

1 **Ground subsidence monitoring with SAR interferometry techniques in**
2 **the rural area of Al Wagan, UAE**

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12 **Abstract**

13 In this work, we investigate the past and present land deformation in Al Wagan
14 area in the United Arab Emirates. The area is primarily an agricultural region
15 where dependence on groundwater is documented. Such a reliance on ground
16 water resources in a region which is characterized by very low precipitation can
17 lead to significant land subsidence as was observed in this study which
18 identified fast and localized deformation trends. The quantification of terrain
19 distortions of large magnitude and small amplitude in this area with SAR
20 Interferometry is a challenging task using moderate resolution data due to the
21 incoherent surface background. Even though SAR acquisitions were sparse

22 over this region, the available ENVISAT, ALOS and Sentinel-1A imagery was
23 analysed with differential interferometry and the Small Baseline Subset
24 technique in order to provide fair estimates about the evolution of the
25 deformation pattern in a limited area. The depletion of the aquifer resources
26 which is confirmed from groundwater level data is speculated to be the most
27 probable cause. However these assumptions require further investigation in
28 order to discover a remediation for this problem.

29 Keywords: SAR Interferometry; SBAS; subsidence; groundwater; UAE

30

31 **Introduction**

32 Excessive groundwater extraction from the subsurface may lead to the depletion of
33 the aquifer's natural resources, which potentially results in ground surface subsidence
34 phenomena. The agriculture sector of the United Arab Emirates (UAE) has been expanding
35 since the 1990s (Statistics Centre of Abu Dhabi, 2014), but since all these farms are situated
36 in an arid region with a hot desert climate and extremely sparse precipitation events, they
37 mostly relied on groundwater for their viability (Environmental Agency of Abu Dhabi,
38 2009). The direct implications of this are the diminishing natural aquifer resources of the
39 UAE, which has already been analysed in another study (Gonzalez et al., 2016), the
40 degradation of the water quality (Vrba, 2003; Al-Naeem, 2014) and the occurrence of
41 localized terrain subsidence (Holzer and Galloway, 2005).

42 Differential Synthetic Aperture Radar (SAR) Interferometry (DInSAR) is one of the
43 most effective remote sensing methods to quantify terrain distortions related to various

44 causes over a large spatial scale (Bürgmann et al., 2000; Ferretti, 2014). The multi-temporal
45 stacking techniques that have been developed in the past two decades, like Persistent
46 Scatterers (PS) (Ferretti et al., 2001; Hooper et al., 2004) and Small Baseline Subset (SBAS)
47 (Berardino et al., 2002; Lanari et al., 2004), are capable of providing accurate estimations of
48 the displacement time series and the deformation velocities, with the latter method being
49 more suitable for rural regions as it is based on identifying Distributed Scatterers (DS) that
50 are more common in natural environments in a stack of small baseline interferograms. There
51 already exist numerous studies where groundwater related subsidence phenomena were
52 successfully assessed with SBAS techniques (Chaussard et al., 2014; Kim et al., 2015;
53 Artese et al., 2016). However, the accuracy and the reliability of the results is affected by
54 temporal decorrelation in rural areas where incoherent surface types, like sand and vegetation
55 are the dominant land covers (Bamler and Hartl, 1998; Massonnet and Feigl, 1998).

56 DInSAR applications are very sparse in the UAE mainly due to the extended presence
57 of the sandy geological background which significantly degrades the performance of the
58 process and the reliability of the final outputs. A previous study had demonstrated the
59 application of SBAS over an extended area of the Eastern part of the UAE focusing on the
60 most coherent parts and providing rough indications about the locations of particular ground
61 distortion zones (Cantone et al., 2013). However, to the authors' knowledge, there are no
62 existing studies attempting to monitor localized surface deformations in the UAE at small
63 scales. The major challenge in this rural area where the subsidence phenomena were
64 extremely fast and localized was to define the optimum data resolution that allowed the
65 detection and depiction of the real deformation magnitude in the presence of high
66 decorrelation noise.

67 This work is part of the study currently being conducted in the UAE aiming to detect
68 regions affected by subsidence phenomena induced by subsurface resources exploitation and
69 to provide quantitative estimations about the spatiotemporal patterns of the terrain
70 displacements. The objective of the present work is to analyse the observations obtained by
71 conventional DInSAR and SBAS techniques and to demonstrate that despite the presence of
72 low coherence that obstructs the interferometric processing, reliable results may still be
73 obtained to provide a comprehensive overview of the subsidence phenomenon in the study
74 area over a larger spatial extent than is possible by sparse permanent GPS stations.

75

76 **Study Area**

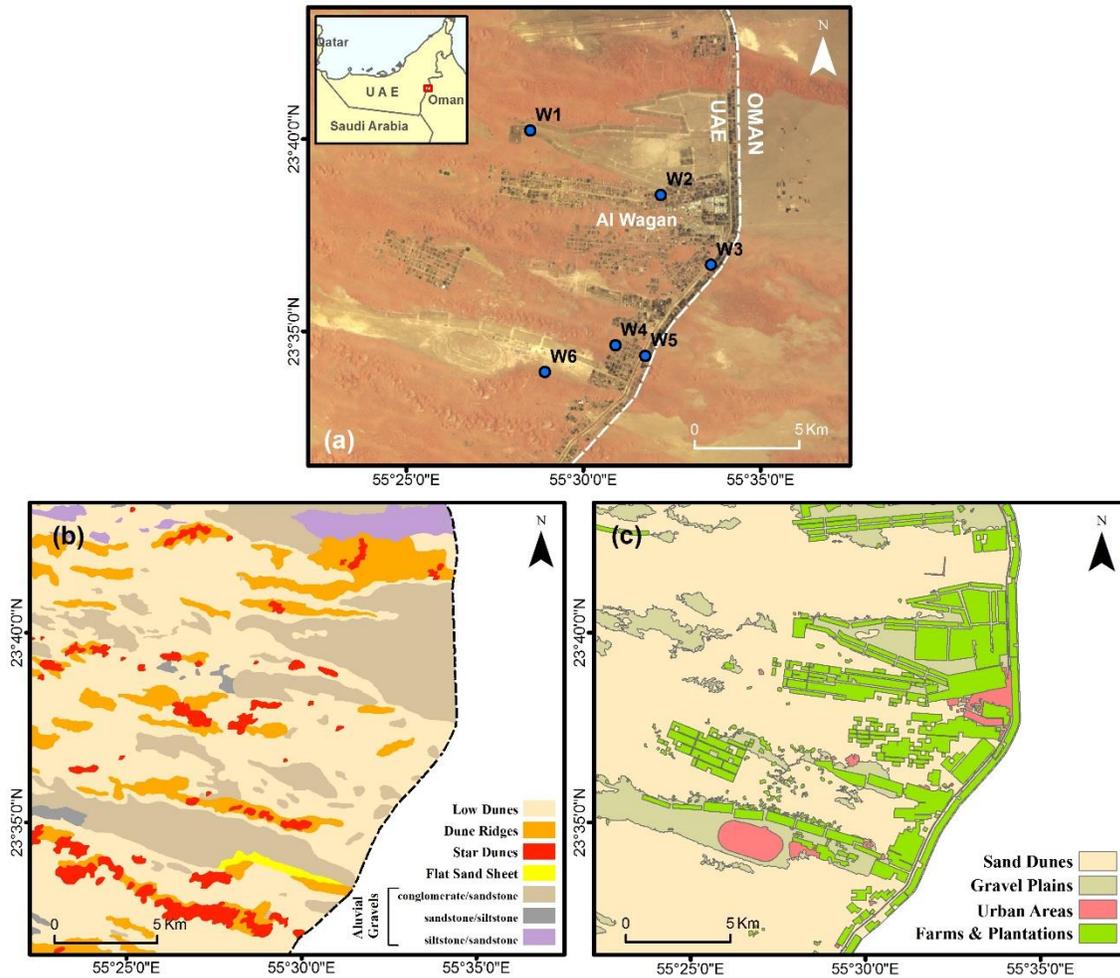
77 Al Wagan is located on the eastern part of the Abu Dhabi Emirate, approximately 65
78 kilometres southwest of the city of Al Ain at the borders with Oman. Figure 1a depicts the
79 study area along with the locations of monitored wells from the Environmental Agency of
80 Abu Dhabi (EAD). It is a hot and arid region at the edge of the Rub' al Khali desert with a
81 generally flat terrain - except for the star dunes - at an average altitude of 170 meters. The
82 surrounding area of this sparsely populated town facilitates many farms at the west side of
83 the borders which unequally spread up to 18 kilometres towards the desert gradually fading
84 out. The farms along with other vegetated areas cover around 71 km² over a total area of
85 approximately 357 km².

86 The borders of the UAE with Oman at the study area roughly depict the transition
87 from the fluvial deposits in the east to the Aeolian deposits to the west. The geological
88 background of the north part of the region where the main settlement is sited consists of
89 fluvial quaternary deposits which are composed of a mixture of sandstone, conglomerates

90 and siltstone (Hili formation), formed by torrent activity originating from the Omani
91 mountains. The central and southern parts are mainly covered by low and star dunes of sandy
92 material transferred by aeolian processes. The heterogeneity of the local geology shown in
93 figure 1b is clearly visible in radar and optical satellite imagery and sporadic shifts between
94 these two main background types are not uncommon. The geological map of the area is not
95 totally reflected on the land use map (figure 1c), implying that the farms are not located
96 exclusively on the alluvial fans but on the sand deposits as well.

97 There were no active faults in the study area observed in the geological maps neither
98 mentioned in literature. The nearest regions that were affected by active tectonism in the
99 studied period were the Fujairah Mountains and the Musandam Peninsula (Yagoub, 2015).
100 Therefore it was considered in this study that the effects of earthquakes on surface
101 deformations were negligible and they were not taken into account.

102



103

104 Figure 1. a) Landsat 8 composite image (bands 4, 3, 2) of the study area dated 26/5/2017. b)
 105 Geological map (BGS) c) Generalized land cover map (EAD).

106

107 **Data**

108 The interferometric analysis was performed with the use of the freely available C-
 109 band and L-band SAR data covering the study area. Two C-band image stacks acquired from
 110 the ENVISAT satellite by the Advanced Synthetic Aperture Radar (ASAR) sensor were
 111 obtained from the European Space Agency (ESA). The descending stack consists of 32
 112 acquisitions covering the period 7/8/2003 – 22/4/2010, while the ascending includes only 10

113 images from the period 3/4/2007 – 14/10/2008. Other C-band data used in this study were
 114 the available Sentinel-1A acquisitions over the area of interest, but the sparsity of the data
 115 did not allow a proper time series analysis. Three SAR acquisitions were used for differential
 116 interferogram generation dated 18/2/2016, 2/12/2016 and 8/3/2017. The L-band data were
 117 also sparse over this region, consisting of 7 ALOS PALSAR acquisitions from the period
 118 5/8/2007 – 10/8/2009 obtained from the Alaska Satellite Facility (ASF). The general
 119 characteristics of these radar data are presented in table 1.

120

121 Table 1. Main characteristics of the SAR data used in the study.

Sensor	Number of images	Pass	Track/Frame	Acquisition Mode	Polarization
ENVISAT/ASAR	32	Descending	435/3123	IW-IM	VV
ENVISAT/ASAR	10	Ascending	13/459	IW-IM	VV
ALOS/PALSAR	4	Ascending	460/567	FBD	HH-HV
ALOS/PALSAR	3	Ascending	460/567	FBS	HH
Sentinel-1A	3	Ascending	130	IW	VV-VH

122

123 The topographic phase removal was performed with the 3 arcsecond Digital Elevation
 124 Model (DEM) of the Shuttle Radar Topography Mission (SRTM). Other data that were used
 125 include the geological map of the area produced by the British Geological Survey (BGS)
 126 under contract from the Ministry of Energy of the UAE, the Terrestrial Habitat map of the
 127 UAE and the groundwater levels from the monitored wells of the study area obtained from
 128 the Environmental Agency of Abu Dhabi (EAD) in the period 5/3/2000 – 9/9/2015.

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130 **Methods**

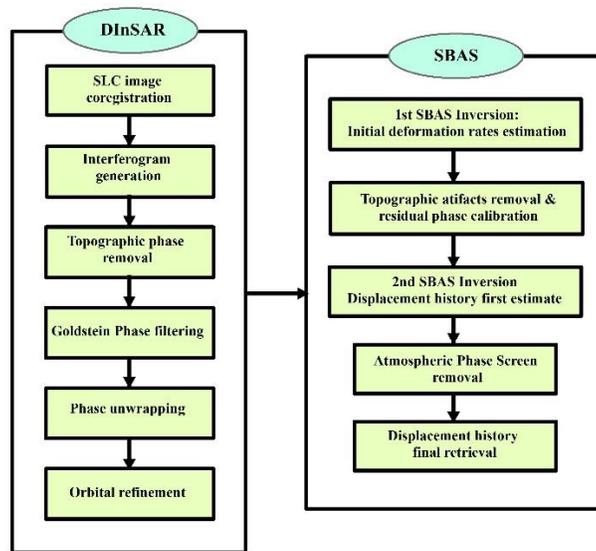
131 DInSAR relies on the isolation of the deformation phase component which
132 corresponds to the terrain movements between the two SAR acquisitions of similar geometry.
133 The Single Look Complex (SLC) images of each stack were coregistered on a single master
134 scene in order to compute the interferograms. Due to the flat relief of the study area the
135 SRTM DEM of 90 m. resolution was considered adequate for subtraction of the local
136 topography, since differences between trials with the ASTER and the 30 m. resolution SRTM
137 DEMs were insignificant. The Goldstein method was used to apply frequency domain phase
138 filtering to the differential interferograms (Goldstein and Werner, 1998) with small windows
139 in order to avoid aliasing fast deformation fringes, at the expense of preserving a lot of
140 decorrelation noise in the final outputs. After the phase unwrapping stage (Pepe and Lanari,
141 2006), orbital refinement was performed for estimation and removal of residual phase ramps
142 from the unwrapped interferometric stack by selecting ground control points common for all
143 the generated interferograms.

144 The SBAS technique is based on generating and stacking a series of interferograms
145 with small perpendicular baselines in order to reduce the spatial decorrelation effects.
146 Therefore, the DInSAR process was applied to interferograms that satisfied certain
147 temporal/perpendicular orbit criteria depending on the monitored surface types, the critical
148 baseline and the relative space-time dispersion of the acquisitions in each stack. At the first
149 inversion of the SBAS algorithm a primary estimate of the mean velocity field was obtained
150 by applying a low degree displacement model to the unwrapped phases. This low pass
151 deformation component along with possible topographic remnants was subtracted from the

152 wrapped interferograms. Subsequently, the residual phase was unwrapped and added back to
153 the subtracted deformation signals (Berardino et al., 2002).

154 In the second step the refined unwrapped interferograms were inverted with a
155 Singular Value Decomposition (SVD) approach in order to derive the time series of the
156 displacements (Lanari et al., 2007). The atmospheric signals were removed with low pass
157 spatial and high pass temporal filtering operations (Ferretti et al., 2000) and the final
158 deformation history and velocity of the multi-temporally coherent pixels was extracted along
159 with the topographic residuals between the DEM and the actual ground surface. The results
160 were geocoded in UTM zone 40N projected coordinate system and the data processing was
161 mainly performed with Sarscape software. The general workflow of the interferometric
162 processes applied in this study is summarized in figure 2.

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Figure 2. Flow chart of the interferometric processes.

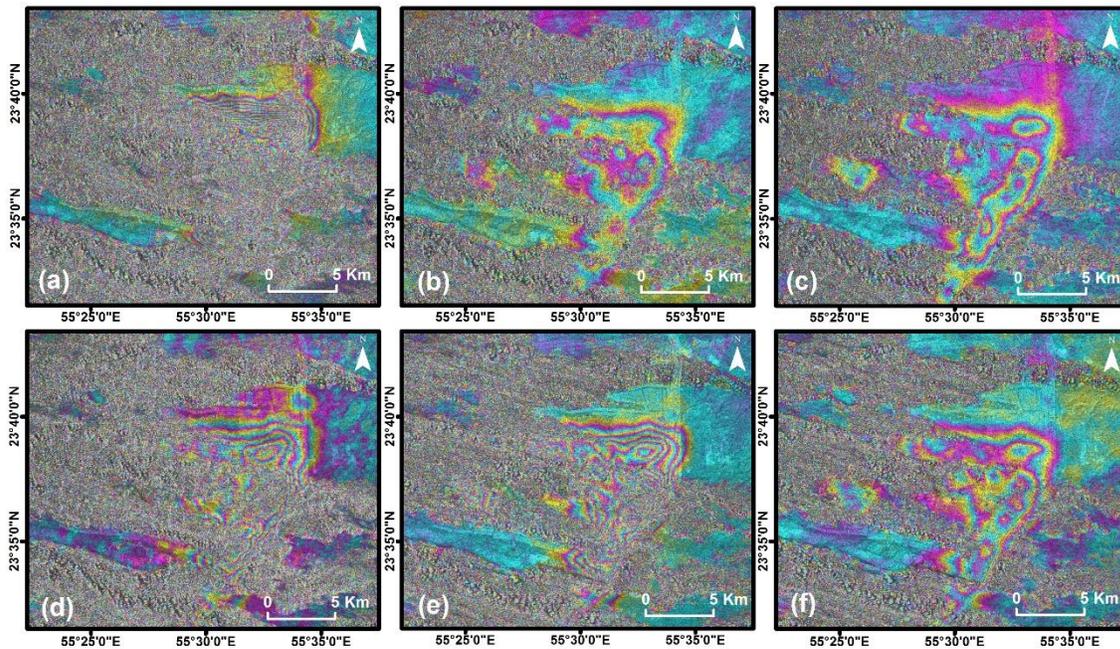
167

168 **Interferometric processing and results**

169 *ENVISAT data*

170 Since the acquisitions of the descending stack were temporally irregular with a gap
171 period from July 2004 to February 2006 (figure 4b), the thresholds for interferogram
172 generation were set to 800 days temporal / 360 m. perpendicular baseline in order to avoid
173 underlaps in the stack. The coherence maps showed that the alluvial gravels were remaining
174 generally coherent for more than two years in short baseline interferometric pairs, but the
175 coherence of the sand dune areas was completely lost within 35 days. The block structured
176 farms located in the sandy areas maintained their coherence at moderate levels for 100 - 150
177 days.

178 Despite the low phase signal quality the differential interferograms revealed a fast
179 subsidence pattern extended throughout the whole farmland area with the deformation
180 magnitude along the Line-Of-Sight (LOS) peaking at several points within the region,
181 especially along the west side of the highway that links the border settlements. In the central
182 and south regions, the deformations were even greater. But this observation relied only on
183 small temporally separated pairs as temporal decorrelation prevented from obtaining long
184 term estimations. A seasonal variation was also observed as the summer period pairs tended
185 to exhibit larger displacement signals than those of the winter season, reaching almost one
186 fringe per month south of the town of Al Wagan. If the observed subsidence was related to
187 groundwater pumping, the increased water demands of agriculture in a desert environment
188 during the hot and dry periods could explain this phenomenon. These primary observations
189 were consistent in all the temporally coherent interferograms of both the ascending and the
190 descending stacks, which reinforced the initial assumption of a fast subsidence phenomenon.



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193

194 Figure 3. Differential interferograms from ENVISAT pairs. a) From descending pair 3/4/2004 -
 195 9/3/2006, where temporal and fast fringes decorrelation are dominant in long time spans. b)
 196 Descending pair 29/11/2007 - 17/4/2008, indicative of the lower fringe rates in the winter seasons. c)
 197 Descending pair of 17/4/2008 – 26/6/2008. d) Descending pair 7/5/2009 - 11/2/2010, where the
 198 deformation pattern is clear despite the decorrelation noise. e) Ascending pair 3/4/2007 – 12/2/2008.
 199 f) Ascending pair 8/1/2008 – 27/5/2008.

200

201 The detection of this fast subsidence pattern depended on the maximum detectable
 202 deformation gradient, which theoretically is one fringe per pixel (Massonnet and Feigl,
 203 1998). Since in reality this threshold is significantly lower due to the presence of
 204 decorrelation noise in low coherent areas (Baran et al., 2005), and fast fringes decorrelation
 205 was observed in all the interferograms of more than two years separation, the images were

206 multilooked 1 x 5 (range x azimuth). For larger factors, the deformation pattern became
207 undetectable or ambiguous even in temporally adjacent pairs.

208 In order to obtain an initial estimation of the deformation velocities, interferometric
209 stacking was applied (Strozzi et al., 2001) after the phase unwrapping stage with pairs of
210 short perpendicular baselines (< 250 m). Due to temporal decorrelation and phase
211 unwrapping failures the reliability of the output linear deformation velocities was
212 compromised, however the areas of stability were delineated for subsequent reference point
213 selections.

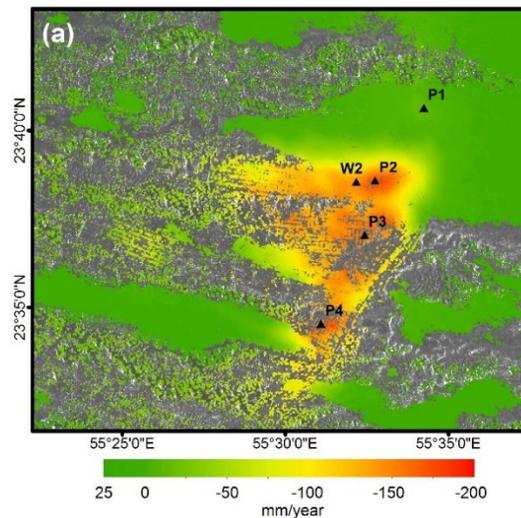
214 The interferograms of the descending stack were then used as inputs for the SBAS
215 algorithm. After discarding the pairs with significant unwrapping errors, 84 interferograms
216 were inverted in order to extract the displacement history and velocities of the multi-
217 temporally coherent points. The relative time-baseline position of the images is shown in
218 figure 4b. The resulting deformation maps showed that cumulative subsidence exceeded 1
219 meter and LOS displacement rates reached -18 cm/year in the most affected areas during the
220 sampling period, with the time series exhibiting a small seasonal effect. However, the
221 coherent pixel grid was sparse at the central and south parts of the study area where
222 subsidence appeared larger in the short temporal baseline interferograms, exhibiting
223 displacements equal or lower than those of the north part. Decorrelation and unwrapping
224 errors prevented from obtaining accurate estimates in those regions.

225 Figure 4a depicts the vertical deformation rates in the study area (valid under the
226 assumption of negligible horizontal component). Point P1 is located in the northernmost part
227 of the region where even though displacement rates were low, an accelerating subsidence
228 trend was observed after 2007. P2 is sited in the fast subsiding area near the town at the local

229 deformation maximum, while W2 is a monitored well within the same area. P3 and P4 are
230 representative points of the central and south subsiding regions respectively. Even though
231 according to the displacement history of figure 4c they exhibit lower or equal deformations
232 with point P2, it is noted that coherence levels were marginal in these areas and a sufficient
233 density of measure points was not achieved.

234 As a general observation, the linear model that was applied to the displacement
235 velocities could roughly capture the seasonal behaviour of the ground deformations. It is also
236 noted that a quadratic model was a more appropriate fit in the southern region proximate to
237 P4 as it was in the area of P1, with the subsidence phenomenon slightly accelerating after
238 2006. In either case, it is not certain whether this was due to the higher frequency of
239 acquisitions after that year (figure 4b), since the large temporally separated pairs of the
240 preceding gap period included many phase inconsistencies in the low coherent areas.

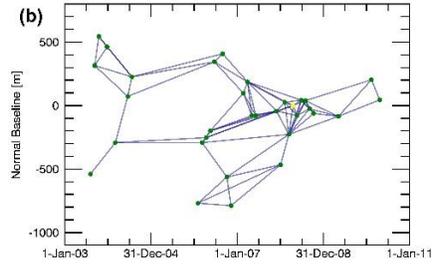
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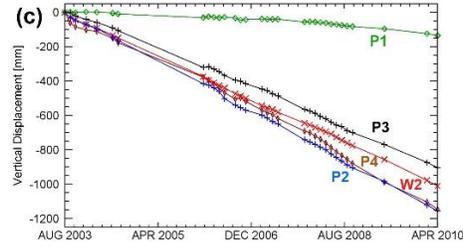
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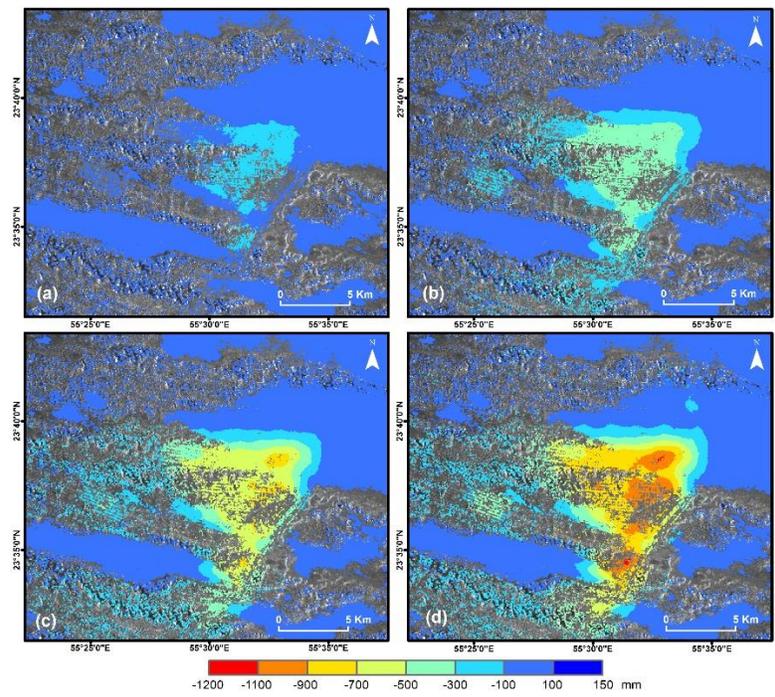


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248 Figure 4. a) Vertical displacement velocities of the period 2003 - 2010. b) Time position plot
 249 of the descending stack images, where nodes and arcs represent scenes and interferograms
 250 respectively. c) Displacement history of selected points in the study area.

251



252

253

254 Figure 5. Evolution of the LOS deformation pattern obtained from the descending
255 acquisitions. a) 7/8/2003 - 22/7/2004. b) 7/8/2003 - 22/6/2006. c) 7/8/2003 - 22/5/2008. d)
256 7/8/2003 - 22/4/2010.

257

258 The SBAS technique was applied with 20 interferograms of the ascending stack
259 within 400 days temporal separation and 330 m. maximum normal baseline for comparison
260 with the results obtained from the descending stack. The generated maps replicated the
261 ground distortion pattern of the previous implementation which verified that the vertical
262 movements were dominant, considering the similar incident angle of the ascending and
263 descending acquisitions. The cumulative LOS displacements derived from the two stacks
264 were comparable in the northern area during similar observation periods.

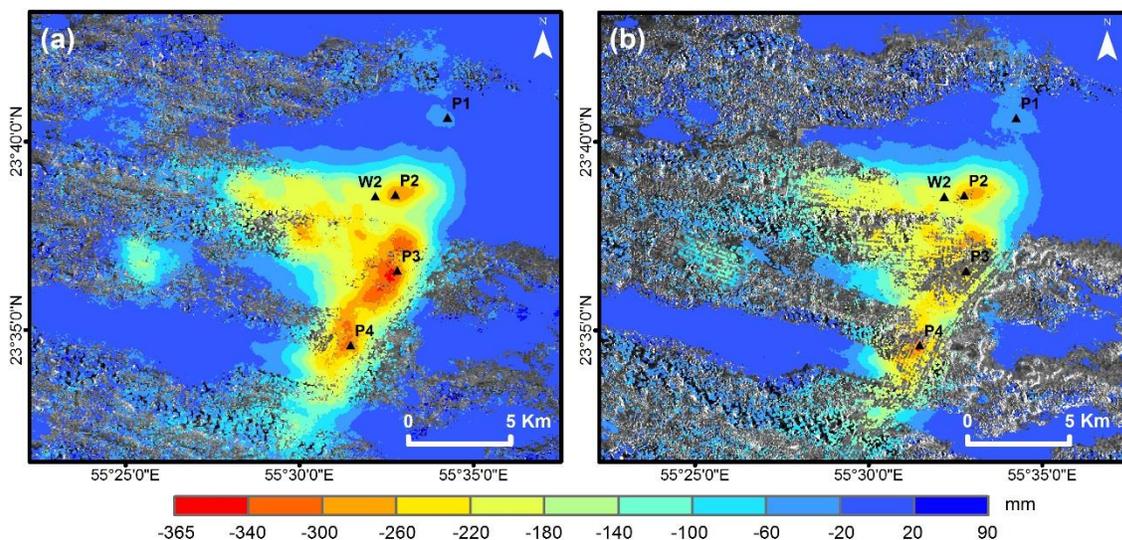
265 However, in the central and southern regions of the study area, where the
266 multitemporal coherence was increased in the ascending stack, the deformations of the
267 commonly coherent pixels appeared significantly larger than those estimated from the
268 descending stack. These local deviations were mostly attributed to phase inconsistencies
269 induced by decorrelation noise in the interferograms of the descending stack rather than to a
270 horizontal component of the terrain movements. The coherence maps of both stacks are
271 shown in figure 6.

272 The comparison of the cumulative LOS displacements between the two stacks in the
273 sampling period of the ascending acquisitions is demonstrated in figures 7a and b. Figure 7d
274 illustrates the vertical displacement history of the selected points in the study area after the
275 application of a linear model to the subsidence rates. Points P1, P2 and W2 follow the same

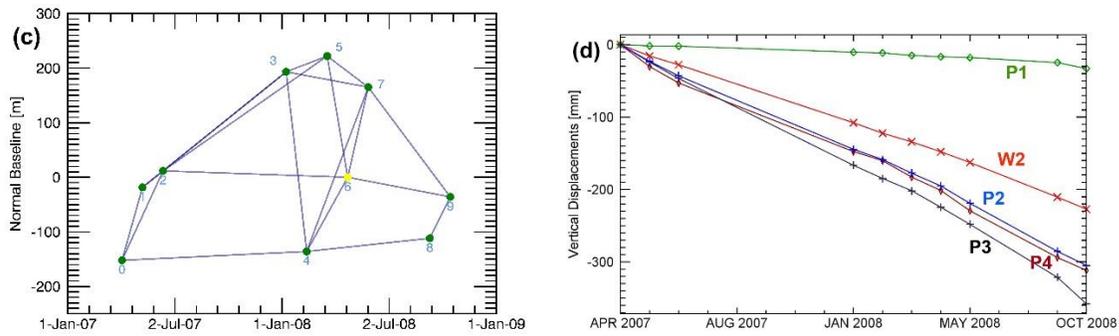
276 pattern as derived from the descending stack. The estimated displacements in these areas
277 were precisely estimated due to the alluvial sediment background, which was adequately
278 coherent in both stacks. Points P3 and P4 exhibited faster subsidence rates than P2,
279 confirming the primary estimate from the differential interferograms of both stacks that these
280 regions were subsiding faster than the northern area.

281 On the other hand, the ascending stack could not provide precise estimates due to the
282 sparsity of acquisitions whereas local phase unwrapping failures in low coherent regions
283 were neither absent in the ascending pairs. Indicatively, the (multitemporal) coherence values
284 of points P3 and P4 were 0.48 and 0.26 respectively contrasted to 0.23 for both points in the
285 descending stack. Even though the time span of the acquisitions included two summer
286 periods, higher order models were not applied due to the extended low coherent regions and
287 the small number of scenes. Nevertheless, this short time interval was considered too short
288 to draw safe conclusions about the exact deformation behaviour of these areas.

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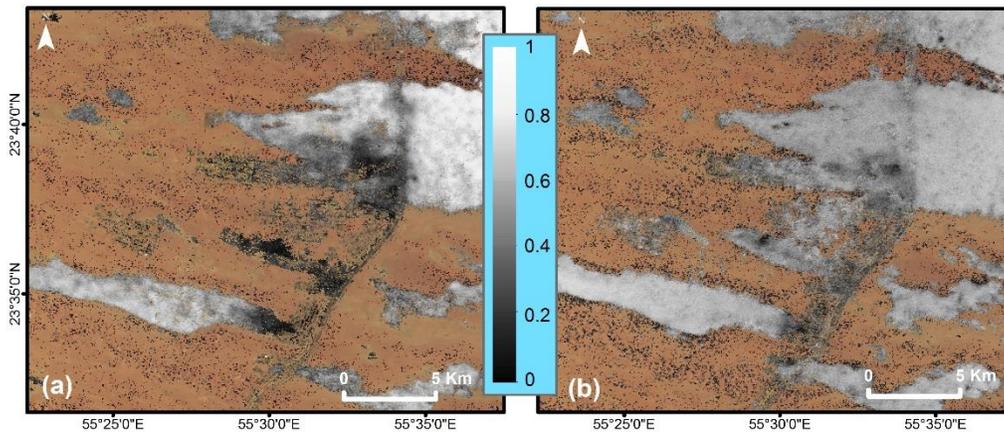
293 Figure 5. Comparison of the LOS displacements obtained from the two stacks with SBAS. a)

294 Ascending stack, 3/4/2007 - 14/10/2008. b) Descending stack, 29/3/2007 - 9/10/2008. c) Time

295 position plot of the descending stack images. d) Vertical displacement history derived from the

296 ascending stack (3/4/2007 - 14/10/2008).

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300 Figure 6. Multitemporal coherence. a) Descending stack. b) Ascending stack.

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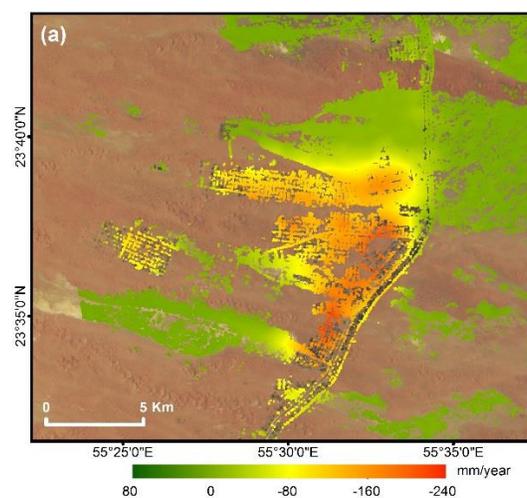
302 *ALOS data*

303 All possible interferograms within a temporal threshold of 700 days were calculated

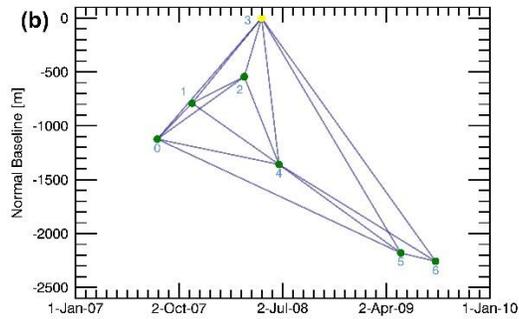
304 from the HH polarization images of the PALSAR acquisitions as all normal baselines were

305 less than 2300 m. Interferograms between the FBD and FBS mode acquisitions were created
306 after oversampling the FBD images in range to the pixel spacing of the FBS scenes (Werner
307 et al., 2007). As it was expected, the L-band interferograms exhibited higher coherence levels
308 and greater spatial coverage providing with a more comprehensive picture of the deformation
309 pattern in longer time spans than the ENVISAT pairs. After the phase unwrapping stage, 16
310 interferograms were inverted with SBAS.

311 The resulting deformation maps demonstrated an almost identical subsidence pattern
312 in terms of extent, even though non-systematic deviations were observed in comparison to
313 the C-band derived displacement maps that locally exceeded 2 cm when compared in similar
314 periods. Nonetheless, these differences were considered insignificant after taking into
315 account the small number of the scenes, the subsidence magnitude and the precision of
316 PALSAR in the achieved coherence levels (Pasquali et al., 2014). During the period 5/8/2007
317 – 10/8/2009, subsidence rates of 22 cm/year in the south, 21 cm/year in the central and 16
318 cm/year in the north farmland areas were observed from the ALOS velocity map (figure 8a).
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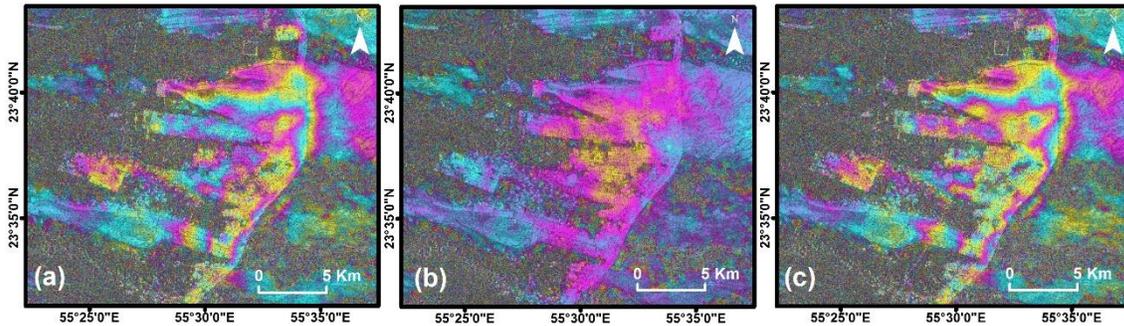
323 Figure 8. a) Subsidence rates obtained from the ALOS data. b) Time-baseline plot of the ALOS
 324 images.

325

326 *Sentinel-1A data*

327 The Sentinel-1A images were multilooked 8 x 2 (range x azimuth) due to the lower fringe
 328 rates and the differential interferograms (generated from the vertical polarization component)
 329 verified the continuation of the subsidence phenomenon in a similar extent but with a lower
 330 magnitude. Deformation fringes are clearly visible in pairs that include the summer season
 331 but they are not as prominent in the winter pair, indicating declining deformation trends that
 332 might have become a seasonal phenomenon, even though this is still an assumption. The
 333 generated wrapped interferograms are shown in figure 9. Results that provide stronger
 334 evidence about the recent evolution of the terrain distortions will be obtained when the
 335 number of acquisitions from Sentinel-1 will be sufficient for a time series interferometric
 336 analysis and the atmospheric artefacts in the interferograms are compensated for. Similar
 337 deformation trends were also observed in GPS data recordings of the Abu Dhabi GRS
 338 Network station of Al Wagan (personal communication with Abu Dhabi Municipality, Town
 339 Planning sector, Spatial Data Division).

340



341

342 Figure 9. Differential interferograms from Sentinel-1A pairs. a) 18/2/2016 - 2/12/2016. b) 2/12/2016
343 - 8/3/2017. c) 18/2/2016 - 8/3/2017.

344

345 **Correlation with groundwater levels**

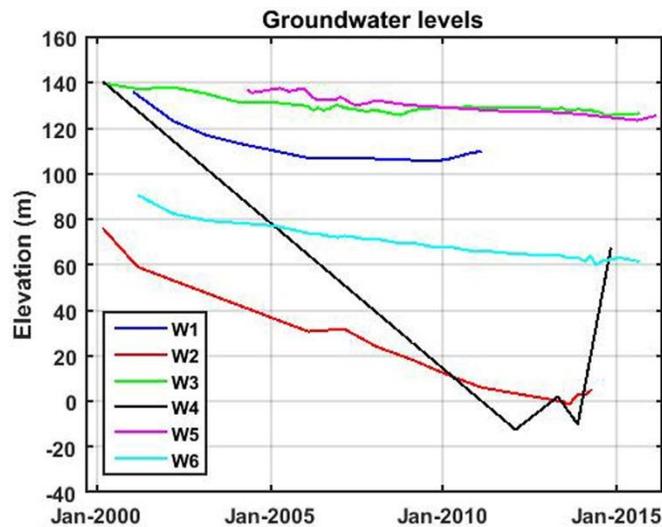
346 Since the subsidence phenomena in Al Wagan were attributed to water pumping from the
347 subsurface, correlations with the groundwater levels were sought. We considered the water
348 level data from local wells monitored by the Environmental Agency of Abu Dhabi (EAD).
349 Even though these measurements were not concurrent and with the same frequency for all
350 the wells, their temporal patterns was examined to define whether a relationship of the
351 groundwater level fluctuations with the terrain deformations exists.

352 The groundwater elevation reduces from east to west verifying that the aquifer surface
353 is inclined towards the west direction as mentioned in a previous study (Al Shahi, 2002). The
354 water level time series of the wells located within the north subsiding area are generally lower
355 contrasted to the surrounding region. This could be possibly related to groundwater extraction
356 which had started before the sampling period, but data prior to 2000 were not available. The
357 number of these wells is not adequate to extract conclusions about the spatial properties of

358 the local aquifer as these should be examined in a wider area level with more similar datasets,
359 but an indication of these spatial characteristics is locally provided.

360 The main observation regards the declining trends exhibited from all of the
361 monitoring wells of the farmland area and the surroundings from 2000 and for the main
362 period of the DInSAR observations 2003-2010. In some of the wells, a recovery trend was
363 observed after 2012 without reaching the levels of 2000-2002, but in other cases the levels
364 were still reducing until September 2015. Inspection of Landsat 5 and 7 imagery from 1984
365 - 2003 showed that the land use scenery changed dramatically during the 1990s, as it was the
366 period of the farmland expansion. However, quantitative evidence about the start of the
367 aquifer resource exploitation was not available. Plots of the time series of the water levels
368 from selected representative wells are illustrated in figure 8.

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371 Figure 10. Groundwater levels of the monitored wells shown in figure 1a.

372

373 **Conclusions**

374 Interferometric processing over the agricultural area of Al Wagan revealed a localized
375 subsidence phenomenon which is speculated to be groundwater related. It was observed from
376 different sensors that subsidence velocities reached 20 cm/year during 2003-2010 with the
377 terrain distortions exceeding -1 m. The land deformations of the study area would not have
378 been detected if the resolution was lowered by large multilooking factors. Even though the
379 area is not suitable for interferometric applications due to temporal decorrelation, DInSAR
380 and SBAS techniques may still provide a terrain distortion metric if no other tools are
381 available. The coherence maps were perfectly correlated with the land cover and the
382 geological maps of the region, an aspect worthy of future investigation in order to define and
383 develop the most appropriate techniques for mapping the surface deformations in the broader
384 area of the UAE. The Sentinel-1A data verified that the surface deformation phenomena are
385 still existent in the region constituting a potential hazard for the future.

386 According to the water level data, a correlation of the deformations with the
387 groundwater resources exists. However, these data do not suffice to explain the seasonality
388 of the subsidence pattern, while the fact that some areas are more affected than others even
389 though the water level declinations appear similar or lower suggests that the local geology
390 controls the phenomenon as well. The evaluation of the subsidence velocity field suggests
391 that the loose sandy material background is more prone to terrain distortions than the
392 compacted gravelly sediments of fluvial origin, but this also depends on the thickness of the
393 formations. This aspect will be further sought in the geotechnical engineering properties of
394 the different soil types of the study area by conducting in-situ sampling in the subsiding

395 farmland area and correlating the different subsidence rates with both the aquifer resources
396 depletion and the geophysical parameters of the varying geological background.

397 Another aspect that could be correlated with the deformation pattern of the farmlands
398 and the groundwater extraction for their maintenance regards the integration of optical
399 imagery within this study. Since the agricultural areas kept expanding during the past 20
400 years and the land cover changes are captured in Landsat images, this type of changes should
401 be reflected in the vegetation indices as well. It is expected that the time series analysis of
402 these indices for this temporal interval will be capable of providing more comprehensive
403 explanations about the excessive use of the aquifer resources in the area and the ground
404 deformations that were induced by this process in the study area. Future work will also
405 include accurate GPS measurements with Global Navigation Satellite Systems (GNSS)
406 equipment to validate the deformation time series that will be derived from the Sentinel-1A
407 data.

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