

## Research Article

# A Regionalization of Downscaled GCM Data Considering Geographical Features in a Mountainous Area

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This study establishes a methodology for the application of downscaled GCM data in a mountainous area having large spatial variations of rainfall and attempts to estimate the change of rainfall characteristics in the future under climate change. The Namhan river basin, which is in the mountainous area of the Korean peninsula, has been chosen as the study area. neural network-simple kriging with varying local means (ANN-SKlm) has been built by combining the artificial neural network, which is one of the general downscaling techniques, with the SKlm regionalization technique, which can reflect the geomorphologic characteristics. The ANN-SKlm technique was compared with the Thiessen technique and the ordinary kriging (OK) technique in the study area and the SKlm technique showed the best results. Future rainfall levels have been predicted by downscaling the data from CNRM-CM3 climate model, which was simulated under the A1B scenario. According to the results of future annual average rainfall by each regionalization technique, the Thiessen and OK techniques underestimated the future rainfall when compared to the ANN-SKlm technique. Therefore this methodology will be very useful for the prediction of future rainfall levels under climate change, most notably in a mountainous area.

## 1. Introduction

Climate change can affect the spatial and temporal distributions of water resources as well as the intensities and frequencies of extreme hydrological events [1]. Climate change causes temperature increases over a long period of time, change in rainfall amounts and patterns, dry spells, and a rise in sea levels. There are many ongoing studies on the evaluation of climate change impacts on water resources such as river flow quantity, water-quality, ecology, groundwater, agriculture, snowmelt, and hydroelectric power generation. The results of these studies are integrated and published by the IPCC in assessment reports [2–6].

One of the important issues for the analysis of climate change impact is the part related to the downscaling of GCM (general circulation model) data. Many researchers

are focusing their efforts on suggesting and improving various downscaling techniques. The studies on downscaling techniques had arisen from the issue on the resolution of GCM. Therefore, the studies on utilizing GCM data from the perspective of hydrology can be largely classified into dynamical and statistical downscaling techniques.

The dynamical downscaling technique is a numerical model based on a computer program that provides variables and fluxes on weather and climate by meteorological equations. Regional climate models (RCMs) and limited area models (LAMs) are the quintessential examples. It is a process of creating weather scenarios through a physical circulation process while having a GCM as the boundary condition [7, 8]. Since RCMs can simulate the study area in high resolution, many studies utilized RCMs [9–13]. Many projects that evaluated the impact of climate change and vulnerability

to climate change in limited regions, such as the UK Climate Impacts Programme [14], the Prudence [15–17], the North American Project NARCCAP, and CRES [18], used dynamical downscaling technique.

Statistical downscaling techniques use the statistical correlation between low resolution GCM and local observed data [19, 20]. Hanssen-Bauer and Førland [21], Hellström et al. [22], and Fowler et al. [10] applied a statistical downscaling technique that used the relationship between atmospheric hydrometeorological variables and rainfall/temperature of the ground surface. Cubasch et al. [23], Kidson and Thompson [24], Hanssen-Bauer et al. [25], and Chu et al. [26] also used statistical downscaling techniques by drawing the main components of pressure fields or geographical-potential heights. There are also studies that applied statistical downscaling techniques, such as the studies using artificial neural networks (ANN), which is mainly used in nonlinear forecasting [27–30], and the studies that applied canonical correlation analysis (CCA) [31–37]. Kyoung et al. [38] applied K-nearest neighbors (KNN) as a statistical downscaling technique.

Recently, there are also the comparison studies of dynamical downscaling techniques and statistical downscaling techniques [39–42]. It is possible to understand the strengths and weaknesses of each technique from these papers. Although both statistical and dynamical downscaling methods have their own merits, statistical downscaling is more widely adopted in hydrological studies because it is less computationally demanding [43]. There are some specific reasons why hydrologists have used statistical downscaling for GCM data with the transfer function method. One reason is that it is difficult to directly utilize RCM data on hydrologic analysis without calibrating RCM data. Another reason is that there is a limitation that all information required for future water resources impact evaluation cannot be easily obtained from RCM data. However, when GCM data is used, there is an issue caused by the resolution of GCM. A GCM for which one grid is 300–500 km represents the Korean peninsula by 2–4 grids. Most of the Korean peninsula (about 70%) is mountainous area, which has fairly large spatial deviations in weather phenomena and is strongly affected by geomorphologic characteristics; therefore, geomorphologic characteristics should be considered when applying downscaling and regionalizing techniques in mountainous areas. It can be said that general downscaling techniques that downscale GCM data to the locations of weather stations in a basin cannot properly reproduce the weather characteristics of the basin.

Therefore, this study intends to improve the spatial limitation of the statistically downscaled GCM data using a regionalization method, which should precede the climate change impact evaluation on water resources in a mountainous area. For this purpose, a methodology accommodating geomorphologic characteristics will be suggested for the downscaling and regionalization. It will then be analyzed and compared with existing methodologies. Lastly, the change in rainfall characteristics caused by climate change will be examined by forecasting future rainfall, which will be downscaled and regionalized from future weather data by the methodology suggested in this study.

## 2. Data Used and Methodology

### 2.1. Climatic Datasets

**2.1.1. Observed Climate Data.** The Namhan river basin in Korea has been chosen as the study area because it represents the meteorological and geomorphologic characteristics of the Korean peninsula fairly well and good quality data with a sufficiently long historical period was available. The total area of the Namhan river basin is 12,577 km<sup>2</sup> and the length of the river is 375 km. The upstream of the basin belongs to highland with altitudes in the 500 m to 1,700 m range, while the downstream belongs to a hilly area with altitudes ranging from 40 m to 500 m. The data from fifteen weather stations under the jurisdiction of the Korea Meteorological Administration (KMA) in the Namhan river basin will be used for this study (Figure 1). In addition, the information from 26 rainfall stations in the Namhan river basin under the jurisdiction of the Ministry of Land, Infrastructure, and Transport (MLIT) is collected to examine the applicability of downscaling techniques suggested by this study (Figure 1). In other words, we use rainfall data obtained from fifteen stations of KMA for the evaluation of the SKlm downscaling technique suggested in this study. For the comparison of the SKlm technique with Thiessen and OK techniques, the rainfall data from 26 rainfall stations of MLIT are used. Here the rainfall stations are more densely distributed than the other fifteen weather stations and so spatial accuracy of rainfall by the SKlm technique can be evaluated by the rainfall data from the 26 stations in the basin.

**2.1.2. NCEP-NACAR Reanalysis Data.** The National Oceanic and Atmospheric Administration (NOAA), Earth System Research Laboratory (ESRL), Physical Sciences Division (PSD) provides the observed data of 29 global units, including NCEP reanalysis data from their homepage. The NCEP reanalysis data, which is mainly used in studies related to climate change, includes temperature, relative humidity, specific humidity, sea level pressure, evaporation, U-wind, and V-wind observations from January 1, 1948, until now for the time scales of month, day, and every four hours. The monthly NCEP data from 1973 to 2008 for the Korean peninsula was downloaded and is utilized in this study.

**2.1.3. Climate Model and Scenario.** Kyoung [44] had chosen the BCM2, CNRM-CM3, FGOLS, and MIHR models, which simulate the Korean peninsula as land, not ocean, out of 24 GCM models given by the IPCC Data Distribution Centre (IPCC DDC) on higher priority and examined the applicability on the Korean peninsula. He finally suggested that the CNRM-CM3 model is the most appropriate model for the Korean peninsula. Therefore, this study adopted CNRM-CM3 model by referring to the results of Kyoung [44]. This study also selected the A1B scenario suggested by IPCC AR4 [5] because A1B is closest to reality, considering the whole world is trying to maximize energy resource efficiency and find alternative energy sources. There is a paper by Kwon et al. [45] in which the investigators performed their analysis using the A1B scenario.

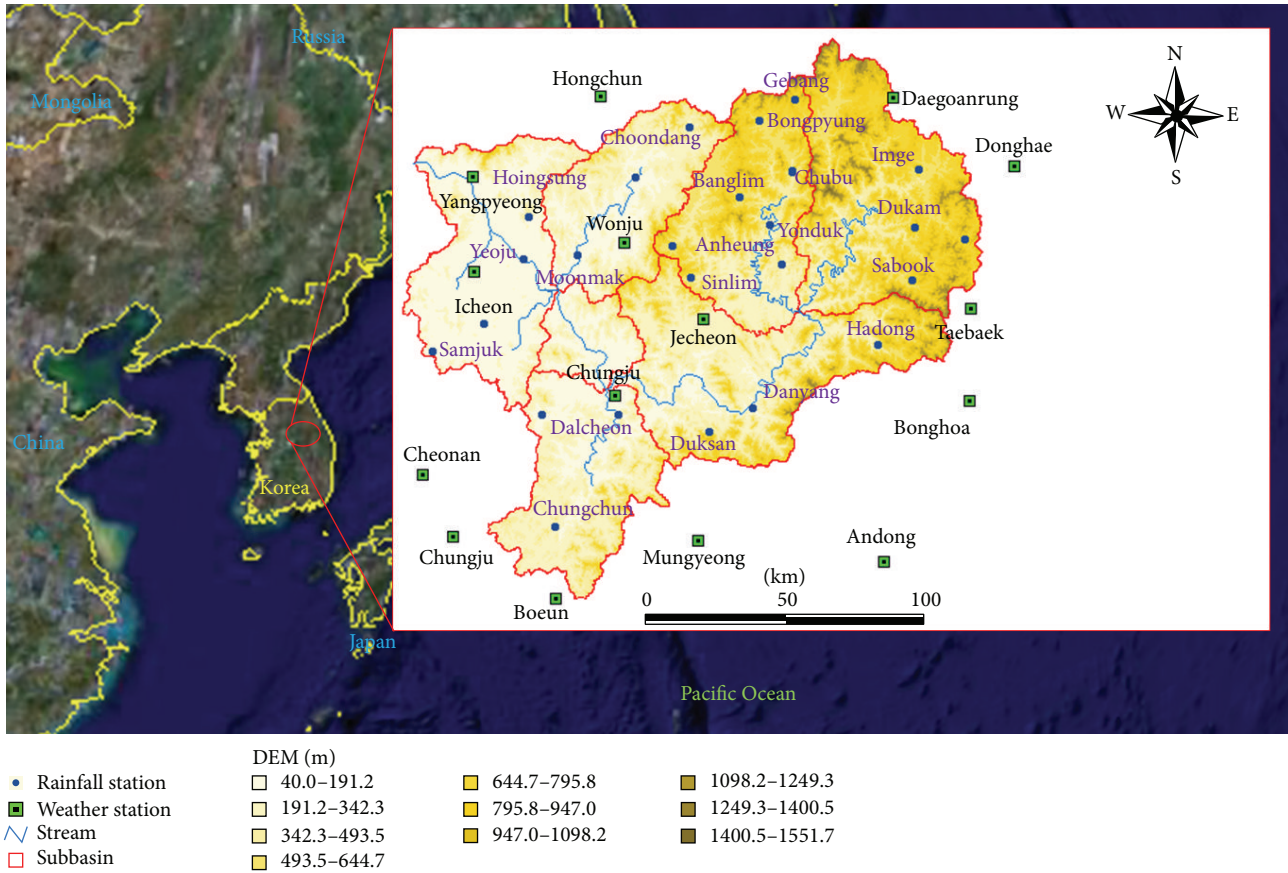


FIGURE 1: Study area and the location of weather and rainfall stations.

**2.2. Methods.** This study examines the artificial neural network (ANN) and simple kriging with varying local means (SKlm) technique as a spatial downscaling technique. Recently, many studies suggested the ANN technique, which is a nonlinear model of the data series, and ANN is better than other techniques by way of systematic evaluation of various techniques [27, 28, 46]. Therefore, this study also applies ANN, which is judged to have superior applicability in the simulation of nonlinear characteristics of weather data series. For the spatial distribution of monthly rainfall series in the application of downscaling technique, the kriging technique, which is applied in various areas for a spatial interpolation, is considered. The SKlm technique has been applied as a downscaling technique because it is an improved version of the simple kriging technique, which is capable of spatial interpolation using geomorphologic characteristics as secondary data.

**2.2.1. Spatial Downscaling Technique Using ANN.** ANN is a model of neurotransmission by a neuron, which is a nerve cell in the human brain. ANN is an empirical pattern search technique that enables the consideration of a nonlinear relationship between input variables and output variables. ANN is used in various areas because of its unique applicability [47, 48]. This includes the field of climate science, where its applicability is proven [49, 50].

In this study, an ANN model is suggested to apply the downscaling technique through GCM data. The spatial downscaling procedure using the ANN model is summarized as shown in Figure 2.

- (1) Select the common factors that can generate desired output values by collecting GCM data, NCEP data, and weather data at weather stations. Create input data for the ANN.
- (2) Create hidden layer considering both the input data (NCEP) and output data. Build a network by connected weights and build multilayer perceptron model.
- (3) Create objective function from output data and comparable observed values at weather stations.
- (4) Create output data from the connected weights of random seed. Compare it with observed values using the objective function.
- (5) Complete the learning by making most optimum connected weights by way of repetitive adjustment of connected weights using error backpropagation algorithm.
- (6) Simulate future weather phenomena series by using learned ANN and GCM data as input data.



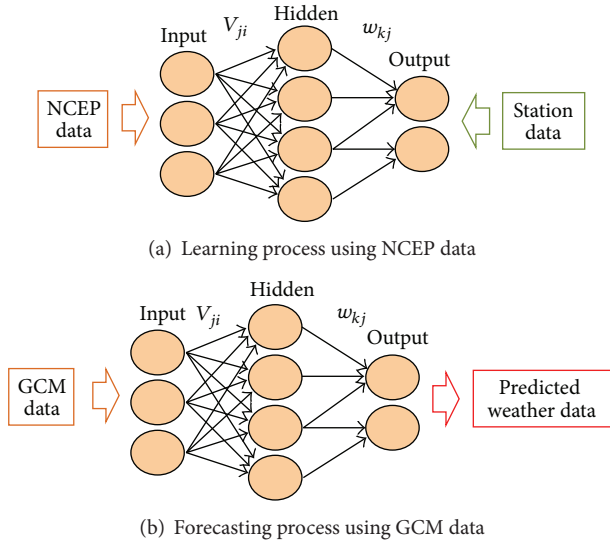


FIGURE 2: Downscaling technique using ANN.

**2.2.2. Considering the Influence of Geomorphologic Characteristics on Downscaled Rainfall.** Kriging is a generic name adopted by geostatisticians for a family of generalized least-squares regression algorithms. The simple kriging (SK) estimate of  $z(u)$  is

$$z_{sk}^*(u) - m = \sum_{\alpha=1}^{n(u)} \lambda_{\alpha}^{sk}(u) [z(u_{\alpha}) - m], \quad (1)$$

where  $m$  is the stationary mean of  $z(u)$  and  $\lambda_{\alpha}^{sk}$  is the weight assigned to datum  $z(u_{\alpha})$ .

Simple kriging with varying local means (SKlm) amounts to replacing the known stationary mean in the SK estimate (1) by known varying means  $m_{sk}^*(u)$  derived from the secondary information [51]. It may be written as

$$Z_{sk}^*(u) - m_{sk}^*(u) = \sum_{\alpha=1}^{n(u)} \lambda_{\alpha}^{sk}(u) [Z(u_{\alpha}) - m_{sk}^*(u)]. \quad (2)$$

The characteristic of SKlm is that the local mean ( $u$ ) of primary data is replaced by the  $m_{sk}^*(u)$  of secondary data.

This methodology can be applied through the following procedures.

- (1) Calculate the local means by a regression relationship between the primary data and the secondary data.
- (2) Calculate the residuals using the relationship between observed value and local mean.
- (3) Make a variogram on the residuals.
- (4) Estimate residuals for ungauged locations by performing kriging on the residuals.
- (5) Obtain the estimated values for ungauged locations by adding residuals to the local mean.

### 3. Model Construction and Estimation of Future Monthly Rainfall

This study suggests a methodology that can account for geomorphologic characteristics when applying downscaling and regionalization techniques as a way to estimate rainfall in a basin by using GCM data. To this end, this study adopts CNRM-CM3 GCM model and SRES A1B for 1973–2100. ANN is built for downscaling of the GCM data by applying learning and verification using monthly data, NCEP data of the NOAA, and weather data of the KMA (of a 36-year period). SKlm is applied for regionalization of the downscaled GCM using a Digital Elevation Model (DEM) and is compared with ordinary kriging (OK) and Thiessen methods in 26 rainfall stations of MLIT. The techniques will be used to predict and assess the monthly rainfall by future GCM data. The study procedure is shown in Figure 3.

**3.1. ANN Model.** In order to construct the ANN model by spatial downscaling of GCM weather data, this study set the average temperature, specific humidity, wind speed, and average atmospheric pressure at sea level as the input layer and the monthly rainfall as the output layer. As seen in Figure 4, a multilayered ANN model consisting of one input layer, two hidden layers, and one output layer has been built.

Monthly NCEP data from 1973 to 2008 (36 years) had been used for the learning period. For the forecast period, data from 2011 to 2100 (90 years) had been used based on CNRM-CM3 weather data (Table 1). Five years from the learning period (2004 to 2008) were set up as examination periods and the applicability of the constructed ANN model was reviewed by comparing it to observed data.

The output results of the 60-month (five-year) period for six weather stations were examined in the Namhan river basin. The result was that though they could not reproduce the extreme values of observed monthly rainfall, they simulated the observed values fairly well. The verification of the ANN model for rainfall series is shown in Figure 5.

In order to examine the applicability of the ANN model constructed in this study, a comparison with the results of the statistical downscaling model (SDSM) [52] was done. SDSM can be considered a general statistical downscaling technique based on multiple regression analysis. According to the model validation measures, such as the coefficient of correlation (CC), coefficient of determination ( $R^2$ ), model efficiency (ME), root mean squared error (RMSE), and prediction error variance (PEV), ANN was better overall. The better results are shown in bold in Table 2.

**3.2. SKlm.** SKlm is a technique that can utilize secondary data when performing spatial interpolation. This study used the Digital Elevation Model (DEM) as the secondary data for the spatial interpolation required for the downscaling technique.

The relationship between elevation and monthly rainfall was analyzed for each month using the monthly averages of fifteen weather stations in the Namhan river basin. Figure 6 depicts the regression analysis results for March, June, September, and December. It was discovered that the elevation and monthly rainfall have strong correlation with  $R^2$

TABLE 1: Input data of ANN.

| Data                          | Learning                     | Prediction                       | Time scale   |
|-------------------------------|------------------------------|----------------------------------|--------------|
| Data period                   | NCEP<br>1973~2008 (36 years) | CNRM-CM3<br>2011~2100 (90 years) | Monthly data |
| Learning (calibration) period | 1973~2003 (31 years)         | —                                | Monthly data |
| Verification period           | 2004~2008 (5 years)          | —                                | Monthly data |

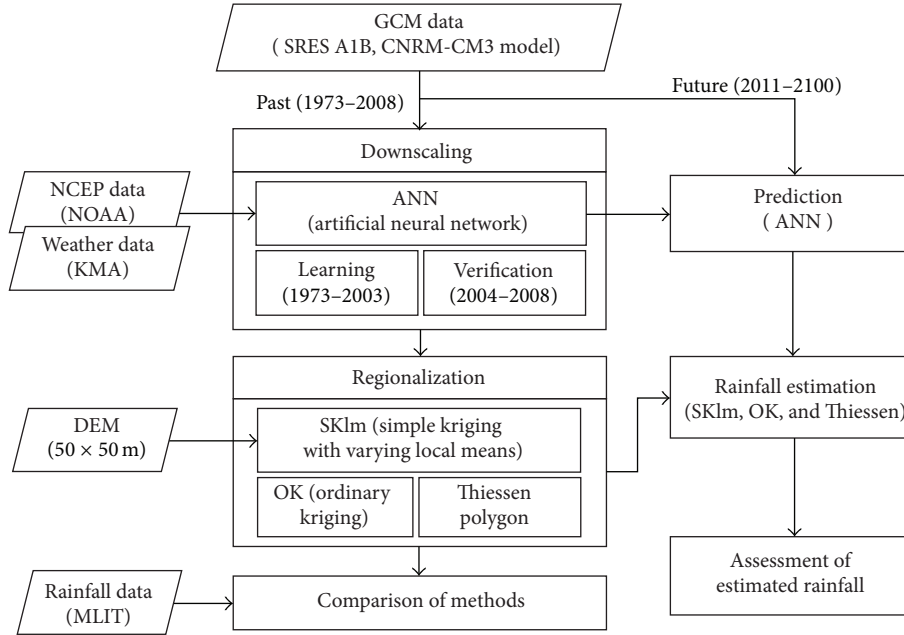


FIGURE 3: Flowchart of the study.

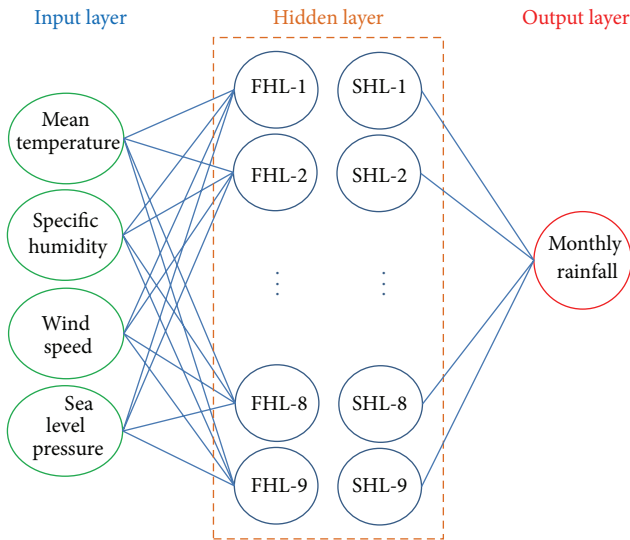


FIGURE 4: Construction of ANN model.

between 0.62 and 0.86. The residuals of rainfall in each weather station were calculated using this primary regression relationship, and the SKlm technique was applied.

In order to examine the applicability of the SKlm technique on 26 weather stations, this study applied and compared the Thiessen method and ordinary kriging (OK) method, which are general techniques in spatial interpolation.

The results of each method created by using fifteen weather stations are shown in Figure 7. Here the result of the Thiessen polygon method was expressed as fifteen values representing fifteen stations. The result of OK was expressed as spatial-interpolated values by the distances of fifteen stations. The result of SKlm was expressed as spatial-interpolated values by DEM and the distances of fifteen stations.

The spatial-interpolated data on two chosen weather stations with different elevations (Dukam (EL. 720 m) and Yeosu (EL. 45 m)) by applying Thiessen, OK, and SKlm techniques in sequence were compared with observed data and suggested in Figure 8. The rainfall of Dukam station, which is located at a high elevation, was estimated differently by each method. However the rainfall of Yeosu station, which is located at a low elevation, was estimated similarly by each method.

According to the results of spatial interpolation from Figure 9 and Table 3, it was found that the SKlm technique reflects the geomorphologic characteristics (elevation) fairly well. CC and  $R^2$  on the forecasted values and observed

TABLE 2: Comparison of CC,  $R^2$ , ME, RMSE, and PEV of ANN and SDSM.

|       |      | Daegoanrung  | Wonju        | Chungju       | Yangpyeong   | Icheon       | Jecheon      |
|-------|------|--------------|--------------|---------------|--------------|--------------|--------------|
| CC    | ANN  | <b>0.689</b> | <b>0.851</b> | <b>0.886</b>  | <b>0.863</b> | <b>0.854</b> | <b>0.834</b> |
|       | SDSM | 0.642        | 0.802        | 0.801         | 0.798        | 0.798        | 0.745        |
| $R^2$ | ANN  | <b>0.475</b> | <b>0.724</b> | <b>0.786</b>  | <b>0.745</b> | <b>0.729</b> | <b>0.696</b> |
|       | SDSM | 0.412        | 0.643        | 0.642         | 0.638        | 0.637        | 0.556        |
| ME    | ANN  | <b>0.963</b> | <b>0.958</b> | <b>0.77</b>   | 0.034        | 0.581        | <b>0.86</b>  |
|       | SDSM | 0.962        | 0.945        | 0.604         | <b>0.064</b> | <b>0.853</b> | 0.797        |
| RMSE  | ANN  | <b>2.208</b> | <b>1.769</b> | <b>13.395</b> | <b>8.503</b> | <b>0.552</b> | <b>1.747</b> |
|       | SDSM | 2.879        | 2.619        | 14.951        | 11.659       | 3.661        | 4.43         |
| PEV   | ANN  | 0.043        | 0.029        | <b>0.402</b>  | <b>0.468</b> | <b>0.582</b> | <b>0.511</b> |
|       | SDSM | <b>0.015</b> | <b>0.026</b> | 0.63          | 0.82         | 0.765        | 0.86         |

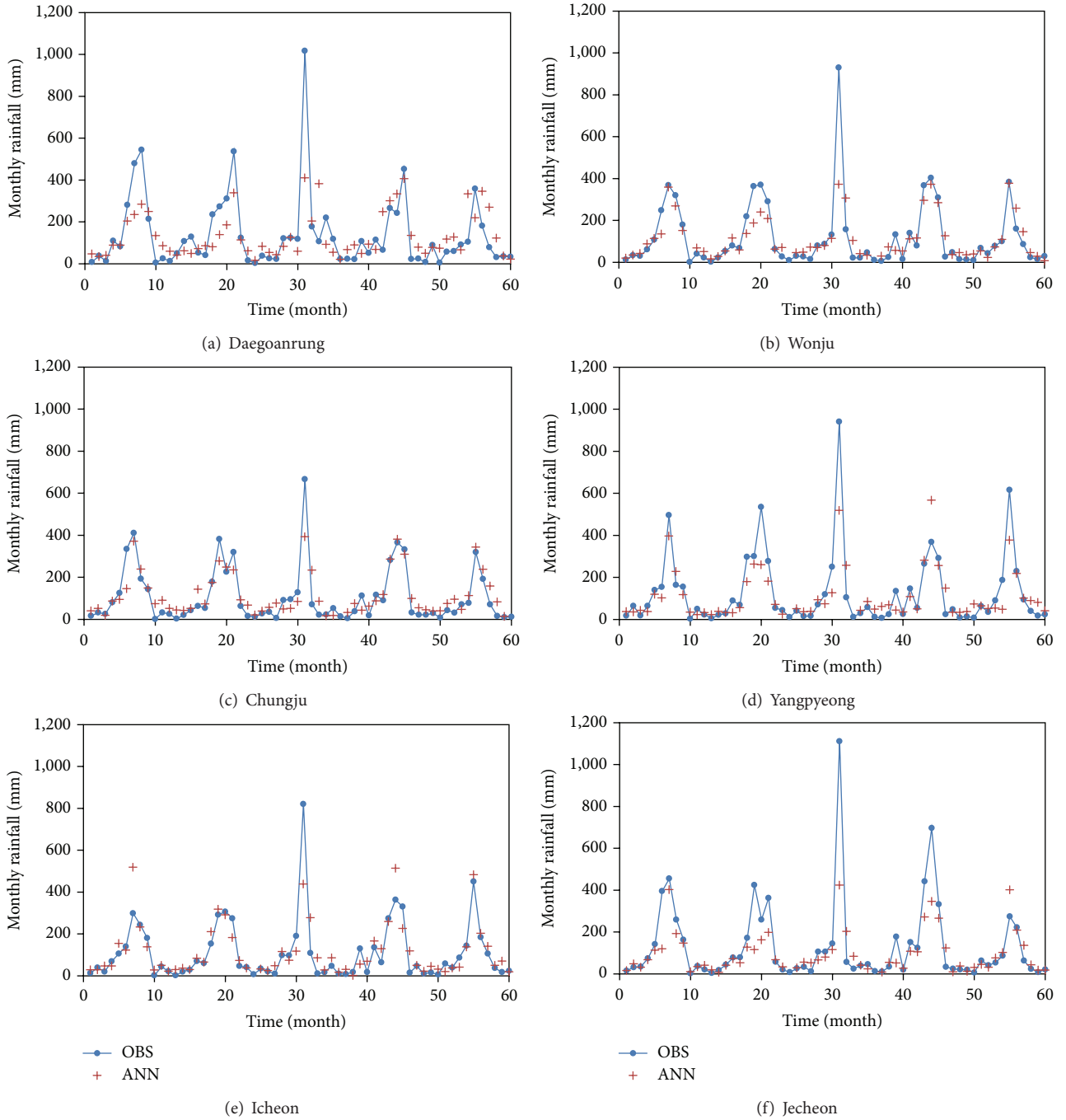


FIGURE 5: Comparison of rainfall series for the verification period in the Namhan river basin.

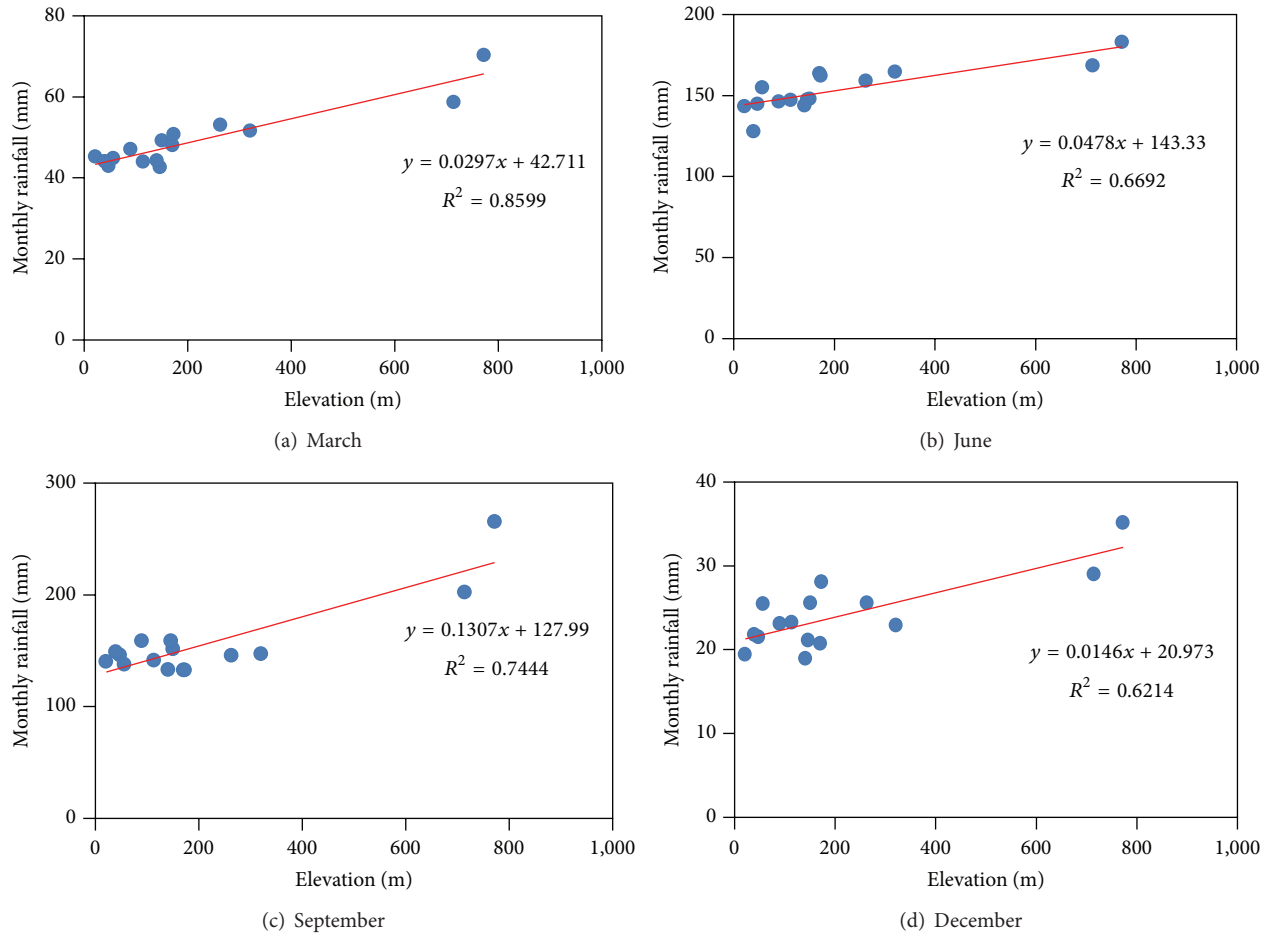


FIGURE 6: Relation between elevation and monthly rainfall.

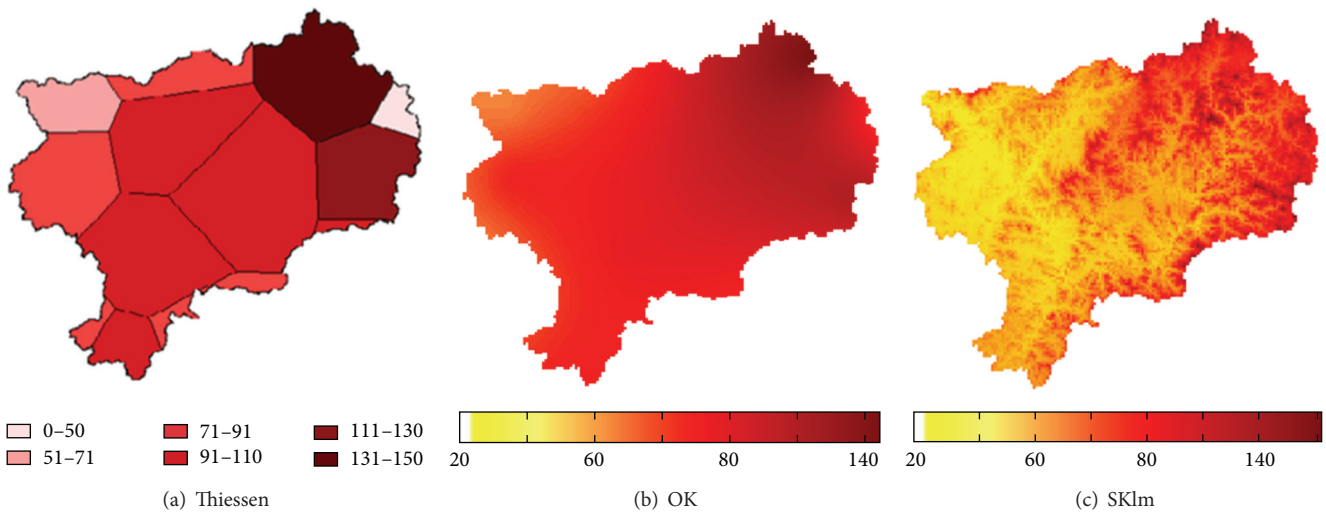


FIGURE 7: Spatial interpolation results by each technique (unit: mm).

values on each technique were calculated for quantitative analysis. Consequently, it was found that the SKlm technique has the best applicability, with  $CC = 0.788$  and  $R^2 = 0.622$ . Characteristics of techniques had been reviewed and it has been found that the SKlm technique had the best result, when

weather stations are located at a high elevation. However, when weather station elevation was low, there was no large deviation among weather stations. The Thiessen and OK techniques were judged as having poor applicability because they could not consider the elevation factor.

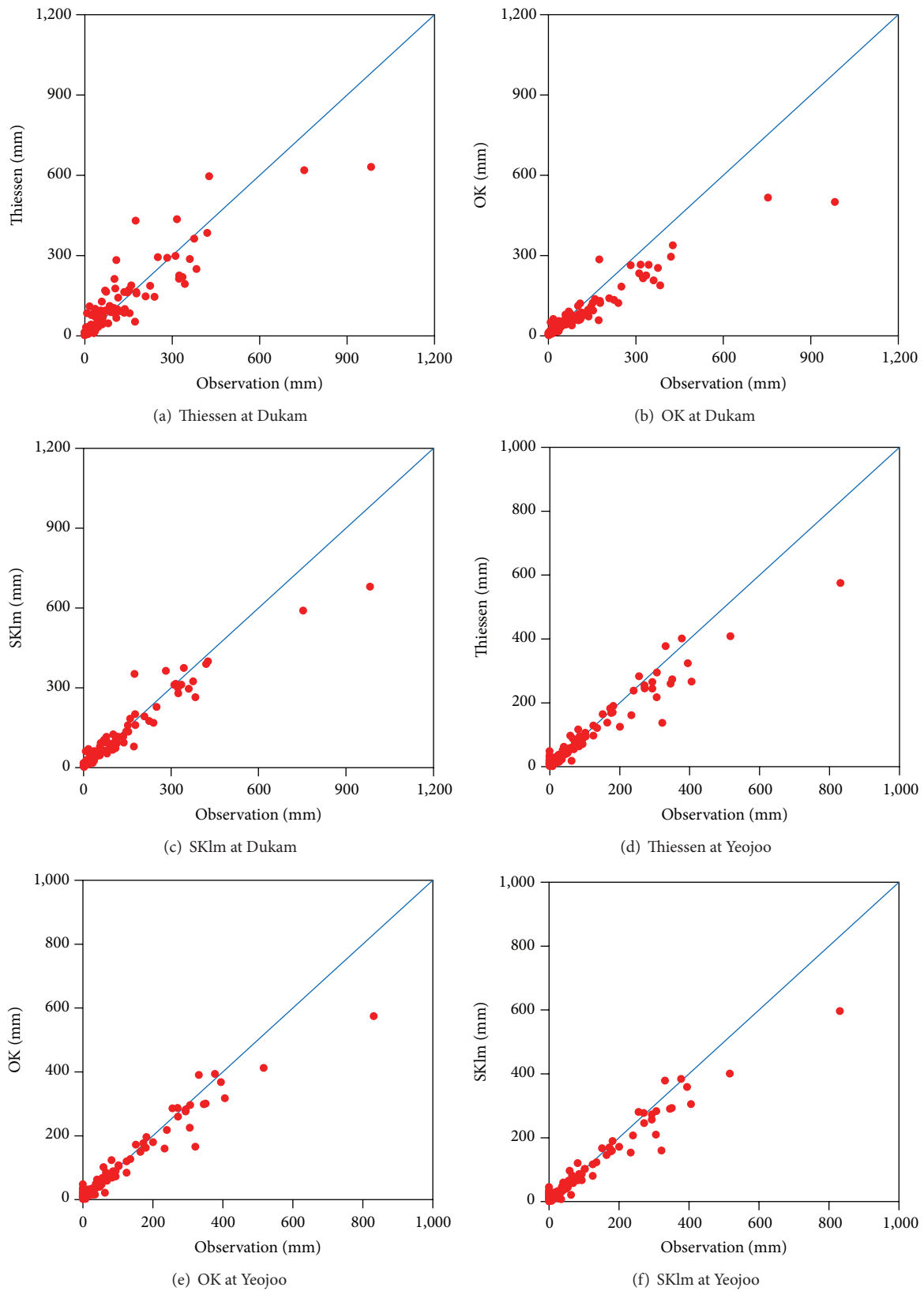
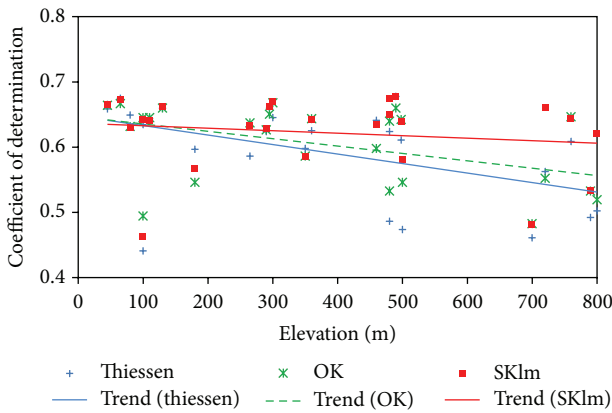


FIGURE 8: Examination results by technique dependent on elevation difference.



TABLE 3: Comparison of spatial interpolation techniques.

| Code     | Rainfall station | Elevation<br>(m) | CC       |       |       | $R^2$    |       |       |
|----------|------------------|------------------|----------|-------|-------|----------|-------|-------|
|          |                  |                  | Thiessen | OK    | SKlm  | Thiessen | OK    | SKlm  |
| 10024060 | Bangrim          | 480              | 0.697    | 0.730 | 0.806 | 0.486    | 0.533 | 0.649 |
| 10014140 | Imge             | 498              | 0.781    | 0.801 | 0.799 | 0.611    | 0.642 | 0.638 |
| 10024240 | Sinrim           | 460              | 0.801    | 0.773 | 0.796 | 0.641    | 0.598 | 0.634 |
| 10024230 | Anheunh          | 480              | 0.790    | 0.800 | 0.821 | 0.624    | 0.640 | 0.675 |
| 10024260 | Yeonduk          | 360              | 0.791    | 0.802 | 0.801 | 0.625    | 0.644 | 0.642 |
| 10024200 | Pyungchang       | 295              | 0.812    | 0.806 | 0.814 | 0.660    | 0.650 | 0.663 |
| 10014150 | Sabook           | 790              | 0.702    | 0.730 | 0.730 | 0.492    | 0.533 | 0.533 |
| 10014240 | Dukam            | 720              | 0.750    | 0.743 | 0.813 | 0.563    | 0.552 | 0.661 |
| 10014100 | Goangdong        | 760              | 0.780    | 0.804 | 0.802 | 0.608    | 0.646 | 0.644 |
| 10064100 | Chundang         | 290              | 0.790    | 0.791 | 0.792 | 0.624    | 0.626 | 0.628 |
| 10034100 | Danyang          | 265              | 0.766    | 0.798 | 0.796 | 0.586    | 0.637 | 0.633 |
| 10034180 | Hadong           | 800              | 0.709    | 0.721 | 0.787 | 0.502    | 0.519 | 0.620 |
| 10024160 | Gebang           | 700              | 0.679    | 0.695 | 0.694 | 0.461    | 0.483 | 0.481 |
| 10064020 | Hoingsung        | 130              | 0.812    | 0.812 | 0.814 | 0.660    | 0.660 | 0.663 |
| 10024210 | Daehoa           | 500              | 0.688    | 0.739 | 0.762 | 0.473    | 0.546 | 0.581 |
| 10074100 | Jooam            | 300              | 0.803    | 0.817 | 0.818 | 0.645    | 0.668 | 0.669 |
| 10074030 | Yeojoo           | 45               | 0.811    | 0.815 | 0.816 | 0.659    | 0.664 | 0.665 |
| 10064060 | Munmak           | 65               | 0.821    | 0.816 | 0.820 | 0.675    | 0.667 | 0.672 |
| 10074090 | Sulsung          | 100              | 0.796    | 0.803 | 0.801 | 0.634    | 0.645 | 0.642 |
| 10074080 | Samjuk           | 110              | 0.802    | 0.803 | 0.801 | 0.644    | 0.645 | 0.641 |
| 10054020 | Mokge            | 80               | 0.806    | 0.795 | 0.794 | 0.649    | 0.632 | 0.630 |
| 10044150 | Eumsung          | 490              | 0.823    | 0.812 | 0.823 | 0.678    | 0.660 | 0.678 |
| 10044130 | Dalcheon         | 100              | 0.664    | 0.703 | 0.681 | 0.441    | 0.494 | 0.463 |
| 10034160 | Duksan           | 350              | 0.773    | 0.766 | 0.765 | 0.598    | 0.586 | 0.585 |
| 10044010 | Chungchun        | 180              | 0.772    | 0.739 | 0.753 | 0.596    | 0.546 | 0.567 |
| Average  |                  |                  | 0.769    | 0.777 | 0.788 | 0.593    | 0.605 | 0.622 |

FIGURE 9: Trend of  $R^2$  dependent on elevation.

When these characteristics were considered, it was judged that the SKlm technique will have better applicability in the Korean peninsula, which has relatively smaller density in the number of weather stations in the area with high elevation. In addition, it is expected that the SKlm technique will be quite useful as a downscaling technique, as it considers the geomorphologic characteristics of the Korean peninsula.

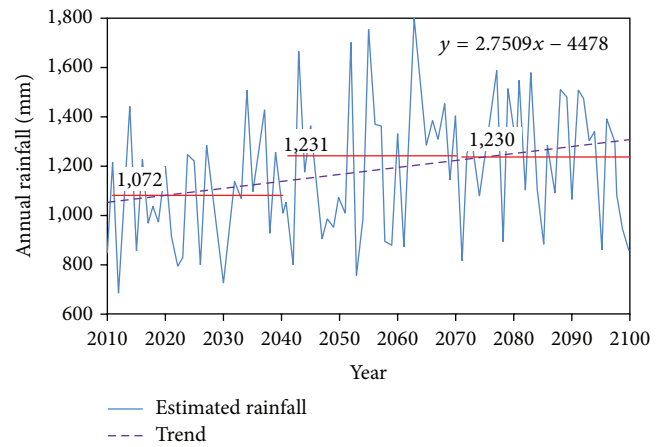


FIGURE 10: Forecast of future annual rainfall levels using climate change scenario data.

### 3.3. Forecast of Future Monthly Rainfall under Climate Change.

In order to use the climate change scenario data in the scale of a basin, this study applied spatial downscaling techniques combining the ANN and SKlm techniques, which are believed to have better applicability, to the Namhan river

TABLE 4: Change of monthly rainfall in each target period.

|           | Reference     | 2011–2040     |              | 2041–2070     |              | 2071–2100     |              |
|-----------|---------------|---------------|--------------|---------------|--------------|---------------|--------------|
|           | Rainfall (mm) | Rainfall (mm) | Increase (%) | Rainfall (mm) | Increase (%) | Rainfall (mm) | Increase (%) |
| January   | 26            | 28            | 8            | 25            | −2           | 21            | −16          |
| February  | 20            | 19            | −4           | 15            | −23          | 17            | −13          |
| March     | 9             | 12            | 29           | 12            | 34           | 14            | 54           |
| April     | 26            | 36            | 38           | 41            | 60           | 40            | 56           |
| May       | 69            | 87            | 26           | 109           | 57           | 128           | 85           |
| June      | 199           | 248           | 25           | 244           | 23           | 262           | 32           |
| July      | 302           | 290           | −4           | 378           | 25           | 379           | 25           |
| August    | 195           | 153           | −21          | 175           | −10          | 151           | −23          |
| September | 90            | 96            | 6            | 119           | 32           | 111           | 23           |
| October   | 52            | 61            | 17           | 64            | 24           | 61            | 17           |
| November  | 25            | 20            | −21          | 27            | 9            | 26            | 4            |
| December  | 24            | 22            | −9           | 21            | −16          | 20            | −18          |

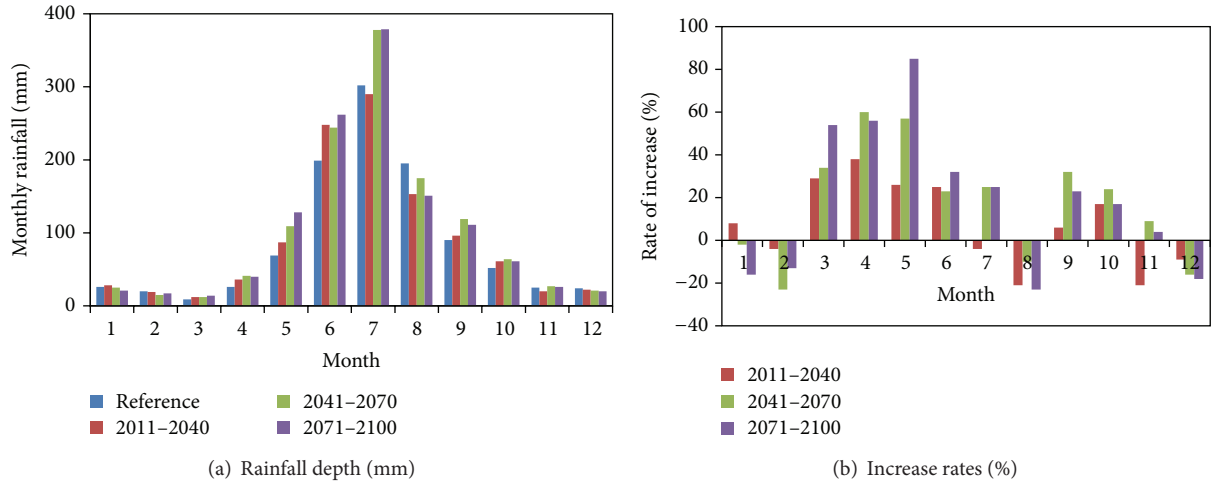


FIGURE 11: Change of monthly rainfall in each target period.

basin. The monthly average rainfall of the Namhan river basin until the year 2100 had been forecasted by getting the average of results downscaled to the spatial resolution of  $50 \times 50$  m. The change of annual average rainfall levels, which had been prepared based on the aforementioned results, is in Figure 10. According to the results of a primary regression, the annual rainfall is in an increasing trend. When the averages of forecasted rainfalls were calculated by segmenting those to 30-year target periods each, the result suggested that the rainfall after the later part of the 21st century will show a converging trend. It is believed that this result reflects the characteristics of the A1B climate change scenario, which reflects human effort to reduce carbon gas emission.

In order to examine the impact of climate change on the Namhan river basin in more detail, the changes in the characteristics of monthly runoffs from 2011 to 2100 were divided into three target periods (reference period, 2011–2040, 2041–2070, and 2071–2100). The analysis results forecasted that rainfall will increase during the rainy season (June and July) and decrease in the dry season (December, January, and February) in Figure 11 and Table 4. The results

also suggested that the big increase of rainfall during the rainy season will cause the future annual rainfall to increase (Figure 11(a)). According to the analysis results of monthly increasing ratios, the rainfall increasing ratios in spring (March, April, and May) will be bigger (Figure 11(b)). The reason for this is believed to be the earlier arrival of the rainy season relative to years past. In other words, the change in the temporal aspect will be an important factor along with the increase in the quantitative aspect.

#### 4. Discussions

When GCM data was used for the impact evaluation of climate change, the downscaling technique was typically applied. However, existing downscaling techniques have limitations in reproducing the rainfall characteristics that have large spatial variability in mountainous areas. As a methodology to address this issue, the ANN and SKlm techniques, both of which can use secondary data when performing spatial interpolation, were applied together and the results were compared to the Thiessen and ordinary kriging techniques,

TABLE 5: Geomorphologic characteristics of the Namhan river basin.

| Elevation (El. m) |       | Area (km <sup>2</sup> ) | Rate (%) | Remark  |
|-------------------|-------|-------------------------|----------|---|
| From              | To    |                         |          |   |
| 1                 | 200   | 3,136.4                 | 25.3     | Cheonan (21.3 m), Donghae (39.5 m), Yangpyeong (47.4 m), Cheongju (56.4), Icheon (90 m), Andong (140.7 m), Hongcheon (146.2), Moonkyung (170.8), Chungju (113.7 m), Wonju (150.7 m) |
| 201               | 400   | 3,462.8                 | 27.9     | Jecheon (263.1 m), Bonghoa (320.9 m)  |
| 401               | 600   | 2,236.6                 | 18.0     |   |
| 601               | 800   | 1,897.1                 | 15.3     | Taebaek (714.2), Daegoanrung (772.4 m)  |
| 801               | 1000  | 1,159.9                 | 9.4      |   |
| 1,001             | 1,600 | 506.8                   | 4.1      |   |
| $\Sigma$          |       | 12,399.6                | 100.0    | (i) Min: 40.0 m<br>(ii) Max: 1,551.7 m<br>(iii) Mean: 438.4 m<br>(iv) Std. dev.: 289.7 m  |

which are general downscaling techniques. According to the results (Figure 8 and Table 3), the SKlm technique using the elevation data (DEM data) in the basin performed better than existing techniques. The reason for this is that rainfall deviations at different elevations have a significant impact on the results, and the Namhan river basin is located in a mountainous area. The elevations of the Namhan river basin were at minimum elevation of 40 m and at maximum elevation of 1551.7 m, where the difference is more than 1,000 m (Table 5). The average elevation is 438.4 m and the standard deviation is 289.7 m, which means that the elevation deviations are significantly large. Therefore, the estimated values could have poor accuracy if existing planar spatial interpolation techniques are applied. In general, it is known that elevation and rainfall have a strong positive correlation. Most of the weather stations in the Namhan river basin are located in a low elevation area (less than elevation of 300 m), excluding the Daegwanryeong weather station, which is at an elevation of 772.4 m. The weather stations are located in such a low elevation area because the station management and maintenance convenience were put on a higher priority than obtaining accurate data. Since most of the Korean peninsula is mountainous, it was judged that this issue was causing an underestimation of rainfall levels.

Based on the above results, the Thiessen, OK, and SKlm techniques were combined in the forecasting and comparison of future aerial average rainfalls of the Namhan river basin (Table 6). Consequently, the SKlm technique produced the biggest average rainfall for the Namhan river basin and the Thiessen technique produced the smallest average rainfall. The difference between the SKlm technique and the Thiessen technique was 70 mm during the reference period and 91 mm for the period of 2071 to 2100. In other words, it was observed that the deviation gets bigger as the time frame progresses into the future.

Based on the results of Section 3.2, if the SKlm technique was the best technique in representing the rainfall spatial distribution characteristics of the Namhan river basin, it can

TABLE 6: Average rainfall from each spatial interpolation technique in the basin (mm).

|           | Thiessen<br>(A) | OK<br>(B) | SKlm<br>(C) | (C)–(A) | (C)–(B) |
|-----------|-----------------|-----------|-------------|---------|---------|
| Reference | 959             | 974       | 1,029       | 70      | 55      |
| 2011–2040 | 994             | 1,015     | 1,072       | 78      | 57      |
| 2041–2070 | 1,144           | 1,153     | 1,231       | 87      | 78      |
| 2071–2100 | 1,139           | 1,146     | 1,230       | 91      | 84      |

be said that the existing spatial interpolation techniques for climate change impact evaluation were seriously underforecasting rainfall. Given the rapid increase of rainfall in the future rainy season as seen in the monthly analysis results (Figure 10, Table 3), if the plans to respond and adapt to climate change are established on the basis of the results produced by existing techniques, there will likely be many issues during the flood seasons in the future.

It can be said that rainfall in a basin may be affected by various complex elements such as meteorological characteristics, geomorphologic characteristics, and land coverage characteristics. This is especially the case for the Korean peninsula, where rainfall events under the impact of the monsoon are usually dominant; however, extreme weather phenomena are heavily impacted by tropical cyclones, such as typhoons. It is also true that there are large deviations between regions due to the geomorphologic characteristics of a mountainous area. These issues are causing many difficulties in the accurate estimation of rainfall in basins when evaluating the impact of climate change.

This study focused on the influence of geomorphologic characteristics of a basin in order to produce more accurate rainfall data using future weather data from GCMs. Among these geomorphologic characteristics, the impact of elevation data was examined as a factor and was found to have the biggest impact on rainfall variability. However, it was believed

that examination on the slope and aspect of basin, which had been made by river flow, was also required for considering the characteristics of the Korean peninsula, which is affected by the monsoon and tropical cyclones. The accuracy of aerial rainfall estimation in the basin would be effectively enhanced if these impacts are comprehensively reflected.

## 5. Conclusions

By themselves, general downscaling techniques that downscale GCM data to a location in a basin cannot properly reproduce the weather characteristics of the region, especially in a mountainous area. In order to address this issue, this study suggested a downscaling and regionalization technique, which is needed for the climate change impact evaluation using GCM data in a basin or region, and examined the changes in the future rainfall characteristics. A methodology which applied the technique considering the basin geomorphologic characteristics and which combined the ANN and SKlm techniques was suggested. It was found that the SKlm technique had better applicability in mountainous areas like the Korean peninsula than the existing techniques such as the Thiessen polygon method or ordinary kriging.

The methodology suggested in this study was applied on the CNRM-CM3 model results, which was a simulation of the SRES A1B scenario. The future monthly average rainfall was forecasted for the Namhan river basin. According to the forecasted results of the monthly average rainfall in the basin, the total rainfall will continuously increase in the future. There were also notable trends, such as the increasing rainfall trend during the rainy season and the decreasing rainfall trend during the dry season. Furthermore, it was discovered that the arrival of the rainy season will shift to an earlier point in time.

By comparing the results generated by our novel approach against the future annual average rainfall forecasted by existing methodologies, we found that there were substantially large deviations that were dependent on the techniques used. It was discovered that the choice of methodology was a significant factor in the success or failure of future water resources planning. Therefore, the methodology suggested by this study, which was found to have good applicability to mountainous areas, could contribute to the establishment of successful water resources planning in the future. It may also provide some information that helps us overcome the limitations that statistical downscaling techniques suffer in low spatial resolutions when compared to their dynamical downscaling counterparts.

## Conflict of Interests

The authors declare that there is no conflict of interests regarding the publication of this paper.

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