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*MANUSCRIPT*

A guideline to select an estimation model of daily global solar radiation between  
geostatistical interpolation and stochastic simulation approaches

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23

**Abstract**

24 This study compares geostatistical interpolation and stochastic simulation approaches for the  
25 estimation of daily global solar radiation (GSR) on a horizontal surface in order to fill in  
26 missing values and to extend short record length of a meteorological station. A guideline to  
27 select an approach is suggested based on this comparison. Three geostatistical interpolation  
28 models are developed using the nearest neighbor (NN), inverse distance weighted (IDW), and  
29 ordinary kriging (OK) schemes. Three stochastic simulation models are also developed using  
30 the artificial neural network (ANN) method with daily temperature (ANN(T)), relative  
31 humidity (ANN(H)), and both (ANN(TH)) variables as predictors. The six models are  
32 compared at 13 meteorological stations located across southern Quebec, Canada. The three  
33 geostatistical interpolation models yield better performances at stations located in a high  
34 density area of GSR measuring stations compared to the three stochastic simulation models.  
35 The guideline suggests an optimal approach by comparing a threshold distance, estimated  
36 according to a performance criteria of a stochastic simulation model, to the distance between a  
37 target and its nearest neighboring station. Additionally, the spatial correlation strength of daily  
38 GSRs and the at-site correlation strength between daily GSRs and the predictor variables  
39 should be considered.

40

41 *Keywords:* artificial neural networks, geostatistical interpolation, global solar radiation, spatial  
42 correlation, temperature, relative humidity.

## 43 1. Introduction

44 Global solar radiation (GSR) on a horizontal surface of the earth is an important variable for  
45 many analyses involving agricultural and plant growth, air and water temperatures,  
46 environmental and biological risk, and solar electric generation. However, instruments  
47 measuring solar irradiation (i.e., Kipp or Eppley pyranometers) are relatively expensive and  
48 difficult to manage [1], compared to those of common meteorological variables such as air  
49 temperature, precipitation, and relative humidity. Therefore, meteorological stations for GSR  
50 are generally less abundant than those for the common meteorological variables. Furthermore,  
51 observed GSR datasets are usually short timeseries and have large gaps of missing values.

52 Geostatistical interpolation approaches can be adopted to fill in missing values and to  
53 extend short record length of the GSR at a station using observed GSR data on the other  
54 stations located near the desired station. Kriging [2-7], nearest neighbor [4], and inverse  
55 distance weighted average [5,8,9] approaches have been applied frequently for the spatial  
56 interpolation.

57 At-site physical and statistical approaches can also be used for GSR simulations.  
58 Physical models (e.g., [10-12]) use complex physical interactions between the GSR and the  
59 terrestrial atmosphere, such as the Rayleigh scattering, radiative absorption by ozone and  
60 water vapour and aerosol extinction. Stochastic simulation models (e.g., [13-20]) use  
61 empirical relationships between GSR and meteorological covariables such as sunshine hours,  
62 temperature, and relative humidity at a desired station. This study considers stochastic  
63 simulation models as they are relatively simple to develop and require fewer input variables  
64 compared to physical models [16,17]. Although linear and non-linear regressions as well as  
65 artificial neural networks (ANNs) can be employed to drive empirical relationships between  
66 the common meteorological variables and the GSR, many studies [16-18,20-22] have shown

67 the superiority of ANN approaches to regression-based approaches.

68 Sunshine duration, one of the most explanatory variables for GSR simulation  
69 [18,21,23], has not been recorded at most meteorological stations in Canada since 1999 due to  
70 its difficulty of measurement [1]. Temperature [13-15,17-19,24-28] and relative humidity  
71 [18,21] are alternative covariables although they have weaker correlation with GSR compared  
72 to sunshine duration [18].

73 Geostatistical interpolation and statistical simulation approaches for GSR estimation  
74 have been applied separately in many studies, however, they have been rarely compared in an  
75 application study. Therefore, this study compares geostatistical interpolation and statistical  
76 simulation approaches to fill in missing values and to extend short record length of daily GSR  
77 timeseries. The spatial interpolation approaches considered include the nearest neighbor, the  
78 inverse distance weighted, and the ordinary kriging methods. The stochastic simulation  
79 models include three ANN-based models with daily temperature and/or daily relative  
80 humidity as input variables. The six models are applied at 13 meteorological stations located  
81 across southern Quebec (45.1~50.3 °N and 64.2~79.0 °W), Canada. Furthermore, a guideline  
82 to choose an approach between the geostatistical interpolation and the statistical simulation  
83 approaches is provided for the estimation of daily GSR on the study area.

84

## 85 **2. Methodologies**

### 86 2.1 Geostatistical interpolation models

87 Three geostatistical interpolation models are developed based on nearest neighbor (NN),  
88 inverse distance weighted (IDW), and ordinary kriging (OK) schemes for daily GSR. The NN  
89 model employs the simplest algorithm among the three models. This model selects the value  
90 of the nearest station to the location of interest and does not consider the other values of

91 neighboring stations in order to yield a piecewise-constant interpolation map.

92 The IDW interpolation algorithm adopts the assumption that the interpolation value at  
 93 a location of interest is inversely proportional to the distances of nearby stations. The  
 94 interpolation value of the model is a weighted average of the values of multiple stations and  
 95 the weight assigned to each nearby station diminishes as the distance from the interpolation  
 96 point to that station increases. The IDW model interpolates the daily GSR value  $R(x_0)$  at an  
 97 ungauged location  $x_0$  from observations  $R(x_i)$  at locations  $x_1, \dots, x_n$  as follows:

98

$$99 \quad \hat{R}(x_0) = \sum_{i=1}^n w_i R(x_i) \quad (1)$$

$$100 \quad w_i = \frac{1/d_i}{\sum_{j=1}^n 1/d_j}, \quad i = 1, 2, \dots, n \quad (2)$$

101

102 where  $\hat{R}(x_0)$  is an interpolated value of  $R(x_0)$  and  $d_i$  represents distance between  $R(x_0)$   
 103 and  $R(x_i)$ .

104 Kriging is a geostatistical interpolation technique based on the linear least square  
 105 estimation algorithm. Ordinary kriging (OK) is the most common among many kriging  
 106 approaches. OK estimates the best linear unbiased estimator based on a linear model. The  
 107 interpolation value of the OK at a location  $x_0$  is given by the following equation:

108

$$109 \quad \hat{R}(x_0) = \begin{pmatrix} w_1 \\ \vdots \\ w_n \end{pmatrix}^T \begin{pmatrix} R(x_1) \\ \vdots \\ R(x_n) \end{pmatrix} \quad (3)$$

110

111 where  $w_1, \dots, w_n$  are the weights of the OK that fulfill the unbiased condition  $\sum_{i=1}^n w_i = 1$ . The

112 weights are obtained by the below OK equation system:

113

$$114 \begin{pmatrix} w_1 \\ \vdots \\ w_n \\ \mu \end{pmatrix} = \begin{pmatrix} \gamma(x_1, x_1) & \cdots & \gamma(x_1, x_n) & 1 \\ \vdots & \ddots & \vdots & \vdots \\ \gamma(x_n, x_1) & \cdots & \gamma(x_n, x_n) & 1 \\ 1 & \cdots & 1 & 0 \end{pmatrix}^{-1} \begin{pmatrix} \gamma(x_1, x_0) \\ \vdots \\ \gamma(x_n, x_0) \\ 1 \end{pmatrix} \quad (4)$$

115

116 where  $\mu = E[R(x)]$  is a Lagrange parameter employed to minimize the kriging error under

117 the unbiased condition, which is assumed to be an unknown constant in the OK.  $\gamma(x_i, x_j)$  is a

118 variogram function to calculate the spatial dependency between  $R(x_i)$  and  $R(x_j)$ . Several

119 variogram functions are available such as exponential, Gaussian, and spherical models. In this

120 study, the spherical variogram function is selected based on trial and error examination. The

121 variogram is estimated for each day, based on observed daily GSR dataset of nearby stations.

122 The detail descriptions of variogram models and ordinary kriging can be found in [29,30].

123 To verify the interpolation performances of the three models, a leave-one-out cross-

124 validation approach is employed. Among the observations at  $n$  stations, GSR values of one of

125 those stations are interpolated using the observations at the remaining  $n-1$  stations. This

126 process is repeated for all the observation stations. The interpolated  $\hat{R}(x_i)$  is then compared

127 to the associated observation  $R(x_i)$  at each station in order to evaluate the performance of

128 the interpolation models.

129 2.2 Stochastic simulation models

130 Three stochastic simulation models are developed to estimate daily GSR using the ANN  
 131 approach as a transfer function and daily maximum and minimum temperatures and/or daily  
 132 mean relative humidity as input variables. Feed forward ANNs have been frequently  
 133 employed to simulate GSR [16-18, 20, 21] from the meteorological input variables. This  
 134 study also employs a three-layer feed forward ANN model, which includes an input layer, a  
 135 single hidden layer, and an output layer of computation nodes. The ANN models are trained  
 136 by the Bayesian regularization backpropagation (BRBP) algorithm, which is a network  
 137 training function that updates the weight and bias values according to the Levenberg-  
 138 Marquardt optimization [31]. An important issue in ANN modelling is the determination of  
 139 the number of hidden nodes. Fletcher and Goss [32] suggested that the optimal number of  
 140 hidden nodes could be within  $(2p^{0.5} + o) \sim (2p+1)$ , where  $p$  and  $o$  are the numbers of  
 141 independent and dependent variables, respectively. The hyperbolic tangent sigmoid function  
 142 is employed for the hidden layer and the linear function is used for the output layer. Detailed  
 143 descriptions of these various activation functions are provided in [31]. The three ANN models  
 144 used to simulate daily GSR series from daily meteorological variables are as follows:

$$145 \hat{R} = ANN(T_{\max}, T_{\min}, R_a) \quad (5)$$

$$146 \hat{R} = ANN(H, R_a) \quad (6)$$

$$147 \hat{R} = ANN(T_{\max}, T_{\min}, H, R_a) \quad (7)$$

148  
 149  
 150 where  $T_{\max}$  and  $T_{\min}$  are daily maximum and minimum temperatures ( $^{\circ}\text{K}$ ) and  $H$  is daily  
 151 mean relative humidity in a given day. The ANN represents the three-layer feed forward  
 152 ANN trained by the BRBP algorithm. The  $R_a$  is the solar irradiation on a horizontal surface

153 at the top of the atmosphere, which is a function of latitude and Julian day of a site. It is  
154 calculated by using the standard geometric method provided by [33]. The details of the  
155 method are also available in [20]. The three models are called ANN(T), ANN(H), and  
156 ANN(TH) hereafter based on employed input variables. The numbers of hidden nodes  
157 selected for the ANN(T), ANN(H), and ANN(TH) are 4, 3, and 5, respectively, based on a  
158 trial-and-error procedure.

159 Daily GSR at ungauged stations that measure other predictor variables can be  
160 simulated using a regional ANN-based model. This model is calibrated based on all available  
161 GSR and meteorological observations for the region of interest, which allows for the  
162 simulation of GSR at ungauged stations where covariables are available. For instance, Fortin  
163 et al. [17] and Jeong et al. [20] tested a regional ANN-based model to simulate daily GSR for  
164 regional areas located in eastern Canada. They calibrated this model using observations  
165 obtained from a set of stations and validated the model using those obtained from a different  
166 station set. Regional ANN(T), ANN(H), and ANN(TH) models are also considered in this  
167 study using a leave-one-out training procedure. In this approach, the regional ANN models  
168 are trained for a given station using observations of all the remaining stations for the  
169 calibration period, which is repeated for all stations. In the regional ANN models, mean GSR  
170 varies according to  $R_a$ , which is a function of the latitude of each station.

171

### 172 2.3 Model evaluation measures

173 Simulation performances are evaluated using the mean bias error (MBE), root mean square  
174 error (RMSE), and R-square (coefficient of determination). The MBE and RMSE are given by  
175 the following equations:

176

$$177 \quad \text{MBE} = \frac{1}{m} \sum_{i=1}^m (\hat{R}_i - R_i) \quad (8)$$

$$178 \quad \text{RMSE} = \left[ \frac{1}{m} \sum_{i=1}^m (\hat{R}_i - R_i)^2 \right]^{0.5} \quad (9)$$

179

180 where  $R_i$  and  $\hat{R}_i$  are observed and simulated daily GSR values and  $m$  is the record length.

181 R-square (coefficient of determination) is the squared value of the (Pearson's product-

182 moment) linear correlation coefficient between observed and simulated values. It can provide

183 the proportion of explained variance of observations by an applied model and is defined by

184 the following equation:

185

$$186 \quad r^2 = 1 - \frac{\sum_{i=1}^m (\hat{R}_i - R_i)^2}{\sum_{i=1}^m (R_i - \bar{R})^2} \quad (10)$$

187

188 where  $\bar{R}$  is the mean of the observed GSR values.

189

### 190 3. Study area and data

191 Daily GSR, maximum and minimum temperatures, as well as mean relative humidity are

192 obtained from 13 meteorological stations of Environment Canada (EC) located between

193 latitude 45.1°N to 50.3°N and longitude 64.2°W to 79.0°W (i.e. Southern Quebec, Canada).

194 The daily GSR and the two predictor variables are obtained for the analysis period from 2003

195 to 2010. Figure 1 shows the locations of the 13 stations across southern Quebec, which have

196 less than 10 % of missing data of daily maximum and minimum temperatures, relative  
197 humidity, and GSR for the analysis period. The figure also distinguishes the GSR stations  
198 excluded from this analysis due to more than 10 % of missing values of any of the three  
199 previously mentioned variables. Stations recording daily maximum and minimum  
200 temperatures and relative humidity are presented when they have less than 50 % of missing  
201 data for the analysis period. The south of Quebec is the most populated and productive area in  
202 the province and has higher density of observation stations than the rest of the province. The  
203 three stochastic simulation models are calibrated and validated on the 2003-2007 and 2008-  
204 2010 periods, respectively. The three geostatistical interpolation models interpolate the daily  
205 GSR for each observation station by using the leave-one-out cross-validation method for the  
206 2008-2010 period. Performances of the six models are finally compared for the 2008-2010  
207 validation period at the 13 selected stations.

208 Table 1 presents the information (station identification numbers, latitudes, longitudes,  
209 and altitudes) of the 13 stations in ascending order of their latitudes. Annual and seasonal (i.e.,  
210 DJF for winter, MAM for spring, JJA for summer, and SON for autumn) averages of daily  
211 GSR for the 2003-2010 period are also provided. In general, it is known that GSR decreases  
212 as latitudes increase; however, the annual or seasonal GSR of the stations do not show a clear  
213 decrease as their latitudes increase because the study area covers a small range of latitude (5.2  
214 degree). Furthermore, some stations are located in complex climate conditions directly  
215 affected by the St-Lawrence River and convections from the Atlantic Ocean (i.e., stations 8, 9,  
216 11, and 13) or from the continent (i.e., station 10). As daily GSR and predictor variables are  
217 not linearly correlated, linear correlation coefficients between the solar transmissivity (i.e., the  
218 ratio of incoming GSR on the surface of the earth to solar irradiation at the top of the  
219 atmosphere) and diurnal temperature range (DTR;  $T_{max} - T_{min}$ )  $\gamma(R/R_a, DTR)$  series as well

220 as between solar transmissivity and daily mean relative humidity  $\gamma(R/R_a, H)$  series are  
221 presented. The solar transmissivity and DTR have positive correlations since a cloudy day has  
222 smaller GSR, and also a smaller DTR due to a lower  $T_{max}$  during the day by blocking sunlight  
223 as well as a higher  $T_{min}$  during the night by preventing radiative cooling, when compared to a  
224 clear day. However, the solar transmissivity and mean daily relative humidity are negatively  
225 correlated since a clear day has less humidity than a cloudy day. Correlations between daily  
226 GSR and DTR and relative humidity of station 11 are weaker than those of the other stations.  
227 This station is located on the south shore of the Lower St-Lawrence valley, which has  
228 complex climate conditions affected by the river and convections from the continent and the  
229 Atlantic Ocean.

230

## 231 **4. Results**

### 232 4.1 Comparison of model performances

233 Table 2 presents performance measures of the geostatistical interpolation models for each  
234 station for the 2008-2010 validation period. The NN, which is the simplest approach, yields  
235 the worst performance, whereas the OK, which is the most sophisticated approach, shows the  
236 best performance, although there is a larger magnitude of MBE for OK than for IDW. The  
237 three models generally produce larger MBE at stations 10, 12, and 13, which have larger  
238 differences in annual mean GSR values compared to the other stations (see Table 1 for values  
239 of annual mean GSR and Figure 1 for station locations). The three models yield small RMSEs  
240 at stations located in the high density area (i.e., stations 1 to 9), whereas they yield large  
241 RMSEs at stations located in the low density area (i.e., stations 10 to 13). In this low density  
242 area, the nearest stations to the stations 10-13 are located within a distance of 482.0, 235.2,  
243 236.1, and 363.8 km, respectively, whereas those to the stations 1-9 are located within 100 km.

244 It is notable that the performance of the geostatistical interpolation models depends on the  
245 density of the network of stations and on the statistical homogeneity of GSR values.

246 Table 3 presents performances of the stochastic simulation models for each station for  
247 the calibration and validation periods. The differences of the performances between the  
248 calibration and the validation periods are modest for each model and for each station,  
249 implying that the three models are calibrated well without overfitting and that they have good  
250 generalization ability for a new data set. Average differences between the two periods are 0.35  
251 MJ/m<sup>2</sup>/day for MBE, 0.20 MJ/m<sup>2</sup>/day for RMSE, and 1.8 % for R-square. Among the three  
252 stochastic simulation models, the ANN(TH) uses both temperature and relative humidity as  
253 input variables and yields the best performance. The ANN(T) and the ANN(H), which employ  
254 either temperature or relative humidity as an input variable, yield similar performances for all  
255 stations, except for the station 11, which showed the weakest correlations between daily GSR  
256 and predictors among the selected stations (Table 1).

257 Figure 2 compares RMSEs of the geostatistical interpolation and the stochastic  
258 simulation models for each station at annual and seasonal scales for the validation period. The  
259 geostatistical interpolation models generally show better performance than the stochastic  
260 simulation models for the stations located in the high density area (i.e., stations 1 to 9).  
261 However, these models perform differently for the stations located in the low density area (i.e.,  
262 stations 10 to 13). The poor performances of the geostatistical interpolation models in the low  
263 density area are expectable as the models use spatial correlations, which exponentially  
264 decrease as distance increase. Especially in spring and summer, RMSEs of the geostatistical  
265 models tend to be larger at stations 10, 12, and 13 than the stochastic simulation models,  
266 indicating that spatial correlation structures of GSR are weaker in spring and summer than in  
267 winter and autumn. However, the stochastic simulation models have similar performances for

268 all stations, except for the station 11, as they only use at-site relationship between the daily  
269 GSR and the input variables.

270 Figure 3 presents scatter plots between observed and simulated daily GSRs for the  
271 validation period and for stations 7 and 13, which are located in the high density and the low  
272 density (i.e., north-eastern boundary) areas, respectively. In Figures 3a to 3f, the geostatistical  
273 interpolation models show better agreement with the 1:1 line than the stochastic simulation  
274 models at the station 7. As shown in Tables 2 and 3, the OK model yields the best  
275 performance among the six models at this station. However, the geostatistical interpolation  
276 models tend to overestimate the observed values at the station because, on average, daily  
277 GSRs at the station are smaller than its neighboring stations (see Table 1). In Figures 3g to 3l,  
278 the geostatistical interpolation models show worse agreement with the 1:1 line than the  
279 stochastic simulation models at the station 13. The ANN(TH) model yields the best  
280 performance among the six models.

#### 281 4.2 Guidelines for model selection

282 RMSEs and R-squares of daily GSR series between a target and a neighboring station versus  
283 their distance for all possible pairs of stations are presented in Figure 4, at an annual and  
284 seasonal scales for the 2003-2010 period. In other words, the RMSEs and R-squares of the  
285 NN method are calculated, under the assumption that the pair of stations includes the target  
286 station and its nearest neighbor. Trend lines of RMSEs and R-squares are estimated by the  
287 logarithmic and exponential functions respectively using the non linear least square algorithm.  
288 Equations and R-squares of the trend lines for annual and season scales are provided in the  
289 figures. Therefore, the trend lines provide approximate RMSEs or R-squares of the NN  
290 method for a target station with its nearest neighbor on the study area. For instance, according  
291 to the equation presented in Figure 4a, if an observed daily GSR value is available at the

292 nearest neighboring station located at a distance of 200km from a target station, the NN  
293 method can approximately simulate the daily GSR at the target station with an expected  
294 RMSE of 4.3 MJ/m<sup>2</sup>/day at an annual scale. Spatial correlation strengths vary between  
295 seasons. For instance, in winter and autumn, the spatial correlation structures are stronger than  
296 those in spring and summer. The study area usually shows more homogenized weather and  
297 solar radiation conditions in winter and autumn compared to spring and summer seasons  
298 because of less convection from Atlantic and/or continental sources.

299 Using the equations presented in Figure 4, a threshold distance (TD) between a target  
300 and its nearest neighboring station can be estimated according to a desired level of  
301 performance (i.e., RMSE or R-square) based on the NN model. In Table 4, estimated TDs of  
302 the NN model are presented based on the RMSEs of each ANN(T), ANN(H), and ANN(TH)  
303 models presented in Table 3. Based on the table, worse performances of the NN models are  
304 expected than the stochastic simulation models at stations 10, 12, and 13 as their nearest  
305 neighboring stations are located further than their TDs. Similarly, the NN model can yield  
306 slightly better performance than ANN(T) and ANN(H), but it can yield a worse performance  
307 than ANN(TH) annually at station 11. This can be explained by the NN model requiring  
308 nearest neighboring stations to be within 263 km for ANN(T), 245 km for ANN(H), and 212  
309 km for ANN(TH) at an annual scale, but the nearest station (i.e., station 9) is actually at a  
310 distance of 235.2 km from station 11.

311 The TDs presented in Table 4 can be used as a guideline to select an approach  
312 between geostatistical interpolation and stochastic simulation models by comparing estimated  
313 TDs to the distances of the nearest neighboring stations when filling in missing values and  
314 extending record length of daily GSR is required at an observation station. There are three  
315 possible cases; (1)  $TD > \text{distance of nearest neighboring station}$ ; (2)  $TD \approx \text{distance of nearest}$

316 neighboring station; (3)  $TD < \text{distance of nearest neighboring station}$ . For the first case,  
317 applying the geostatistical interpolation models is recommended. For instance, on average,  
318 better annual performances of the NN model can be expected than the ANN(T), ANN(H), and  
319 ANN(TH) when a nearest neighboring station is within 162, 164, and 121 km, respectively.  
320 However, the availability of predictor variables (i.e., temperature and/or humidity) of  
321 statistical simulation models and the seasonal spatial correlation strengths of geostatistical  
322 interpolation models should be considered to select an optimal approach. As ANN(TH) yields  
323 better performance than ANN(T) or ANN(H), the former's TD is shorter than the latter's.  
324 Shorter TDs are estimated in summer compared to the winter and autumn seasons due to a  
325 weak spatial correlation structure in summer. In the second case, applying more sophisticated  
326 geostatistical interpolation models (e.g., IDW and OK) than the NN model is recommended.  
327 As an example, at station 7, the IDW and OK models yield better performances, whereas the  
328 NN model yield a worse performance compared to the ANN(TH) annually (see Figure 3a).  
329 Finally, in the third case, applying stochastic simulation models is recommended as they  
330 generally can perform better than geostatistical interpolation models. Since the best  
331 performance model cannot always be applied for a specific period at a selected station due to  
332 a lack of available predictor variables and observed GSR values of neighboring stations, the  
333 TD criterion of the proposed guideline can be used to suggest an optimal approach. The  
334 guideline and TD can also be used for other GSR stations that were excluded in this analysis  
335 due to short record-length (Figure 1).

336 Under the assumption that a target station has only predictor variables, regional  
337 stochastic simulation models are developed using the GSR and predictor variables measured  
338 at the other stations. Table 5 presents annual performances of regional models for the 13  
339 stations and their TDs to the nearest neighboring stations to produce similar RMSEs to the

340 regional models. Among the regional ANNs, ANN(TH) yields the best performance while  
341 regional ANN(T) and ANN(H) yield similar performances to each other. Again, station 11  
342 shows the worst performance among the 13 stations. The RMSEs of the regional models are  
343 0.21~0.27 MJ/m<sup>2</sup>/day larger than those of the at-site models. The worse performances of  
344 regional ANNs are reasonable compared to the at-site ANNs as the regional ANNs at each site  
345 do not use the observed GSR data of that site for the model calibrations. Consequently, the  
346 TDs of the regional models are also 14.9~29.1 km longer than those of the at-site models.  
347 These TD values and the ones presented in Table 5 can thus be used to select an appropriate  
348 approach between geostatistical interpolation and regional ANN simulation approaches in  
349 order to estimate daily GSR at ungauged (or short-record) stations.

350

## 351 **5. Concluding remarks**

352 Geostatistical interpolation and stochastic simulation approaches are compared in this study to  
353 fill in missing values and to extend short record length of the daily global solar radiation  
354 (GSR). However, it is notable that the comparison is only based on the performances of two  
355 approaches because they have different application constraints and algorithms to each other.  
356 For instance, geostatistical interpolation approaches provide interpolated values at any point  
357 in a region including a target station; however, they need observations of daily GSR on the  
358 other stations located near the target station to estimate the spatial correlation structure.  
359 Stochastic simulation approaches provide estimated values only at the target station using  
360 observed daily GSR series as a dependant variable, and daily temperatures as well as  
361 humidity series as independent variables.

362 The simplest nearest neighbor (NN) model yields the worst performance, whereas the  
363 most sophisticated ordinary kriging (OK) model shows the best performance among the three

364 geostatistical interpolation models. The three geostatistical interpolation models generally  
365 yield smaller RMSEs at stations located in the high density area (i.e., stations 1-9) than those  
366 located in the low density area (i.e., stations 10-13). The difference of the performances of  
367 geostatistical interpolation models between the high and low density areas can be explained  
368 by the exponential decrease of the spatial correlations between stations as the distance  
369 increase. Among the three at-site stochastic simulation approach models, the ANN(TH) yields  
370 better performance than the ANN(T) and ANN(H), while the ANN(T) and ANN(H) yield  
371 similar performances to each other. The three stochastic simulation models produce similar  
372 performances for all stations, except for the station 11, which is exposed to a complex climate  
373 and showed weaker relationships between GSR and predictors. Regional stochastic simulation  
374 models can simulate daily GSR series at stations, where only predictor variables are available;  
375 however, the performances of the regionalized models are worse than the at-site models.

376 In the comparison between the geostatistical interpolation and the stochastic  
377 simulation models, the geostatistical models perform better at stations located in the high  
378 density area, but they perform worse at stations located in the low density area, compared to  
379 the stochastic simulation models. Equations that can approximately estimate the RMSE and  
380 R-square based on the NN model using the distance between a target and its nearest  
381 neighboring station are presented. By using these equations, a guideline is suggested to select  
382 an approach between the geostatistical interpolation and the stochastic simulation approaches.  
383 A stochastic simulation approach is recommended when the distance between a target and its  
384 nearest neighboring station is longer than the threshold distance (TD) estimated according to  
385 the RMSE of a stochastic simulation model. In the opposite case, when the TD is longer than  
386 the distance between a target and its nearest neighboring station, a geostatistical interpolation  
387 approach is recommended. When the TD is similar to the distance between a target and its

388 nearest neighboring station, more sophisticated geostatistical interpolation models (e.g., IDW  
389 and OK) have generally proven to perform better than a stochastic simulation model in this  
390 study.

391 Although, this study suggests a guideline to select an appropriate simulation approach  
392 for daily GSR between geostatistical interpolation and stochastic simulation approaches, the  
393 guideline is dependent on the spatial correlation strength of daily GSRs and the at-site  
394 correlation strength between daily GSRs and the predictor variables. It is proved that spatial  
395 correlation strengths for seasonal scales have stronger in winter and autumn compared to  
396 those in spring and summer in the study area. Simulation of sub-daily GSR will be considered  
397 in future work as it is generally more important than daily GSR to estimate solar energy  
398 output due to the non-linear relationship between the radiance and the energy output.

399

#### 400 **Acknowledgements**

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404 his help.

405

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492  
493

494 Table 1

495 Station identification number, location information (latitude, longitude, and altitude) as well  
 496 as annual and seasonal averages of daily GSR of the selected stations for the 2003-2010  
 497 analysis period. Linear correlation coefficients between solar transmissivity and diurnal  
 498 temperature range (DTR)  $\gamma(R/R_a, DTR)$  as well as between solar transmissivity and relative  
 499 humidity  $\gamma(R/R_a, H)$  are also provided.

No.	Station #	Lat (°N)	Lon (°W)	Altitude (m)	Average GSR (MJ/m <sup>2</sup> /day)					$\gamma(\frac{R}{R_a}, DTR)$	$\gamma(\frac{R}{R_a}, H)$
					Annual	Winter	Spring	Summer	Autumn		
1	7022579	45.05	-72.86	152.4	11.61	4.97	14.58	18.39	8.51	0.53	-0.63
2	702FQLF	45.12	-74.29	49.1	12.83	6.21	16.00	19.91	9.21	0.54	-0.66
3	702LED4	45.29	-73.35	43.8	13.07	6.44	16.49	20.27	9.09	0.50	-0.63
4	7024280	45.37	-71.82	181.0	11.63	5.46	14.53	18.04	8.49	0.58	-0.65
5	702327X	45.72	-73.38	17.9	12.79	6.45	16.10	19.49	9.11	0.52	-0.68
6	7025442	46.23	-72.66	8.0	12.72	6.28	16.07	19.65	8.89	0.52	-0.62
7	7011983	46.69	-71.97	61.0	11.67	5.72	14.99	17.83	8.13	0.63	-0.68
8	701Q004	46.78	-71.29	91.4	11.50	5.44	15.26	17.77	7.54	0.53	-0.72
9	7041JG6	47.08	-70.78	6.0	11.97	5.56	15.29	18.57	8.48	0.51	-0.65
10	7086716	48.25	-79.03	318.0	11.83	5.39	15.96	18.49	7.49	0.54	-0.73
11	7056068	48.51	-68.47	4.9	12.16	4.99	16.02	19.49	8.15	0.29	-0.47
12	7065639	48.84	-72.55	137.2	12.63	6.21	17.09	18.78	8.44	0.52	-0.68
13	7044328	50.27	-64.23	11.0	11.28	4.18	15.46	17.92	7.56	0.58	-0.59
Avg.		46.86	-72.05	83.2	12.13	5.64	15.68	18.82	8.39	0.52	-0.65

500

501

502 Table 2  
 503 Performance measures of the three geostatistical interpolation models for the 2008-2010  
 504 period.

station #	MBE (MJ/m <sup>2</sup> /day)			RMSE (MJ/m <sup>2</sup> /day)			R-square (100 <sup>-1</sup> %)		
	NN	IDW	OK	NN	IDW	OK	NN	IDW	OK
1	-0.67	-0.33	-0.45	2.53	2.34	2.15	0.91	0.92	0.94
2	0.08	0.51	0.16	2.24	2.53	2.38	0.92	0.90	0.91
3	0.63	0.40	0.15	2.44	2.03	2.17	0.92	0.94	0.93
4	-0.26	-0.50	-0.37	2.34	2.60	2.50	0.92	0.89	0.90
5	-0.24	0.18	-0.14	2.36	1.80	1.86	0.91	0.94	0.94
6	1.46	0.61	0.71	2.99	2.01	1.98	0.89	0.94	0.95
7	-1.03	-1.24	-1.14	2.32	2.17	2.01	0.93	0.94	0.95
8	-0.37	-0.13	0.09	2.24	1.91	1.89	0.92	0.95	0.94
9	0.26	0.22	0.26	2.60	2.68	2.49	0.90	0.90	0.91
10	-0.78	-0.42	-0.48	5.50	5.49	5.10	0.58	0.55	0.60
11	-0.48	-0.32	-0.20	4.55	4.18	3.79	0.73	0.76	0.80
12	0.34	0.55	0.67	4.93	4.22	3.75	0.67	0.73	0.79
13	-0.45	-0.82	-0.66	5.51	5.80	5.33	0.62	0.55	0.61
avg.	-0.12	-0.10	-0.11	3.27	3.06	2.88	0.83	0.84	0.86

505

506

507 Table 3  
 508 Performance measures of the three stochastic simulation models during the 2003-2007  
 509 calibration and the 2008-2010 validation periods.

	MBE (MJ/m <sup>2</sup> /day)			RMSE (MJ/m <sup>2</sup> /day)			R-square (100 <sup>-1</sup> %)		
	ANN(T)	ANN(H)	ANN(TH)	ANN(T)	ANN(H)	ANN(TH)	ANN(T)	ANN(H)	ANN(TH)
Calibration period (2003~2007)									
1	0.03	0.31	0.02	3.76	3.82	3.19	0.76	0.76	0.83
2	-0.04	0.05	0.00	3.94	3.88	3.18	0.78	0.78	0.86
3	-0.06	-0.04	-0.04	4.48	4.49	3.86	0.74	0.74	0.81
4	-0.03	0.05	0.01	3.77	3.77	3.06	0.75	0.75	0.84
5	0.03	0.05	-0.01	3.77	3.72	3.14	0.79	0.79	0.85
6	0.00	-0.07	0.07	3.88	4.12	3.26	0.77	0.74	0.84
7	0.09	0.02	-0.04	3.77	3.94	3.01	0.78	0.76	0.86
8	-0.04	0.00	0.00	4.02	3.62	3.16	0.75	0.80	0.85
9	0.00	-0.01	0.01	4.39	4.35	3.64	0.72	0.72	0.80
10	-0.01	0.03	-0.02	3.62	3.14	2.78	0.80	0.85	0.88
11	0.04	0.00	-0.04	4.87	4.53	4.35	0.68	0.72	0.75
12	-0.07	0.10	0.03	3.83	3.66	3.20	0.78	0.80	0.85
13	-0.10	0.03	0.00	3.75	3.76	3.37	0.80	0.80	0.84
avg.	-0.01	0.04	0.00	3.99	3.91	3.32	0.76	0.77	0.83
validation period (2008~2010)									
1	0.66	0.78	0.53	3.92	4.08	3.37	0.79	0.77	0.84
2	-0.36	-0.09	-0.28	3.89	3.84	3.25	0.77	0.77	0.84
3	-0.75	-0.56	-0.66	3.65	3.80	3.22	0.81	0.79	0.85
4	-0.01	-0.27	-0.21	3.62	3.81	3.03	0.79	0.76	0.85
5	-0.34	0.21	-0.08	3.40	3.36	2.94	0.80	0.81	0.85
6	-0.07	-0.32	-0.09	3.81	3.89	3.12	0.77	0.77	0.85
7	-0.86	-0.42	-0.67	3.42	3.63	2.84	0.81	0.77	0.87
8	0.40	-0.40	-0.06	3.76	3.75	3.28	0.79	0.78	0.83
9	0.26	0.29	0.28	4.24	4.19	3.58	0.74	0.74	0.81
10	0.33	0.50	0.49	3.81	3.46	3.08	0.76	0.81	0.85
11	-0.52	-0.53	-0.60	4.76	4.63	4.32	0.69	0.71	0.75
12	0.36	0.08	0.25	3.77	3.57	3.18	0.78	0.79	0.84
13	0.18	0.70	0.51	3.79	4.11	3.54	0.79	0.77	0.82
avg.	-0.05	0.00	-0.05	3.83	3.85	3.29	0.78	0.77	0.83

510 Table 4

511 Threshold distances (TDs; in km) between the target and nearest neighboring station for the NN model to produce same RMSEs as the  
 512 ANN(T), ANN(H), and ANN(TH), respectively. The TDs are calculated by equations presented in Figure 4 with the RMSEs of the  
 513 three stochastic simulation models at each station and each time scale for the validation period presented in Table 3.

	ANN(T)					ANN(H)					ANN(TH)				
	annual	winter	spring	summer	autumn	annual	winter	spring	summer	autumn	annual	winter	spring	summer	autumn
1	167.0	356.1	209.9	109.5	217.7	181.9	212.4	199.2	145.8	249.9	123.7	194.3	146.3	88.5	168.2
2	164.0	303.6	196.6	127.9	183.9	160.1	185.4	224.0	104.7	214.1	115.8	158.6	158.8	75.2	154.6
3	144.0	231.5	176.4	101.1	207.8	156.2	157.3	221.6	94.4	248.4	114.0	121.0	163.2	66.3	183.8
4	142.0	205.5	170.9	108.8	161.6	156.9	130.0	201.3	112.1	211.1	103.0	120.7	133.3	72.1	129.2
5	125.6	155.1	142.1	102.6	154.7	123.4	126.4	145.0	94.7	178.6	98.2	99.2	126.1	71.8	118.8
6	157.2	183.8	197.3	116.7	184.5	163.9	106.6	183.2	134.3	250.4	108.2	116.5	127.8	85.4	135.6
7	126.9	208.6	145.5	84.5	128.6	142.2	119.7	170.5	97.1	132.7	92.7	99.3	120.4	59.3	80.2
8	153.2	260.3	195.0	115.8	168.3	152.1	137.2	207.1	122.2	143.4	117.8	137.5	155.5	96.2	102.8
9	198.4	349.5	231.0	148.1	274.5	192.8	231.4	200.0	186.0	190.6	138.8	207.0	153.5	123.2	138.4
10	157.3	78.1	212.1	113.0	201.8	129.9	112.6	151.9	101.1	184.1	106.1	64.2	135.2	81.0	135.6
11	263.0	162.5	242.8	270.5	313.2	245.6	125.1	253.6	261.4	237.8	206.8	116.2	212.3	216.8	212.2
12	153.8	145.4	174.3	114.8	194.1	137.9	110.0	158.1	109.0	187.5	111.6	107.5	132.1	89.0	124.1
13	155.2	99.4	172.2	153.5	142.3	184.7	102.2	162.6	218.4	134.1	136.1	79.7	138.4	145.7	102.5
avg.	162.1	210.7	189.7	128.2	194.9	163.7	142.8	190.6	137.0	197.1	121.0	124.7	146.4	97.7	137.4

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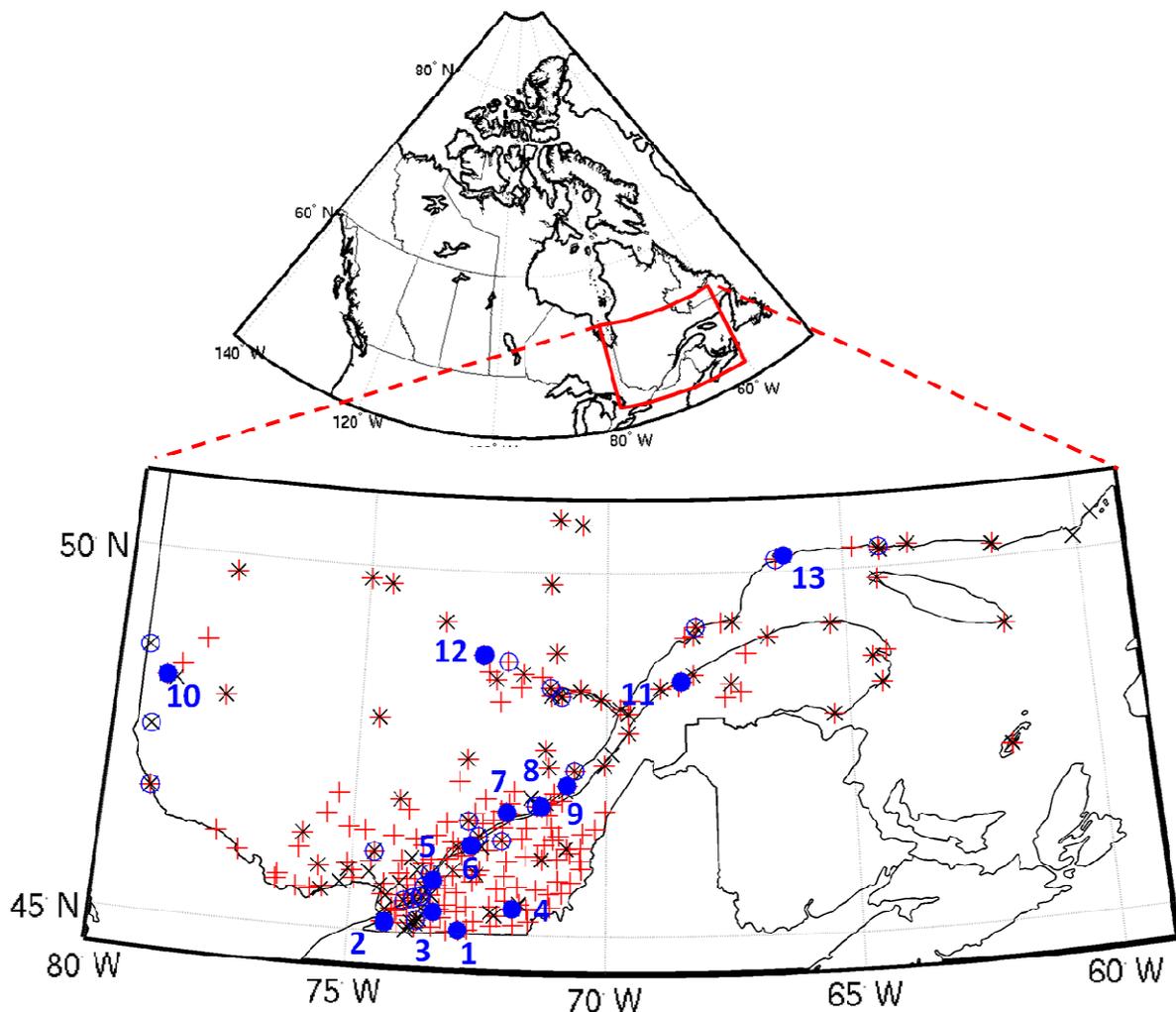
515 Table 5

516 Annual RMSEs of regional ANN(T), ANN(H), and ANN(TH) and their threshold distances  
 517 (TDs) of the nearest stations for the NN model to produce same RMSEs as the regional  
 518 models. The TDs are calculated by equations presented in Figure 4 with the RMSEs of the  
 519 three regional models at each station during the validation period.

	regional ANN(T)		regional ANN(H)		regional ANN(TH)	
	RMSE	TD	RMSE	TD	RMSE	TD
	(MJ/m <sup>2</sup> /day)	(km)	(MJ/m <sup>2</sup> /day)	(km)	(MJ/m <sup>2</sup> /day)	(km)
1	3.89	164.0	4.14	187.7	3.48	131.4
2	3.97	171.1	3.91	165.8	3.30	119.4
3	3.65	144.3	3.77	153.9	3.19	112.4
4	3.98	172.2	3.93	168.1	3.38	124.6
5	3.46	130.3	3.37	124.0	2.97	99.6
6	3.89	164.4	3.84	160.1	3.16	110.5
7	4.47	225.1	3.64	143.7	3.39	124.8
8	3.75	152.1	4.10	183.4	3.23	114.8
9	4.28	202.7	4.19	193.0	3.63	142.3
10	3.81	157.4	3.72	149.9	3.20	112.7
11	5.66	428.8	5.43	377.4	5.63	421.5
12	3.97	171.2	3.82	158.3	3.44	128.7
13	3.78	154.6	4.92	286.3	4.34	209.2
avg.	4.04	187.6	4.06	188.6	3.56	150.1

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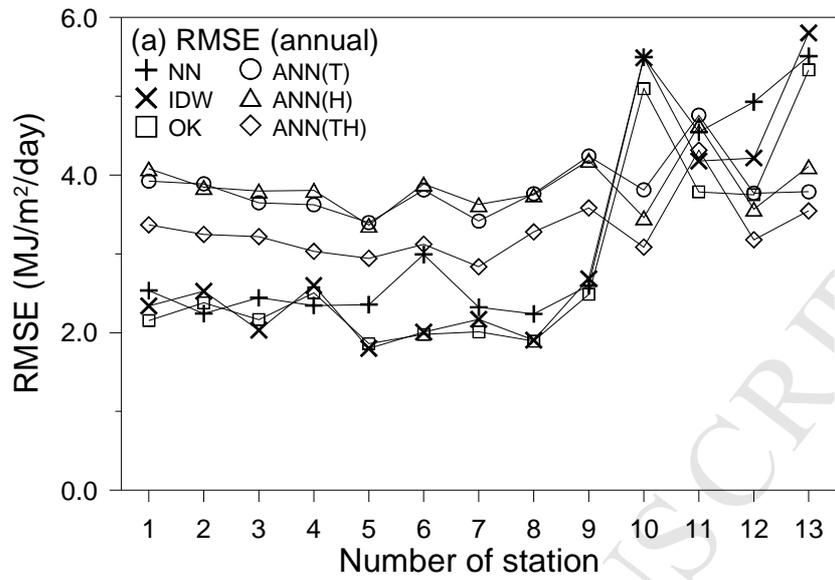


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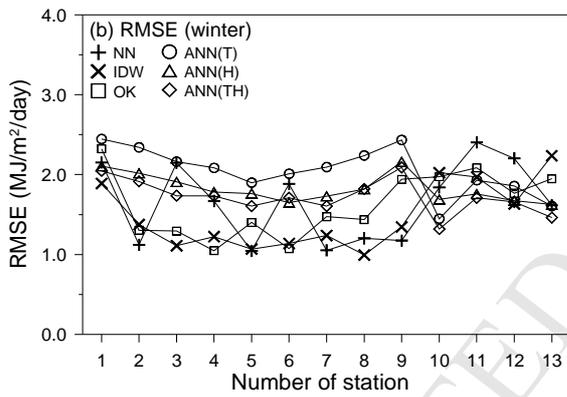
523 Fig. 1. Map of southern Québec, Canada. Stations are presented by red '+' and black 'x' when  
 524 they have observed daily temperature and relative humidity respectively, when they have less  
 525 than 50 % of missing data for the common analysis period (from 2003 to 2010). Blue filled  
 526 circles represent the selected meteorological stations, which have less than 10 % of missing  
 527 data of daily temperature, relative humidity, and GSR for the common analysis period. Blue  
 528 open circles represent the GSR stations excluded in this analysis due to more than 10 %  
 529 missing values of any of the three previously mentioned variables.

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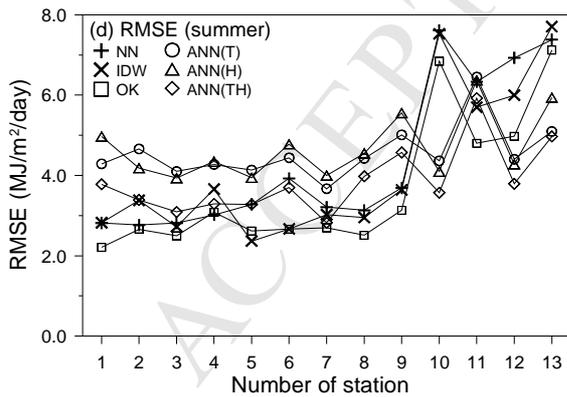
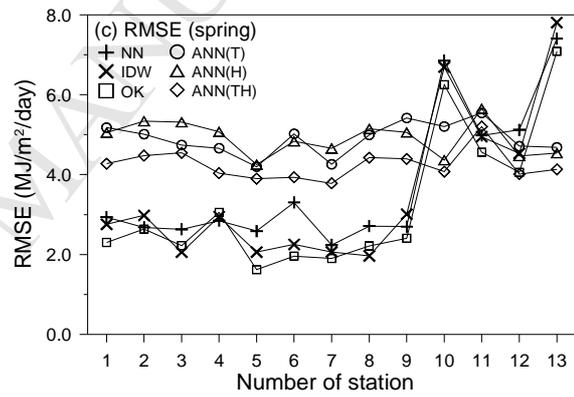
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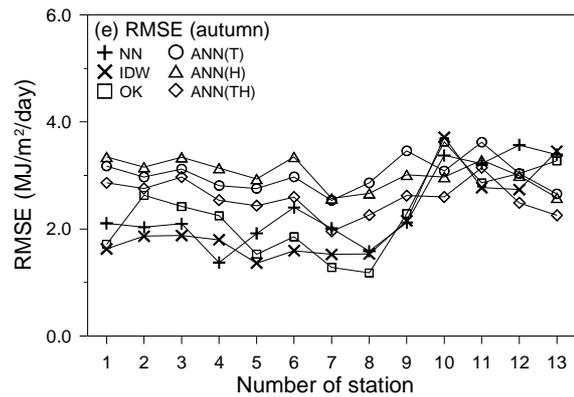
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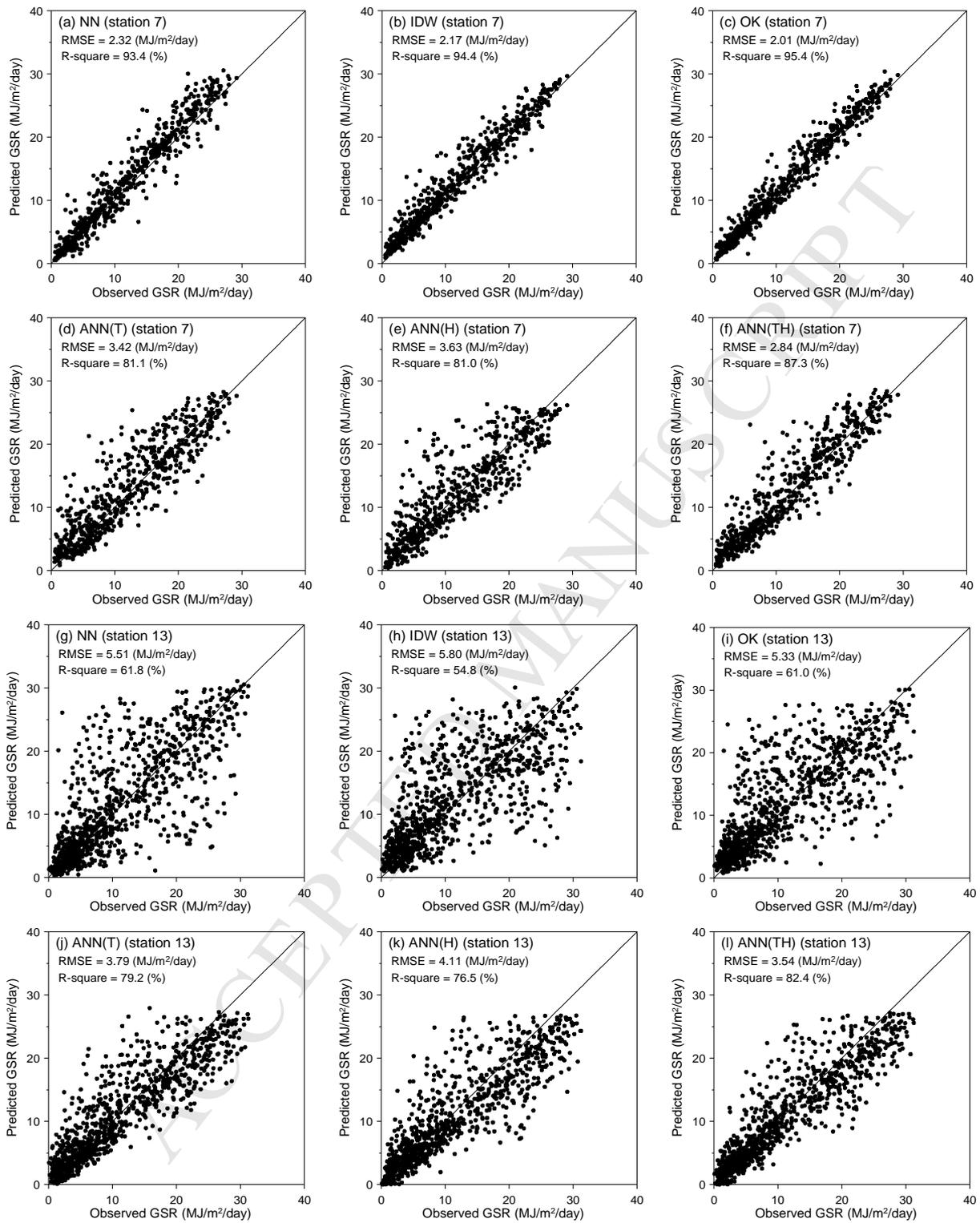


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535 Fig. 2. RMSEs of the three geostatistical interpolation (NN, IDW, and OK) and the three  
 536 stochastic simulation (ANN(T), ANN(H), and ANN(TH)) models for each stations at (a)  
 537 annual and (b-e) seasonal scales during the 2008-2010 validation period.

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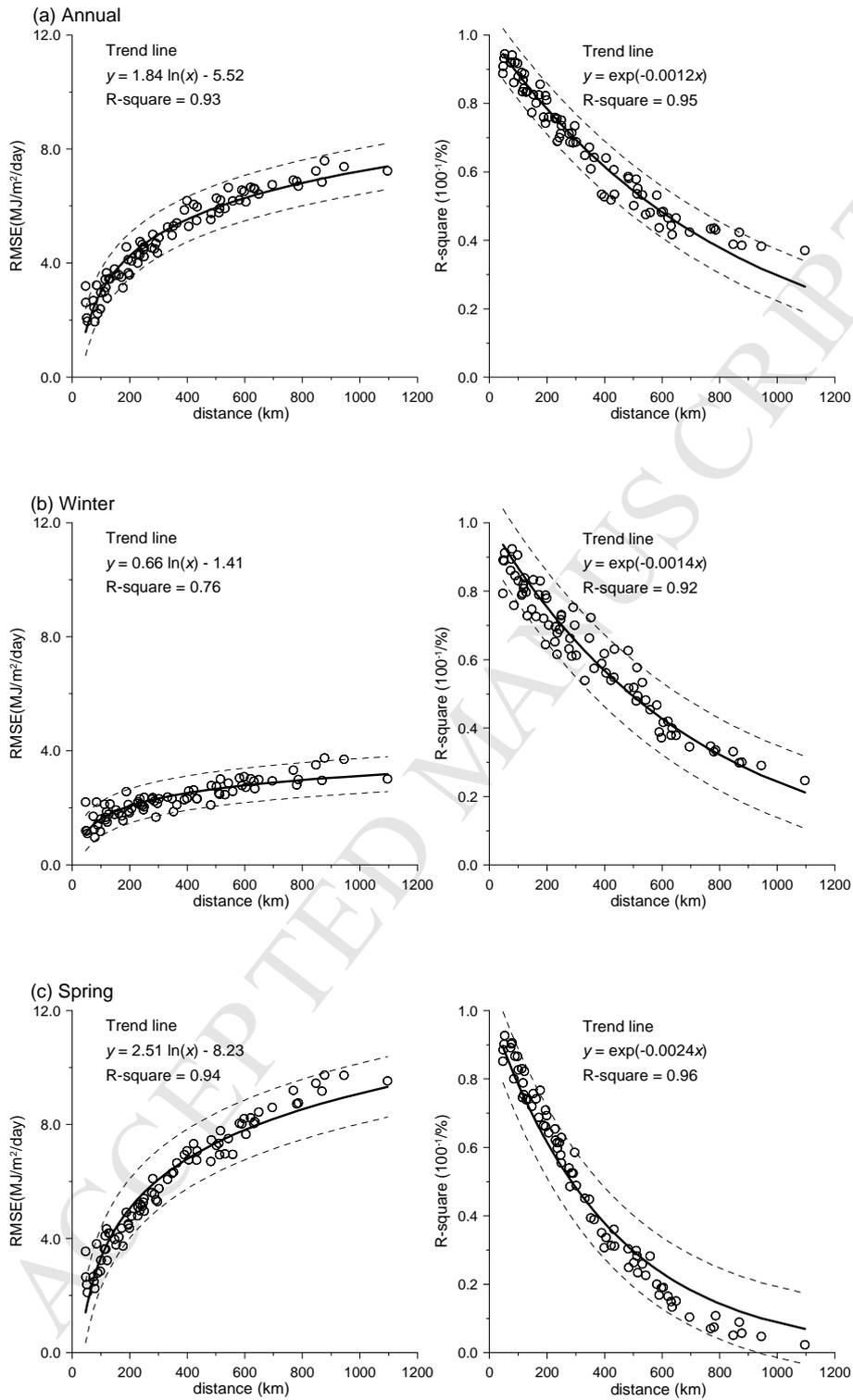
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543 Fig. 3. Scatter plots of daily GSRs between observation and predictions by the three  
 544 geostatistical interpolation and the three stochastic simulation models at the stations 7 (a-f)  
 545 and 13 (g-l).

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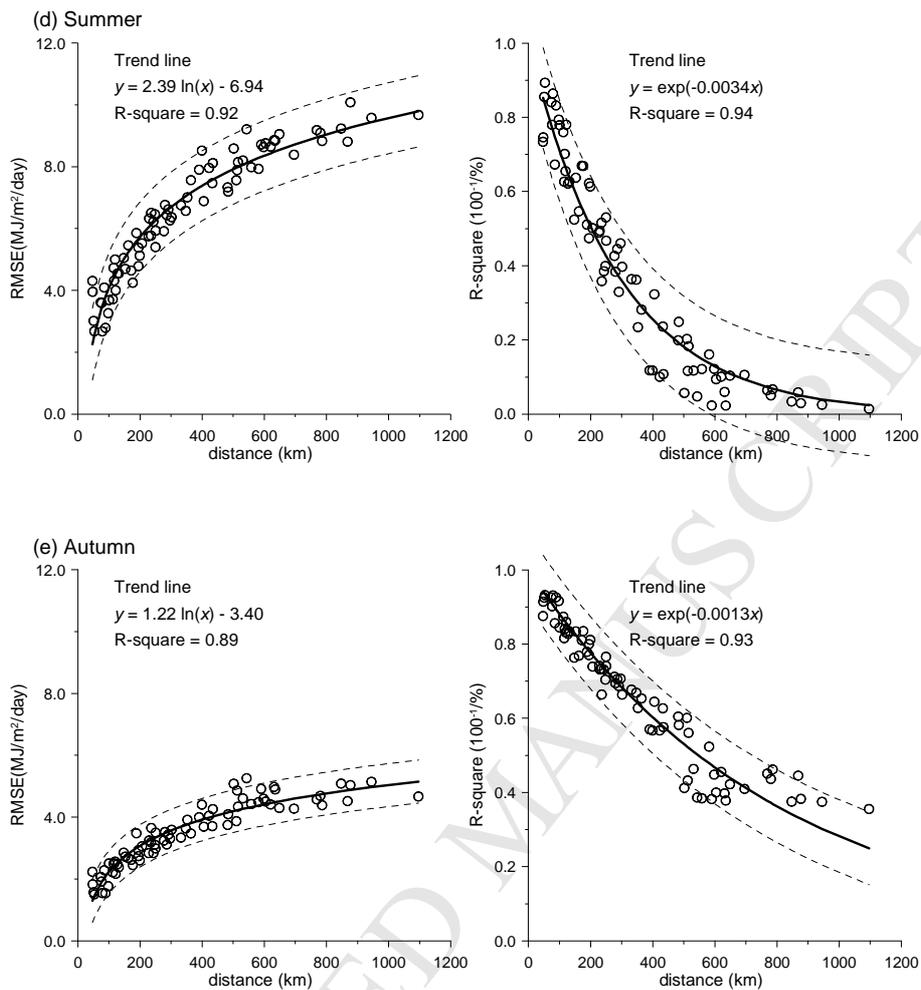
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556 Fig. 4. RMSE and R-square of daily GSR series between a target and its neighboring stations  
 557 versus the distance between the two stations for all possible combinations during the analysis  
 558 period from 2003 to 2010. Trend lines of RMSEs and R-squares are estimated by logarithmic  
 559 and exponential functions, respectively. Equations and R-squares of the trend lines are  
 560 presented on the figures. The dotted lines represent the 95 % confidence interval of the trend  
 561 lines.

562

- Models for estimating daily global solar radiation are investigated.
- Geostatistical interpolation and stochastic simulation approaches are compared.
- Geostatistical models yield better performance at a high density measurement area.
- Stochastic models show better performance at a low density measurement area.
- A guideline to select an optimal estimation approach is then suggested.