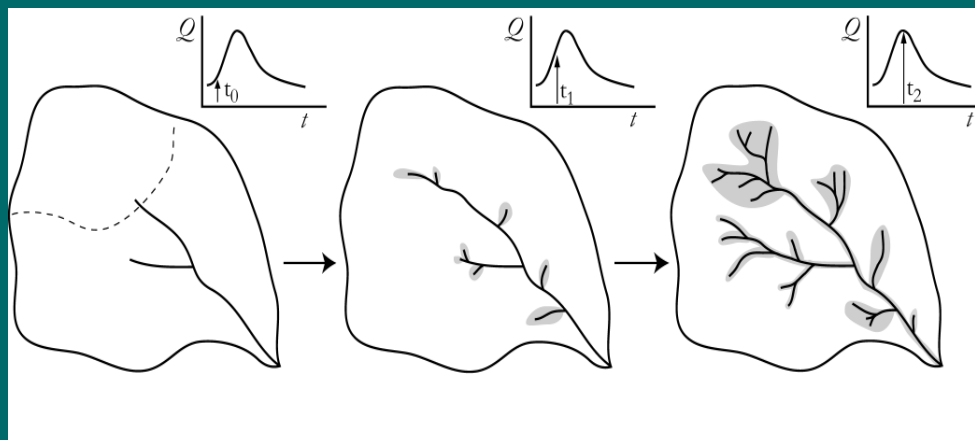

Coupled modeling of groundwater/surface water interactions: successes and challenges from recent applications



Claudio Paniconi

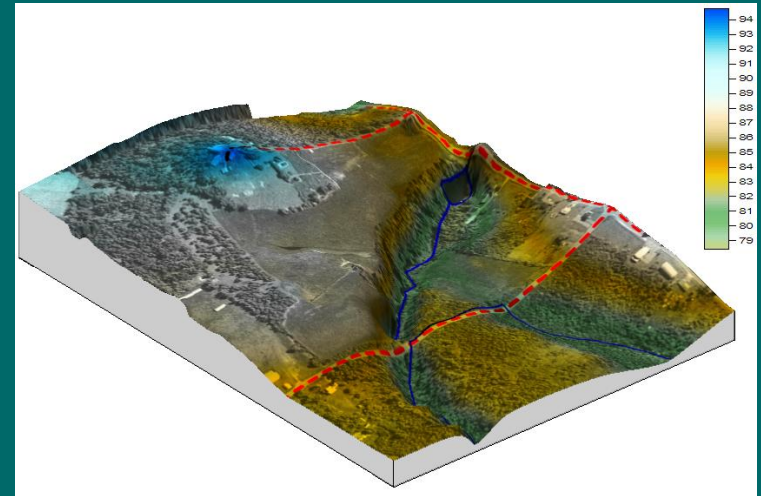
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Context

Proper understanding and representation of hydrosphere interactions (between the atmosphere, land surface, soil zone, aquifers, rivers/lakes, and vegetation) is increasingly relevant to climate prediction, environmental protection, and water management

We are at a crossroads in hydrological modeling:

- models (of all flavors) are being integrated across many disciplines and over multiple scales, and they are being intercompared
- better datasets are increasingly being made available (for hypothesis testing and model validation) that provide observations (on the ground, airborne, and from space) of more processes, in more detail, and at higher accuracy
- computational boundaries are continually being pushed (cost and capabilities of systems, efficiency and robustness of algorithms), for easier and more effective data analysis and process simulation

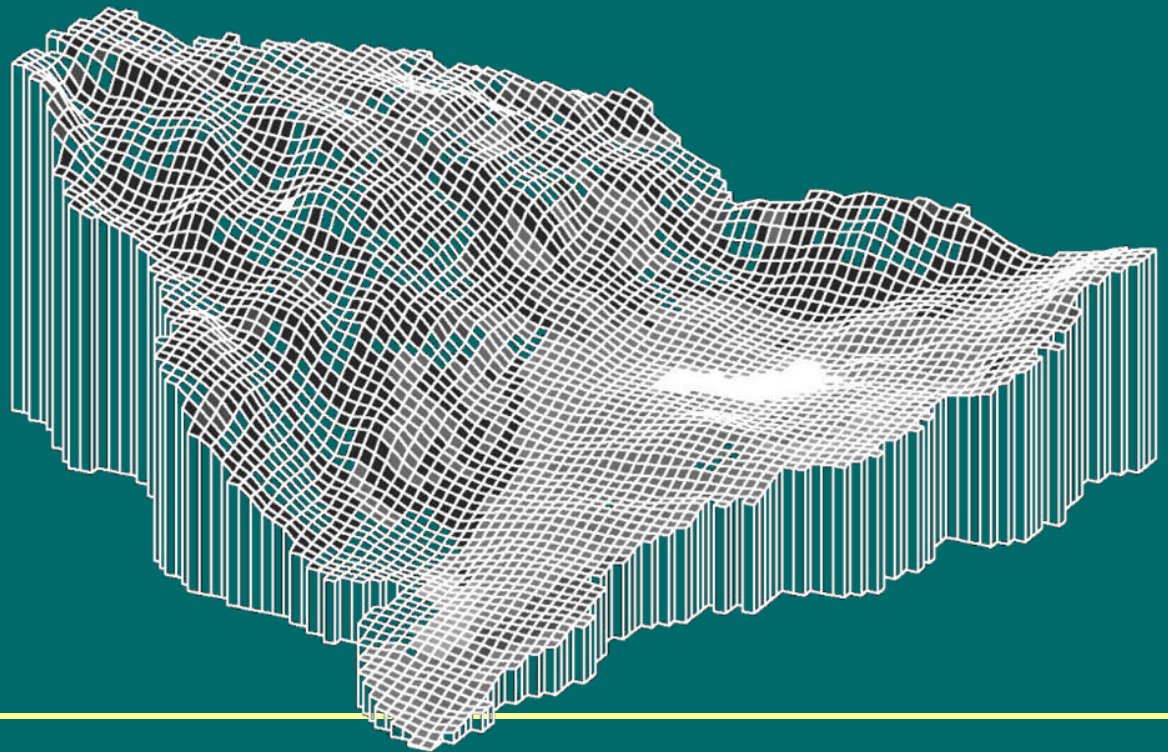


Outline

CATHY (CATchment HYdrology) model description

Some recent studies (successes and challenges)

Extensions and evolution of the model



CATHY (CATchment HYdrology) model description

$$\sigma(S_w) \frac{\partial \psi}{\partial t} = \nabla \cdot [K_s K_{rw}(S_w) (\nabla \psi + \eta_z)] + q_s(h) \quad (1)$$

$$\frac{\partial Q}{\partial t} + c_k \frac{\partial Q}{\partial s} = D_h \frac{\partial^2 Q}{\partial s^2} + c_k q_L(h, \psi) \quad (2)$$

σ	general storage term [1/L]: $\sigma = S_w S_s + \phi(dS_w/d\psi)$	z	vertical coordinate +ve upward [L]
S_w	water saturation = θ/θ_s [1]	q_s	subsurface equation coupling term (more generally, source/sink term) [L ³ /L ³ T]
θ	volumetric moisture content [L ³ /L ³]	h	ponding head (depth of water on surface of each cell) [L]
θ_s	saturated moisture content [L ³ /L ³]	s	hillslope/channel link coordinate [L]
S_s	specific storage [1/L]	Q	discharge along s [L³/T]
ϕ	porosity (= θ_s if no swelling/shrinking)	c_k	kinematic wave celerity [L/T]
ψ	pressure head [L]	D_h	hydraulic diffusivity [L ² /T]
t	time [T]	q_L	surface equation coupling term (overland flow rate) [L³/LT]
K_s	saturated conductivity tensor [L/T]		
K_{rw}	relative hydraulic conductivity [1]		
η_z	zero in x and y and 1 in z direction		

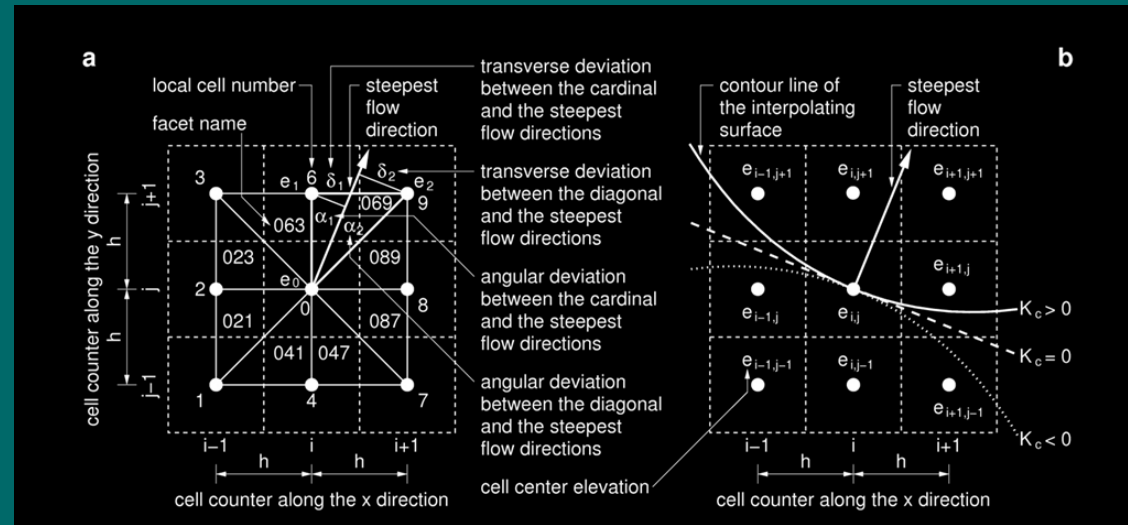
(1) Paniconi & Wood, *Water Resour. Res.*, 29(6), 1993 ; Paniconi & Putti, *Water Resour. Res.*, 30(12), 1994

(2) Orlandini & Rosso, *J. Hydrologic Engrg.*, ASCE, 1(3), 1996 ; Orlandini & Rosso, *Water Resour. Res.*, 34(8), 1998

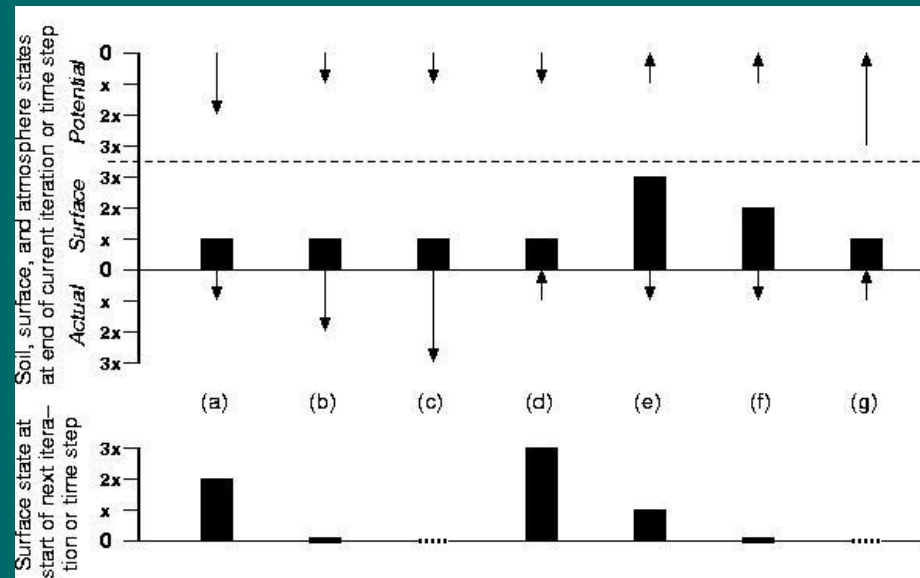
(1)+(2) Putti & Paniconi, *CMWR Proceedings*, 2004; Camporese, Paniconi, Putti, & Orlandini, *Water Resour. Res.*, 46(W02512), 2010

Main features of the model

Path-based description of surface flow across the drainage basin;
 several options for identifying flow directions, for separating channel cells from hillslope cells (same governing equation), and for representing stream channel hydraulic geometry.



The coupling term for the model is computed as the balance between atmospheric forcing (rainfall and potential evaporation) and the amount of water that can actually infiltrate or exfiltrate the soil. This threshold-based boundary condition switching partitions potential fluxes into actual fluxes and changes in surface storage.



Subsurface flow module

Various functional forms for $S_w(\psi)$ and $K_{rw}(\psi)$

Heterogeneities (K_{sx} , K_{sy} , K_{sz} , S_s , ϕ) by "zone" and by layer

DEM-based (uniform) grid or user-defined (nonuniform) surface grid input

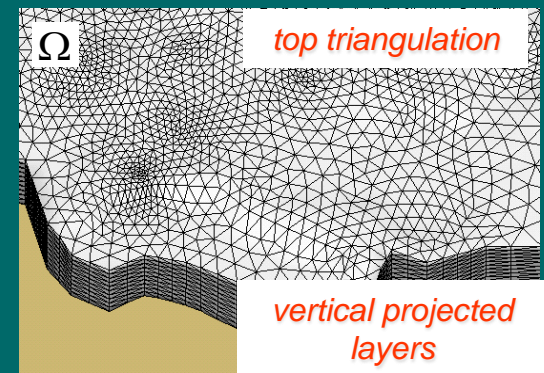
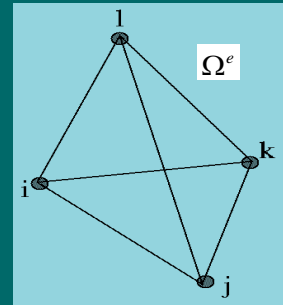
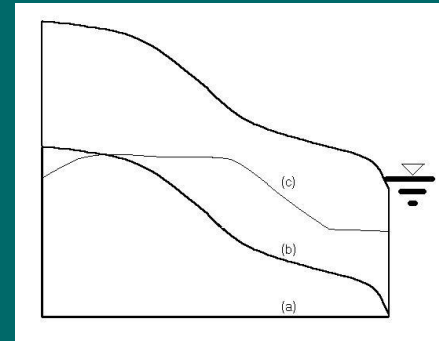
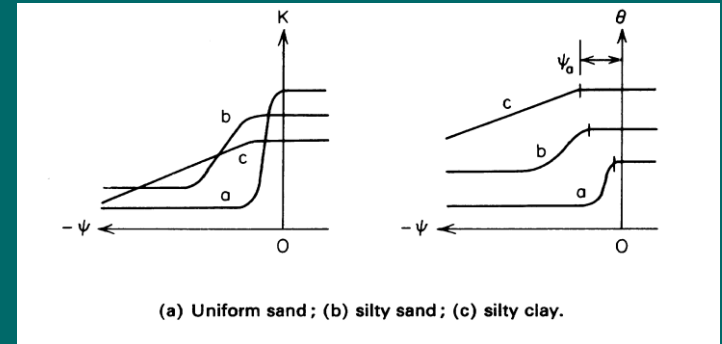
3D grid automatically generated with variable layer thicknesses and different base ("bedrock") shapes

Finite element spatial integrator (Galerkin scheme, tetrahedral elements, linear basis functions)

Weighted finite difference discretization in time

Time-varying boundary conditions: Neumann, Dirichlet, source/sink terms, seepage faces, and atmospheric fluxes

Adaptive time stepping; Newton and Picard linearization; selection of CG-type linear solvers; etc



Surface flow module (cell differentiation, lake handling, other features)

Overland (hillslope rills) and channel flow along s

DEM pre-analysis for definition of cell drainage directions, catchment drainage network and outlet, etc

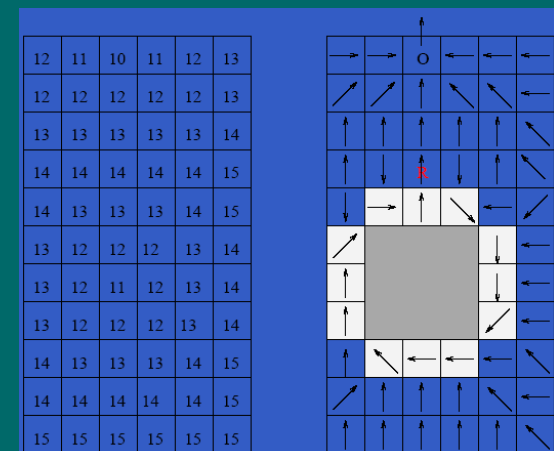
"Constant critical support area": overland flow \forall cells with upstream drainage area $A < A^*$; else channel flow (2 other threshold-based options also implemented)

Leopold & Maddock scaling relationships; Muskingum-Cunge solution scheme (explicit and sequential); etc

"Lake boundary-following" procedure to pre-treat lakes

Storage and attenuation effects of lakes and other topographic depressions are accounted for by transferring with infinite celerity all the water drained by the "buffer" cells to the "reservoir" cell; level pool routing calculates the outflow from this cell:

$$\frac{\partial V}{\partial t} = I(t) - O(h^*)$$



Surface flow module (drainage network flow characteristics)

*

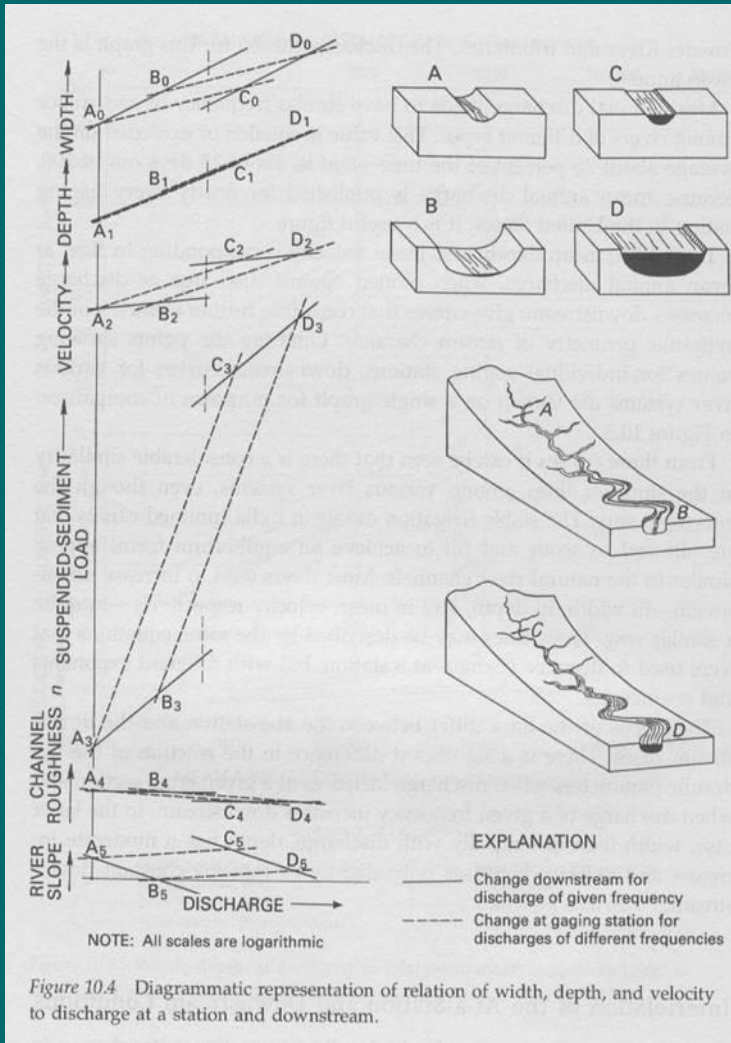


Figure 10.4 Diagrammatic representation of relation of width, depth, and velocity to discharge at a station and downstream.

Surface runoff propagated through a network of rivulets and channels automatically extracted from the DEM.

Spatial (term I) and temporal (term II) variations of flow characteristics of the drainage network (stream channel geometry W and conductance coefficient k_s) derived from application of downstream (according to upstream drainage area) and at-a-station (according to flow discharge) fluvial relationships:

$$W(A, Q) = \underbrace{W(A_s, Q_f)}_{\text{I}} Q_f \underbrace{(A_s)^{-b'}}_{\text{I}} \underbrace{(A/A_s)^{w(b''-b')}}_{\text{I}} \underbrace{Q^{b'}}_{\text{II}}$$

$$k_s(A, Q) = \underbrace{k_s(A_s, Q_f)}_{\text{I}} Q_f \underbrace{(A_s)^{-y'}}_{\text{I}} \underbrace{(A/A_s)^{w(y''-y')}}_{\text{I}} \underbrace{Q^{y'}}_{\text{II}}$$

* From L. B. Leopold and T. Maddock Jr. (1953), "The hydraulic geometry of stream channels and some physiographic implications", U. S. Geological Survey, Professional Paper no. 252

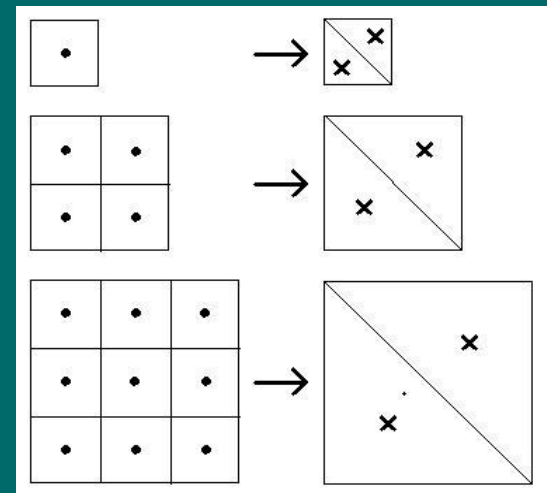
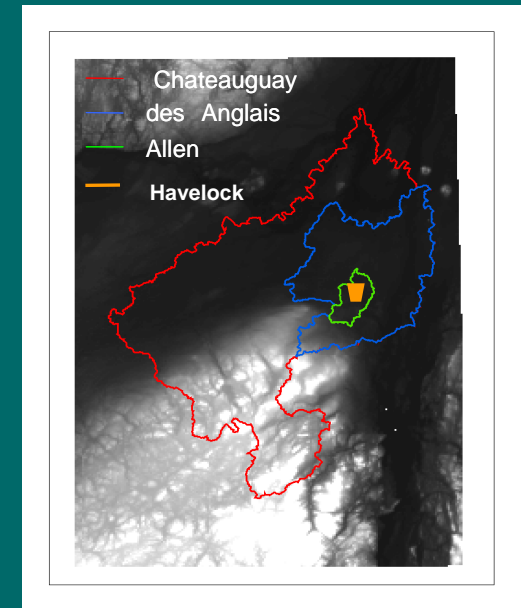
Coupling, time stepping, and iteration

"Pond_head_min" threshold parameter accounts for microtopography

Coupled system solved sequentially*: surface first, for Q^{k+1} and h^{k+1} ; then subsurface, for ψ^{k+1} ; finally overland flow rates q_L^{k+1} are back-calculated from subsurface solution [**sequential solution procedure but with iterative BC switching during subsurface resolution to resolve the coupling*]

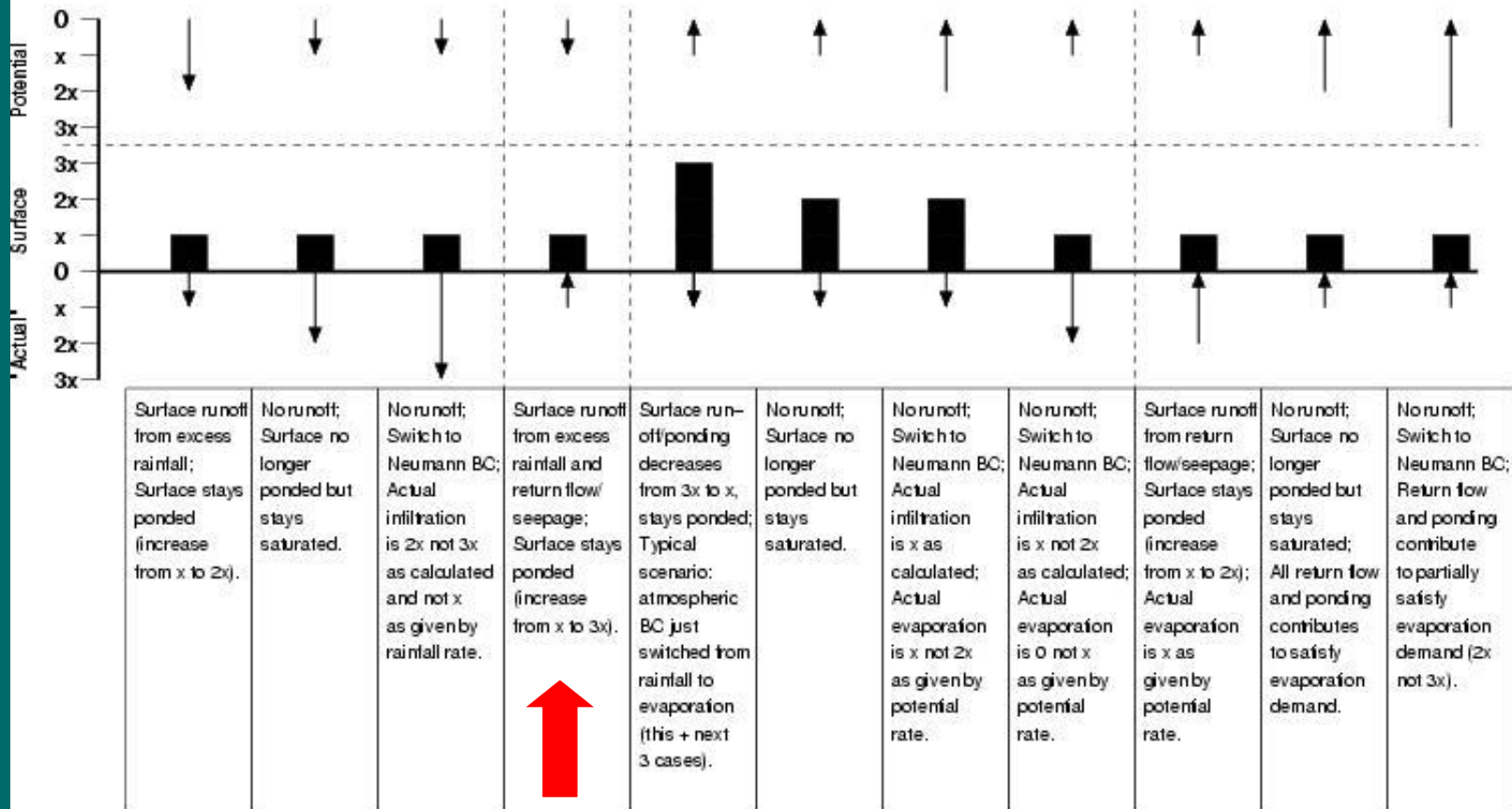
Nested time stepping: one or more surface solver time steps for each subsurface time step (based on Courant and Peclet criteria for the explicit surface routing scheme; also reflects typically faster surface dynamics compared to subsurface)

Interaction between cell-based surface grid and node-based subsurface grid includes input option for coarsening of latter grid. Allows us to exploit slower subsurface dynamics and looser grid constraints (implicit scheme), and can lower CPU and storage costs of 3D module



Boundary condition-based coupling (surface BC switching procedure)

Case I: Pondered surface



Case II: Saturated but not ponded

Case III: Unsaturated

Case IV: Dry (stage-two drying)



Analogous, but more straightforward
(as treated in subsurface-only mode)

Some recent studies (successes and challenges)

Recharge estimation (impact of heterogeneity)

Hydrograph separation (model coupling approaches)

Bedrock leakage (and the importance of boundary conditions)

Seepage faces (more on BCs)

Predicting near-surface soil moisture (a “too-wet” bias?)

Storage–discharge hysteresis (and other nonlinear phenomena)

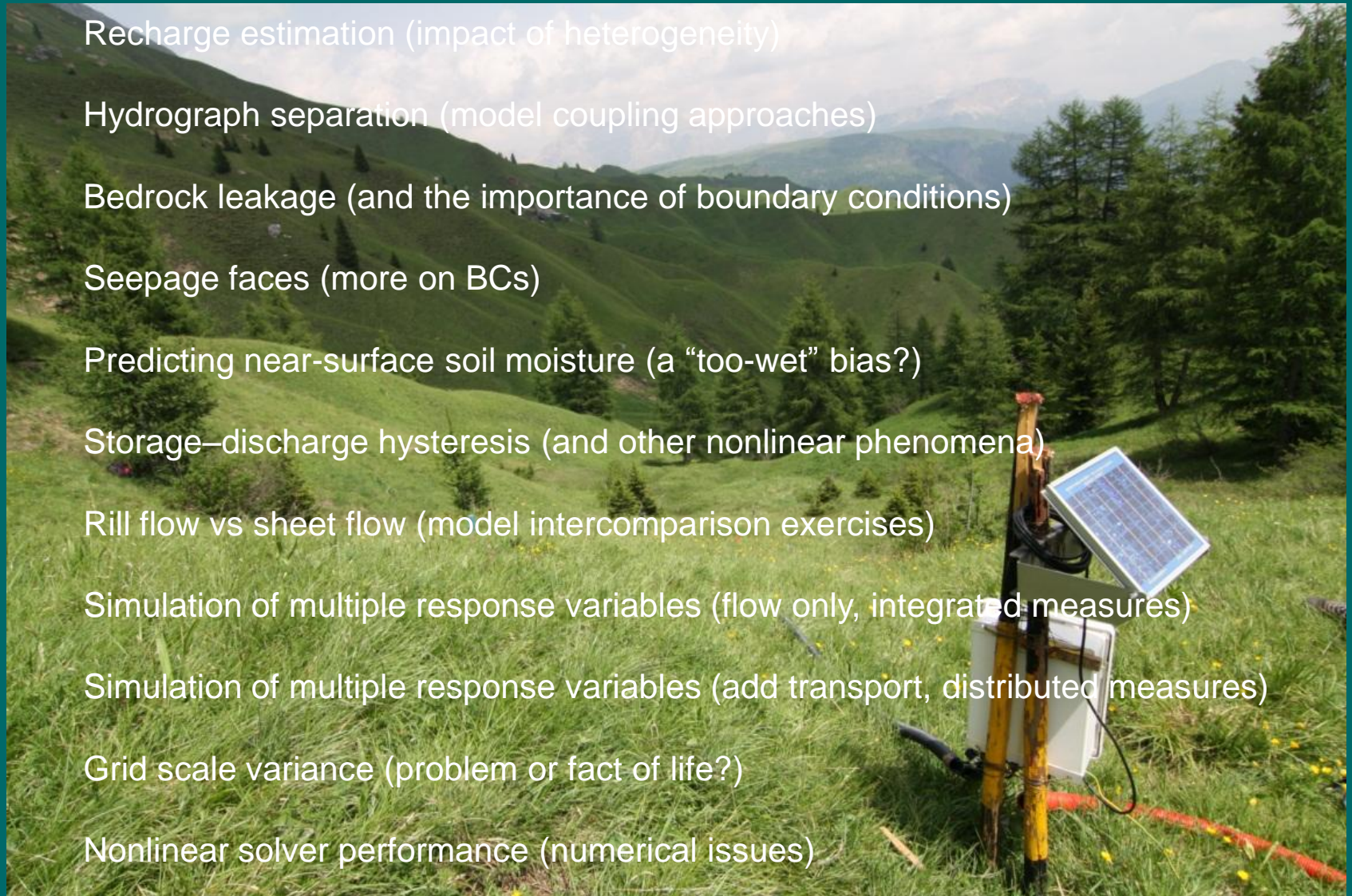
Rill flow vs sheet flow (model intercomparison exercises)

Simulation of multiple response variables (flow only, integrated measures)

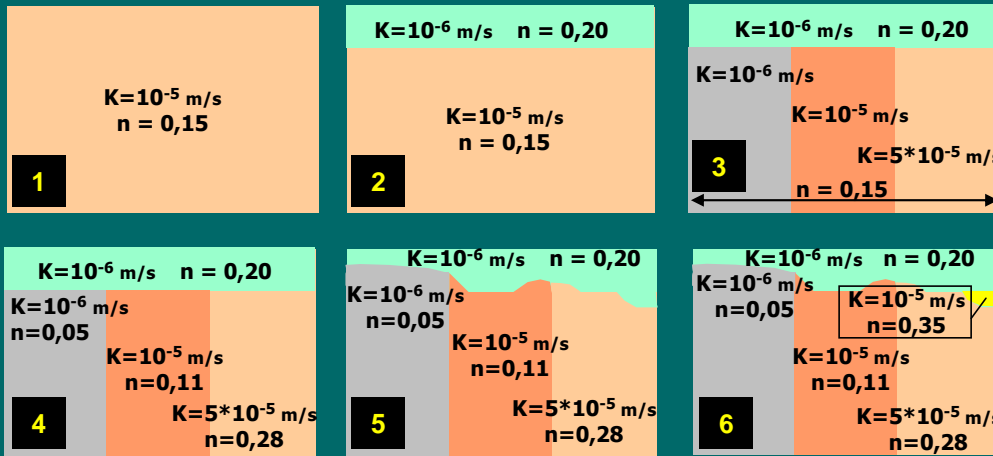
Simulation of multiple response variables (add transport, distributed measures)

Grid scale variance (problem or fact of life?)

Nonlinear solver performance (numerical issues)



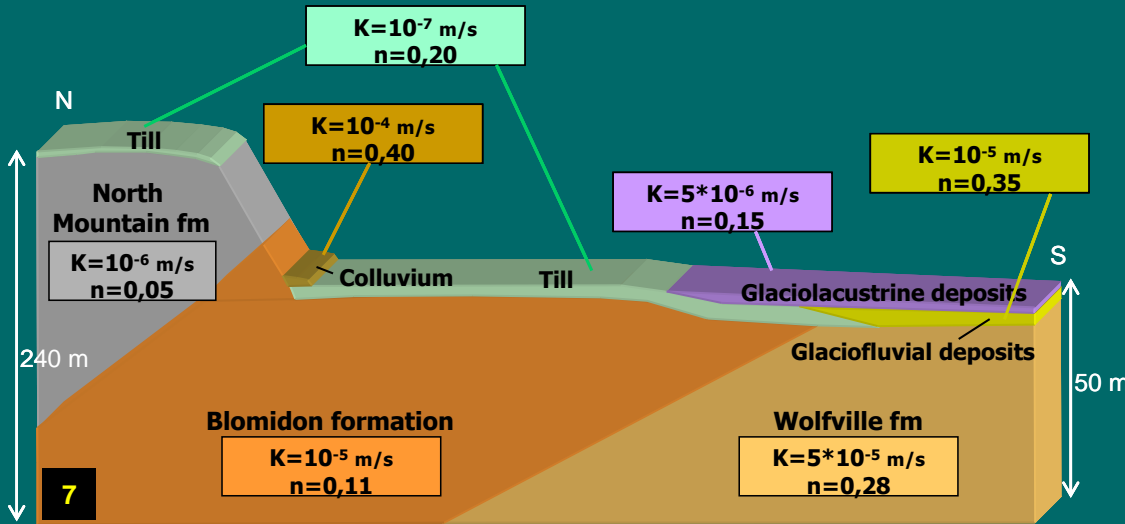
Recharge estimation (impact of heterogeneity): Thomas Brook catchment, Annapolis Valley, Nova Scotia



Higher-K colluvium and glacial deposits; smaller proportion of lower-K North Mtn fm

Increasing thickness (and nonuniformity) of surface layer

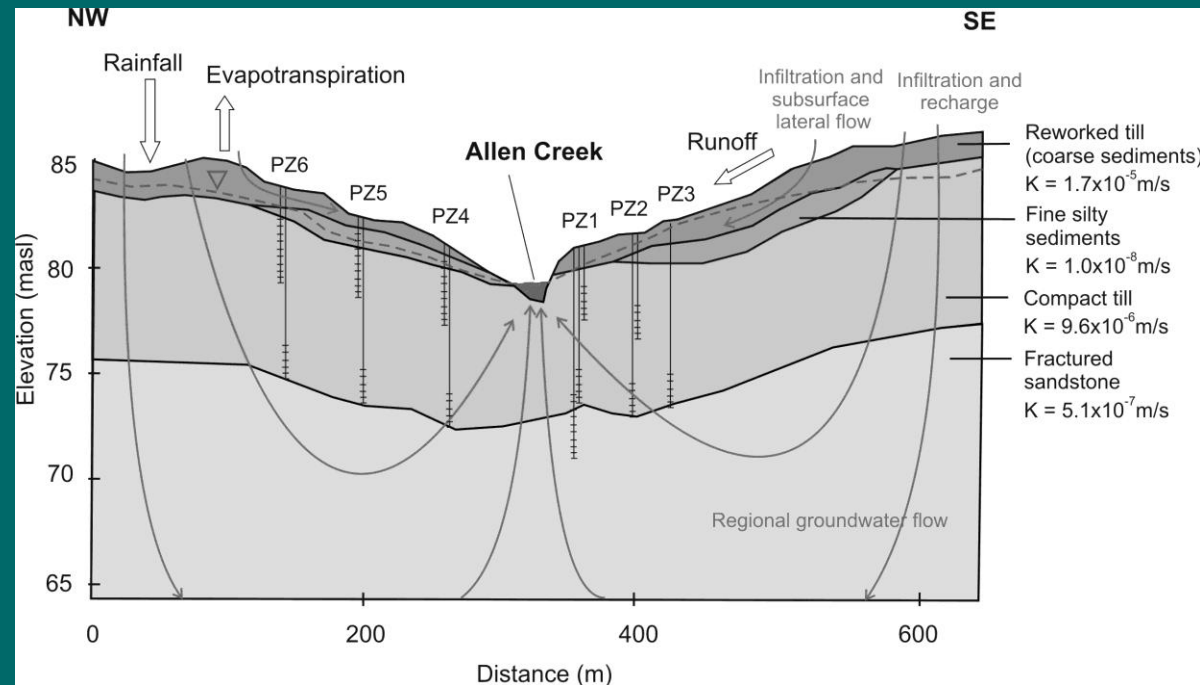
Lower-K surface layer



Configuration	Average streamflow (m ³ /s)	Annual Recharge (mm)
1	0,252	567
2	0,243	439
3	0,180	457
4	0,177	449
5	0,167	375
6	0,158	514
7	0,223	675
8 (lower K 3 surf. fms.)	0,180	368
9 (snow accumulation)	0,176	349
Mean	0,195	466
Std. dev.	0,0347	105

How do model-simulated values compare to other techniques?
 What is the role of mechanisms such as reinfiltration or fill-and-spill?

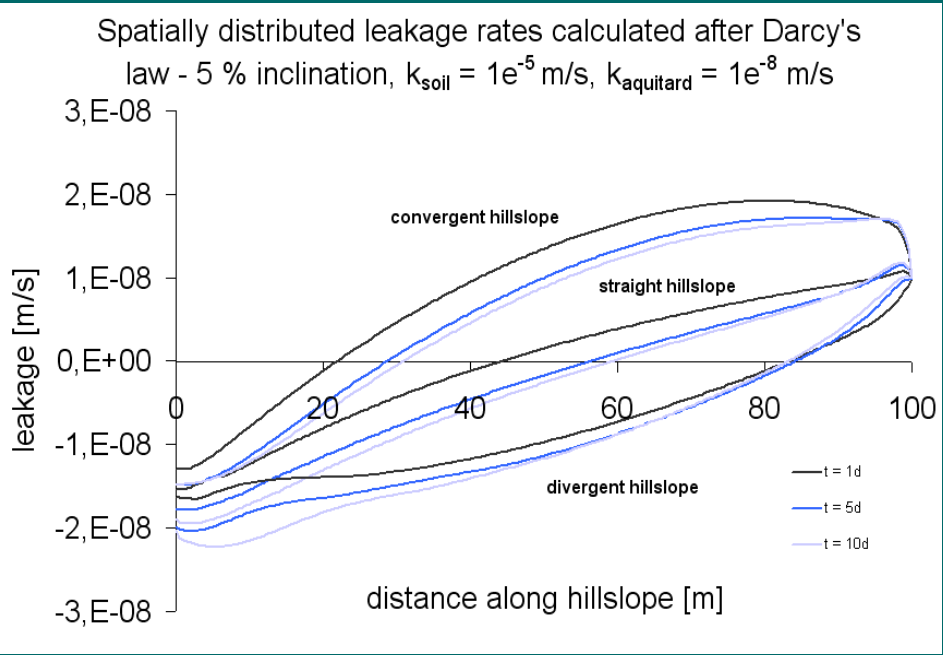
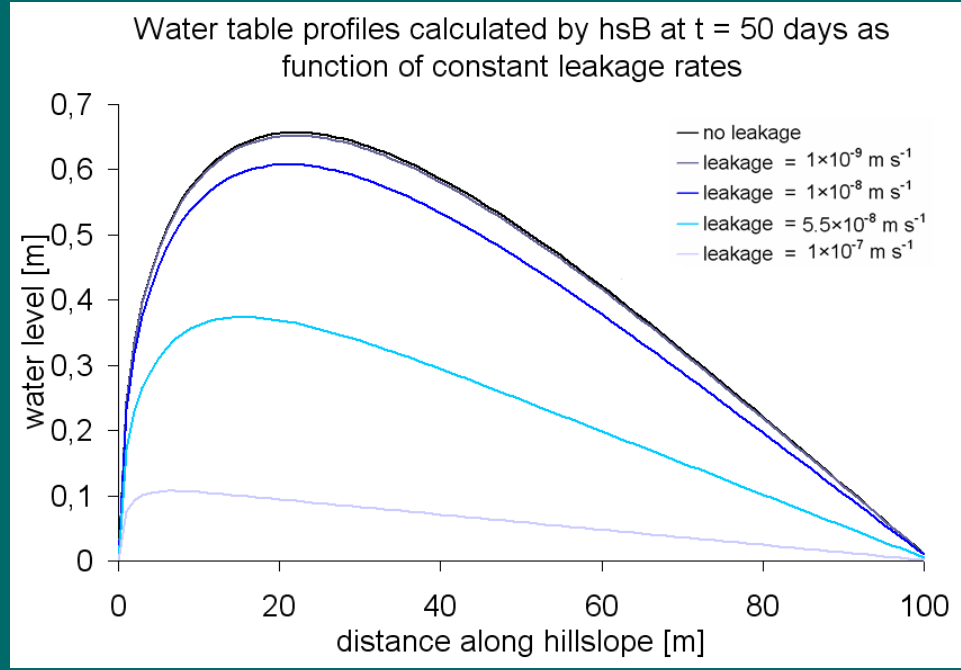
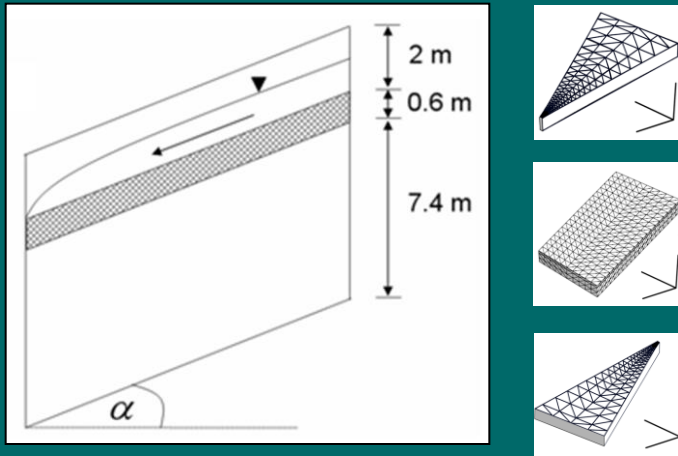
Hydrograph separation (model coupling approaches): Havelock hillslope, southwestern Quebec



Water budget component (mm/y)	HELP + FEFLOW	CATHY
Precipitation	1038	1038
Evapotranspiration	556	556
Recharge	214	233
Total Discharge	456	500
Surface runoff	231	/
Subsurface runoff	36	/
Baseflow	189	/
Exchange with regional fractured aquifer		
+ve (reg.aq. to hillslope)	4	77
-ve (hillslope to reg.aq.)	17	4
Storage change	14	55

Loose coupling (simplified model) vs CATHY: is hydrograph separation really so straightforward?

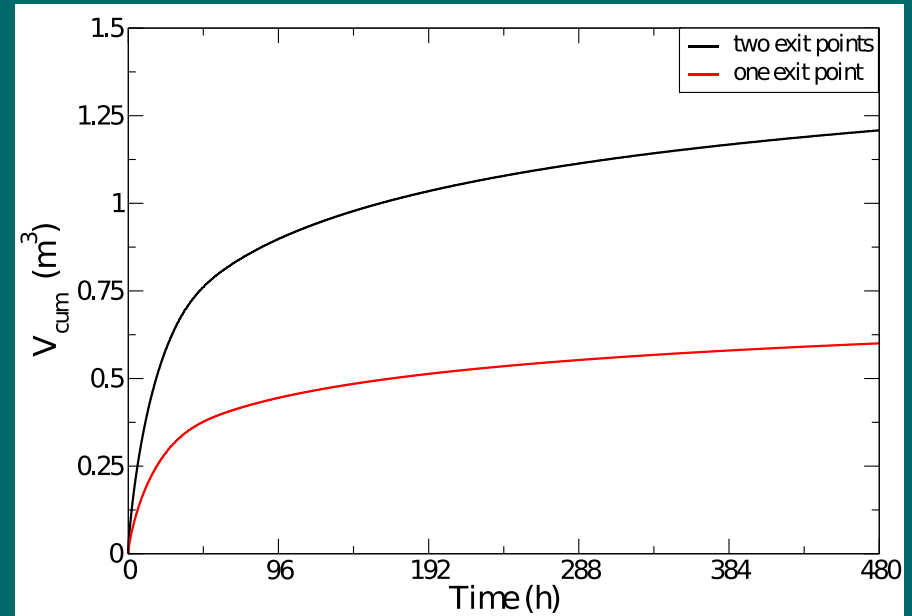
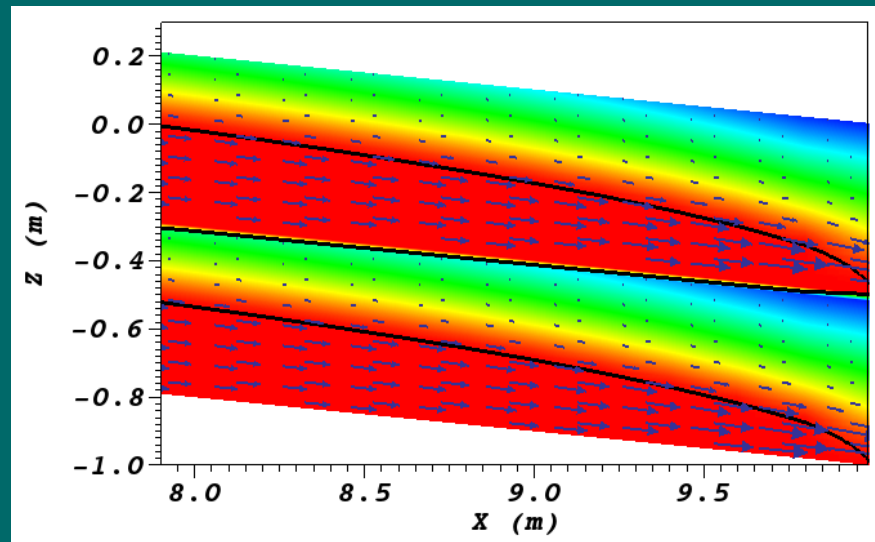
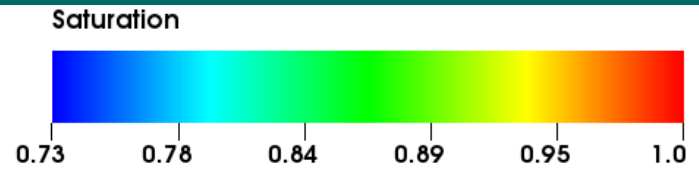
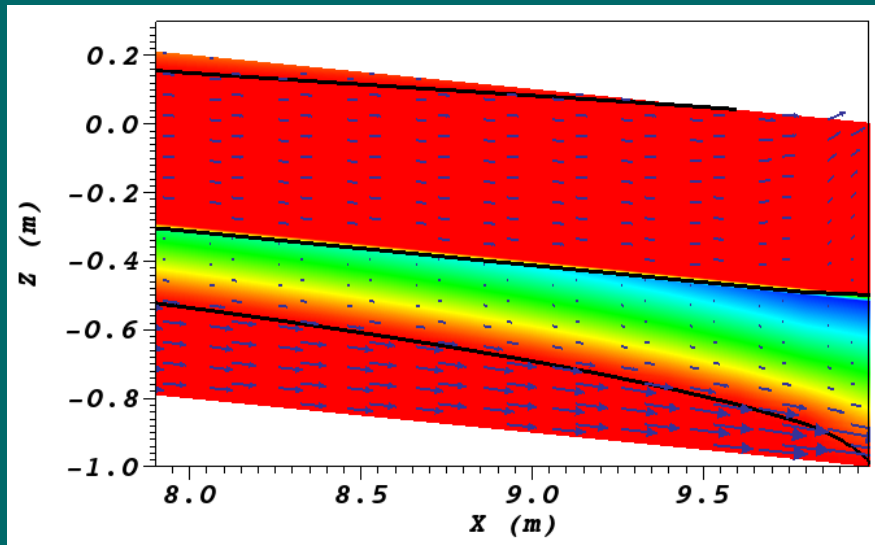
Bedrock leakage (and the importance of BCs): idealized hillslopes / sloping unconfined aquifers



Questioning a fundamental paradigm in hillslope hydrology.

Highly dependent on downslope BC treatment – not just a numerical issue.

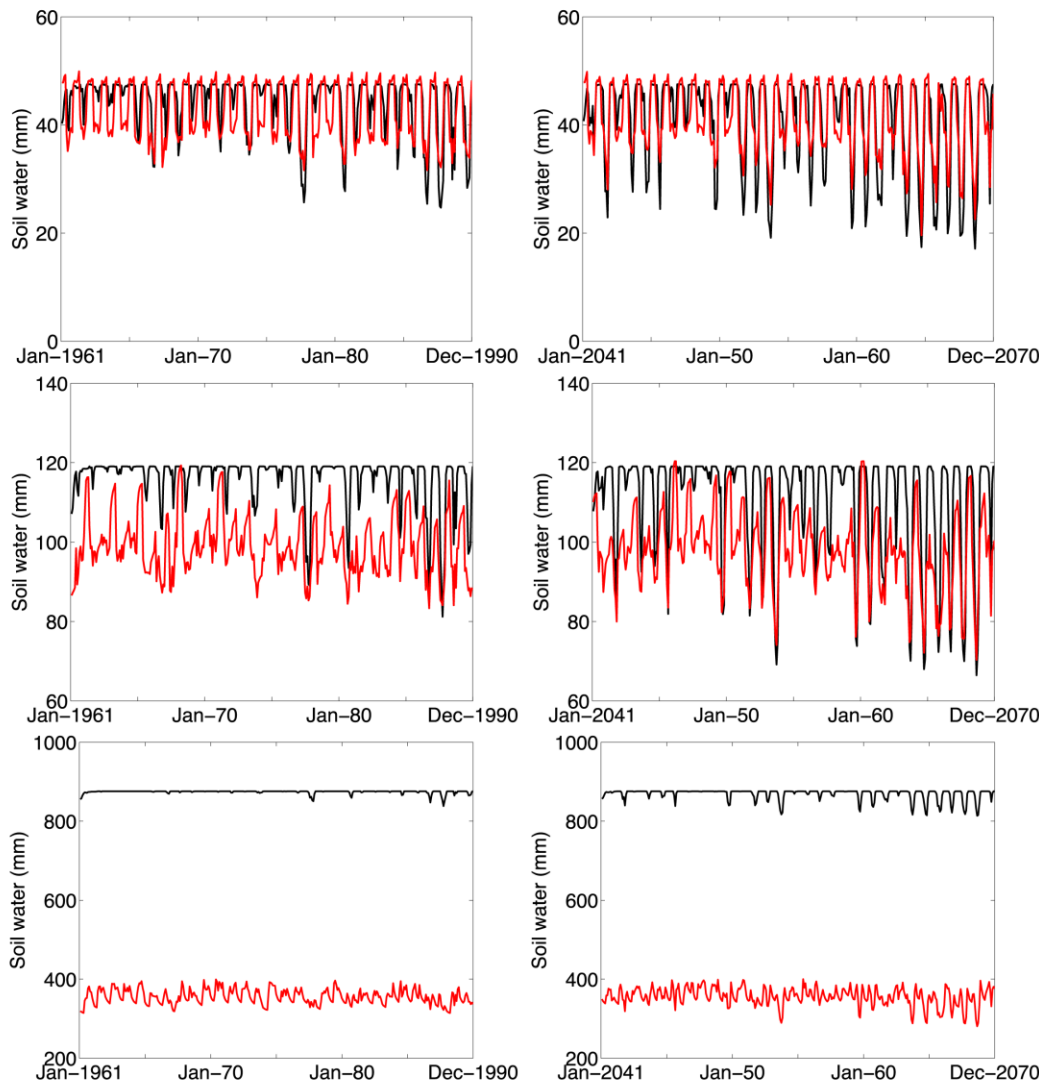
Seepage faces (more on BCs): idealized hillslopes, Landscape Evolution Observatory (LEO)



Issues:

- seepage face vs Dirichlet BC treatment
- interaction between seepage face and catchment outlet
- SF algorithm for full (random) heterogeneity

Predicting near-surface soil moisture (a “too-wet” bias?): des Anglais river basin, southwestern Quebec



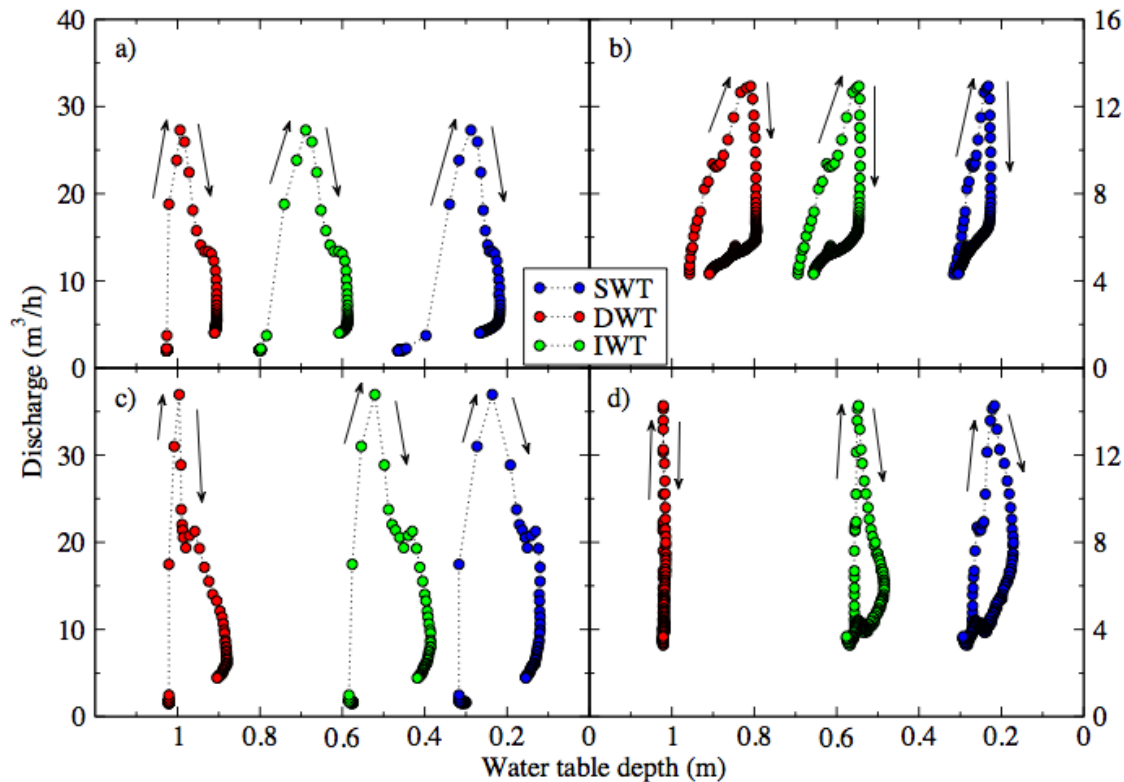
CLASS (red) and CATHY (black) results for monthly soil water content at different depths (shallow to deep from top to bottom) and for past (left) and future (right) climate projections.

Is there a bias in the model?

Possible causes:

- surface BC handling (eg, need seepage faces along stream banks?)
- too-coarse temporal rainfall resolution (peak rain rates get smoothed out → more infiltration, less surface runoff)
- missing transpiration
- too-coarse grid around steep terrain (eg, Covey Hill) misses important dynamics
- missing agricultural (eg, tile) drainage
- ...

Storage–discharge hysteresis (and other nonlinear phenomena): Larch Creek catchment, northern Italy



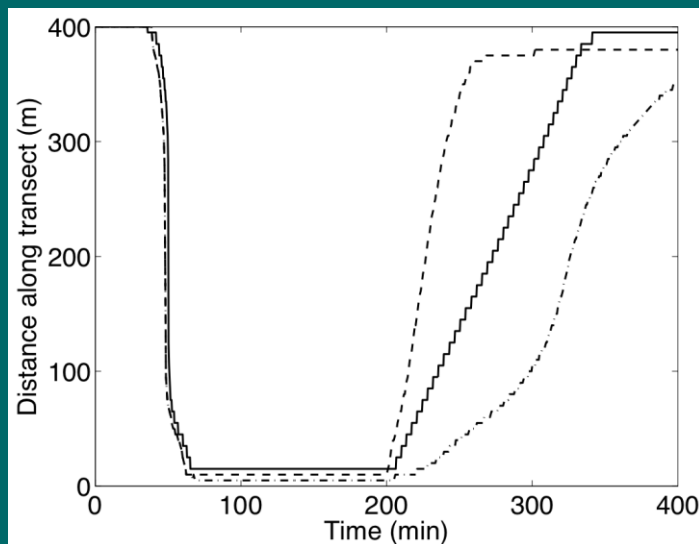
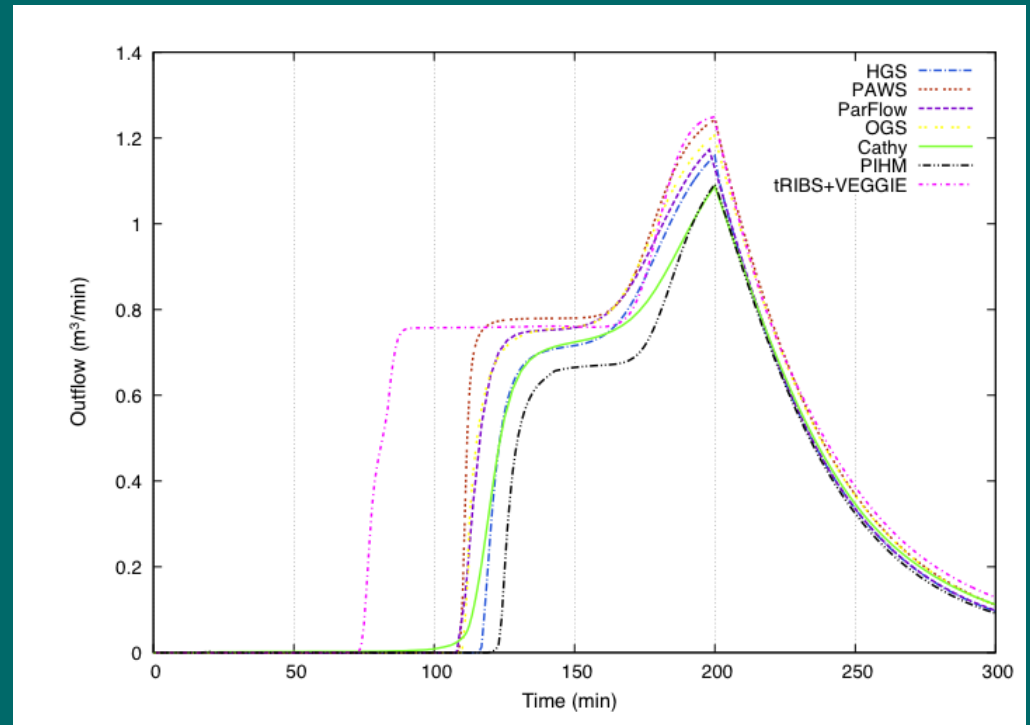
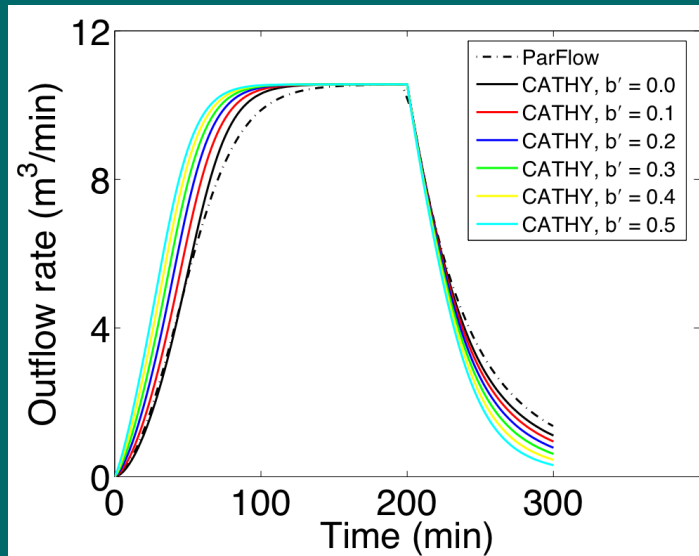
CATHY can reproduce hysteresis and thresholding behavior observed in the relationship between the subsurface storage and discharge responses of a small catchment. No ad hoc parameterization is needed.

Is there any link to or contribution from unsaturated zone hysteresis?

Nature and role of nonlinear phenomena in atmosphere–land surface–soil–aquifer interactions and feedbacks are poorly understood.

Simulated (top) and observed (bottom) responses in shallow, deep, and intermediate observation wells for 7–8 August 2009 (left) and 16–18 August 2009 (right) rainfall events.

Rill flow vs sheet flow (model intercomparison exercises): benchmark problems

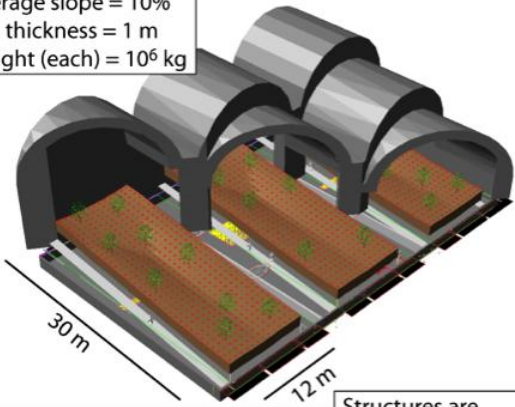


Benchmarking is a complicated business even for synthetic test cases ... Why and how do different models (even based on the same equations) perform differently? And what to do about it??

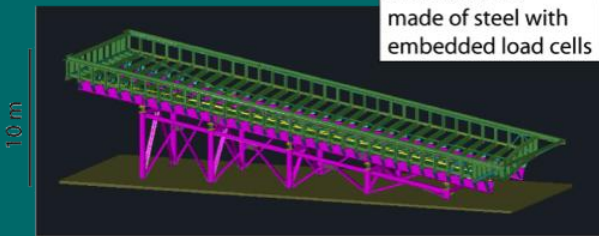
Evolution of the point of intersection between the water table and the land surface for the sloping plane test case. The outlet face is at $x = 400$ m. ParFlow: solid line; CATHY: dashed-dotted (sheet flow) and dashed (rill flow).

Simulation of multiple response variables (flow only, integrated measures): Biosphere 2 LEO

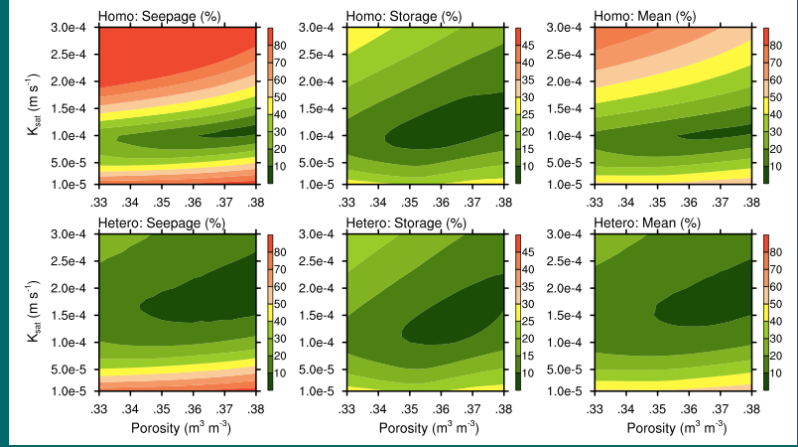
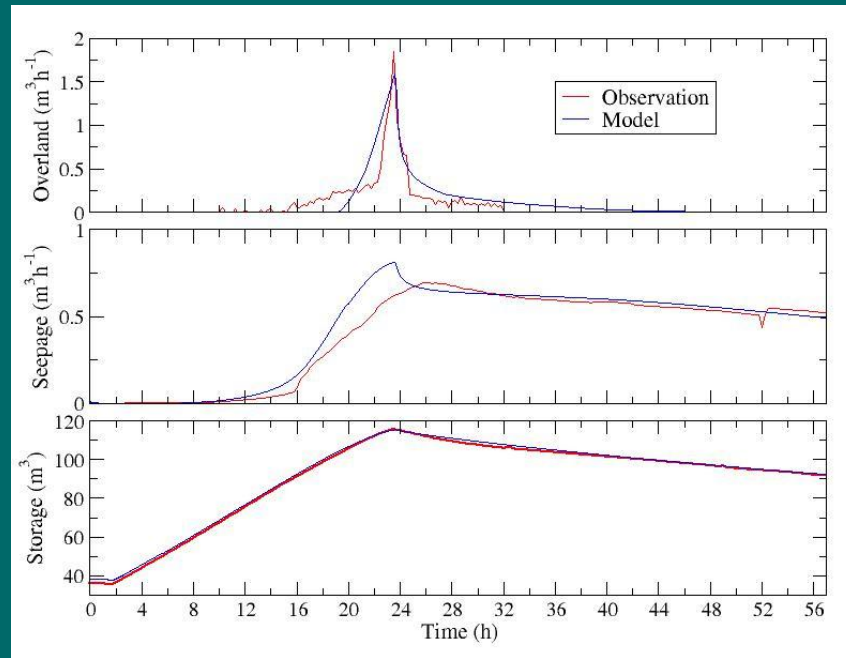
Average slope = 10%
Soil thickness = 1 m
Weight (each) = 10^6 kg



Structures are made of steel with embedded load cells



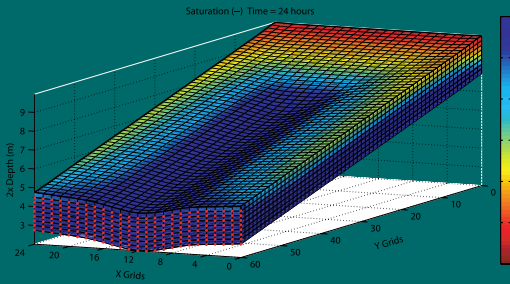
10 m



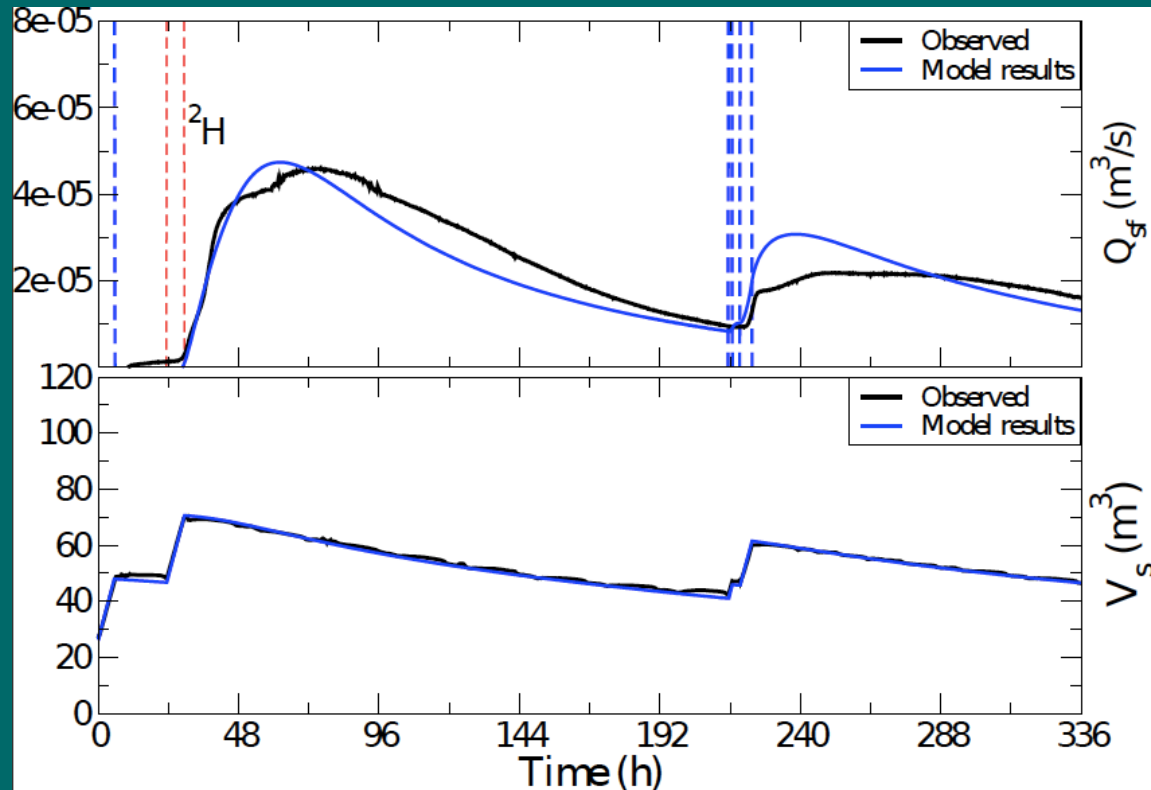
Issue of equifinality: does the mechanism we invoke (incipient downslope heterogeneity) imply (sole) causation?

“Perfect knowledge” of the bottom BC ... how much does this help?

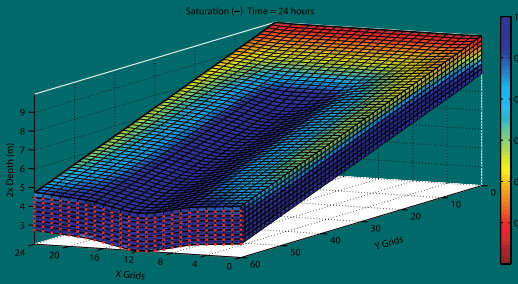
Simulation of multiple response variables (add transport, distributed measures): Biosphere 2 LEO



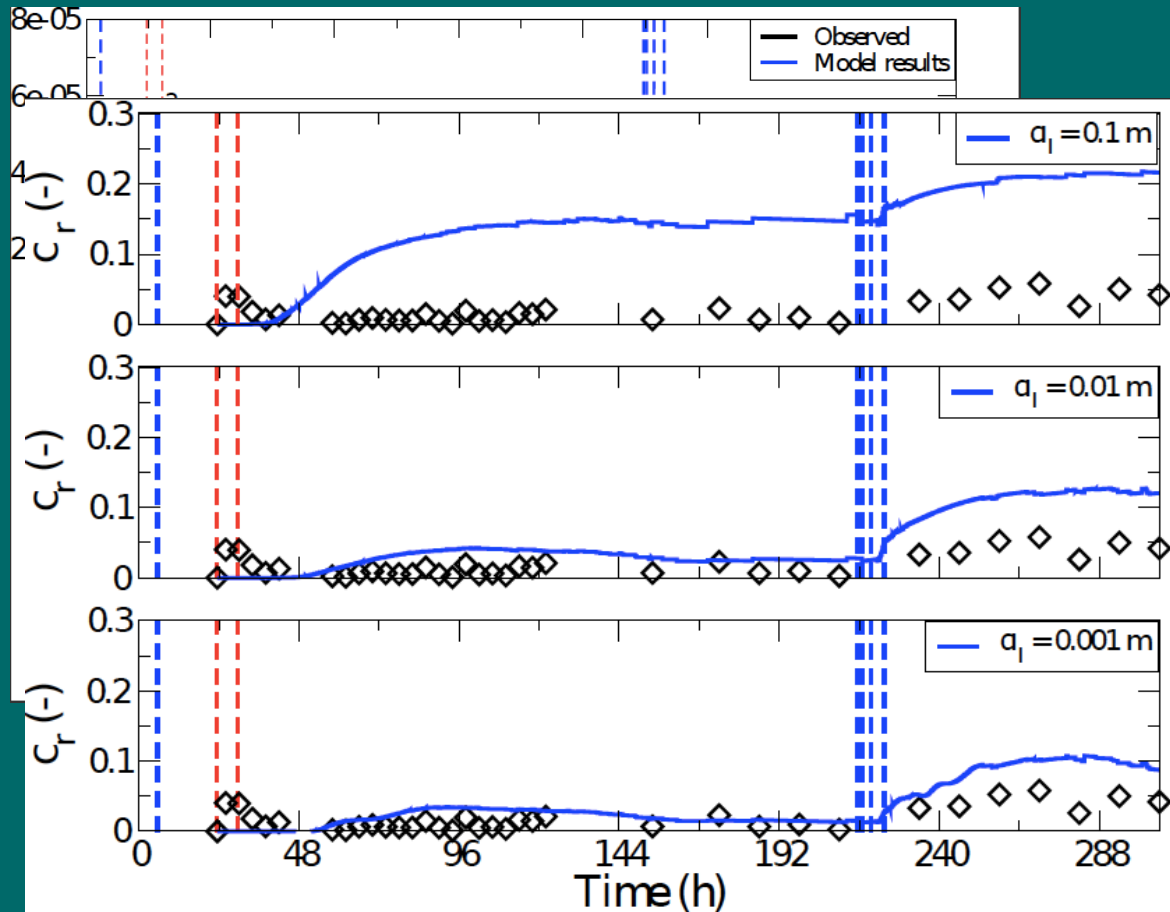
All three variables on previous slide are integrated measures of the hillslope response. How does the model perform when we examine distributed responses? And what happens when we include solute transport? ...



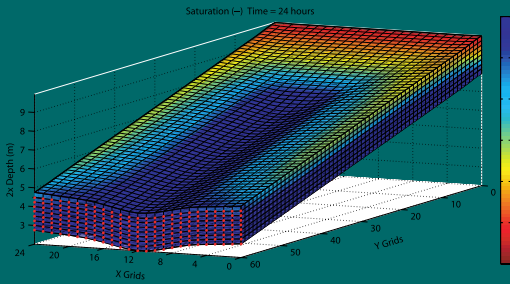
Simulation of multiple response variables (add transport, distributed measures): Biosphere 2 LEO



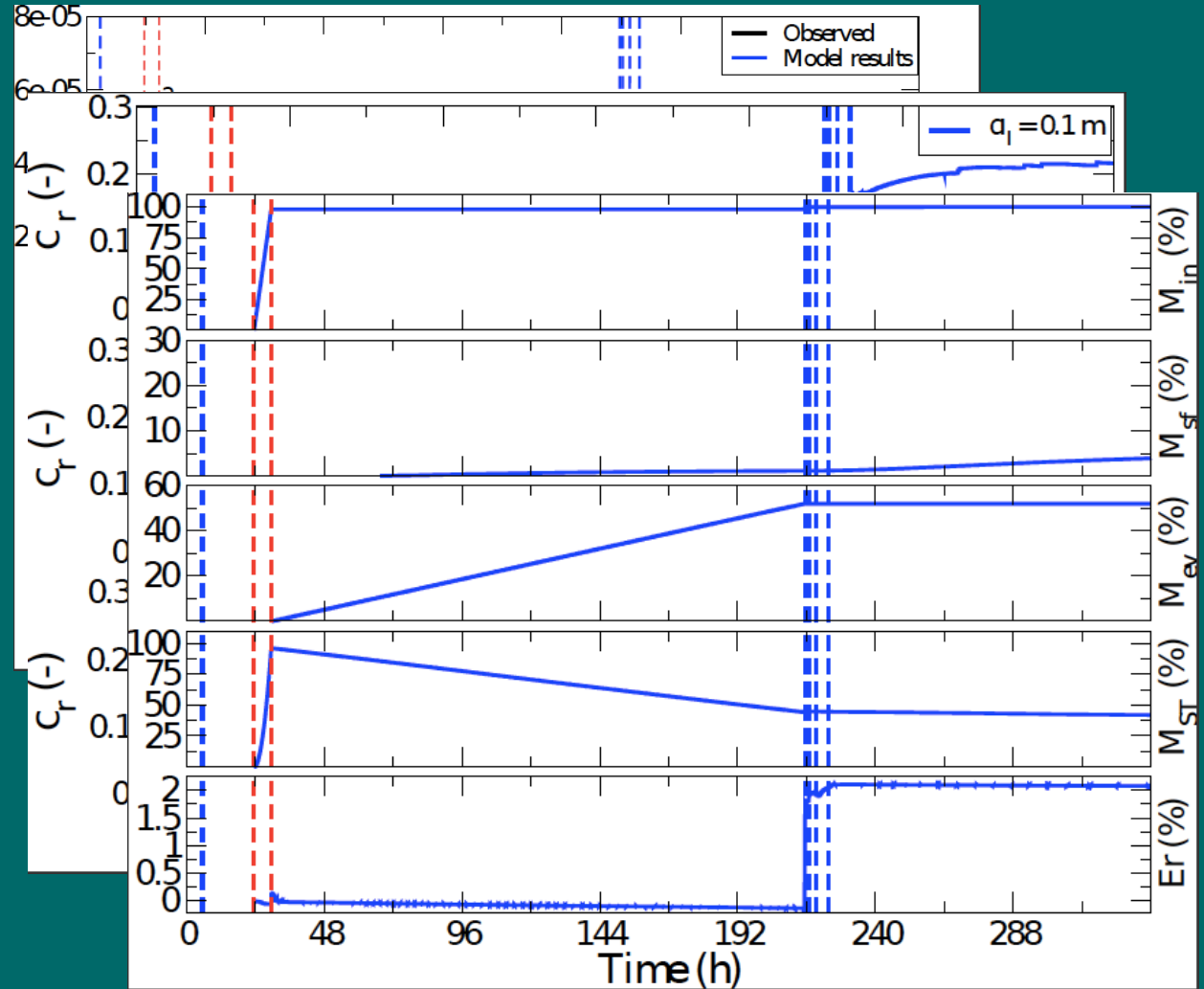
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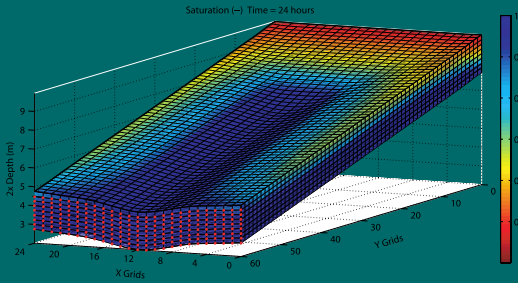
Simulation of multiple response variables (add transport, distributed measures): Biosphere 2 LEO



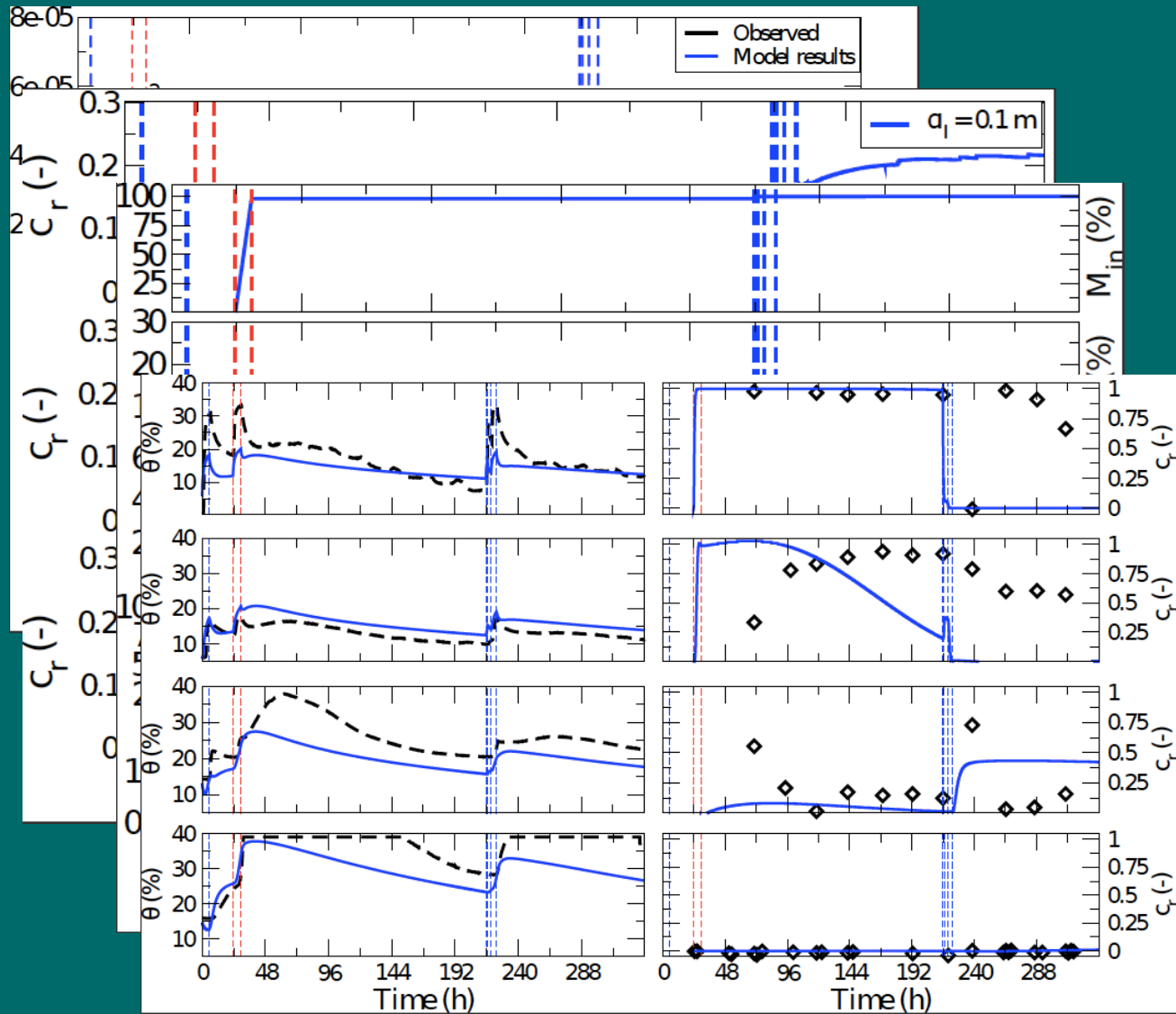
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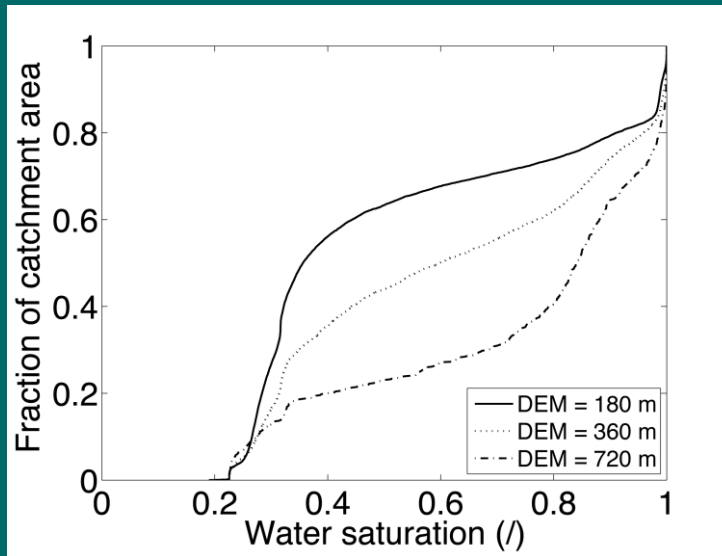
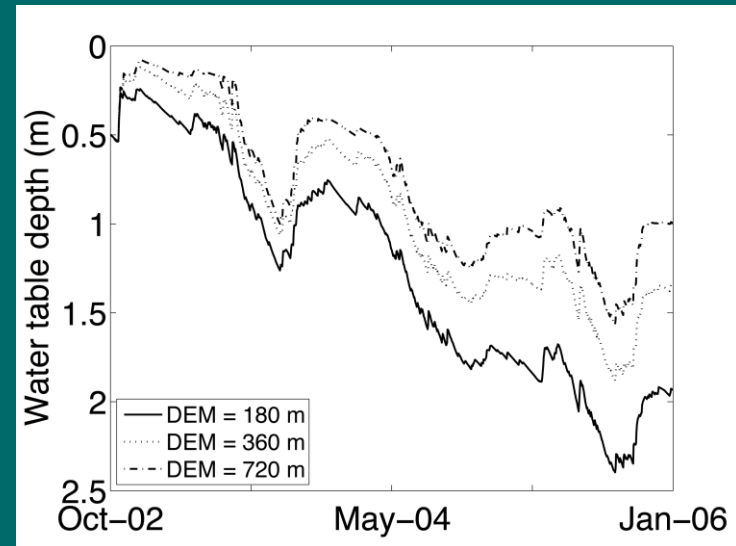
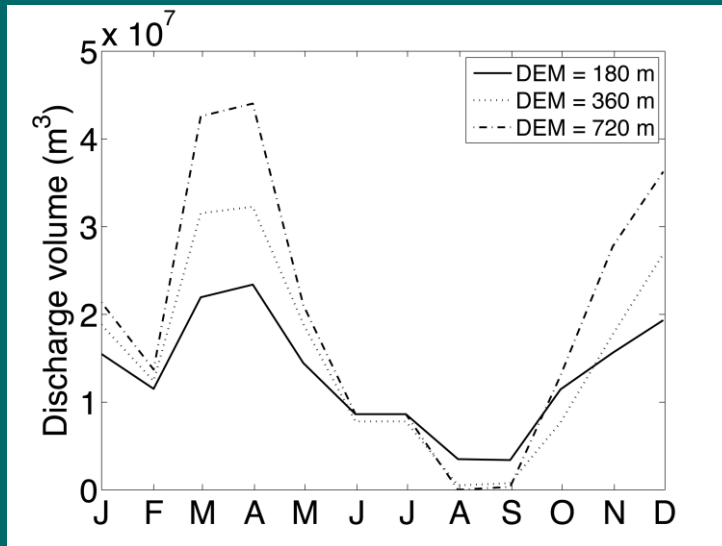
Simulation of multiple response variables (add transport, distributed measures): Biosphere 2 LEO



All three variables on previous slide are integrated measures of the hillslope response. How does the model perform when we examine distributed responses? And what happens when we include solute transport? ...



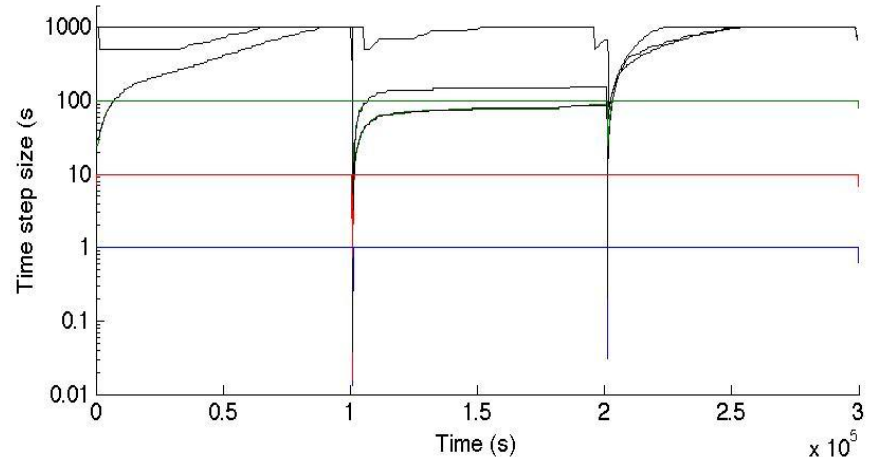
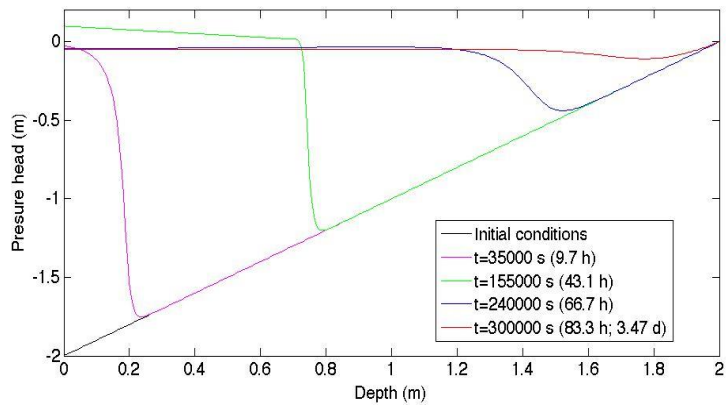
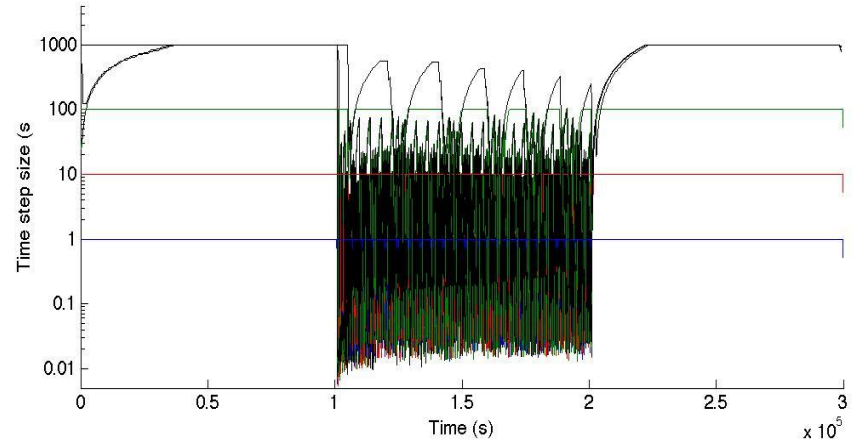
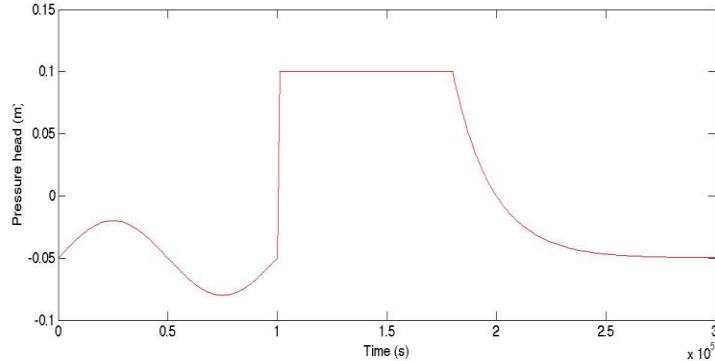
Grid scale variance (problem or fact of life?): des Anglais river basin, southwestern Quebec



Comparison of simulation results at 3 different DEM resolutions: average monthly streamflow discharge, catchment-averaged daily water table depth, and cumulative frequency distribution of surface soil saturation after a 10-day rain period.

There are many reasons (causes) for grid scale variance (and not limited to just the CATHY model). One of the most serious challenges in catchment-based hydrological / ecological modeling ...

Nonlinear solver performance (numerical issues): synthetic test case with dynamic ponding



Iterative schemes (Picard, Newton, nested Newton, ...), mass conservation (including data assimilation considerations), accuracy of velocity calculations (especially important for gw recharge and for solute transport), etc – all still ongoing research topics!

Adaptive time stepping behavior for Picard (top) and Newton (bottom) at four different vertical discretizations (10, 50, 250, and 500 layers for black, green, red, and blue curves, respectively).

Extensions and evolution of the model (flow and transport; other processes)

Flow (water quantity and distribution)

Surface

$$\frac{\partial Q}{\partial t} + c_k \frac{\partial Q}{\partial s} = D_h \frac{\partial^2 Q}{\partial s^2} + c_k q_s$$

Subsurface

$$\sigma(S_w) \frac{\partial \psi}{\partial t} = \nabla \cdot [K_s K_r(S_w) (\nabla \psi + \eta_z)] + q_{ss}$$

Transport (water quality and interactions with other substances)

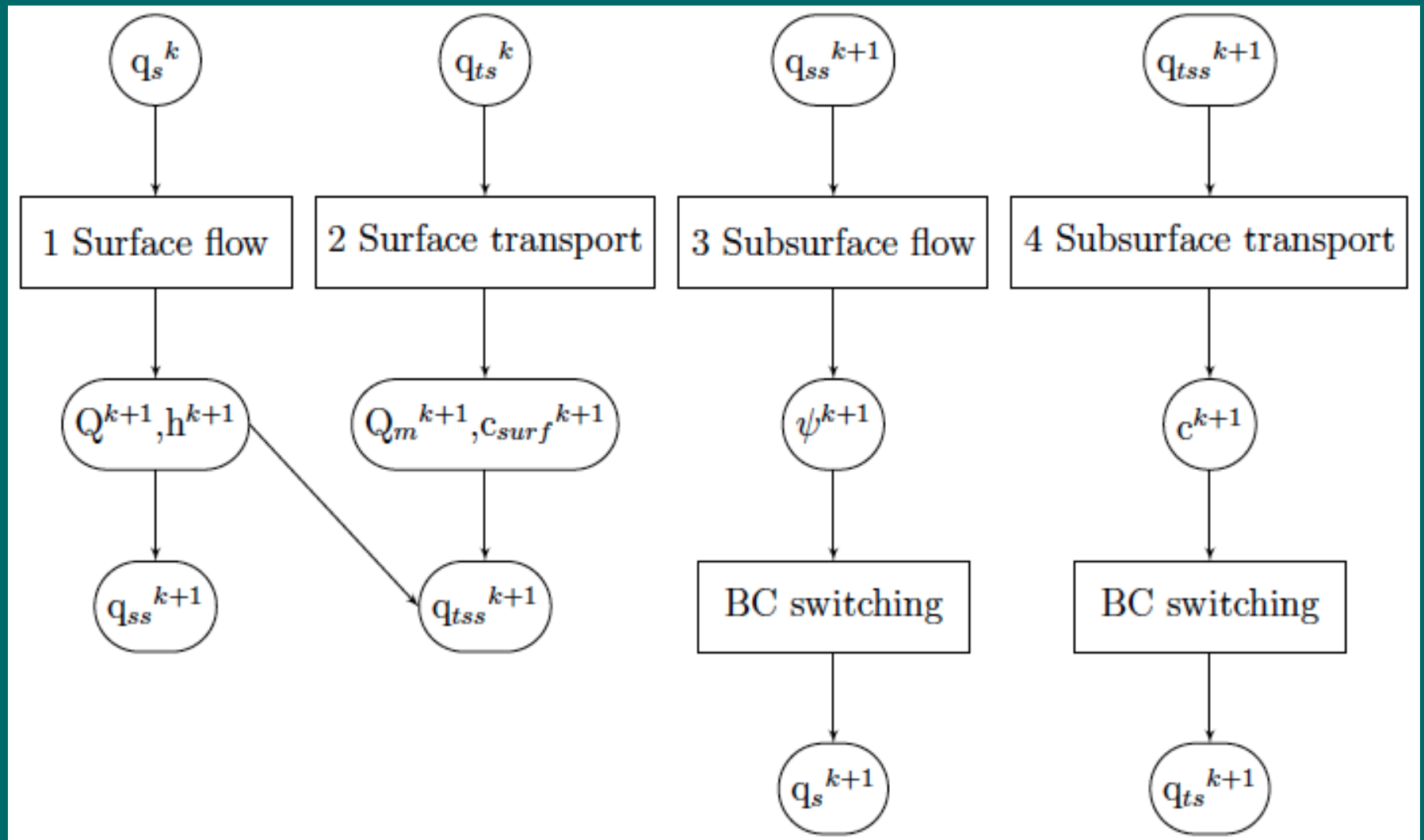
Surface

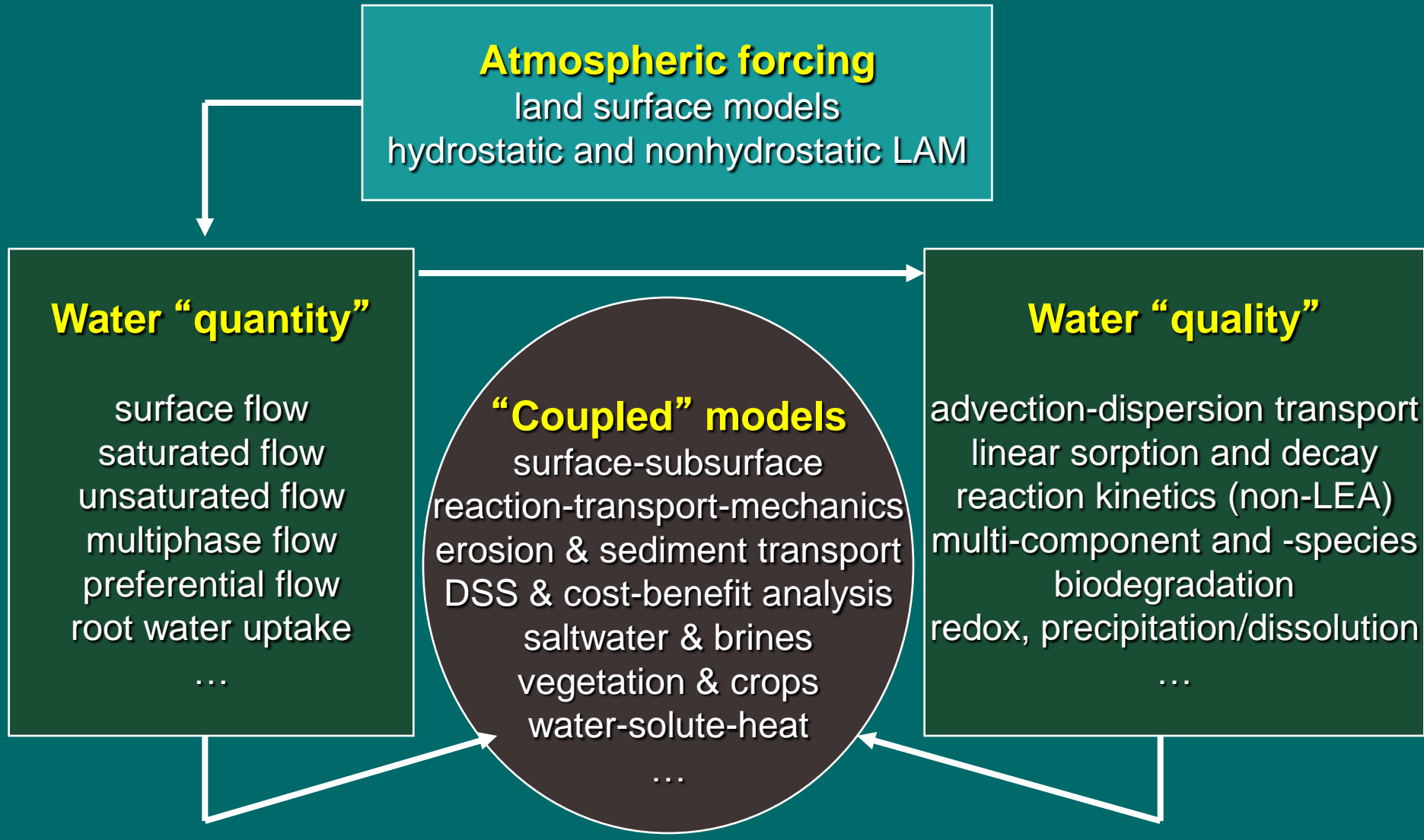
$$\frac{\partial Q_m}{\partial t} + c_t \frac{\partial Q_m}{\partial s} = D_c \frac{\partial^2 Q_m}{\partial s^2} + c_t q_{ts}$$

Subsurface

$$\frac{\partial \theta c}{\partial t} = \nabla \cdot [-qc + D \nabla c] + q_{tss}$$

Coupling scheme (sequential; iterative BC switching provides updated information)





Evolution of the model

Catchment/DEM-based subsurface flow modeling



soil freezing & snowmelt, preferential flow, unstructured grids, RCM/NWP coupling (CATHY-NoahMP-WRF?), ...

Collaborators

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many others (sensitivity tests, case studies, ...)
