## 1750 YEARS OF LARGE RAINFALL EVENTS INFERRED FROM PARTICLE SIZE AT EAST LAKE, CAPE BOUNTY, MELVILLE ISLAND, CANADA

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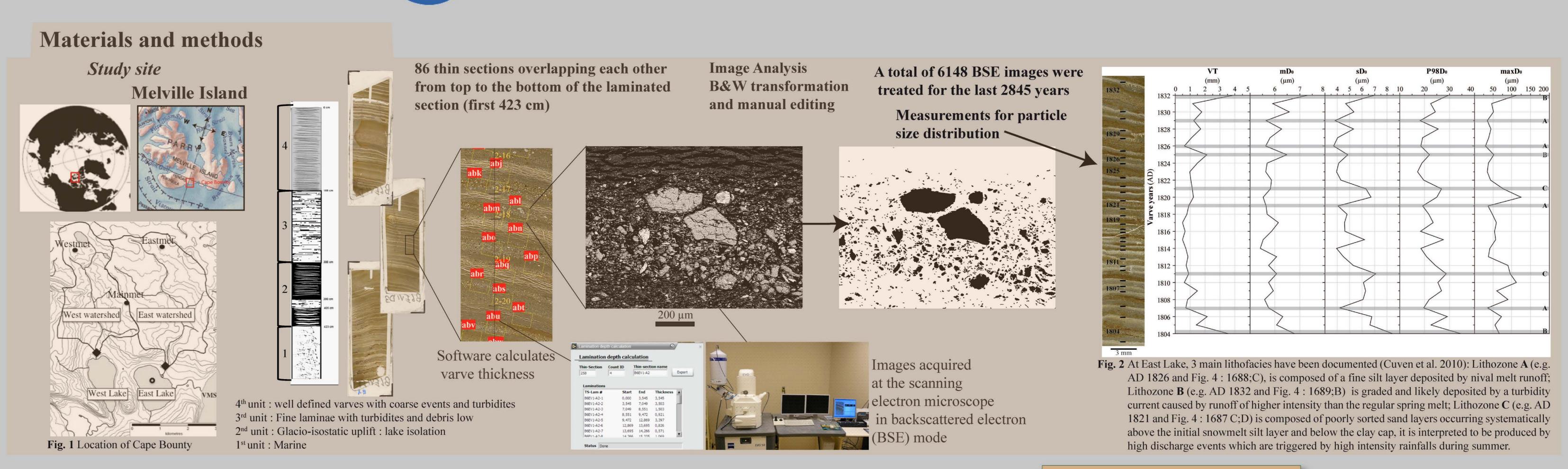






#### Introduction

Major changes occurred in Arctic climate during the 20th century, including an important increase of surface air temperature particularly between 1970 and 2000 (Moritz et al. 2002). To better understand these changes beyond the available short instrumental record, there is a need to obtain long-term proxy climate records across a wide geographical area. In the western Canadian High Arctic, annually laminated sediments from lakes are the only available records with precise chronological control in an area that lacks tree-ring and ice core records. However, caution is needed when calibrating instrumental data with varve thickness (VT), since sediment accumulation can result from different hydroclimatic and geomorphic processes, such as snowmelt, rain events and landslides. In addition to VT, grain size measured at the annual scale is a sedimentary parameter that may be used as a potential variable to reconstruct paleo-hydroclimatic condition. Here, we use the methods pioneered by Francus (1998) together with a new image acquisition and analysis software developed at INRS (Francus and Nobert 2007), to obtain grain size measurements within each varve in a long Arctic lake sequence in order to demonstrate its potential for paleoenvironmental reconstructions.



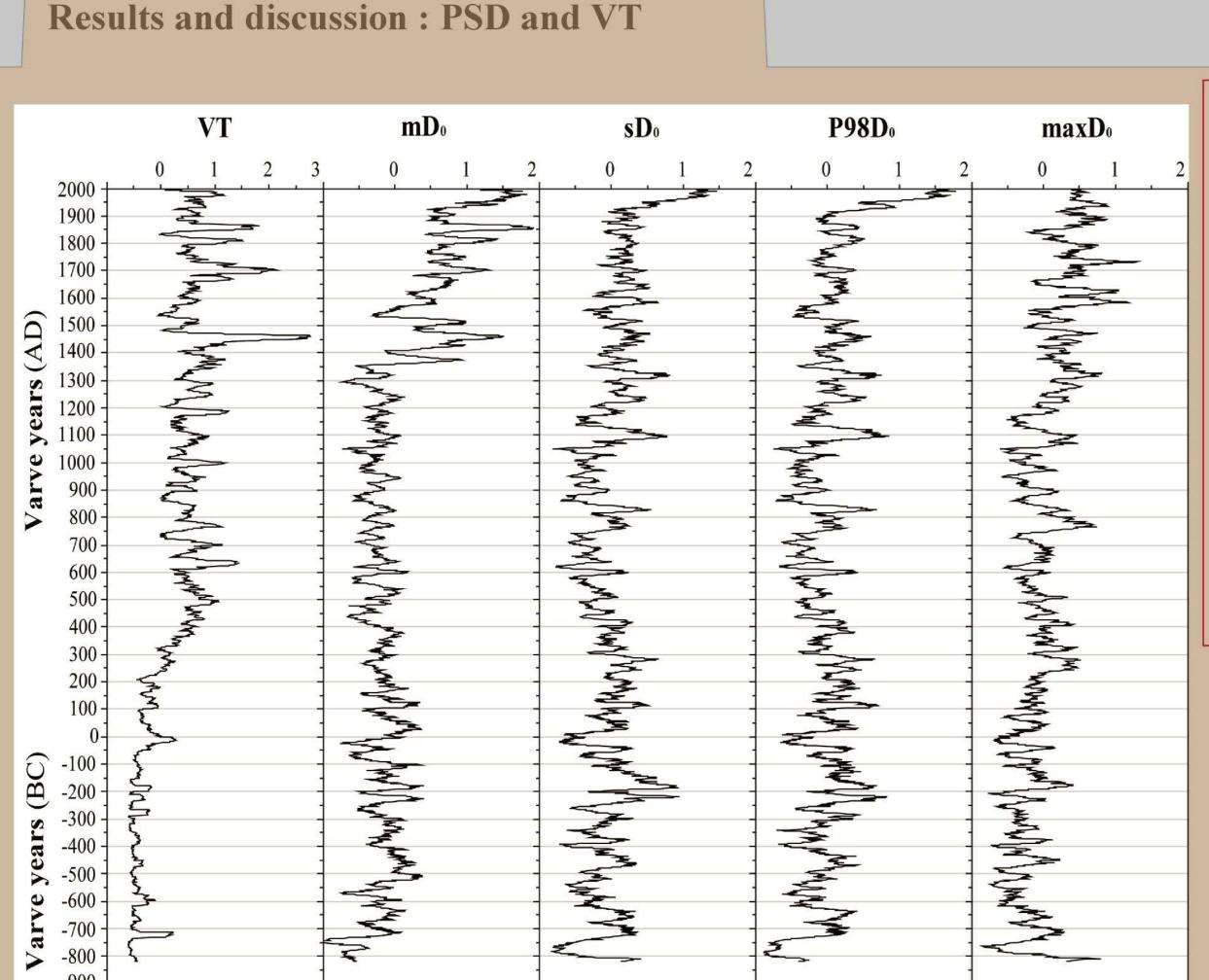


Fig. 3 East Lake varve thickness (VT) and four particle size indices for the last 2,845 years. Data are standardized relative to the mean and standard deviation. A 20-year running mean is applied on the series

## **Main observations**

- The long-term evolution of VT and PSD indices at Cape Bounty is strikingly different.

- P98D<sub>0</sub> and sD<sub>0</sub> vary similarly through time and they are strongly correlated (r = 0.83, p < 0.0001). They increase to maximum at the beginning of the 20th century with highest values during the 1990s.

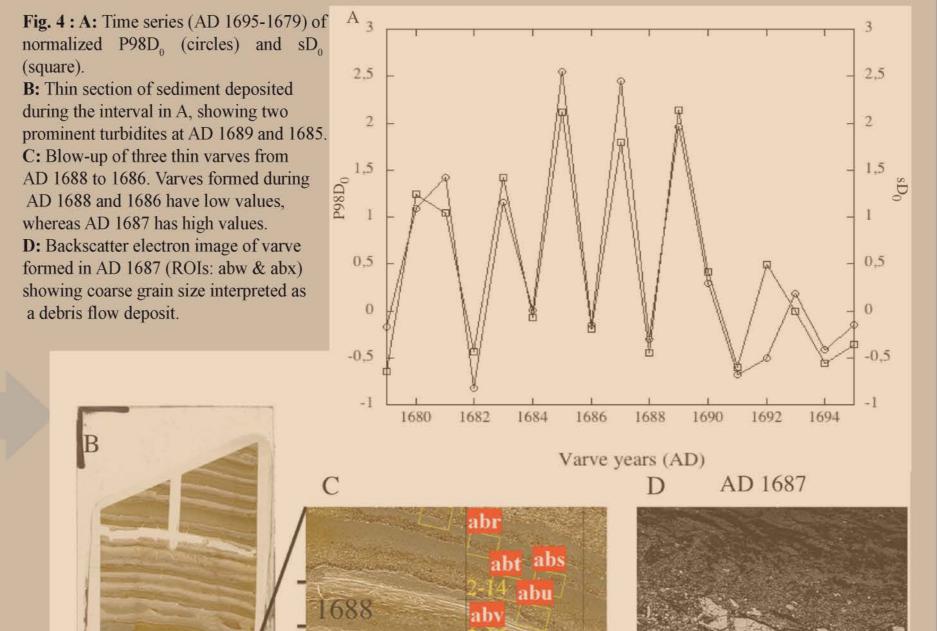
- The clear step change of mD<sub>0</sub> around AD 1350 does not correspond to a detected change in known boundary conditions in the watershed.

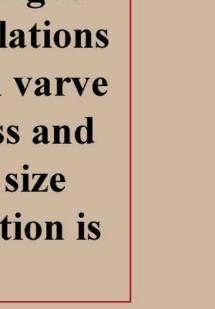
Table 1: the Pearson's (r), Spearman's ( $\rho$ ), and Kendall's ( $\tau$ ) correlations between varve thickness and annually resolved particle-size distribution indice for the raw (top half) and decorrelated (bottom half) time series, and their

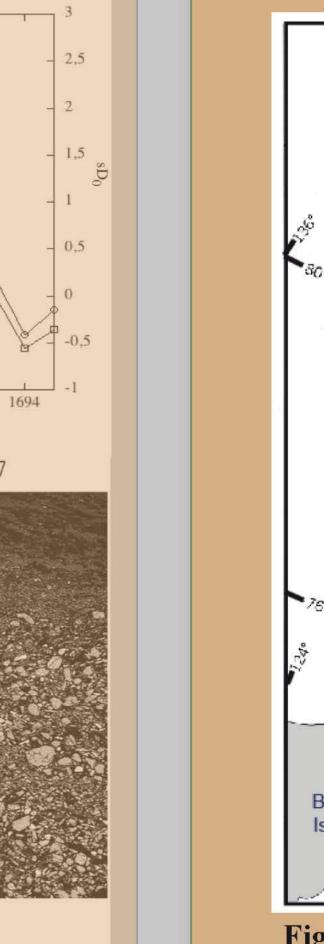
Wicthous	Original time series			
r	0.5343	0.2153	0.2999	0.0426
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(0.0753)
ρ	0.3228	0.1790	0.2618	0.1186
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(< 0.0001)
τ	0.2238	0.1217	0.1784	0.0792
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(< 0.0001)
Methods	Decorrelated time series			
r	0.5503	0.2092	0.3149	0.0166
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(0.4879)
ρ	0.3967	0.1765	0.2745	0.0724
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(0.0025)
τ	0.2745	0.1194	0.1868	0.0483
	(< 0.0001)	(< 0.0001)	(< 0.0001)	(0.0025)

**High-energy** facies can be found inside relatively thin laminae

The strenght of correlations between varve thickness and particle size distribution is weak







Nearby meteorological stations

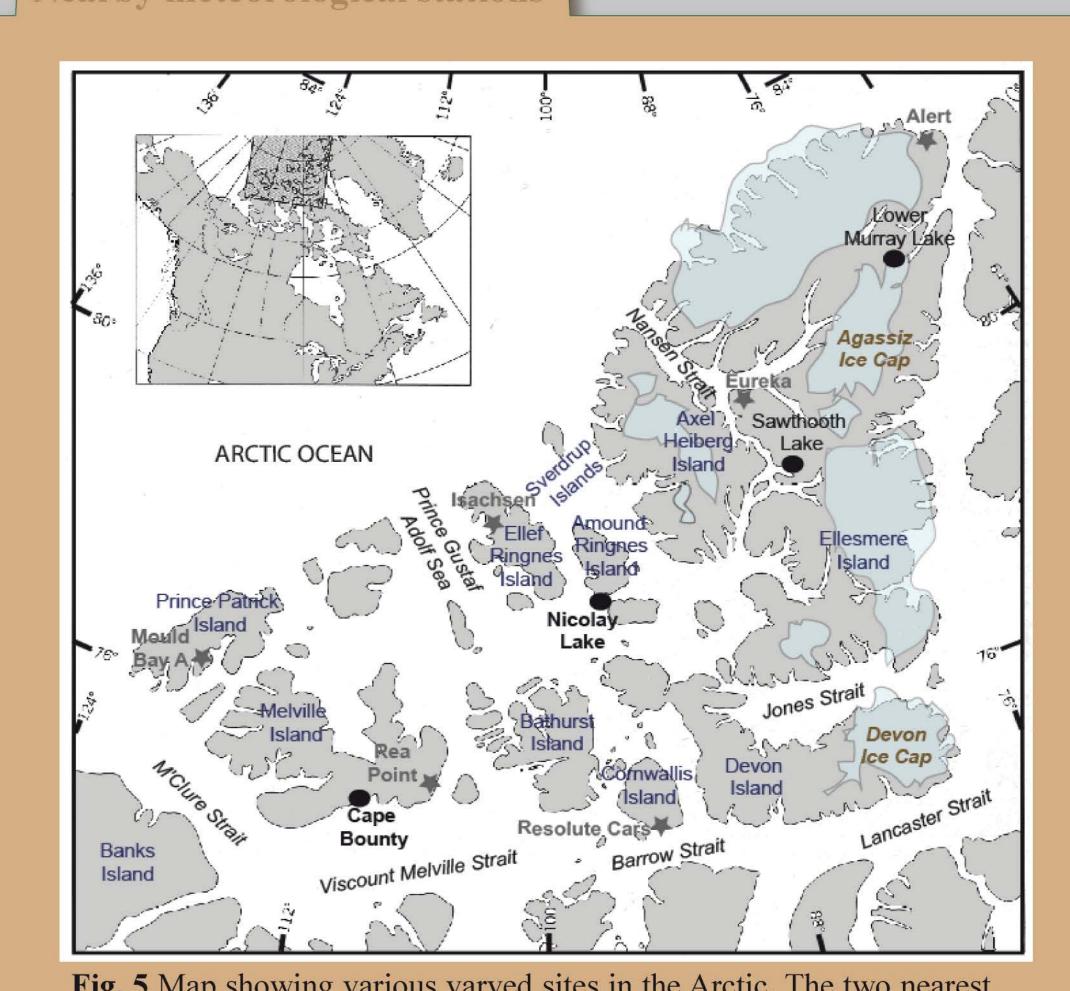


Fig. 5 Map showing various varved sites in the Arctic. The two nearest meteorological stations from Cape Bounty are Rea Point (100 km east) and Mould Bay (300 km north west).

### Calibration with meteorological

### able 2: Pearson's (P), Spearman's (S), and Kendall's (K) correlations between particle-size distribution index P98D, and instrumental

Rea Point (1985–1971) Mould Bay (1996–1971) Variables 0.3608 0.2000 0.5165 June temp. (°C) (0.2424)(0.3282)(0.0109)(0.0069)(0.0111)0.3679 0.2762 0.2459 0.2113 0.1373 July temp. (°C) (0.1779)(0.2260)(0.3420)(0.1686)(0.3000)0.2492 0.4536 0.3737 MDD (May-0.5893 0.4286 (0.0528)(0.0232)(0.0275)(0.0199)(0.0609)(0.0777)Largest rainfall 0.8479 0.7668 0.5742 0.1046 -0.0041 0.0123 (0.0009)(0.0035)(0.9473)(0.0001)(0.6110)(0.9841)0.0337 0.1246 0.0867 Annual snowfall

(0.5590)

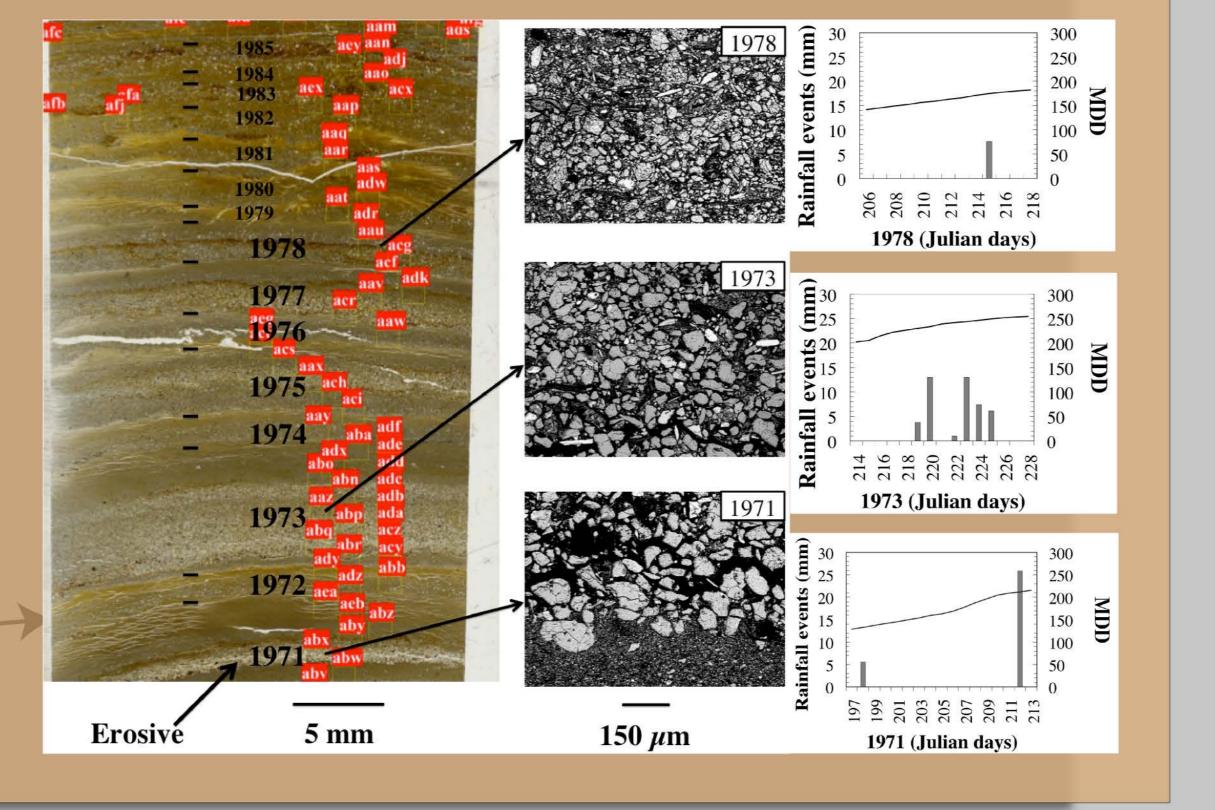
(0.8729)

(0.5513)

(0.5631)

Fig. 6 Thin section (right) with regions of interest and varve boundaries. The right panel shows backscatter electron (BSE) images of sediment deposited during years AD 1971, 1973, and 1978 (ROIs abx, aaz, and aau respectively) that contain coarse grain size. Beside each BSE image is the corresponding year with rainfall events and the melting degree days above 0°C (MDD). In those years, warm conditions prevail before the precipitation events, especially for AD 1971 and 1973, where coarsest particle are observed. Maximum temperatures reached uncommun values during the 4 days (> 13°C) before the major rainfall event that occurred on July 30 (212 Julian day), 1971, the year with the coarsest particle size during the calibration record.

(0.4266)



# Climatic

### reconstruction

 $\log (rainfall + 1) = -9.5337 + 3.2706 \log P98D$ 

Fig. 7 Linear regression between log-transformed (Box and Cox 1964) P98D and largest rainfall event/year at Rea Point using robust estimation (Maronna et al. 2006), and the 90% confidence interval (dotted lines).

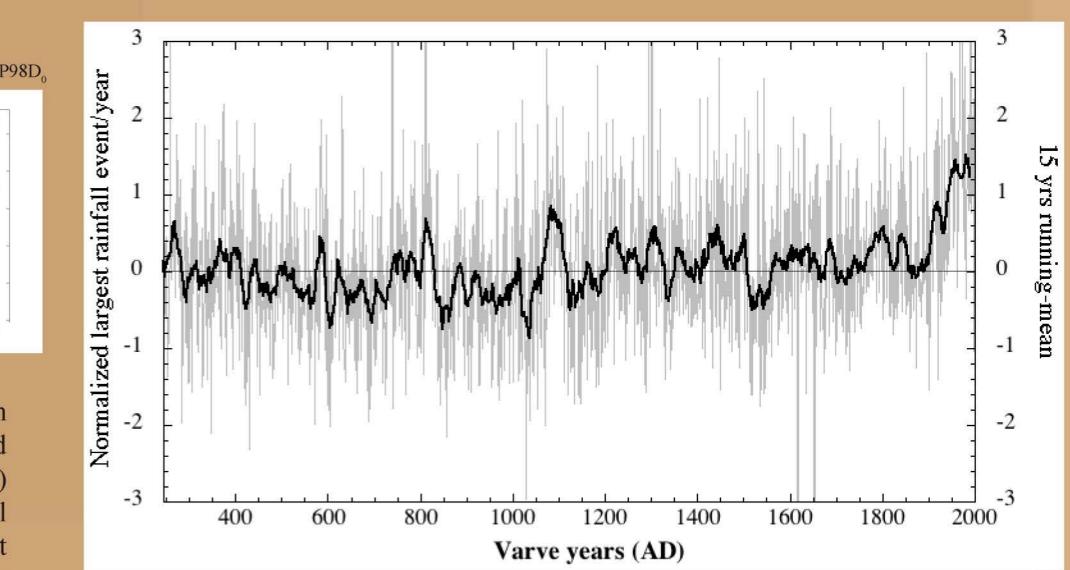


Fig. 8 Large rainfall events increase at ~AD 1920 with highest levels after ~ AD 1950. This increase is similar to the increase of temperatures documented by Kaufman et al. (2009). The peak at ~AD 1100 appears to be a strong anomaly similar to what is seen at the beginning of the 20th century. This could also be linked to warmer temperatures recorded in the Northern Hemisphere according to Moberg et al. (2005).

### Conclusion

Grain size signal is different compared to thickness signal.

Our sedimentary record is sensitive to large rainfall events, particulraly when warm conditions predate the rain events. This is supported by process work at Cape Bounty (Dugan et al. 2009; Lewis et al. 2011)

Rain events at East Lake increased to unprecedented levels in the 20<sup>th</sup> century compared to the last 1,750 years.

Lapointe F, Francus P, Lamoureux F S, Saïd M, Cuven S. (2012) 1,750 years of large rainfall events inferred from particle size at East Lake, Cape Bount, Melville Island, Canada. Journal of paleolimnology: 47 (4) doi: 10.1007/s10933-01209611-8