergens* Zone of the upper Katian (Zhang 2011) (Fig. 3).

### **Organic-rich interval on Akpatok Island, the Hudson Strait Basin**

A bituminous, brown, argillaceous limestone with a thickness of 4.5 m was described above the high sea level at section II, i.e. locality 3A in Zhang (2018, fig. 3B), in a gully to the south of the Premium Homestead Akpatok L-26 well location; and it was correlated to the “Boas River shale” between the Bad Cache Rapids and the Churchill River groups by Workum et al. (1976). Recently, field investigation at the same locality proved that this bituminous, brown, argillaceous limestone is not an outcrop, but a pile of rubble, and that the bituminous, brown, argillaceous limestone only occurs at the bottom of the pile (Zhang and Mate 2015; Zhang 2018). The rubble contains maximum and average TOC 4.19% and 3.11% (Zhang and Mate 2015). According to the conodonts and graptolite preserved in the organic-rich limestone rubble, Zhang (2018) concluded that the organic-rich rubble is most likely from the Foster Bay Formation, rather than from between the Amadjuak and Akpatok formations, and biostratigraphically the Foster Bay Formation is correlated with the conodont *Aphelognathus divergens* Zone of the upper Katian (Fig. 3).

### **Organic-rich interval on southern Baffin Island, the Foxe Basin**

Macauley (1987) described an organic-rich interval from the Jordan River section on southern Baffin Island. The interval was designated as the Boas River Formation without formal diagnosis and description, and its stratigraphic position was suggested to be between the Amadjuak and Akpatok formations by Sanford and Grant (1990, 1998, 2000) (Fig. 2). However, Riva (2000) in the same volume (McCracken and Bolton 2000) as Sanford and Grant (2000) demonstrated that the stratigraphic position of this organic-rich interval is near the base of Amadjuak Formation. Zhang (2012) visited the Jordan River section and reported a brown-black

shale with a thickness of less than two meters and containing maximum and average TOC 12.97% and 7.8% in the lower Amadjuak Formation. Thus, both Riva (2000) and Zhang (2012) concur that the organic-rich interval on southern Baffin Island in the Foxe Basin is in the lower Amadjuak Formation instead of between the Amadjuak and Akpatok formations (Fig. 2). In the limestone beneath this organic-rich interval at the Jordan River section, the typical Amadjuak *Fisherites-Maclurites* Fauna was found (Zhang 2012), and one specimen of conodont *Amorphognathus superbus* Rhode was recovered (McCracken 2000), suggesting a correlation of the organic-rich interval in the Foxe Basin with the lower Katian conodont *Amorphognathus superbus* Zone (Fig. 3).

## **Graptolites in the Upper Ordovician organic-rich intervals on Southampton, Akpatok and southern Baffin islands**

### **Graptolites on Southampton Island, the northern Hudson Bay Basin**

As mentioned above, the lower of the three organic-rich intervals identified in unit 1 of the lower Red Head Rapids Formation at Cape Donovan section on Southampton Island was correlated to the “Boas River shale” (Zhang 2008, 2011). This correlation was based on the unique laminated nature of organic-rich argillaceous limestone and the occurrence of the trilobite *Pseudogygites hudsoni* Ludvigsen in the lower organic-rich interval at the Cape Donovan section and the “Boas River shale” at the Boas River section (Zhang 2008, fig. 9), which is the only interval where graptolites occur.

Jackson (1971) described a new species of graptolite, *Glyptograptus hudsoni* (subsequently transferred to *Anticostia* Stewart and Mitchell 1997), from a slab of limestone of

unknown location and stratigraphic position on Southampton Island, but its stratigraphic position was referred to as the “oil shale horizon” of latest Ordovician age (Nelson and Johnson 1966).

More recently, on Southampton Island, graptolites were found in limestone interbeds within the “Boas River shale”, i.e. the lower of the three organic-rich intervals (Zhang, 2008, 2011). The limestone interbeds are rich in graptolites, but contain only *Anticostia lata* (Elles and Wood 1906) (Figs. 4A, 4B) (senior synonym of *Climacograptus inuiti* Cox, 1933a; see Štorch et al. 2011 for details of genus transfer). *Anticostia hudsoni*, however, is not present among the recovered specimens.

Riva (1988) noted that *Anticostia lata*, *An. prominens*, and *An. cf. An. hudsoni* occurred together in the *An. prominens* Zone in the Vauréal Formation on Anticosti Island. Thus, the slab of limestone from an unknown locality on Southampton Island containing *An. hudsoni* (Jackson 1971) was most likely from the same stratigraphic level as the limestone interbeds yielding the graptolite *An. lata* in the lower organic-rich interval of the Red Head Rapids Formation reported by this study. Moreover, the presence of both *An. lata* and *An. hudsoni* provided a biostratigraphic connection between the lower organic-rich interval on Southampton Island and the *An. prominens* Zone on Anticosti Island.

### **Graptolites on Akpatok Island, the Hudson Strait Basin**

A “chocolate-coloured limestone” bed yielding the graptolite *Anticostia lata* described by Cox (1933a) occurs at an elevation between 143 and 144 m in a creek at the southern end of Akpatok Island (Cox 1933b). Unfortunately, there was no chance to visit this locality during the field investigations (Zhang and Mate 2015, Zhang 2018).

The graptolites reported here from Akpatok Island were found in the bituminous, argillaceous limestone rubble at the same locality described as section II by Workum et al. (1976). Most graptolites found in the samples are very fragmentary, and only few good specimens are identifiable. This graptolite fauna includes *Amplexograptus decipiens* Riva (subsequently transferred to the genus *Anticostia* Stewart and Mitchell 1997) (Figs. 4E, 4F) and *Rectograptus socialis* (Lapworth) (Figs. 4C, 4D). On Anticosti Island, the former is associated with *Anticostia lata* in the *An. prominens* Zone (Riva 1988; fig. 1), and the latter's highest occurrence is in the lower part of the zone (McLaughlin et al. 2016; fig. 2). Therefore, the bituminous, argillaceous limestone rubble recovered by this study could be correlated with the "chocolate-coloured limestone" bed of Cox (1933b), an interval in the Upper Ordovician Foster Bay Formation on Akpatok Island (Zhang 2018). As a result, both the bituminous, argillaceous limestone rubble reported herein and the "chocolate-coloured limestone" band of Cox (1933b) can be correlated with the *Anticostia prominens* Zone on Anticosti Island (Riva 1988).

The *Anticostia prominens* Zone on Anticosti Island was related to the *Dicellograptus anceps* Zone of Scotland and the *Paraorthograptus pacificus* Zone of Siberia, Kazakhstan and China (Riva 1988). This correlation is partially followed by this present study, i.e. the lower organic-rich interval of the Red Head Rapids Formation on Southampton Island and the organic-rich interval of the Foster Bay Formation on Akpatok islands could be correlated to an uncertain interval within the *Dicellograptus anceps* and *Paraorthograptus pacificus* zones, most likely within the *Paraorthograptus pacificus* Zone (Fig. 3). This is based on 1) the fauna reported above, particularly the presence of *Anticostia lata* and *An. hudsoni* in the lower organic-rich interval of the Red Head Rapids Formation on Southampton Island and of *An. lata* and *An. decipiens* in the organic-rich interval of the Foster Bay Formation on Akpatok islands, and 2) the

correlation of these organic-rich intervals to the conodont *Aphelognathus divergens* Zone (Zhang 2011, 2018), which is, in turn, broadly correlated to the *Paraorthograptus pacificus* Zone (Cooper and Sadler 2012).

Reports on the occurrence of *Anticosti hudsoni* and *An. decipiens* are rare, and their worldwide stratigraphic distribution is unclear. Globally, the stratigraphic distribution of *An. lata* varies from section to section. *Anticostia lata* is generally seen in the *Dicellograptus anceps* Zone, such as in north-central Nevada, USA (Štorch et al. 2011) and on the Canadian Arctic Islands (Melchin 1987), but its highest occurrence has been found to be in the *Normalograptus extraordinarius* Zone of the Hirnantian in north-central Nevada, USA (Štorch et al. 2011) and China (Chen et al. 2000a, 2005). In some cases *Anticostia lata* is restricted to the *Dicellograptus complexus* Subzone of the *Dicellograptus anceps* Zone, as seen in Scotland (Williams 1982a, 1982b, 1987), or the *Dicellograptus ornatus* Zone in Idaho, USA (Mitchell et al. 2003). Therefore, based on the known graptolite species from Southampton and Akpatok islands, the biostratigraphic position of the organic-rich interval in the lower Red Head Rapids and Foster Bay formations could extend above the *Paraorthograptus pacificus* Zone, but it is definitely much higher than that in the lower Amadjuak Formation on southern Baffin Island (see discussion below).

### **Graptolites on southern Baffin Island, the Foxe Basin**

Two graptolite species, *Climacograptus (Diplacanthograptus) spiniferus* Ruedeman and *Amplexograptus praetypicalis* Riva, were reported from the organic-rich interval in the lower Amadjuak Formation on southern Baffin Island; based on the graptolite fauna, the organic-rich interval was correlated with lower *Diplacanthograptus spiniferus* Zone of latest Mohawkian–early Edenian age (Riva 2000).

The graptolites found during the recent field investigation (Zhang 2012) are well preserved on the weathered bedding plane of black shale in outcrops of organic-rich interval of the lower Amadjuak Formation. All specimens belong to one of the two species reported by Riva (2000), i.e. *Amplexograptus praetypicalis* (Figs. 4G, 4H). Therefore, this present study agrees with the stratigraphic correlation of the organic-rich interval in the Amadjuak Formation with the lower *Diplacanthograptus spiniferus* Zone suggested by Riva (2000); the zone is correlated to the lower Katian by the GTS of 2012 (Fig. 3).

Regionally, the graptolites found in the organic-rich interval in the Amadjuak Formation on southern Baffin Island suggest a biostratigraphic connection to the lower Utica Shale in St. Lawrence Lowlands of Quebec (Riva 1969) and the Indian Castle Member of the Utica Shale in New York State (Goldman et al. 1999), in which *Diplacanthograptus spiniferus* is omnipresent.

Outside North America, *Amplexograptus praetypicalis* was not only reported from the *Diplacanthograptus spiniferus* Zone, and possibly from a zone below, the *Diplacanthograptus lanceolatus* Zone, in Xinjiang, China; however, the distributions of *Di. spiniferus* and *Di. lanceolatus* do not overlap in that region (Chen et al. 2000b). Therefore, it is only natural to correlate the organic-rich interval in the lower Amadjuak Formation on southern Baffin Island yielding both *Am. praetypicalis* and *Di. spiniferus* to the *Diplacanthograptus spiniferus* Zone.

### **Proposal to abandon the name Boas River Formation**

As mentioned above, the name “Boas River shale” was introduced by Heywood and Sanford (1976) for a thin unit of organic-rich brown shale and laminated argillaceous limestone in a stream cut of the Boas River on Southampton Island. Afterwards, it was formally named Boas River Formation (Sanford and Grant 1990). Since then, the name Boas River Formation has

been widely used for the Upper Ordovician organic-rich units in the northern Hudson Bay, Foxe Basin and Ungava Bay areas (Sanford and Grant 1990, 1998, 2000), as well as in Hudson Bay Lowlands (Armstrong and Lavoie 2010a, 2010b; Lavoie et al. 2013; McCracken et al. 2013; Hahn et al. 2016, 2017; Nicolas and Armstrong 2017). All the known localities with the exposed Upper Ordovician organic-rich intervals in these areas, except for Hudson Bay Lowlands, were visited by Zhang (2008, 2011, 2012, 2018; Zhang and Mate 2015). Various stratigraphic problems have become apparent in using the Boas River Formation in the region, and the name was not used in recent studies (Zhang 2008, 2011, 2012, 2018). It is here suggested to abandon this stratigraphic term in the areas covered by this study. The main points are:

- Type locality

At the type locality of the Boas River Formation, the exposure is very poor (Zhang 2008, figs. 8a and 8b). The strata above the Boas River Formation are eroded off, and those below not exposed. The type section does not show a stratigraphic contact between the Bad Cache Rapids and the Churchill River groups as interpreted by Heywood and Sanford (1976).

- Lithostratigraphy

- Three oil shale intervals were identified in unit 1, thin-layered and laminated argillaceous limestone interbedded with oil shales, of the Red Head Rapids Formation in a well exposed stream cut section in the Cape Donovan area on Southampton Island (Zhang 2008, 2011). Based on its unique laminated lithology and distinctive trilobite and conodonts, the Boas River Formation was correlated to the lower of the three organic-rich intervals in unit 1 of the Red Head Rapids Formation (Zhang 2008, 2011).

- The stratigraphic position of the Boas River Formation is in the lower Amadjuak Formation, right above the lowest occurrence of the distinct Amadjuak *Fisherites-Maclurites* Fauna on southern Baffin Island (Zhang 2012).
- The outcrop of the Boas River Formation on Akpatok Island was not found by the present authors. However, the brown organic-rich argillaceous limestone rubble, representing the Boas River Formation, can be correlated to the “chocolate-coloured limestone” described by Cox (1933a) at an elevation between 143 and 144 m on the island (Cox 1933b). This elevation is within the lower Foster Bay Formation, as the elevation can be used to identify the stratigraphic unit in the area where the strata are horizontally distributed (Zhang 2018). Thus, the Boas River Formation on Akpatok Island lies within the lower Foster Bay Formation (Zhang 2018).

Therefore, lithostratigraphically the Boas River Formation is an organic rich interval occurring in different mappable stratigraphic units in different areas; and it only represents a brief facies change in the homogeneous unit. As initially defined, it cannot be mapped as an independent stratigraphic unit.

- Biostratigraphy

The Boas River Formation or the organic-rich interval is typically diachronous in the northern Hudson Bay, Hudson Strait and Foxe basins, although a mappable unit can be with diachronism. Based on the graptolite and conodont faunas, the Boas River Formation on Southampton and Akpatok islands can be related to the upper Katian graptolite *Dicellograptus anceps* and *Paraorthograptus pacificus* zones and conodont *Aphelognathus divergens* Zone, but on southern Baffin Island, it can be correlated to the lower Katian graptolite *Diplacanthograptus spiniferus* Zone and conodont *Amorphognathus superbus* Zone (Fig. 4).

- Hudson Bay Lowlands

In the Hudson Lowlands (northern Ontario and northeastern Manitoba), although beyond the coverage of this paper, the stratigraphic control of the Boas River Formation needs a discussion here. It has been located between the Bad Cache Rapids and Churchill River groups for a long time (Sanford and Grant 1990, 1998; Armstrong and Lavoie 2010a, 2010b; Lavoie et al. 2013; Nicolas and Armstrong 2017), but most recently, it has been correlated to the lower Bad Cache Rapids Group (Hahn et al. 2016, 2017). This seems a reasonable correlation, because some conodonts, such as *Periodon grandis* (Ethington) and *Phragmodus undatus* Branson and Mehl were reported from the Boas River Formation in drilling cores in Winisk River area in northern Ontario (McCracken 1990); and these species are restricted to the Bad Cache Rapids Group of the lower Katian in the northern Hudson Bay, Hudson Strait and Foxe basins (Zhang, 2011, 2013, 2018). However, the conodonts recovered from the Boas River Formation at Asheweig River section in northern Ontario include *Rhipidognathus symmetricus* Branson, Mehl and Branson, *Amorphognathus duftonus* Rhodes and other relatively long-lasting species (McCracken et al. 2013). *Rhipidognathus symmetricus* was typically recovered from the dolomitic facies Red Head Rapids Formation (Zhang 2011) and its equivalent Foster Bay Formation (Zhang 2013, 2018) in the northern Hudson Bay, Foxe and Hudson Strait basins, and the local *Rhipidognathus symmetricus* Zone was correlated with North American Midcontinent conodont *Aphelognathus divergens* Zone of the upper Katian (Zhang 2011, 2013, 2018) in these areas. Thus, based on the known conodonts from northern Ontario (McCracken 1990; MacCracken et al. 2013), the Boas River Formation would be at two different stratigraphic levels.

## Conclusion

The organic-rich intervals previously named the Boas River Formation, interpreted as being of the same stratigraphic position and the same age, are found in the Upper Ordovician succession exposed on Southampton (northern Hudson Bay Basin), Akpatok (Hudson Strait Basin) and southern Baffin (Foxe Basin) islands. They are, however, at the different stratigraphic position and of the different age, as proven by field observations and the conodonts from these intervals (Zhang 2008, 2011, 2012, 2018) and graptolites reported herein. Lithostratigraphically, the organic-rich intervals are not mappable units, but are intercalated within different mappable stratigraphic units in the areas cited above. In the northern Hudson Bay and Hudson Strait basins, the intervals are in the lower Red Head Rapids Formation, or in its equivalent lower Foster Bay Formation, although in the Foxe Basin the interval is found in the lower Amadjuak Formation. Therefore, the name Boas River Formation is not deemed appropriate for usage in the three basins of northern Canada. In addition, biostratigraphically, the organic-rich intervals are diachronous. In the northern Hudson Bay and Hudson Strait basins, the organic-rich intervals can be correlated to the upper Katian *Paraorthograptus pacificus* Zone; and that in Foxe Basin to the lower Katian *Diplacanthograptus spiniferus* Zone.

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## Figure Captions

**Figure 1.** Northern part of the Phanerozoic Hudson Bay Basin and adjacent Hudson Strait and Foxe basins (modified from Wheeler et al. 1997). Red triangles point the locations of organic-rich intervals in the Upper Ordovician succession.

**Figure 2.** Stratigraphic nomenclature of Hudson Bay, Foxe, and Hudson Strait basins. Black strips represent organic-rich intervals (unscaled).

**Figure 3.** The Upper Ordovician organic-rich intervals in northern Hudson Bay, Hudson Strait and Foxe basins and their correlation with the Ordovician chronostratigraphic and biostratigraphic framework adopted from Cooper and Sadler (2012). The black bars represent the stratigraphic position of the Upper Ordovician organic-rich intervals (unscaled). Ea: Eastonian; Bo: Bolindian.

**Figure 4.** Graptolites from Southampton, Akpatok, and southern Baffin islands. (A) and (B) *Anticostia lata* (Elles and Wood) from lower organic-rich interval in the lower Red Head Rapids Formation on Southampton Island with (B) showing proximal end, GSC139805 and GSC139806. (C) and (D) *Rectograptus socialis* (Lapworth) from brown organic-rich argillaceous limestone rubble of the lower Foster Bay Formation on Akpatok Island with (D) showing proximal end, GSC139807 and GSC139808. (E) and (F) *Anticostia decipiens* (Riva) from brown organic-rich argillaceous limestone rubble of the lower Foster Bay Formation on Akpatok Island, with (F) showing enlargement of black box in (E), GSC139809. (G) and (H) *Amplexograptus*

*praetypicalis* Riva, (G) full rhabdosome from the organic-rich interval of the lower Amadjuak Formation on southern Baffin Island, with (H) showing enlargement of black box in (G), GSC139810. Black bars represent 1 mm.



Figure 1

174x139mm (300 x 300 DPI)

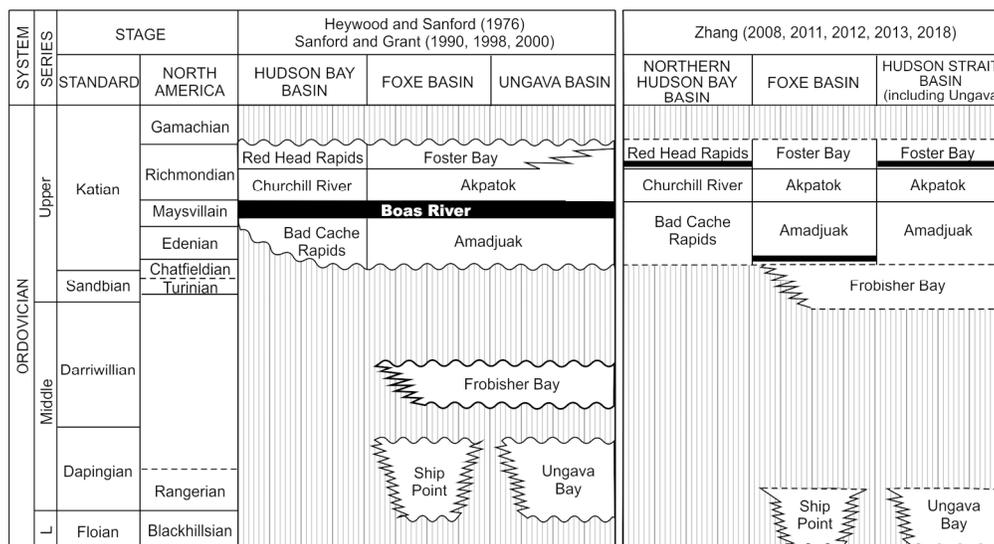


Figure 2

160x87mm (300 x 300 DPI)

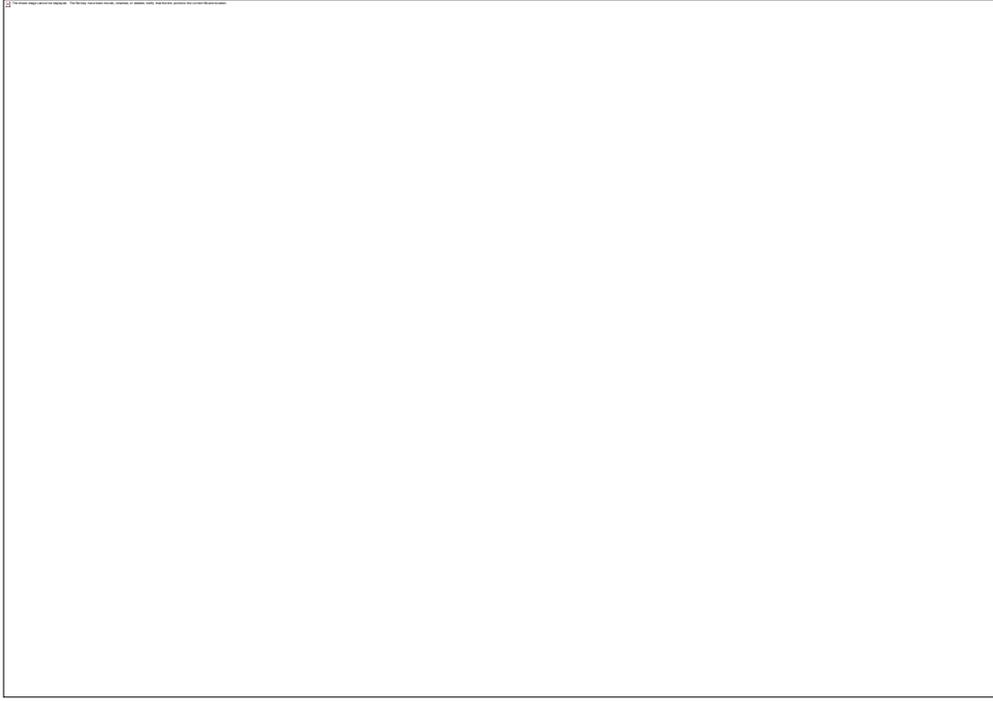


Figure 3

137x96mm (300 x 300 DPI)



Figure 4

159x195mm (300 x 300 DPI)