

1 **Evaluating pixel-based versus object-based image analysis approaches for lithological**
2 **discrimination using WorldView-3 VNIR Imagery**

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8 **Abstract**

9 The object-based against pixel-based image analysis approaches were assessed for lithological
10 mapping in a geologically complex terrain using the VNIR bands of WorldView-3 (WV-3)
11 satellite imagery. The study area is Hormuz Island, southern Iran, a salt dome composed of
12 dominant sedimentary and igneous rocks. When performing the object-based image analysis
13 (OBIA) approach, the textural and spectral characteristics of the lithological features were
14 analyzed by the use of support vector machine (SVM) algorithm. However, in the pixel-based
15 image analysis (PBIA), the spectra of lithological end-members, extracted from imagery, were
16 used through the spectral angle mapper (SAM) method. Several test samples were used in a
17 confusion matrix to assess the accuracy of classification methods quantitatively. Results
18 showed that OBIA was capable of lithological mapping with an overall accuracy of 86.54%,
19 which was 19.33% greater than the accuracy of PBIA. OBIA also reduced the salt-and-pepper
20 artifact pixels and produced a more realistic map with sharper lithological borders. This
21 research showed limitations of the pixel-based method due to relying merely on the spectral
22 characteristics of rock types when applied to the high-spatial-resolution VNIR bands of
23 WorldView-3 imagery. It is concluded that the application of an object-based image analysis

24 approach obtains a more accurate lithological classification when compared to a pixel-based
25 image analysis algorithm.

26

27 **Keyword:** object-based image analysis, pixel-based image analysis, lithological mapping,
28 WorldView-3, Hormuz Island, spectral angle mapper, support vector machine

29 **1. Introduction**

30 Producing the lithological maps has undergone continuous evolution associated with
31 technological improvements in related fields. At the current time, advances in sensor
32 technology and developments in image processing approaches are the two main improvements
33 in collecting geological data and lithological mapping. Many researchers have recently used
34 multispectral data such as thematic mapper (TM), operational land imager (OLI), and advanced
35 spaceborne thermal emission and reflection radiometer (ASTER) to extract information about
36 rocks and alterations as well as their spatial distribution (e.g., Naghadehi et al., 2014; Ducart
37 and Silva, 2016; Ibrahim et al., 2018; Noori et al., 2019; Bolouki et al., 2020). Although the
38 pixel size of 30 m in Landsat and ASTER SWIR imagery is not appropriate for producing a
39 large scale and accurate geological map, they are beneficial for reconnaissance mapping to
40 guide geologists for more detailed field observations and mappings (Sun et al., 2017; Testa et
41 al., 2018; Bedini, 2019; Rajendran and Nasir, 2019). However, the Worldview-3 (WV-3)
42 satellite has recently provided alternative operational data that could efficiently be applied for
43 large-scale mapping of terrestrial features, including lithological units.

44 WV3 benefits from significant improvements such as high spatial resolution (1.24 m in VNIR
45 and 3.7m in SWIR bands), more spectral bands (16 multispectral bands), and high geometric
46 and radiometric accuracies associated with high radiometric resolution (11-bit in VNIR and

47 14-bit in SWIR bands) than the ASTER data. As a result, WV-3 data have been recently utilized
48 by remote sensing geologists in various disciplines.

49 Mars (2018) applied band ratios and Logical Operator Algorithms (LOAs) on data of WV-3 to
50 map goethite, calcite and dolomite, epidote-chlorite, and muscovite, using the absorption
51 features of Fe^{3+} , CO_3^{2-} , Fe- Mg-OH, and Al-OH, respectively, in Mountain Pass, California.

52 Ye et al. (2017) assessed the capabilities of WorldView-3 data compared to the ASTER and
53 OLI imagery for lithological mapping using a support vector machine (SVM) algorithm. They
54 estimated higher accuracies of 17% and 14% for WV-3 data outputs than, respectively, ASTER
55 and OLI data, and attributed it to the higher spatial resolution of WV-3 bands. Sun et al., (2017)
56 enhanced the alteration minerals in the Pobei area of Xinjiang Uygur Autonomous Region,
57 China, using short wave infrared data of WorldView-3. These authors proposed five principal
58 component analysis (PCA) models and ten mineral indices for enhancing the alteration
59 minerals. The WV-3 and ASTER TIR data were applied by Bedini (2019) for mineral mapping
60 in the Rodalquilar deposits, Spain. He expressed that the geographic dispersal of goethite was
61 successfully enhanced by combining all VNIR bands and band-1 of the SWIR region of WV-
62 3 and suggested that ASTER TIR data could map quartz-rich zones.

63 The primary remote sensing contexts such as training data and statistical assumptions are used
64 to classify images by running algorithms such as supervised vs. unsupervised, parametric vs.
65 non-parametric, per-pixel vs. sub-pixel, and pixel-based image analysis (PBIA) vs. object-
66 based image analysis (OBIA) (Thapa and Murayama, 2009). To date, most geologists have
67 used pixel-based methods to map rock units, in which, they classified lithology based on per-
68 pixel or sub-pixel formats without considering the contextual data for neighboring pixels (e.g.,
69 Hewson et al., 2017; Ayoobi and Tangestani., 2018; Liu et al., 2018). In per-pixel classification
70 algorithms, each image pixel is independently assigned to a unique lithology if the spectra of
71 pixel and the lithological end-member are highly suited. (Elnagheeb and Bromley, 1994). Two

72 well-known algorithms being used for per-pixel mapping of geological targets are spectral
73 angle mapper (SAM) (Kruse et al., 1993) and spectral feature fitting (SFF) (Clark and Roush,
74 1984). However, they lead to ignoring the spatial correlations between pixels of the imagery.
75 Moreover, spatial information can supply extra information related to the shape and size of
76 different structures, which could help identify and classify surface features with high accuracy.
77 Blaschke (2010) has concluded that the object-based image analysis (OBIA) approach
78 delineates a remarkable classification method for remote sensing objectives. In OBIA, several
79 attributes or features are associated with each of the image objects, and these attribute values
80 can be derived from the imagery. The selection of an optimal set of features for the
81 classification of unknown image objects is a crucial step and is very important for designing a
82 useful classification system (Cai et al., 2018).

83 Recently, the OBIA approach has extensively been applied to enhance and map the Earth's
84 surface features. For instance, Petropoulos et al. (2012) investigated OBIA and SAM methods
85 for land use/land cover mapping in a heterogeneous Mediterranean land using Hyperion
86 imagery. They estimated a higher overall accuracy and Kappa coefficient for OBIA results.
87 Additionally, the forest waste due to the gold excavation in Guyana was evaluated by
88 Mengisteab et al. (2014) using OBIA on the Landsat data, during which, they effectively
89 enhanced and specified the minor mining activities at the area.

90 Moreover, few articles have already been published on the geological utilizations of OBIA.
91 Van der Werff et al. (2007) applied Observatoire pour la Mineralogie, l'Eau, la Glace et
92 l'Activite (OMEGA) data for geological mapping on Mars using an object-based processing
93 method. Grebby et al. (2016) illustrated that the object-based image analysis method could
94 successfully map the rock types in an area covered by vegetation. They applied the Airborne
95 LiDAR (Li) and Airborne Thematic Mapper 9 (ATM9) data and discriminated chalky marl,
96 pillow lava, dyke, and alluvium-colluvium deposits. Aufferman et al. (2017) mapped the

97 Krafla volcanic rocks of the Icelandic volcanic zone by the use of OBIA and spectral angle
98 mapper (SAM) methods on Landsat 8 and SPOT-5 images. They revealed that SAM was
99 successful in producing detailed lava surface morphology maps; however, it partly led to a salt-
100 and-pepper effect. They concluded that despite the more efficient results of the OBIA approach,
101 it is sensitive to the objects derived from image segmentation. The mapping of geological
102 structures such as lineaments and faults was analyzed by the OBIA method in southwest
103 England (Yeomans et al., 2019) using the high-resolution airborne geophysics and LiDAR data.
104 They suggested that the OBIA method is highly effective for lineament detection.

105 An overview of the published articles indicated that geologists have conventionally used PPIA
106 methods for enhancement and identification of rock types, methods that are generally
107 performed based on the spectral characteristics of desired features. Unlikely, in the object-
108 based image analysis approach, the segmentation of image data into homogeneous and
109 consistent segments is a prerequisite for classification (Hay and Castilla 2008; Lang et al.,
110 2008; Blaschke 2010). The spatial dimensions, including parameters such as distances,
111 neighborhoods, and topologies, are essential in the OBIA approach, which is a primary reason
112 for an increase in its usage in recent years (Benz et al. 2004; Blaschke et al. 2004).

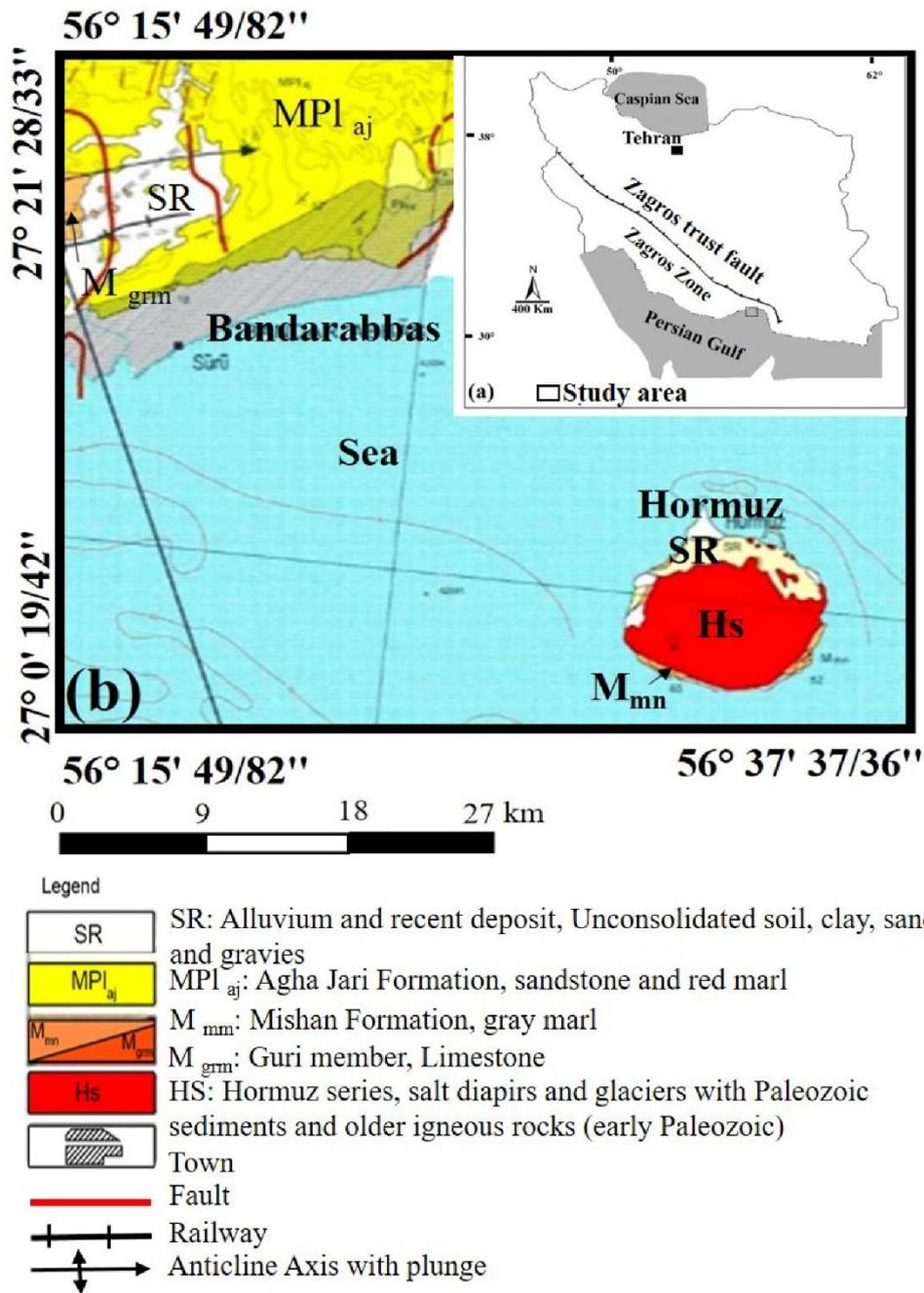
113 Despite the advantages reported for the OBIA approach (Castillejo-González et al., 2009;
114 Petitjean et al., 2012; Matton et al., 2015), rare publications are available on its performance
115 on the WV-3 data for discriminating lithological feature. This paper investigated the potential
116 of an object-based approach (support vector machine) and compared it to a pixel-based
117 approach (spectral angle mapper) for classification and information extraction of lithological
118 units in Hormuz Island, southern Iran. This island is a geologically salt dome, well-known for
119 of its particular setting and varying types of exposed rocks and minerals. Considering that the
120 VNIR bands of WorldView-3 can detect the dominant spectral features of rock outcrops of the
121 study area, this data set was applied for this research. The classification accuracies were

122 subsequently analyzed and compared using the parameters of confusion matrices and the
123 Kappa coefficients.

124 **2. Geological Setting**

125 The study area, Hormuz Island, is an Iranian island in the Persian Gulf with an oval shape and
126 a total area of about 45 km² (Fig. 1). A concentric structure shown at the central part of the
127 island contains salt, gypsum, and anhydrite (Elyasi et al., 1975), surrounded by salt rocks. The
128 salt rocks contain abundant fragments of black shale, black and white dolomite, limestone to
129 sandy limestone, iron oxide-rich strata, as well as outcrops of igneous rocks dominantly
130 consisting of tuff, rhyolite, and trachyte (Sadat Faramarzi et al., 2015). Stocklin (1972 and
131 1974) suggested that the Hormuz salt plug's diapirism has moved the vast enclaves of igneous
132 rocks to the surface, now occurring as isolated outcrops.

133 An iron oxide-rich band surrounded by young sediments wraps around the island. Alluvial
134 deposits that have been demolished from upstream formations are dominantly outspread in the
135 northern half of the Island (Fig. 1), and expand as small patches in other parts. The Hormuz
136 ochre is the most significant mine on the island, with a reservoir of about 390,000 tons (Yazdi
137 et al., 2014). In terms of quality and applications in industry, this red-colored earth pigment is
138 considered a unique raw material (Aqanabati, 2006).



139 Figure 1. The study area in Iran (a), and in 1:250,000 geological map (Fakhari (1988) (b).

140 **3. Materials and Methods**

141 **3.1. Overview**

142 The recently launched WV-3 is a high spatial and spectral resolution satellite that operates at a
 143 height of near 617 km. This satellite provides one panchromatic and eight multispectral bands

144 in the VNIR region, eight bands in SWIR region, and 12 CAVIS (Clouds, Aerosols, Vapors,
145 Ice, and Snow) bands with pixel sizes of, respectively, 0.31 m, 1.24 m, 3.7 m, and 30 m.
146 The VNIR data of WorldView-3, utilized in this study, was acquired on June 16, 2016
147 (www.worldview3.digitalglobe.com). These data were firstly corrected for likely geometric
148 and atmospheric errors, and subsequently, were applied in PBIAs and OBIA approaches by the
149 use of SAM and SVM algorithms for classifying the lithology of Hormuz Island.
150 The WV-3 level 2-A data have already been calibrated and corrected for radiometric and
151 geometric inaccuracies. The datum WGS-84 was used to geo-referencing applied data to UTM
152 zone 40-north projection. The data were also atmospherically corrected using the FLAASH
153 model, available in ENVI software, version 5.3. The effects of seawater and tidal zone on the
154 images were eliminated by applying a masking method. Data processing was supported by
155 extensive field sampling combined with petrographic and spectroscopic studies to identify
156 mineralogy and lithology of rock types. Finally, the accuracy of results was assessed by the use
157 of field criteria and confusion matrices.

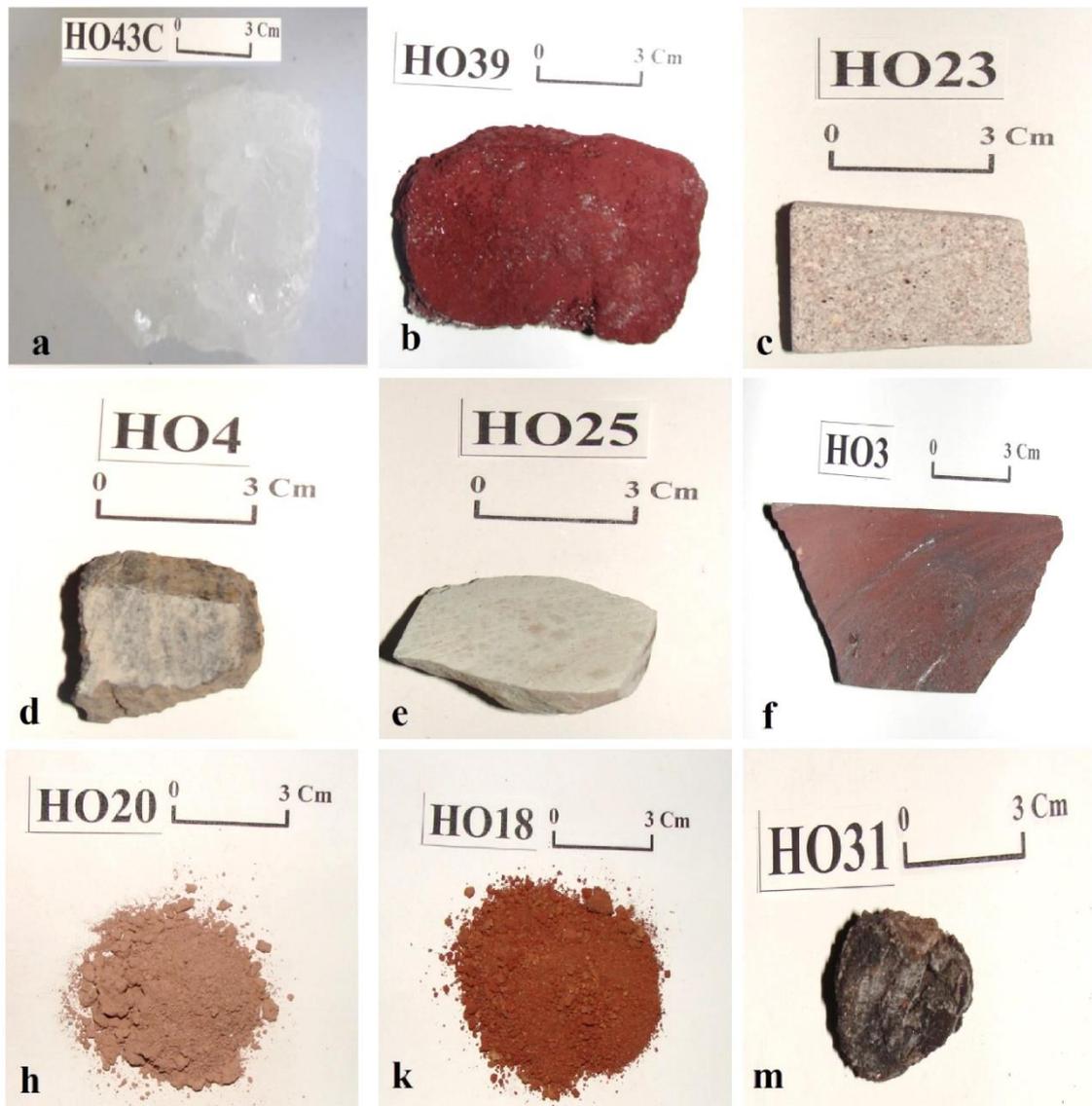
158 **3.2. Field sampling and laboratory studies**

159 According to the field observations, spectroscopy, petrography, and X-ray Diffraction (XRD)
160 studies, the rock units were classified into five groups, including 1) mixture of red soil, gypsum,
161 and anhydrite, 2) mixture of red soil, tuff and anhydrite, 3) white rhyolite tuff, 4) diabase and
162 volcanic tuff, and 5) marl. The validation sites of these lithological features were identified,
163 and 5-10 spectra in the range of 400 nm to 2500 nm were measured for each collected sample
164 using an ASD FieldSpec spectrometer, in the Department of Geography, Bowling Green State
165 University, the USA, which were subsequently averaged for each lithology. The pictures of
166 hand samples and their averaged spectra, resampled to the VNIR bands of WV-3, are shown in
167 Figures 2 and 3. Hormuz Island is dominantly formed of red soil and salt rock (Figs. 2 (a-b-
168 k)). The red color of soil is due to the extensive occurrence of hematite, which reduces the

169 center's amount and the extent to the margins of the island. The major absorption features of
170 red soil and gypsum in their high-resolution spectra are in 1900 nm attributed to the H₂O
171 vibration in anhydrite and gypsum, and 800 nm, because of the charge-transfer effect of ferric
172 iron (Hunt, 1980) (Figs. 3 (a-b)). The second most dominant rock unit is red soil with large
173 amounts of tuff and less anhydrite (Figs. 2 (b-e-k)). The tuffaceous rocks include rhyolite tuff,
174 alkaline rhyolite tuff, and dacite tuff (Figs. 3 (a-b)). Microscopic studies showed that tuffaceous
175 rocks consist mainly of quartz, alkaline feldspar, muscovite, chlorite, and rare epidote and
176 goethite, which are the results of degradation of ferromagnesian minerals (Mahyari, 2016).

177 The measured spectra of white rhyolite tuff displayed an absorption in 800 nm for charge
178 transfer effect of Fe³⁺ (Hunt, 1980) and additional features in 2160 nm and 2330 nm attributed
179 to vibrational modes of Al–OH and Mg–OH (Salisbury and Hunt, 1974) (Figs. 3 (a-b)).
180 Moreover, the diagnostic absorption features of diabase in 400-500 nm and 650-800 nm could
181 be attributed to the charge transfer effect of Fe–O (Hunt, 1980). Similar spectral properties of
182 this rock type in 2200 nm and 2210 nm are due to Mg–OH vibrational processes (Segal, 1983)
183 (Figs. 3 (a-b)). The carbonate interlayers are observed in marl outcrops of the Mishan
184 Formation (Fig. 2 (h)) and also are dispersed western and southwestern the island within a
185 sequence of salt and gypsum. The high-resolution spectra of marl showed significant
186 absorptions in 2000 nm and 2130 nm (Figs. 3 (a-b)) due to Al–OH (Huang and Kerr, 1960) and
187 an insignificant absorption feature in 1900 nm, for H₂O (Hunt, 1980) (Figs. 3 (a-b)).

188



189 Figure 2. Hand samples of; a) salt rock, b) iron oxide, c) rhyolite, d) diabase, e) green tuff, f)
 190 basalt, h) marl, k) iron soil, and m) volcanic tuff

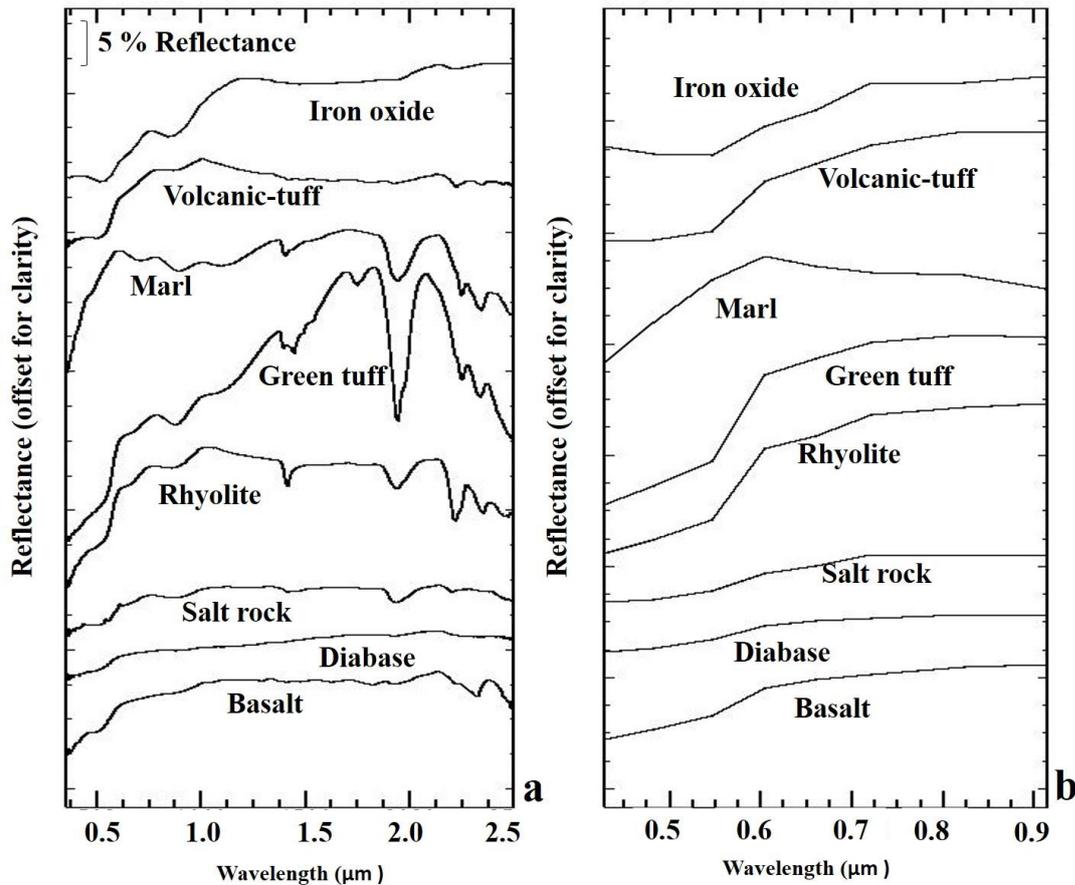
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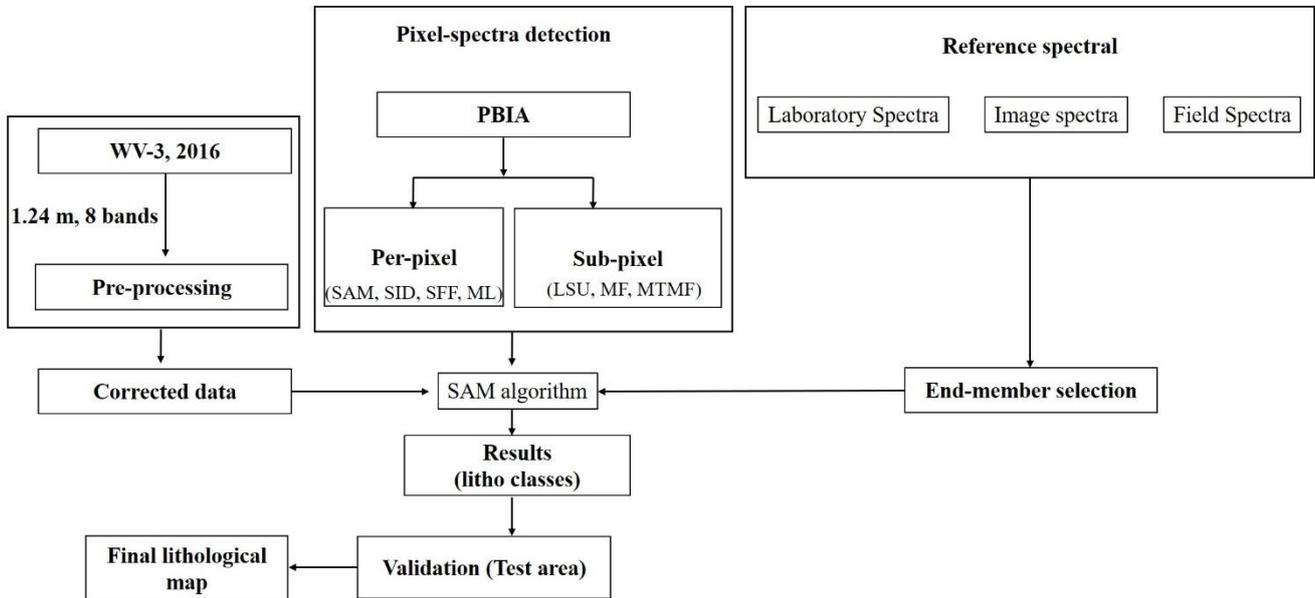
196 Figure 3. a) High-resolution spectra of rock samples, b) spectra of rocks resampled to the
 197 VNIR bands of WV-3.

198 3.3. Pixel-Based Image Analysis (PBIA)

199 PBIA is a spectrum space method that classifies the imagery by finding the analogy of a
 200 reference spectrum to that of a target (Richards, 1993). Spectral characteristics of desired
 201 materials play an essential role in their detection, identification, and classification. The
 202 appropriate spectra are typically selected from spectral libraries or field samples and are
 203 imported to an algorithm. Figure 4 shows a general workflow of the PBIA approach; its
 204 practical procedure is described in subsections “end-member selection” and “classification.” In
 205 cases where no information is available for a class, the spectral measures could be examined
 206 on a single signature vector basis to determine the spectral similarity between the target and
 207 the reference. This commonly is applied for discrimination and identification of specific

208 features, but not for classifying an imagery (Kruse et al., 1993). Moreover, these references are
 209 efficient only if compared with the spectral features are true characteristics of desired materials.

210



211 Figure 4. Flowchart of the PBI approach (PBI A = pixel-based image analysis; WV-3 =
 212 worldview-3; SAM = spectral angle mapper; SID = spectral information divergence; MF =
 213 matched filtering; MTMF = mixture tuned matched filtering, ML = maximum likelihood, SFF
 214 = spectral feature fitting, LSU = linear spectral un-mixing)

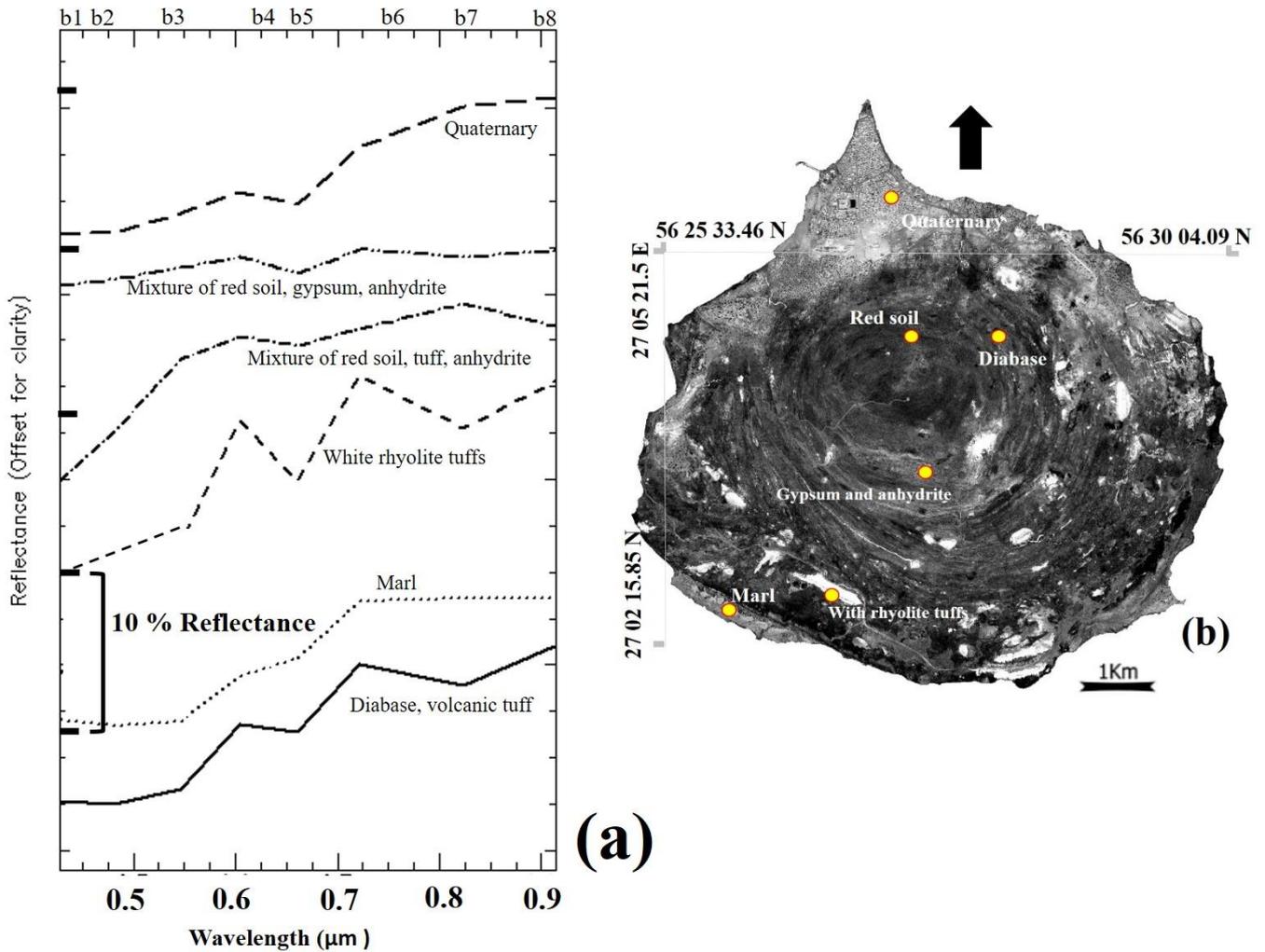
215 3.3.1. End-member selection

216 A reference spectrum or end-member, which represents the known spectral class, should
 217 typically be selected and put into the SAM algorithm when analyzing the desired satellite data.

218 The end-members are generally selected from spectral libraries or are extracted from applied
 219 imagery. Since the image spectra involve the atmospheric conditions of the applied data set, it
 220 is suggested that reference spectra from imagery are usually more valid for detecting the targets
 221 than those selected from libraries (Wang et al., 2004). On the other hand, the spectra extracted
 222 from imagery do not show subtle spectral features, as is evident in reference to spectra of
 223 spectral libraries (Wang et al., 2004). In this study, six reference spectra of lithological features

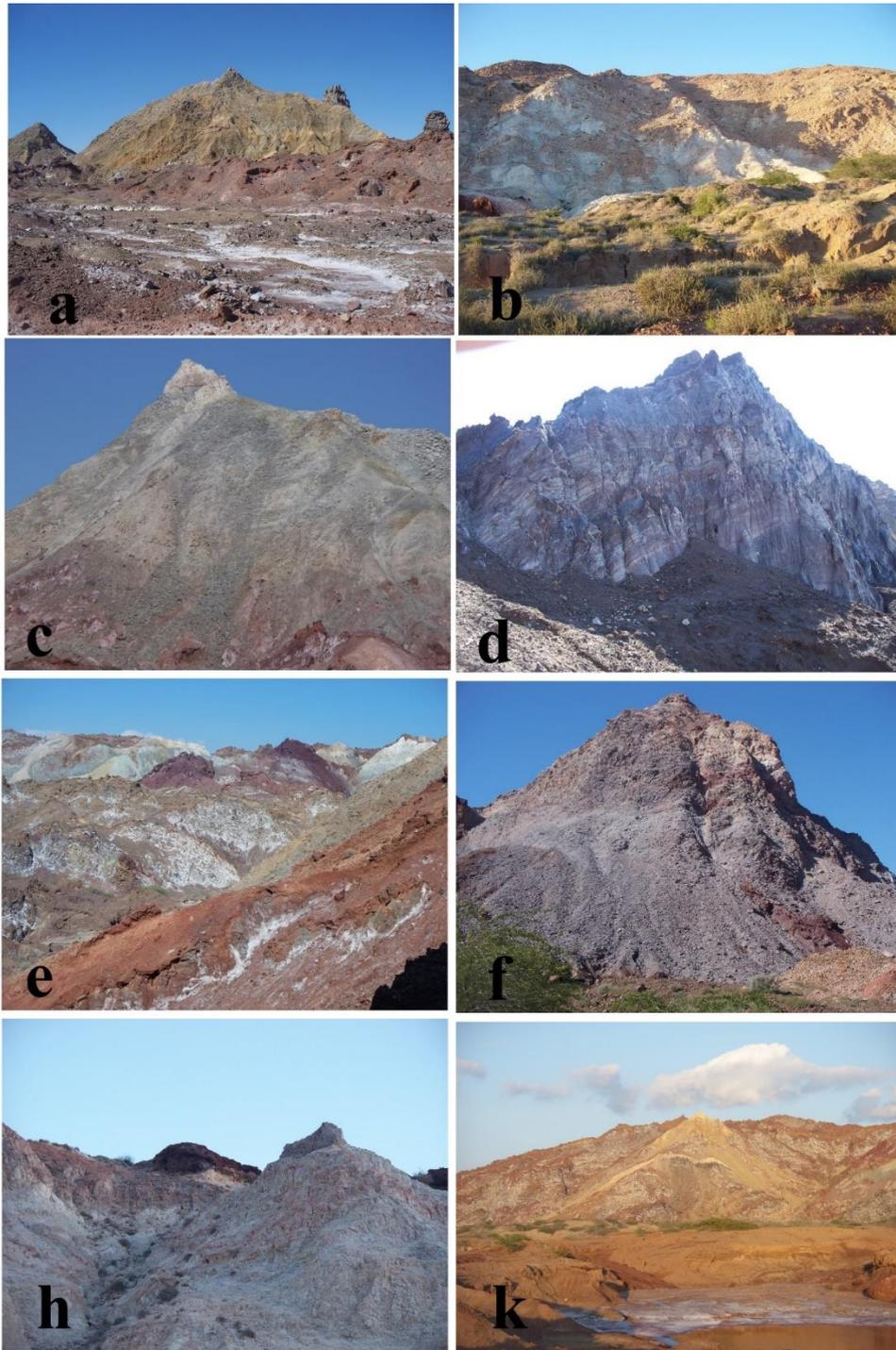
224 were directly extracted from WV-3 imagery (Fig. 5) using the Z-profile tool available in ENVI
 225 software. The representative sites of desired pixels were identified during field observations
 226 (Fig. 6), and appropriate rock samples were collected for further investigations.

227



228 Figure 5. a) Spectral curves of lithological groups, extracted from VNIR bands of WV-3, and
 229 b) locations and the names of collected end-members in a gray image.

230



231 Figure 6. Field photos of a) red soil and marl, b) anhydrite and marl, c) tuff, d) anhydrite and
 232 rhyolite, e) red soil, gypsum and anhydrite, f) diabase, h) gypsum, k) marl.

233 **3.3.2. Spectral angle mapper (SAM) algorithm**

234 This algorithm is categorized as a pixel-based image analysis technique and has extensively
 235 been applied by the remote sensing geologists (e.g., Qiu et al., 2006; Rajendran et al., 2013;

236 Markoski and Rolim, 2014). Kruse et al. (1993) indicated that this algorithm could identify the
237 similarity between a pixel of a data set and the reference spectra by calculating a spectral angle
238 (“ α ” in Eq. 1) between them. They suggested two n-dimensional spectral vectors for this
239 algorithm [Eq. 1], coincided with the spectrum of each pixel (r) and the spectrum of desired
240 end-member (t), in which the number of dimensions is equal to the number of applied bands.

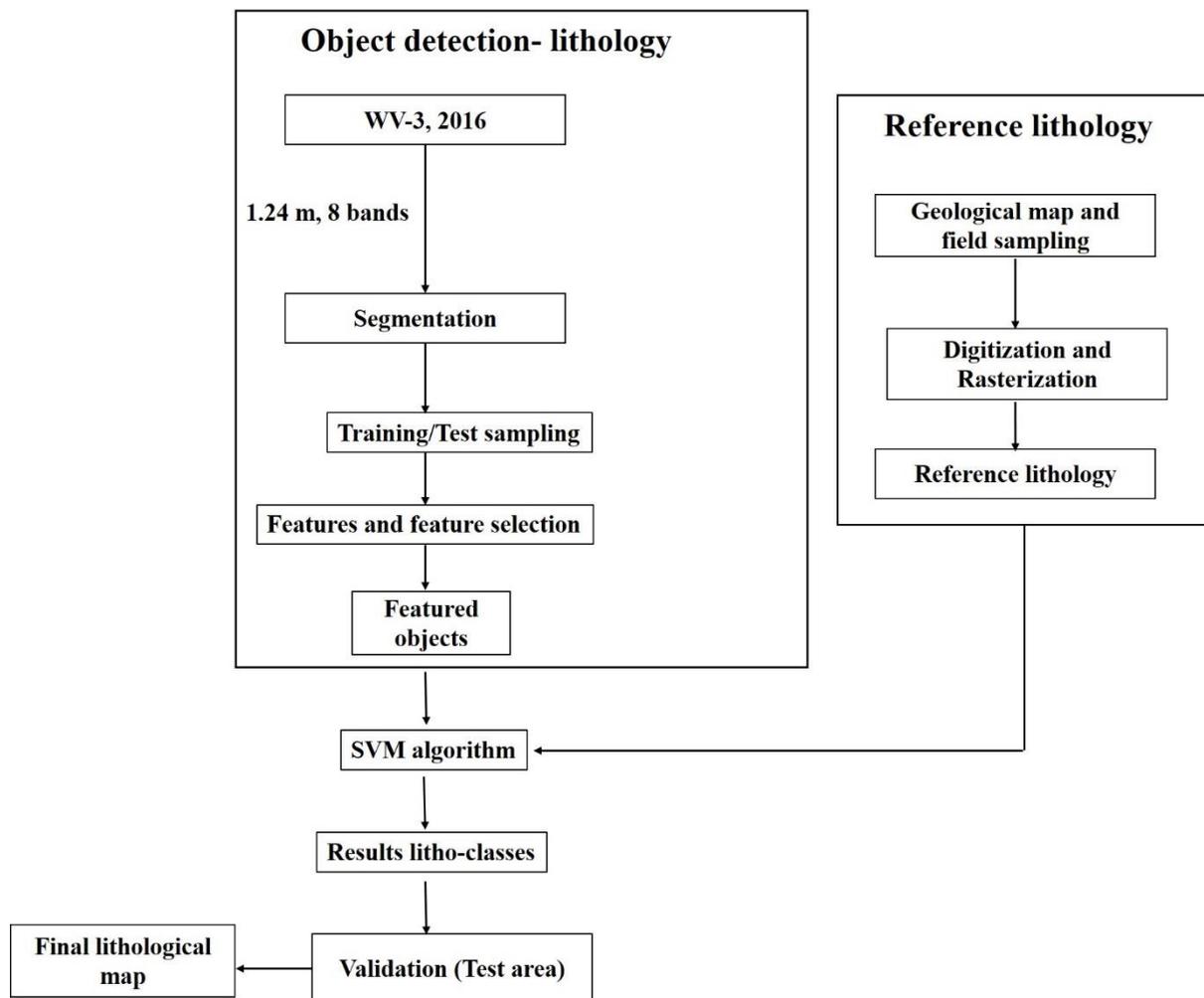
$$\alpha = \cos^{-1} \left(\frac{\sum_{i=1}^n t_i r_i}{\left(\sum_{i=1}^n t_i^2\right)^{\frac{1}{2}} \left(\sum_{i=1}^n r_i^2\right)^{\frac{1}{2}}} \right) \quad [1]$$

243 The pixels with lower spectral angles represent closer similarity to the reference spectrum and
244 appear darker (Research Systems, Inc., 2002; Jensen, 2005). For the propose of lithological
245 mapping, eight bands of WV-3 and the spectra of six previously identified lithological groups
246 were put into the SAM algorithm. Subsequently, the different threshold values per end-member
247 spectrum were examined, and finally, the appropriate pixels attributed to particular lithology
248 were realized based on the lowest spectral angles for the desired end-member.

249 **3.4. Object-Based Image Analysis (OBIA)**

250 The objects in the scale of satellite imagery are various sets of similar pixels that provide the
251 necessary information for the object-based image analysis method. These are similar groups of
252 pixels based on their spectral characteristics such as texture, shape, color, and conditions of
253 surrounding pixels (Tormos et al., 2012). The general workflow for this approach is presented
254 in Fig. 7, includes: 1) segmenting the image, 2) sample selection by the use of a stratified
255 random scheme (Mason et al. 1988), 3) feature selection for scale using correlation-based
256 feature selection (CFS) method (Dorren et al., 2003), 4) classifying the image using SVM
257 classifier (Hsu et al., 2007).

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Figure 7. Flowchart of the OBIA approach.

261

This algorithm defines the decision borders by giving priority to margins between support

262

vectors that spatially contain a minor geometric error (Borges 1998; Melgani and Bruzzone

263

2004). An essential issue for performing this algorithm is selecting a suitable kernel function,

264

which works with two other parameters, including gamma and C. Hsu et al. (2007), indicating

265

that the kernel function re-projects the varying space. C-factor controls the degree of

266

misclassification by SVM. They expressed that this algorithm randomly fixes the complex

267

decision borders with particular spatial specifications using the C-factor. On the other hand,

268

Hsu et al. (2007) also specified that the Gamma factor adapts the spread of the kernel function

269

and that the factor that both determines the spread of the kernel function and controls the

270 susceptibility of the decision boundary to confused support vectors, is set by C parameter.
271 Geologists such as Bahrambeygi and Moeinzadeh (2017) and Gasmi et al. (2016) have already
272 used the object-based image analysis approach in lithological mapping.

273 **3.4.1. Segmentation**

274 The primary step before performing the objected-based image analysis for lithological
275 classification is image segmentation, which leads to generating non-overlapping polygons. The
276 critical factor for defining the lithological segments is a scale that determines the accuracy of
277 image segmentation. Marceau (1999) suggested that exert of different scales in imagery could
278 be possible if the dimensions of the desired object are more significant than the pixel size. To
279 create the appropriate lithological segments in applied imagery, several scales were tested from
280 scale parameters of 5 to 20, considering the pixel size of 1.24 m for WV-3 data and the sizes
281 of objects. Finally, a scale parameter of 10 was selected to be applied to segmenting the image
282 at one scale for any lithological class. If the value of the scale parameter is high, it obtains
283 larger objects.

284 Blaschke and Burnett (2004) indicated that imagery could be categorized into comparatively
285 analogous and essential classes of pixels by an appropriate segmentation algorithm. These
286 segments are subject to be identified using a competent processing technique and converted
287 into relevant objects. Parameters of color and shape control the homogeneity criteria, in which
288 the summation of factor values is equal to 1 for each couple.

289 The degree of analogy in texture is determined by shape, which is a combination of smoothness
290 and compactness and helps extract the desired objects (Trimble, 2015). Considering that in
291 lithological segmentation, we would instead give the most crucial role to spectral information,
292 the ratio of 0.9/0.1 was set to color/shape. Moreover, the ratio 0.5/0.5 was set for

293 smoothness/compactness because we were reluctant to support the smooth or rough segments,
294 and the value 1 was assigned as the weight of the image layer to prevent any prejudice.

295 **3.4.2. Training and Test sampling**

296 A significant phase after segmentation of the image and before implementing an object-based
297 image analysis algorithm is selecting the various rock units that are going to play the role of
298 training samples. Sampling schemes such as systematic, cluster, simple random, and stratified
299 random have already been used by various algorithms (Congalton and Green, 2009). In this
300 study, the selection of lithological training samples was based on the stratified random
301 sampling with the purpose of selection of enough number of polygons for each distinguished
302 lithology group. When performing this sampling method at the scale parameter of 10, the visual
303 and field interpretations of lithological features were used as references for obtaining deduced
304 knowledge of areas. The main issue in stratified random sampling strategy is the correct
305 lithological interpretation of the area for which we used the previous studies (e.g., Mahyari,
306 2016), and field observations. The geological map and GPS points of field observations were
307 then overlapped on the segmented layer to assign a class label to each segmented object.

308 **3.4.3. Features and Feature Selection**

309 The object-based method produces more features than a pixel-based approach due to its logic
310 in engaging the segmented objects. The frequently used features of eCognition software,
311 version 9.0 (Trimble, 2015), including spectral measure, shape, and texture, were directly
312 calculated using WV-3 bands. The spectral measures including mean, max, mode, difference,
313 standard deviation, and brightness were calculated for each lithology using aerosol, blue, green,
314 yellow, red, red edge, NIR1, and NIR2 bands. The shape measures consisting of area,
315 roundness, main direction, density, compactness, rectangular fit, elliptic fit, border index, shape
316 index, and asymmetry were also calculated. Furthermore, the texture measures, including Gray-

317 Level Co-occurrence Matrix (GLCM), homogeneity, and Gray-Level Difference Vector
318 (GLDV) were estimated on the basis of pixels of every lithology. Based on WV-3 bands, other
319 texture parameters consisted of contrast, dissimilarity, entropy, standard deviation, correlation,
320 mean, GLDV angular second moment, entropy, mean, and contrast measured.

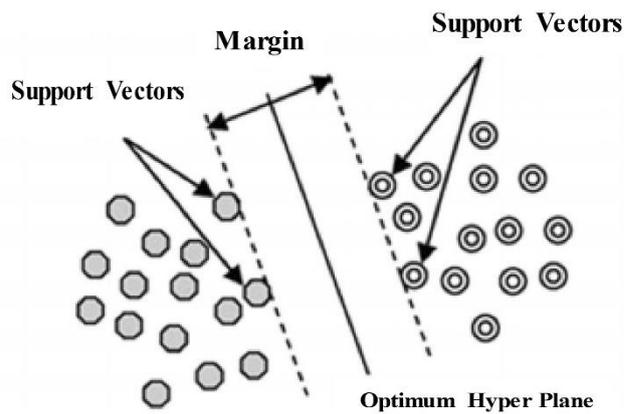
321 **3.4.4. Support Vector Machine (SVM)**

322 The classification method applied in this study (support vector machine) works based on the
323 hypothesis of machine learning through a supervised learning process. This algorithm isolates
324 two desired classes and enlarges the space between them by creating a hyperplane (Kavzoglu
325 and Colkesen, 2009). This method is based on the belief of maximum margin (Fig. 8), which
326 is the distance between identified boundary for classes and the closest samples, and the idea of
327 transforming extent of depiction on the applied data set to the extent of excessive size. The
328 support vectors are intended samples positioned adjacent to the borders of decision (Fig. 8)
329 (Oommen et al., 2008). Four types of kernels, including linear, sigmoid, polynomial, and radial
330 basis function (RBF), execute the concept of transformation of the SVM algorithm (Hsu et al.,
331 2007).

332 The SVM algorithm defines the decision borders by giving priority to margins between support
333 vectors that contain a minor geometric error in space (Borges 1998; Melgani and Bruzzone
334 2004). An essential issue for performing this algorithm is the selection of suitable kernel
335 function, which works with two other parameters including gamma and C. Karatzoglou et al.
336 (2004) stated that when performing SVM algorithm, the varying space is re-projected by kernel
337 function, and C-factor directs the level of misclassification. They described that this algorithm
338 randomly fixes the complex decision borders with particular structural specifications by using
339 the C-factor, in which these structures are a basis of support vector positions in varying space.
340 Karatzoglou et al. (2004) also expressed that the extent of the kernel task is balanced by the

341 Gamma factor, and the C parameter sets the item that determines the extent of the kernel task;
342 this parameter also modifies the susceptibility and of the decision boundary to those support
343 vectors which are confused.

344 Although many types of kernels are available for the SVM, the RBF was applied in this study
345 because it is suggested as a suitable primary option (Karatzoglou et al. 2004).



346 Figure 8. The optimum hyperplane, margin, and support vectors in the SVM algorithm
347 (Kavzoglu and Colkesen, 2009).

348 3.5. Accuracy Assessment

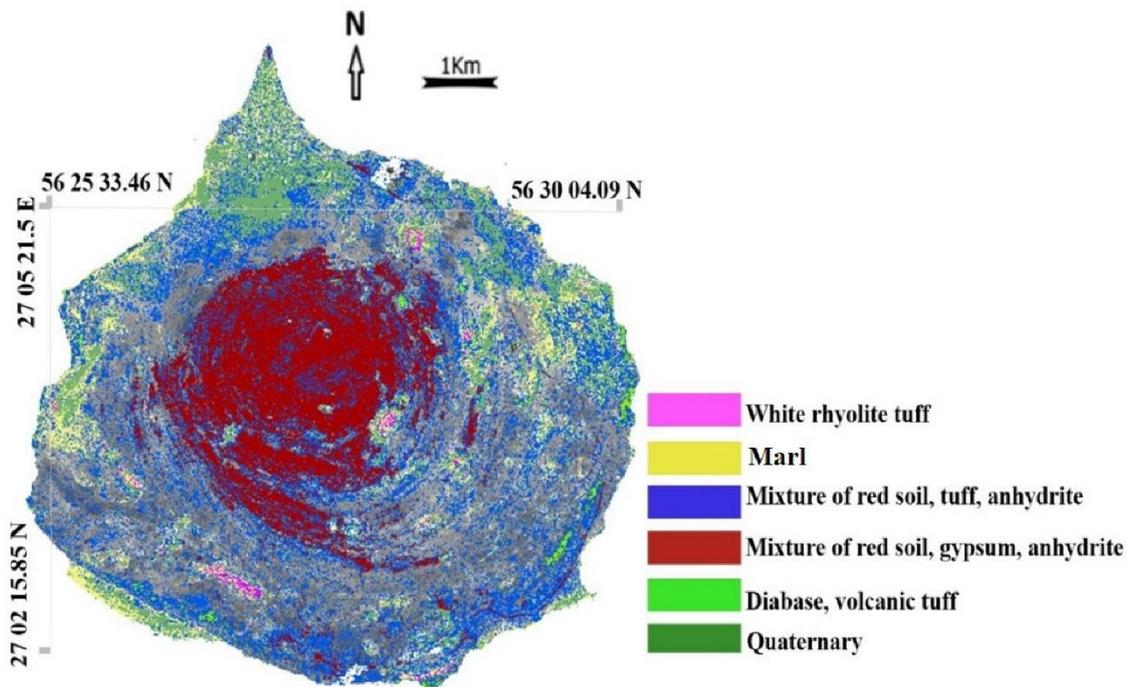
349 The degree of correspondence between PBIA and OBIA classification results and the field and
350 laboratory evidence were assessed to evaluate the lithological plausibility of each classifier
351 output. A random sampling of rock types provided a set of samples that were then spectrally
352 and petrographically analyzed to verify their lithology. The overall accuracy of results was
353 estimated based on a confusion matrix (Congalton and Green, 2009). In this regard, the
354 reference sites with 5023 pixels for SAM and 5949 pixels for SVM algorithms were selected
355 through visual interpretation of the images associated with the general in situ and laboratory
356 validations.

357

358 **4. Results and Discussion**

359 **4.1. PBIA results**

360 To enhance the desired lithological units by spectral angle mapper, various threshold values
361 were examined for spectral angles, and the maximum angle of 0.1 in the range of 0.0-1.0 was
362 set acceptable. In order to produce a lithological map (Fig. 9) from SAM output images, a
363 classification code was assigned to each pixel based on its closest match to the reference
364 spectrum (Kruse et al., 1993; Boardman and Kruse, 1994). The output image shows spatial
365 overlaps for rhyolite and marl units (green and brown pixels). The marl units are more extended
366 eastern and northern the study area in the output image, than what was observed in the field.
367 The mixtures of red soil, tuff, and anhydrite are not well discriminated thought the area. The
368 Quaternary deposits that are mostly extended at the northern parts of the island are
369 misclassified as a salt and pepper mixture. Furthermore, a great circular area peripheral to the
370 central yellow class (Fig. 9) is not attributed to any lithological unit (gray pixels). The overall



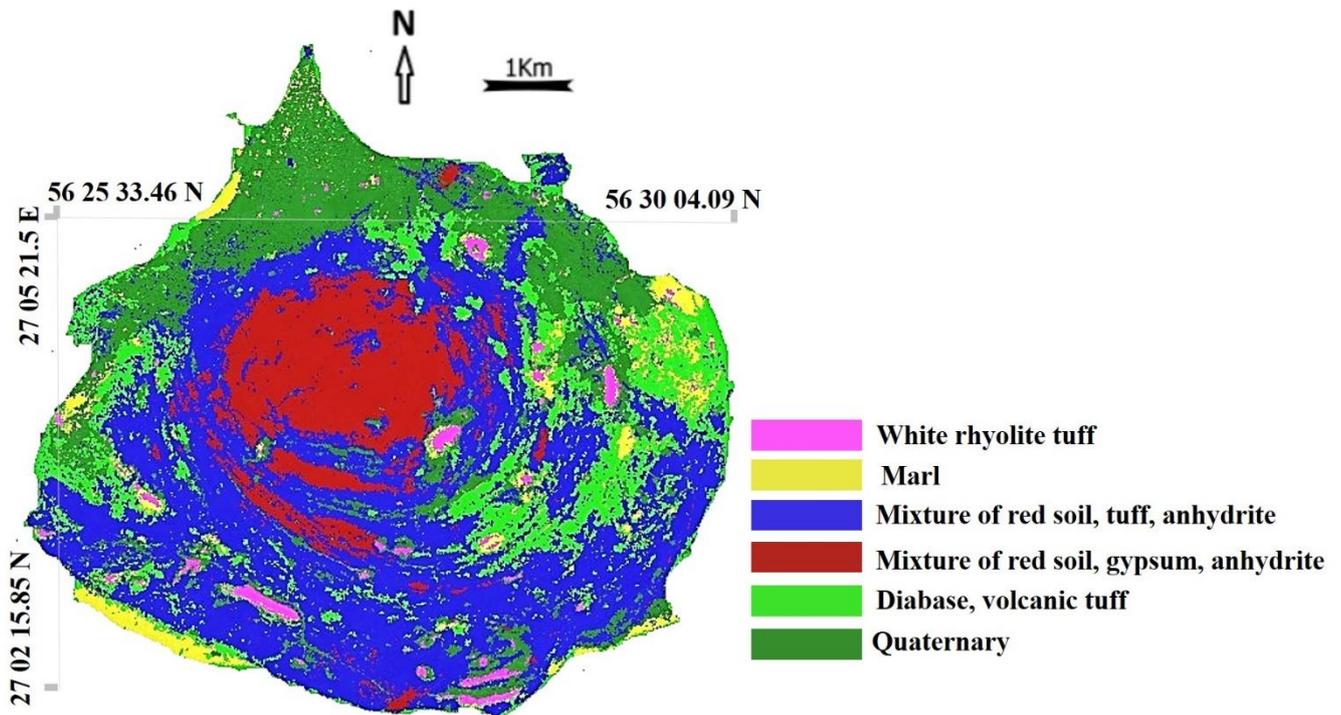
371 accuracy parameter of a confusion matrix was used to assess the validity of the lithological
372 map achieved by this classifier.

373 Figure 9. Classification map of lithological units in Hormuz Island using the Spectral Angle
374 Mapper algorithm.

375 **4.2. OBIA results**

376 In order to discriminate the lithological units, six lithological codes were assigned to extracted
377 objects in SVM classifier. The gamma and C parameters were set to 0 and 2, and the final
378 classification map was produced by the use of the original WV-3 dataset (Fig. 10). In general,
379 this output map showed the lithological extensions and borders more clear and transparent than
380 SAM output, and the whole area was successfully divided into assigned lithological classes.
381 The highest similarity in shape and extent of the classes in output images of the two methods
382 (Figs. 9 and 10) belongs to the mixed class "red soil, gypsum and anhydrite" with a circular
383 shape at the central part of the Hormuz Island. However, the shape, size, borders, and structures
384 of other lithological units are significantly different in two output maps considering that they
385 are highly explicit and recognizable in maps produced by SVM. This method was capable of
386 achieving reliable results by considering the specific spectral absorption and reflection features
387 of desired objects and their textures and spatial relationships. The various lithological types
388 and the Quaternary sediments were also successfully discriminated via the segmentation
389 process performed based on appropriate training areas, even in few numbers.

390 Furthermore, field observations and controls showed that the lithological units extracted by the
391 SVM method are more precise than the results obtained by the SAM technique. Results
392 obtained by SVM demonstrated that lithology classification based on the texture features and
393 spectral characteristics of a high spatial resolution data such as WV-3 outstandingly
394 outperforms the pixel-based image analysis approaches such as SAM technique. The accuracy
395 of lithological classes obtained by this classifier was assessed and presented in the next section
396 using the overall accuracy of a confusion matrix.



398 Figure 10. Classification map of lithological units in Hormuz Island as produced by the SVM
 399 algorithm.

400 4.3. Accuracy assessment

401 Table 1a shows the confusion matrix of results obtained by the SAM. It showed that the SAM
 402 has accurately mapped the diabase with a volcanic tuff class that occurs in central and eastern
 403 parts of the study area, with a user's accuracy of 82.35% and producer's accuracy of 78.87%
 404 (Table 1a). A mixture of red soil, gypsum, and anhydrite was classified with a producer's
 405 accuracy of 70.20% and the user's accuracy of 58.82%. However, the moderate user's accuracy
 406 of a mixture of red soil, gypsum, and anhydrite, white rhyolite tuff, and Quaternary deposits
 407 could be attributed to the similarity in spectral characteristics this lithological classes in applied
 408 bands of WV-3. Table 1b reveals the confusion matrix of the SVM output. Results showed that
 409 the object-based mapping method has been magnificently more accurate in mapping the white
 410 rhyolite tuff and mixtures of red soil, tuff and anhydrite with user's accuracies of 84.14% and

411 92.13% and producer's accuracies of 89.16% and 75.02%. The overall accuracies for SAM and
 412 SVM results were 70.16 and 86.03, respectively (Table 1).

413 Table 1. Confusion matrices for SAM (a) and SVM (b) classification methods (MRGA: Mixing of red
 414 soil, gypsum and anhydrite, MRTA: Mixing of red soil, tuff and anhydrite, DVT: Diabase with volcanic tuff,
 415 WRT: white rhyolite tuffs, M: MARL, and QD: Quaternary deposits).

a)	SAM	MRGA	MRTA	DVT	WRT	M	QD	Total (Pixels)	User.ac.
MRGA	490	153	0	110	0	80	833	58.82	
MRTA	88	510	0	60	180	130	968	52.68	
DVT	120	0	560	0	0	0	680	82.35	
WRT	0	0	150	530	110	130	920	57.60	
M	0	0	0	188	650	119	957	67.92	
QD	0	0	0	0	220	475	695	68.34	
Total (Pixels)	698	663	710	888	1160	934	5023	64.61	
Prod.ac.	70.20	76.92	78.87	59.68	56.03	50.85	65.42	75.71	

416 Overall accuracy = 70.16%

b)	SVM	MRGA	MRTA	DVT	WRT	M	QD	Total (Pixels)	User.ac.
MRGA	870	130	0	0	0	0	1000	87.00	
MRTA	70	820	0	0	0	0	890	92.13	
DVT	0	50	1150	90	0	0	1290	89.14	

WRT	0	0	59	741	78	0	878	84.39
M	0	93	0	0	667	71	831	80.26
QD	0	0	0	0	180	880	1060	83.01
Total (Pixels)	940	1093	1209	831	925	950	5949	85.98
Prod.ac.	92.55	75.02	95.11	89.16	72.10	92.63		86.09

417 Overall accuracy = 86.03%

418 **4.4. Discussions**

419 This study's primary purpose was to compare the performance of a pixel-based image analysis
420 approach versus an object-based approach in the lithological mapping of complex terrain. The
421 capability of approaches in discriminating lithological units was evaluated by the use of
422 confusion matrix parameters (Table 1). Results showed that the object-based approach
423 outperformed the pixel-based method with an average difference of 19.33% in overall
424 accuracy.

425 The spectral-based techniques involve two drawbacks: 1) extraction of spectra from known
426 pure materials, 2) calibration of the pixel spectrum. These techniques are performed based on
427 an approach in which the pixel spectrum is compared to the spectra of a known pure material.
428 Spectra of these materials are generally extracted from imagery or measured of field-collected
429 samples, and if needed, they are selected from known spectral libraries. Methods for extraction
430 of spectra from imagery typically search for pure pixels. Although these methods depend on
431 the size of pixels, such pixels might be rare on the surface. Therefore, the numerical values of
432 these spectral end-members may commonly be associated with noise. This noise shows that

433 spectral discrimination is devalued when a pixel is a mixture of two rock types occurring next
434 to each other.

435 With recent advances in capabilities of the satellite data such as in WV-3, more studies are
436 focusing on the texture of images and extraction of contextual information that is a measure of
437 association between the values of neighboring pixels (Marceau et al., 1990; Hay and Niemann,
438 1994). In comparison to the pixel-based methods that only rely on the DN values of pixels,
439 segments in object-based approaches obtain extra information on the spatial behavior of the
440 objects, which makes it more advantageous (Blaschke and Strobl, 2001; Darwish et al., 2003).
441 Consequently, it is suggested that the object-based methods are more efficient than per-pixel
442 algorithms for mapping the various rock units because the decrease of intra-class variability
443 happens when averaging the DN values of all nearby pixels within objects such as rock classes.
444 Depending on the type of classifier and the input dataset, the efficiency of object-based image
445 analysis methods for target enhancement, such as in the case of a lithology, could be different,
446 although, in general, this approach outperforms the pixel-based methods. Another fact for such
447 variability is that a unique value of scale is not perfect for segregating all the lithological
448 categories.

449 The improved enhancement and discrimination of Quaternary deposits, marl, rhyolite, a
450 mixture of red soil, tuff and anhydrite, and a mixture of red soil, gypsum and anhydrite in this
451 study (Figs. 9 and 10), revealed that the object-based image analysis method is superior over
452 that of the pixel-based approach. In the OBIA classification map, correct spatial distribution is
453 displayed for rhyolite, marl, Quaternary deposits, a mixture of red soil, tuff and anhydrite, and
454 a mixture of red soil, gypsum, and anhydrite. The detection of Quaternary deposits has not
455 always been as easy as with other units. This is mainly because it is a combination of various
456 products of weathering and erosion of upstream outcrops. The visual interpretation of output

457 results in Figure 11-a confirms the improvements of the OBIA method in enhancement and
458 discrimination of lithological units and attributing all pixels to desired classes. Moreover,
459 Figure 11-b displays that considering the importance of intra-class discrepancies for OBIA and
460 similarity in the spectral properties in the PBIA method, the white rhyolite class is efficiently
461 classified by OBIA; however, it is misclassified by PBIA with marl unit. This drawback is also
462 observed for other classes in outputs of the SAM algorithm. Besides, Figure 11-c shows that
463 the object-based method was successful in decreasing the salt-and-pepper pixels associated
464 with spectral-based mapping. The common issues in low-resolution pixels are the
465 heterogeneity in spectral properties of rock units and spectral differences between rocks,
466 vegetation, and Quaternary deposits, in cases that all are present in one pixel.

467 When generating the attributes of objects, such as rock types, in the OBIA approach, the
468 spectral characteristics of all the pixels of a given object are averaged. This leads in decreasing
469 the mapping confusion by reducing the variations within an object. An essential disadvantage
470 of a pixel-based mapping method is that it does not use the data of neighboring pixels to support
471 more correctly recognition of a target class for a pixel. Consequently, if pixels of a class of
472 lithology exhibit local spectral heterogeneity, they may be labeled as different classes.
473 Therefore, the pixel-based methods could obtain a high rate of misclassification such that
474 specific regions of a class of rock might wrongly be classified as another rock unit.
475 Furthermore, if per-pixel methods are applied, usage of imagery with high spatial resolution,
476 such as WV-3, which is needed to separate the small areas of specific rock units, may lead to
477 increased errors in classifications.

478 This study aimed to classify the lithological groups of a geologically complex terrain by
479 focusing on the high spatial resolution of WV-3 VNIR data. This advantage of WV-3 led to
480 the successful enhancement of lithological boundaries in the study area, including places where
481 lithology is not homogeneous, and the outcrops are small. Another investigation has also

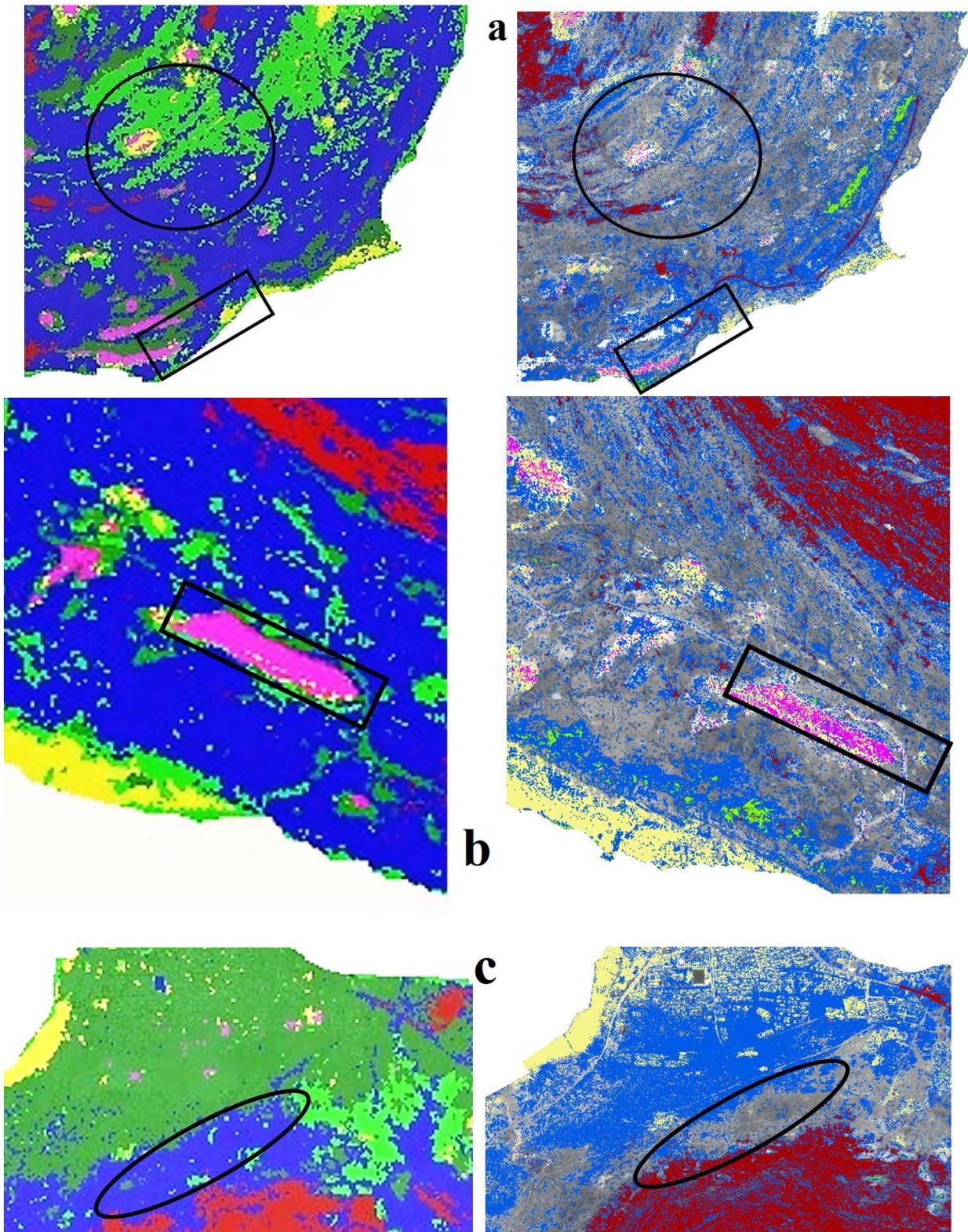
482 suggested the advantageous usage of the spatial resolution of WV-3 in mapping the geological
483 targets (Sun et al., 2017). Bedini (2019) showed that integration of WV-3 and ASTER TIR data
484 could successfully recognize the lithological units by using spectral properties and indices
485 calculated from WV-3 imagery.

486 Although in this study, we proved the advantage of the OBIA method in comparison to the
487 PBIA approach for an improved lithological mapping using only the VNIR bands of WV-3,
488 further works to consider the SWIR bands of this satellite is recommended.

489 Although it is suggested that considering absorption features in the segmentation process gives
490 satisfactory results (Grebby, 2016), their application could be more various than region
491 growing. This research emphasizes the efficiency of object-based image analysis in reducing
492 the spectral variability within an object (here, lithology) and the conjunction of supplementary
493 information extracted from structural and contextual image/object properties to improve the
494 enhancement of rock units.

Objects-based approach-SVM

Pixel-based approach-SAM



495 Figure 11. Comparison of results achieved by object-based and pixel-based approaches: a)
496 improvement in lithological mapping outcrops, b) omission of the ambiguous mixture, and c)

497 filtering salt and-pepper pixel in a mixture of red soil, gypsum, and anhydrite.

498 **5. Conclusions**

499 This study was a comparative approach to show the capabilities of pixel-based and object-
500 based methods and their representative algorithms, SAM and SVM, in discriminating the
501 lithological classes using VNIR data of WorldView-3 of Hormuz Island, southern Iran. Results
502 obtained by these two approaches revealed that the OBIA method was superior compared to
503 the PBIA method. The OBIA could lead to an improved discrimination of lithological groups,
504 clear detection of geological units with complex lithology such as Quaternary deposits, and
505 successful decrease or remove of salt-and-pepper pixels, which were common in the spectral-
506 based output map.

507 Comparing the degree of efficacies of applied methods illustrated that the OBIA conforms to
508 a type of expert interpretation aiming to determine the internal relationships among
509 neighboring pixels. This advantage leads to relatively perfect classification of features, while
510 that of the pixel-based approach is segregated. Furthermore, the WV-3 data, because of its high
511 spatial resolution, is notably suitable for the OBIA approach aiming at the discrimination and
512 classification of lithological units in a geologically complex district. Moreover, realizing the
513 same ideal texture groups by SVM method is a basis for lithological mapping and classifying.
514 It was shown that the OBIA approach produces a more improved and contiguous lithological
515 map than the PBIA method. Overview of the criteria mentioned above showed that the pixel
516 size of 1.24 m for VNIR bands of WV-3 is particularly advantageous for lithological mapping
517 by using the OBIA method rather than the PBIA method.

518

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