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3 1 **Complete Multivariate Flood Frequency Analysis, applied to**
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7 2 **Northern Algeria**
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Abstract

Extreme hydrologic events are commonly described by several dependent characteristics, such as duration, volume and peak flow for floods. Traditionally in Algeria and North Africa, flood frequency analysis (FFA) is conducted as a univariate approach focusing separately on each single of flood characteristics. On the other hand, elsewhere, multivariate FFA studies have been conducted focusing on some FFA steps (especially modelling). The current study aims to consider complete multivariate FFA at-site case studies in Northern Algeria using 11 hydrometric stations. It is also among the first studies dealing with multivariate FFA in a complete way by considering all the required steps of the analysis (multivariate outliers detection, multivariate assumptions testing and copula fitting) and on datasets from Algeria. Multivariate stationarity, homogeneity and independence assumptions have been well verified before modelling. The Weibull distribution is mostly selected as margin distribution for the duration, volume and peak flow series. Frank, Clayton and Gumbel copulas are commonly selected to describe the dependence structure on the three flood pairs of variables. These findings should be interesting in water management and flood risk assessment in these regions. Combining these flood characteristics enables the design of more efficient hydraulic structures.

Keywords: Copula; daily flow; flood characteristics; multivariate assumptions; multivariate frequency analysis; Northern Algeria

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1. Introduction

Hydrological frequency analysis (HFA) aims to relate the magnitude of extreme events to their frequency of occurrence through the use of probability distributions (Chow et al. 1988 in Rao & Hamed 2000). HFA is mainly composed of the following steps: (a) exploratory analysis and outlier detection; (b) checking the basic assumptions of stationarity, homogeneity and independence; (c) modelling and parameter estimation; (d) risk evaluation and analysis (e.g. Chebana 2012).

Flood events can be described as a multivariate stochastic phenomenon by three main variables, that is its duration (D), volume (V) and peak flow (Q_p). Hence, the associated risk analysis should be treated by an effective multivariate probabilistic approach (Reddy & Ganguli 2012). In this context, univariate HFA (based on a single variable) has a number of limitations and drawbacks which leads to considering the multivariate HFA. The latter is the object of a growing number of studies (e.g. Zhang 2005; Grimaldi & Serinaldi 2006b; Kao & Govindaraju 2007; Chebana & Ouarda 2009; Chowdhary 2009; Chebana & Ouarda 2011a, Serinaldi & Kilsby 2013; Requena et al. 2013; Genest & Chebana 2017). The above HFA steps are commonly employed to study flood risk in a univariate setting (e.g. Cunnane 1987 and Rao & Hamed 2000) which is not the case in the multivariate one (see e.g. Chebana et al. 2013).

The copula approach overcomes the restriction of univariate HFA by allowing a combination of different marginal types. It also provides a wider choice of admissible dependence structure as compared to the conventional approach (Chowdhary 2009). The use of copulas in multivariate HFA becomes the appropriate tool to study the dependence structure between the different variables.

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3 67 Hence, several studies were considered copulas in multivariate HFA, e.g. De Michele &
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5 68 Salvadori (2003) identified the relation between intensity and duration of storm rainfall in Italy
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8 69 by using Frank copula. Renard & Lang (2006) suggested applications of Gaussian copula on
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10 70 flood mitigation in France. Grimaldi & Serinaldi (2006a) have proved the adequacy of two
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12 71 copulas (Frank and Gumbel) on the flood characteristics analysed for Kanawha River in West
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14 72 Virginia (USA). Chebana & Ouarda (2007, 2009) presented regional multivariate flood analysis
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17 73 using copula and multivariate L-moments. A Gumbel copula was used by Leonard et al. (2008)
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19 74 to couple the seasonal rainfall maxima marginal distributions on the Murray–Darling Basin,
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21 75 Australia. El Adlouni & Ouarda (2008) proposed the application of copula to analyse the
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24 76 dependence of the water level of Saint-Louis Lake on the maxima flow on the Chateaugay
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26 77 River in Quebec (Canada). Leite Rosa (2011) presented that Frank and Clayton copula fit good
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28 78 in studying a relationship between maximal flow and volume in Portugal. To establish the
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31 79 relation between the different flood characteristics Salarpour et al. (2013) applied the t-copula on
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33 80 the Johor River in Malaysia. A Gumbel copula was selected as the most appropriate model for
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35 81 trivariate frequency analyses of peak discharges, hydrograph volumes and suspended sediment
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38 82 concentrations in Bezak et al. (2014). Drought frequency analysis in Medjerda River basin
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40 83 (Tunisia) was carried using copula in Hamdi et al. (2016) which can be considered as the first
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42 84 multivariate study in North Africa.

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45 85 In Algeria, to our knowledge, there are no published works dealing with multivariate HFA,
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47 86 especially for floods. However, flood frequency analysis (FFA) has always been carried on
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50 87 different approaches involving univariate context, either on flood flows (e.g. Ketrouci & Meddi
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52 88 2012; Rezak 2014), on maximum daily flow in Boutoutaou et al. 2011, on maximum daily
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54 89 rainfall (e.g. Benhattab et al. 2014; Meddi & Belhadj Bouchaib 2010); on maximum annual by
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56 90 Benameur et al. 2017 under arid climate of Biskra in southern Algeria; and on Intensity-

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91 Duration-Frequency curves (IDF) in order to describe the rainfall threshold generating flood (e.g.
92 Yamani et al. 2016).

93 Owing to numerous tragic flood events in Algeria (Table 1) a number of 575 projects against
94 flood with a budget of 2 billion Dollars have been launched since 1999, covering 33 Northern
95 provinces over 42. Unfortunately, the desired object is not reached as the floods continue to
96 record serious problems and considerable disasters. Each year, an average of seven to ten large-
97 scale floods is affecting mainly the big cities. A relevant issue has arisen, therefore, looking for
98 new techniques and procedures with better performances to mitigate such hazard event seems a
99 promising solution. In this vision, the present work aims in studying the floods in Northern
100 Algeria within the multivariate framework to provide a blanket solution for flood risk
101 assessment. This study focuses on the regions most affected by repeated floods (as summarized
102 in Table 1). These basins require adequate facilities planning to mitigate flood risk. Multivariate
103 FFA constitutes a powerful tool allowing the dimensioning of suitable installations based on a
104 more accurate flood risk assessment (e.g. Callau Poduje et al. 2014; Balistrocchi et al. 2017;
105 Jiang et al. 2019; Zhou et al. 2019; Liu et al. 2019).

106 Hence, the objective of this study is to consider the multivariate FFA for Northern Algerian
107 basins based on the most recent techniques with a complete analysis. This work is the first
108 contribution on analysing floods with a multivariate approach considered in Algeria. It is applied
109 to 11 different selected hydrometric stations in Northern Algeria. These data are the subject of
110 the first application on multivariate FFA. On the other hand, the current work is considered as a
111 complete HFA study including, detection of multivariate outliers, multivariate hypothesis testing
112 which are rarely reported combined in the hydrological literature, modelling and risk evaluation.
113 Indeed, Chebana et al. (2013) indicated that ignoring the testing step in multivariate HFA may
114 lead to inaccurate or wrong results and hence to inappropriate decisions regarding, for instance,
115 the design of hydraulic infrastructures.

2. Area and data

The study area covers 11 basins in Northern Algeria investigated to analyse flood (Figure 1).

This region extends from the Tunisian border on the East to the Moroccan border on the West

with an area of 11400 km². It is located between 8° East and 2° West longitudes and 37° and 33°

North latitudes. It has an average annual temperature of 17°C, and the average annual rainfall is

about 900 mm on the East and 350 mm on the West. The basins selected are the most affected by

repeated floods since the last fifty years. Important damages are observed mainly caused by

watercourses (wadi) flooding. The average annual hydrological yield can reach 3250 Hm³ on the

extreme East to 5 Hm³ on the extreme West.

The study area covers 11 basins in Northern Algeria investigated to analyse flood (Figure1). The

characteristics of the selected stations are presented in Table 1. Daily average flow is the raw

data applying in this study. In some areas, data access is very restricted and hence limits the

availability to the daily flow data. The available database from the National Agency of Hydraulic

Resources (ANRH) had 30 hydrometric stations. However, after preliminary analysis of this

database, we found that the majority of these stations either have short record periods (less than

15 years), or their hydrological regime is disrupted (e.g. by dams), or their series present many

gaps. The presence of such anomalies forced us to select only 11 hydrometric stations to conduct

the multivariate FFA. To overcome some these limitations, reconstruction of the historical data

can be considered for future studies, which may allow extending the FFA to other stations

(Grimaldi et al. 2016). In addition, filling the data gaps in a multivariate context using copula is

also a promising option (Aissia et al. 2017).

Indeed, the hydrological regime of the studied basins shows that these wadis flow only during

the wet season (October to May); while the dry season is marked by the lowest or even zero

flows. The dates of the floods show that more than 50% of the analysed floods occur during the

winter season (December, January and February), the spring (40 to 20%) then the fall and rarely

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3 141 in the summer (0 to 19%). These observations are valid at hydrometric stations in the east and
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5 142 centre of the country. However, the stations in the far west (Ghazaouet, Sidi Ali Ben Youb and
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7 Taria) indicate different behaviour, where autumn and spring dominate in the flood dates. The
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10 144 summer season registers 11 to 26% of the total floods in this region. The transition between the
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12 145 Mediterranean and semi-arid climate create a mixed hydrological regime on the extreme north-
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14 western regions (Ghazaouet station). Soultès wadi (Timgad station) is characterised by a
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17 147 complex hydrological regime under a semi-arid climate. A simple hydrological regime
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19 148 dominates the other studied basins (Meddi et al. 2017).

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21 149 The annual average yields of the studied wadis presented in Table 1 are very low ranging from
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23 2.9 to 304.1 10^6 m³/year. This is usually in the Algerian context where the flows record the value
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26 151 of zero m³/s during the dry season (June to September). In Algeria, the hydrological regime of
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28 152 the stream is characterized by temporary flows following the rainfall contributions with the
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30 appearance of the systematic seasonal fluctuation of flows. Indeed, the stream flow is very
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33 154 irregular and low flow is insignificant in the dry period, as it is discussed in different previous
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35 155 papers (e.g. Bouanani, 2004; ABH-OCC, 2006; Touazi et al. 2011; Hallouz et al. 2013; Meddi et
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37 al. 2017).

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40 157 Each series of daily flow makes it possible to extract three series of variables (D, V, Qp)
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42 characterizing each flood as described below. The missing values in the extracted variables are
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44 due to the presence of one or more missing values in the crude database. The Lakhdaria station
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46 shows an example with four missing values for each extracted variable. The obtained values are
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48 verified with the associated hydrograph while the extracted data series are presented in Appendix
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3. Methodology

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57 164 The considered methodology covers all appropriate aspects of multivariate HFA, including data
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59 extraction, preliminary analysis, checking assumptions, modelling as well as risk assessment.
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3 166 Indeed, the selection of flood characteristics is the most important step in this work since all the
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5 167 subsequent results depend on the availability and quality of the input data. It would be more
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8 168 reasonable to separate the fast response from the slow response to estimate adequately the
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10 169 average daily flow (q). However, the dataset obtained from the ANRH presents the average daily
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12 170 flow (q) in m^3/s . This format does not allow separating the fast response from the slow response.
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15 171 The method applied to estimate the flood features is based on a daily average value of flow (q) to
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17 172 construct the yearly hydrograph which corresponds to the hydrological year starting on
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19 173 September 1st of the year (i) and finishing on August 31st of the year ($i+1$). Then, the associated
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21 174 flood characteristics are extracted. The peak flow (Q_p) corresponds to the highest q over the
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24 175 year. The base flow index method has been employed to separate base flow and flood flow (e.g.
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26 176 Yang et al. 2019). In this method, the starting date of the flood coincides with the rise on the
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29 177 annual hydrograph. The ending date refers to the recession curve and the back to the low flow
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31 178 regime. These dates are determined by adjusting the slopes with a linear approximation of the
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33 179 annual hydrograph (e.g. Ben Aissia et al. 2011). Flood duration is the number of days between
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35 180 the start date and the end date of the flood. The volume is calculated by summing the daily flow
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38 181 values recorded during the specified time (here flood duration). The extracted multivariate series
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40 182 (Q_p , D , V) is verified year by year for each station.

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42 183 This data extraction is done similarly to previous studies applied in North America (Canada) by
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44 184 Ben Aissia (2009); Jeong et al. (2013) in Northeastern Canada; Karmakar & Simonovic (2007)
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47 185 in USA; Bačová Mitková & Halmová (2014) in Slovakia; Singh et al. (2015) in India; Hamdi et
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49 186 al. (2016) in Tunisia on droughts.

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52 187 As a first phase, univariate HFA is carried out on the extracted series for each variable. A
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54 188 descriptive statistical analysis is performed on each of the three flood variables (D , V and Q_p);
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57 189 by determining the principal statistical features, such as the mean, the standard deviation, the
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59 190 coefficient of variation (C_v), of skewness (C_s) and kurtosis (C_k). Then, the Pearson correlation
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191 coefficient is calculated to identify the correlation between flood variables. The t-test is
192 considered to check the significance of the correlation.

193 Detection of outliers from the analysed variables is checked. Hunt et al. (1981) defined an *outlier*
194 as "an observation that does not conform to the pattern established by other observations".
195 Rosner test is employed to detect univariate outliers from the analysed series. Rosner's (1983)
196 method assumes that the main body of data is from a normal distribution. Rosner's approach is
197 designed to avoid masking of one outlier by another (in Gilbert 1987).

198 Testing the basic assumptions on hydrologic data must be verified before starting any HFA. In
199 the univariate framework, Mann-Kendall (Mann 1945; Kendall 1975), Wald-Wolfowitz (1943)
200 and Wilcoxon are respectively selected to test stationarity, independence and homogeneity of
201 hydrological series. These tests are nonparametric and widely applied in univariate HFA (e.g.,
202 Bobée & Ashkar 1991; Yue et al. 2002; Kundzewicz et al. 2005 and Khaliq et al. 2009).

203 The most important step in HFA is a selection of the appropriate distribution model. Identifying
204 the fitted probability distribution allows predicting the probability of exceedance for a specified
205 magnitude (quantile) or the magnitude associated with a specific exceedance probability.
206 Hydrologic literature suggests several methods to estimate the parameters associated with these
207 distributions, i.e., the maximum likelihood method (e.g., NERC 1975; Clarke 1994); method of
208 moments (e.g., Chebana et al. 2010) and L-moment method (e.g., Hosking 1990). The choice of
209 a subset of appropriate distribution is performed by the goodness-of-fit test of Pearson (Chi-
210 square, χ^2), see e.g. Benkhaled et al. 2014 and Benameur et al. 2017. The selection of the best
211 fit distribution for each variable is based on the criteria mainly Akaike and Bayesian information
212 criterion (AIC, BIC) proposed, respectively by Akaike (1974) and Schwartz (1978). These
213 criteria are considered on the above subset of the accepted distribution. The smallest criterion
214 AIC or BIC values identify the best fit distribution (Rao & Hamed 2000).

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215 In the final step, a multivariate HFA is conducted to identify the relationship and dependence
216 between the different characteristics of the flood through the pairs (D, V), (D, Qp) and (Qp, V).
217 Firstly, multivariate outlier detection is carried out using depth-based tests (e.g. Chebana &
218 Ouarda 2011b; Chebana et al. 2017). Multivariate outliers can have negative impacts on the
219 selection of the appropriate distribution as well as on the estimation of the associated parameters
220 (Chebana 2012). Mainly three depth functions are considered to detect multivariate outliers:
221 Mahalanobis, Spatial and Tukey. Then, checking the assumptions in multivariate HFA is an
222 important step in the analysis. Multivariate versions of Mann–Kendall and Spearman tests are
223 applied to check the multivariate trend on the flood pairs (Chebana et al. 2013). Covariance
224 Inversion Test statistic (CIT), Covariance Sum Test statistic (CST), Covariance Eigenvalue Test
225 statistic (CET) for both Mann-Kendall and Spearman tests are analysed at 5% of the significance
226 level. Cramer test, M-test, Wilcox test and Zhang test are selected to check the homogeneity for
227 multivariate series and multivariate copula-based test for serial independence at a significance
228 level of 5%.

229 In terms of modelling, the dependence structure between two or more random variables is
230 described by using a copula. It is used to join two different marginal distributions in common
231 multivariate distribution. The notion of copula was introduced in Sklar (1959) to decompose a d-
232 dimensional distribution function H into marginal distribution functions F_1, \dots, F_d and a copula C
233 describing the dependence part of the distribution. A copula is a multivariate cumulative
234 distribution function with univariate uniform margins. Given a random vector $X = (X_1, \dots, X_d)$
235 with joint distribution function H and marginal cumulative distribution functions F_1, \dots, F_d . Sklar
236 showed that there always exists a copula C such that, for all $x_1, \dots, x_d \in \mathbb{R}$

$$\Pr (X_1 \leq x_1, \dots, X_d \leq x_d) = H(x_1, \dots, x_d) = C \{F_1(x_1), \dots, F_d(x_d)\} \quad (1)$$

238 If F_1, \dots, F_d are continuous, then C is unique.

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In this work, five copulas commonly used in hydrology are considered for the multivariate flood analysis namely: Ali-Mikhail-Haq, Clayton, Frank, Gumbel, and Joe (Zhang & Singh 2006; Genest & Chebana 2017). These copulas belong to the Archimedean family. In this work, the copula fitting is based on pseudo-observations in order to avoid parametrically modelling the margins, which is conducted in two steps. Adequate marginal distribution is chosen firstly for each of the studied variables. In the second step, the copula parameter is estimated from the selected copula. This parameter is estimated using the maximum pseudo-likelihood method as the most appropriate method (Kim et al. 2007 and Kojadinovic & Yan 2010). Goodness-of-fit test for copula (Sn and Tn tests) are applied to select potential copula, then among the latter, Akaike information criterion (AIC) is applied to select the most appropriate copula (see Genest & Chebana 2017).

Once the joint distribution is found (including the margins, copula and their parameters), it can be used to calculate multivariate return periods and multivariate quantiles. In multivariate FFA, the return period estimation has an important impact on the designing of the hydraulic structure. Therefore, a given flood event return period may have different occurrence combinations of flood variables (Qp, V, D) and vice versa (Chebana & Ouarda 2011a). Such conditions make the choice of the selected return period an ambiguous and difficult step to establish the joint return period (e.g. Corbella & Stretch 2012). This latter represents, for instance, the event in which x or y or both are exceeded ($X > x$ or $Y > y$) (e.g. Vittal et al. 2015). Multivariate return period issues have been defined by more authors (see. e.g. Salvadori et al. 2011; Chebana & Ouarda 2011a; Salvadori et al. 2013; Requena et al. 2013; Gräler et al. 2013).

Multivariate quantile is presented as a quantile curve describing a specific event for a giving return period. In multivariate HFA, four events are characterized by four quantile curves among many other combinations: simultaneous exceedance $\{X \geq x, Y \geq y\}$, exceedance-non-exceedance $\{X \geq x, Y \leq y\}$, non-exceedance-exceedance $\{X \leq x, Y \geq y\}$ and simultaneous non-exceedance $\{X \leq x, Y \leq y\}$.

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3 264 $\leq x, Y \leq y\}$ where the variables X and Y represent the characteristics of a hydrological
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5 265 phenomenon. Therefore, the simultaneous exceedance $\{X \geq x, Y \geq y\}$ and simultaneous non-
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8 266 exceedance $\{X \leq x, Y \leq y\}$ would be of interest in studying the hydrological phenomenon. For
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10 267 droughts investigating and on the quantiles evaluation, the event $\{X \leq x, Y \leq y\}$ is considered.
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12 268 However, the event $\{X \geq x, Y \geq y\}$ is most important in flood consider and for evaluating a return
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15 269 period (Chebana & Ouarda 2011a).

17 270 4. Results

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20 271 In this section, we present the obtained results by applying the methodology in section 3 on data
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23 272 of section 2.

25 273 4.1. Univariate frequency analysis

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28 274 Descriptive analysis is performed on the verified series and the results are shown in Table 2. It is
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31 275 observed that each station expresses a specific character in describing flood event in each region.
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33 276 V and Q_p values are widely different from station to another. However, the mean value of V and
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35 277 Q_p are highest in the East region presented by the stations Ain El Assel and Khemakhem,
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37 278 compared to the Central region (El Affroun and Baraki stations) and the West region (Taria and
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40 279 Sidi Ali Ben Youb) with the smallest values. The observed decrease in V and Q_p values in East-
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42 280 West orientation can be explained by the natural hydrological regime in Northern Algeria, where
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44 281 the rainfall blade is most important in the East region than the West. Lakhdaria station marks the
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46 282 exception having the maximum V value ($340.9 \cdot 10^6 \text{ m}^3$) and maximum Q_p value ($1489.5 \text{ m}^3/\text{s}$)
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49 283 even though it is located in Central region; because of its large drained area (3620 km^2) and very
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51 284 rough terrain. These features lead to explain that flood at Lakhdaria station is important in term
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54 285 of V and Q_p values. While, flood duration is characterized by slight variation, where its
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56 286 maximum values range between 11 and 22 days over the study area. The variability of flood
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58 287 characteristics between the different stations can be explained by the watershed structure, the
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60 288 relief nature, drainage density, soil moisture and the characteristics of the watercourse, or related

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3 289 to the type of flood itself. It is noted that V of the flood follows the same behaviour as D and Q_p
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5 290 at the studied stations. This can provide information on the existence of a possible connection
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8 291 between these three variables that may be identified in the multivariate analysis. According to
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10 292 Meddi et al. (2017) in a univariate regional study, Ghazaouet station fits in region II which is
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12 293 characterized by rugged terrain that favours the formation of rapid and catastrophic floods.
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15 294 Timgad station belongs to the third region where the flash floods are dominated importantly on
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17 295 autumn season. The other stations appear in the region I characterized by rugged terrain that is
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19 296 very favourable to flow and the rapid advance of a flood.

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22 297 The Rosner test identifies the detected values as no outliers for all the stations and all variables
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24 298 (Appendix A). Furthermore, correlation analysis (Table 3) between the three flood characteristics
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26 299 shows a significant correlation between the majorities of the analysed pairs. The T-test indicates
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29 300 a weak correlation between D and Q_p . Pearson coefficient expresses a strong correlation (> 0.78)
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31 301 between Q_p and V at all the studied stations. This correlation means that flood Q_p is positively
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33 302 related to flood V .

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36 303 The results of checking the basic assumptions using the appropriate tests on all the margins and
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38 304 all stations are summarized in Table 4. We observe that all assumptions can be considered as
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41 305 satisfied for all variables and all stations at the significance level 5%. Except for stationarity
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43 306 (Azzefoun), homogeneity (Azzefoun, Khemakhem and Sidi Ali Ben Youb) and independence at
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45 307 Taria can be accepted at level 1%. This decision is reasonable (see e.g. Benameur et al. 2017),
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48 308 especially when the length dataset is relatively short (25 and 26 years in these cases). According
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50 309 to the Chi-square test of goodness-of-fit and the lowest AIC value, the selected marginal
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52 310 distributions are shown in Table 5. These later allowed the determination of return periods for
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54 311 flood characteristics (Figure 2). Most of the quantiles estimated at 50 or 100 years of return
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56 312 period exceed the maximum values of V and Q_p on the analysed series. For instance, the flood of
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58 313 50 years at Baraki station marks a larger V of $143.7 \cdot 10^6 \text{m}^3$ than already observed. While the
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314 flood of 100 years notes Q_p of $930.8 \text{ m}^3/\text{s}$ higher than the maximum value ($876.4 \text{ m}^3/\text{s}$). At Ain
315 El Assel, the quantiles calculated of V and Q_p for a 100 years return period are smaller than
316 observed in the analysed data. A return period of 100 years of a flood on wadi Mekkera records
317 V of $21.2 \cdot 10^6 \text{ m}^3$ and Q_p of $177.6 \text{ m}^3/\text{s}$ exceeding the maximum observed values. At Timgad
318 station, wadi Soultès revealed that a 50 years return period is characterized by Q_p ($91.8 \text{ m}^3/\text{s}$)
319 and V ($10.1 \cdot 10^6 \text{ m}^3$) which are outside the range of observation. Wadi Isser at Lakhdaria station
320 records V exceeding the observed values for a 100 years return period. However, the
321 bicentennial flood notes Q_p more than $1489.4 \text{ m}^3/\text{s}$.

4.2. Multivariate frequency analysis

322 Regarding the outliers detected by the depth functions, are mostly retained for all variables. In
323 another case, some values are considered as outliers even are not identified by the depth
324 function, because these values are accepted according to the natural flow regime in these regions
325 and the stream flow characteristics: For instance, the detected outliers are not retained in their
326 totality because their consideration may affect the chronology series length, example with (V ,
327 Q_p) series for El Affroun station (Figure 2, Appendix C) where five years are not considered
328 (1972, 1974, 1979, 1980, 1983). Sometimes, the identified outliers are replaced by other ones
329 which allowed the best verifying for the assumption tests, example with (D , V) series for
330 Azzefoun station where the values corresponding to the years of 2000, 2003 and 2010 are
331 replaced by those of 1988 and 2006; or because it seems due to an error in curded data. The other
332 stations give similar results (see Figures 1 to 10 in Appendix C). In other cases, the defined
333 outliers are not retained in order to keep a long chronological series (Timgad, Ghazaouet, and
334 Baraki).

335 Further, the multivariate trend tests of the studied pairs show that all tests accept the no trend
336 assumption for all the stations at level 5%. Except (D , V) and (D , Q_p) series for Azzefoun station
337 where CET and CIT (Mann-Kendall) and CET (Spearman) for (D , Q_p) and (Q_p , V), and for
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3 339 Taria station where CIT (Mann-Kendall) accept at significance level of 1% (Table 6). For the
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5 340 multivariate serial independence test, the obtained results (Table 7) indicate that all pairs for all
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8 341 station are independent. Also, the multivariate homogeneity tests express that all studied series
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10 342 for all the stations are homogeneous (Table 7). Except, El Affroun (D, V) and Sidi Ali Ben Youb
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12 343 (D, Qp) where M-test is accepted at the significance level of 1%, and C-test at Sidi Ali Ben
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15 344 Youb for the pair (D, V). These satisfied finding made it possible to run the final step of the
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17 345 multivariate analysis: copula fitting.

20 346 ***Copula Fitting and Parameter Estimation***

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22 347 The selected as best fit copulas (corresponding to the p-value of the Sn test and Tn test and the
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25 348 smallest AIC value) for flood characteristics pairs are indicated in Table 8. Frank is the
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27 349 predominant copula (11 over 33 cases) followed by Clayton (9 over 33) and Gumbel (8 over 33).
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29 350 Each pair of flood characteristics is described by the appropriate copula which leads to estimate
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32 351 the corresponding quantiles. The occurrence of a flood event is obtained from the contour line of
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34 352 pairs according to the selected copula. Figure 4 gives a graphic presentation of quantiles curves
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36 353 corresponding to the event $X \geq x \cap Y \geq y$. Exceeding probability is applied to calculate the
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39 354 corresponding return period (i.e. 2, 5, 10, 20 and 50 years for all the studied station). The
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41 355 different shapes of the quantile curves from pairs and stations are mainly because of the different
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43 356 copulas and strength of the dependence For instance, the curves lines of (D, Qp) at the stations of
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46 357 Timgad and Taria are likely to be linear since the dependence is very low. The curves lines at
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48 358 Ammi Moussa station (Figures 4-a1, 4-a2, 4-a3) of (D, V) and (D, Qp) pairs are condensed for
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50 359 return period superior than 2 years for flood that last between 10 to 16 days, while the
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52 360 corresponded V and Qp have spaced curves which indicates that a single duration flood could
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55 361 have numerous joint V or Qp value. In studying flood structure dependence, it is important to
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57 362 care an interest considering to (D, V) and (D, Qp) pairs where it is most significant to identify a
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59 363 flood which happens with a short D and the highest Qp or highest V because it does not have the
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364 same impact, especially on hydraulic structures (e.g. dams, bridges). Therefore, (Q_p, V) pair can
365 describe well a flood at stations presented a similar structure.

366 Quantile curves of the pairs (D, V) and (D, Q_p) present the same shape of contour lines at the
367 same station for all the analysed stations. For instance, the curves lines of the flood return period
368 of 5 to 50 years at Khemakhem station show that flood duration ranges between 11 to 18 days.
369 The experimental data are distributed between the contour line of the selected return periods and
370 mostly condensed for the period return inferior to five years. The figures (4-k1, 4-k2) on Wadi El
371 Harrach show that the experimental data present flood with a duration from 10 to 18 days. For
372 the figures (4-g1, 4-g2) for wadi Isser, the maximum V and Q_p are recorded for the medium
373 duration (8 to 9 days) for the short return period (2 and 5 years) against the long duration (>17
374 days) for 50 years return period. These observations justify the interest to give in studying flood
375 at Lakhdaria station based on a combination between D and V because the rough relief favours a
376 large V over a small period of time. The curves lines at El Affroun station (figures 4-c1, 4-c2)
377 indicate that the flood with a duration between 14 to 16 generates a large V and important Q_p ,
378 which is recorded at the short return period (2 to 5 years).

379 Dependence structure designed for the given return period (2 to 50 years) by the quantiles curves
380 of the pair (Q_p, V) on the analysed flood show similarity in the joint structure. The contour lines
381 have a shape of a right angle since the correlation is greater than 0.9, which present the linear
382 form of the dependence structure of the analysed data. For instance, figure 4-b3 indicates that the
383 analysed data are condensed in the short return period (inferior de 5 years). However, the
384 maximum value of V and Q_p have a joint return period of 20 years, which is really important in
385 designing the hydraulic development in this basin. Ain El Assel (figure 4-e3) shows that most
386 flood events have a return period between two to five years. This indicates the interest to give in
387 planning water management at a short return period in this watershed. The figure 4-g3 at
388 Lakhdaria notes that the analysed floods are mostly dominant for the short return period.

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The similar joint structure of (Q_p, V) is identified on the other stations, which indicate the necessity to focus on the short and medium return period and the behaviour of this couple, in identifying the values of V and Q_p in water structure designing and a flood mitigation programs. According to the analysed pairs, these stations show a slight difference in behaviour from the Eastern to the Western regions in responding to flood event. These spatial trends may be explained by the basin structure and nature, or the length of the corresponding wadi, or in the human activities in using the soil surface.

5. Discussion

The main purpose of this study is to give an appropriate method to analyse flood in the regions affected by repeated flood in northern Algeria. For that, complete multivariate frequency analysis on the flood characteristics (all steps not only modelling) is applied. Eleven basins are selected according to the availability of data (time series length and continued observation). It is noticed that these watersheds present different physiographic and climate characteristics from the East (Eastern Constantine Coastal basin) to the West (Oran Coastal basin); which give an appropriate description of the hydrology variability in northern Algeria. On the other hand, in Algeria FFA has been always carried as univariate by using a single variable (usually Q_p); while in this study, three variables are extracted (Q_p, V and D). We use these 11 stations because they present continuous series without gaps compared to other stations; which is relatively the longest among those made available to us by the services of the ANRH. For future work, thinking on extrapolating this method to the ungauged sites should be considered where no or less hydrological information is available. Therefore, this finding will be used on the hydraulic structures design and the protection of towns and villages as well as the agricultural lands against floods. Hence, this first step allows us to conclude on the reliability of such an approach in our regions.

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The daily flow series allows the extraction of three main variables to characterize the flood event risk. This latter is described by the joint dependence structure defined by the associated pairs of flood variables. The dependence structure modelling is performed by using a copula approach which provides important information about flood risk assessment. In Algeria, flood risk issues have been treated in the univariate approach by identifying the best (univariate) distribution which allows estimating quantile for an associated return period. In the presented study, multivariate FFA is conducted in Northern Algeria where 11 watersheds are analysed. A complete multivariate FFA is carried on the extracted variables (D , V , Q_p). The correlation of flood variables shows the link that an important flow can generate a large volume and vice versa. In the same sense, Yue et al. (1999); Kim et al. (2003) and Ben Aissia et al. (2009) revealed a strong dependence between Q_p and V . Most of the identified change point dates belong on the 1970's decade for three mains flood variables (Table B.2). The same period was detected by Bakreti et al. (2013) and Halouz et al. (2013) in analysing the change in the hydrological regime on Western Algeria. The detected change points are not retained because they are located on the series boundaries which affect strongly the ability of the applied test (see e.g. Xie et al. 2014; Xiong et al. 2015; Nayak et al. 2016). Indeed, as a first step, for univariate quantile estimation, Weibull is mostly selected for the duration series (7 out of 11), also for peak flow series (6 out of 11). For volume series (6 out of 11) and Log-Normal type 2 (4 out of 11) are the selected marginal distributions. This manuscript presents a study of at-site FFA where the selected multivariate models for the studied stations are not transferable to other stations (even though it is the same methodology). Hence, there is no need to validate on other stations. In addition, since the aim of FFA is not the forecasting, it is not appropriate to validate on sub-samples (e.g. Salvadori & De Michele, 2004; Zhang & Singh, 2006; Reddy & Ganguli, 2012; Requena et al. 2013; Callau Poduje et al. 2014). The depth functions are used in identifying multivariate outliers. Multivariate tests are used to check the basic assumptions of trend, homogeneity and

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3 438 serial independence. Then, the modelling step is performed based on copulas to model the
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5 439 dependence structure. Based on selection criteria (AIC and BIC), as well as on goodness-of-tests
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8 440 and graphical tools, mostly three copulas (Frank, Clayton and Gumbel) are retained for the flood
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10 441 variables to describe the joint structure of the analysed pairs. A total of 33 pairs are analysed
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12 442 (three pairs for each of the 11 stations), where Clayton (4 out of 11) and Gumbel (4 out of 11)
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14 443 are selected for (D, V); Frank (5 out of 11) and Clayton (3 out of 11) are retained for (Qp, V) and
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17 444 Gumbel (4 out of 11) and Frank (4 out of 11) describe the couple (D, Qp). It is noted that a given
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19 445 return period for a flood event may have various occurrence of flood variables and vice versa.
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21 446 Such findings are valuable tools for water surface management and engineers for hydraulic
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23 447 structures design (dams, sewage systems, bridges, and watercourses developments). Regarding
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26 448 multivariate HFA studies in Mediterranean region, e.g., Leite Rosa (2011) and Hamdi et al.
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28 449 (2016), we notice the similarity in the behaviour of the hydrological variables to assess the
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30 450 extreme event; where Frank and Clayton were selected to model the dependence of the studied
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33 451 variables.

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36 452 The results presented in this manuscript of a multivariate analysis of flood variables (D, V, Qp)
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38 453 offer an important opportunity to develop a new strategy of water management (water supply,
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40 454 irrigation), sewage systems, flood control and reservoir operation, where the several
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42 455 combinations of flood features should be involved to enhance the efficiency and sustainability of
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44 456 water structure. A better knowledge of the hydrological behaviour of flood variables
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46 457 combinations (D, Qp), (D, V), (V, Qp) in multivariate approach is a promising step in Algeria, as
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48 458 it is already in other parts of the world. However, the water demand is increasing especially in its
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50 459 northern part where the majority of the population is located. Therefore, managing the flood to
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52 460 reuse the water in irrigation, or recharge the groundwater may help to minimize the damage
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54 461 caused by flood events and contribute to providing an important quantity of water when the
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56 462 water supply-demand is rising (e.g. summer season).
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3 463 For instance, the return period of 100 years is required in watercourses facilities. Through the
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5 464 combination of two or more hydrological variables, mainly Q_p and V , provide more reliability
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8 465 for hydraulic structure designing, water reservoir management, flood risk assessment and support
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10 466 flood mitigation in the studied basins (e.g. Salvadori & De Michele, 2004; Zhang & Singh,
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12 467 2006; Reddy & Ganguli, 2012; Requena et al. 2013; Callau Poduje et al. 2014; Balistrocchi et al.
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14 468 2017; Jiang et al. 2019; Zhou et al. 2019; Liu et al. 2019).

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18 469 This work presents a case of study of at-site FFA and not regional analysis. It is important to
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20 470 mention that a regional FFA is not a simple collection of a number of at-site FFA cases (see e.g.
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22 471 Ouarda 2017 for more details). In addition, to the best of our knowledge, there are very limited
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24 472 studies simultaneously on multivariate and regional FFA (see Requena et al. 2016). To do a
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27 473 regional or spatial study, it goes beyond the scope of the present project since it requires first an
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29 474 entire project dealing only with data preparation and extraction. In addition, in the case of the
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32 475 Algerian context, the situation is more complex since the data when they are available they do
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34 476 not fulfil the FFA assumptions. Consequently, we need to consider advanced approaches to deal
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36 477 with missing multivariate data, non-standard multivariate models (e.g. multivariate no-
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38 478 stationary). These elements may be the subject of more than one additional manuscript, not to
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41 479 mention the fact that some approaches are not even available in the literature.

44 480 **6. Conclusions and perspectives**

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47 481 The work presented in this paper examines multivariate frequency analysis of flood features in
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49 482 Northern Algeria, in order to add to the current knowledge of hydrological systems and flood
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52 483 management in Algeria. The multivariate HFA framework is conducted on flood event for
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54 484 eleven stations (in seven large watersheds) chosen from the extreme East to the extreme West of
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56 485 the country. This data and this analysis are considered for the first time in Algeria. This can
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59 486 provide an interesting illustration for surface water management and identifying the behaviour of
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3 487 flood risk in the watershed. The methodology considered is complete since it includes all steps of
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5 488 the multivariate HFA. Copula concept is applied to identify the structure dependence of the main
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8 489 flood features (D, V, and Qp). As in other studies, a significant correlation between Qp and V is
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10 490 identified. This result shows the necessity to establish the dependence structure between these
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12 491 variables and to conduct multivariate FFA based on copulas. Copulas selection, among three
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14 492 variable pairs and five stations, shows that the Frank copula is most often selected followed by
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17 493 the Clayton and Gumbel copula. The results obtained can be used by water managers to obtain
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19 494 information for given variables of flood in the studied regions. Through these results, it is
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21 495 demonstrated that applying copulas to analyse multivariate correlated flood characteristics can
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24 496 lead to a joint structure which describes the flood event well. The multivariate FFA study,
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26 497 including the dependence structure between the flood characteristics, is considered as efficient
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28 498 tools in managing water supply in the studied watersheds. These findings are an interesting tool
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31 499 to understand the dependence between the multivariate flood variables; to establish more
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33 500 efficient planning strategies against flood risk; providing support for the decision in water
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35 501 management, designing the hydraulic structure, flood control and mitigation and optimize the
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38 502 water reservoir operation according to the local conditions of each basin. Indeed, it provides
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40 503 flexibility for engineers, useful to fill the gaps in the data and the results can be applied for future
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42 504 regional studies. This kind of analysis is often neglected in FFA in Algeria, where usually only
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44 505 one variable is treated (Qp). On the other hand, these results may be involved in the design of
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47 506 hydraulic structures projects such as dams, bridges, and watercourse developments and
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49 507 protection of cities subject to flood problems (e.g. Karmakar & Simonovic 2007; Bačová
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51 508 Mitková & Halmová 2014; Balistrocchi et al. 2017). This study can be used to motivate the
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54 509 authorities to continue recording data in the existing stations so that they can be used in the
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56 510 future to provide better risk assessment. On the other hand, FFA, especially in multivariate
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58 511 approach, requires a specific type and amount of data, and conditions of stationarity,
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3 512 homogeneity and serial dependence compared with those used in another hydrological study e.g.
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5 513 prevision, rainfall-runoff modelization. Behind the specific results obtained for the considered
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8 514 case studies, the paper has presented a full methodological framework for a complete
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10 515 multivariate frequency analysis which includes all steps of the analysis, i.e. extracting flood
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12 516 features from daily stream flow, descriptive analysis, hypotheses testing, model selection and
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15 517 estimation, risk assessment. This framework is transferable to other locations, flood variables or
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17 518 hydrological events (e.g. droughts, precipitations). Provided the availability of appropriate data
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19 519 sets, it is well recommended to extend this analysis to cover the entire country by adding more
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22 520 dataset and involving more variables. Such a national study would require an additional step of
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24 521 classifying stations in homogenous regions in order to provide a model for each group of
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26 522 stations. The national study also requires a huge work regarding data preparation and databases
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29 523 construction. The present study can be seen as a first and necessary step toward a national study.

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45 46 47 530 **References**

- 48
49
50 531 - Agence de bassin hydrographique Constantinois-Seybous-Mellegue, Constantinois,
51
52 532 ABHCSM (1999-2000), Les Cahiers de l'Agence. Les Côtiers constantinois, N°4, 75 pages.
53
54 533 - Agence de Bassin Hydrographique Oranie-Chott Chergui (ABH-OCC) (2006) Mission 1,
55
56
57 534 Inventaire des Ressources en Eau et en Sols et des Infrastructures de Mobilisation, Cadastre
58
59 535 Hydraulique du Bassin des Hauts Plateaux Oranais
60

- 1
2
3 536 - Agence Nationale d'Intermédiation et de Régulation Foncière. (2015) Consulted on
4
5 537 November 1st 2017. <http://www.aniref.dz/index.php/fr/pourquoi-l-algerie/les-zones->
6
7
8 538 industrielles.
- 9
10 539 - Aissia, M. A. B., Chebana, F., & Ouarda, T. B. (2017). Multivariate missing data in
11
12 540 hydrology—Review and applications. *Advances in water resources*, 110, 299-309.
- 13
14
15 541 - Akaike, H. (1974) A new look at the statistical model identification. *IEEE Trans. Automatic*
16
17 542 *Control* 19(6):716–723
- 18
19 543 - Bobée, B & Ashkar, F. (1991) The gamma family and derived distributions applied in
20
21 544 hydrology. Water Ressources Publications.
- 22
23
24 545 - Bačová Mitková, V.B. & Halmová, D. (2014) Joint modeling of flood peak discharges,
25
26 546 volume and duration: a case study of the Danube River in Bratislava. *J. Hydrol. Hydromech.*,
27
28 547 62, 3, 186–196. DOI:10.2478/johh-2014-0026.
- 29
30
31 548 - Bakreti, A., Braud, I., Leblois, E., Benali, A. (2013) Analyse conjointe des régimes
32
33 549 pluviométriques et hydrologiques dans le bassin de la Tafna (Algérie Occidentale).
34
35 550 *Hydrological Sciences Journal*, 58:1, 133-151, DOI: 10.1080/02626667.2012.745080.
- 36
37
38 551 - Balistrocchi, M., Orlandini, S., Ranzi, R., & Bacchi, B. (2017). Copula-based modeling of
39
40 552 flood control reservoirs. *Water Resources Research*, 53. [https://](https://doi.org/10.1002/2017WR021345)
41
42 553 doi.org/10.1002/2017WR021345
- 43
44
45 554 - Ben Aissia, M.A., Chebana, F., Ouarda, T. B. M. J., Roy, L., Desrochers, G., Chartier, I., &
46
47 555 Robichaud, É. (2011) Multivariate analysis of flood characteristics in a climate change
48
49 556 context of the watershed of the Baskatong reservoir, Province of Québec, Canada.
50
51 557 *Hydrological Processes*, 26(1), 130–142. DOI:10.1002/hyp.8117
- 52
53
54 558 - Ben Aissia, M.A., Chebana, F., Ouarda, T.B.M.J., Bruneau, P., Barbet, M. (2015) Bivariate
55
56 559 index-flood model: case study in Québec, Canada. *Hydrological Sciences Journal*, 60, 2, 247–
57
58 560 268. DOI:10.1080/02626667.2013.875177.
- 59
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41
42
43
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47
48
49
50
51
52
53
54
55
56
57
58
59
60

- 561 - Benameur, S., Benkhaled, A., Meraghni, D., Chebana, F., Necir, A. (2017) Complete flood
562 frequency analysis in Abiod watershed, Biskra (Algeria). *Nat Hazards* (2017) 86:519–534
563 DOI 10.1007/s11069-016-2703-4.
- 564 - Bezak, N., Mikoš, M., Šraj, M. (2014) Trivariate Frequency Analyses of Peak Discharge,
565 Hydrograph Volume and Suspended Sediment Concentration Data Using Copulas, *Water*
566 *Resources Management*, 28 (8), pp. 2195-2212.
- 567 - Bouanani, A. (2004) Hydrologie, transport solide et modélisation, étude de quelques sous-
568 bassins de la Tafna (NO-Algérie), Tlemcen, Algérie: Thèse de doctorat d'état, Université
569 Abou Bekr Belkaid.
- 570 - Callau Poduje, A. C., Belli, A., & Haberlandt, U. (2014) Dam risk assessment based on
571 univariate versus bivariate statistical approaches: a case study for Argentina. *Hydrological*
572 *Sciences Journal*, 59(12), 2216–2232. DOI:10.1080/02626667.2013.871014
- 573 - Chebana, F., & T. B. M. J. Ouarda. (2007) Multivariate L-moment homogeneity test, *Water*
574 *Resour. Res.*, 43, W08406, DOI:10.1029/2006WR005639.
- 575 - Chebana, F., & T. B. M. J. Ouarda. (2009) Index flood-based multivariate regional frequency
576 analysis, *Water Resour. Res.*, 45, W10435, DOI:10.1029/2008WR007490
- 577 - Chebana, F., El Adlouni, S., Bobée, B. (2010) Mixed estimation methods for Halphen
578 distributions with applications in extreme hydrologic events *Stochastic Environment*
579 *Resources Risk Assessment* 24: 359–376. DOI: 10.1007/s00477-009-0325-z
- 580 - Chebana, F., & T.B.M.J. Ouarda. (2011a) Multivariate quantiles in hydrological frequency
581 analysis, *Environmetrics*, 22 (1), 63-78.
- 582 - Chebana, F., Ouarda, T.B.M.J. (2011b) Depth-based multivariate descriptive statistics with
583 hydrological applications. *J. Geophys. Res.* 116 (D10), D10120. HTTP://
584 dx.doi.org/10.1029/2010JD015338.

- 1
2
3 585 - Chebana, F. (2012) Multivariate analysis of hydrological variables in Encyclopedia of
4
5 586 Environmetrics Second Edition, A.-H. El-Shaarawi and W. Piegorisch (eds), John Wiley &
6
7 Sons Ltd, Chichester, UK, pp. 1676-1681, DOI: 10.1002/9780470057339.vnn044.
8 587
9
10 588 - Chebana, F., Ouarda, T.B.M.J & Duong, T.C. (2013) Testing for multivariate trends in
11
12 589 hydrologic frequency analysis. Journal of Hydrology 486; 519–530.
13
14 <http://dx.doi.org/10.1016/j.jhydrol.2013.01.007>
15 590
16
17 591 - Chebana, F., Ben Aissia, M. A., Ouarda, T.B.M.J. (2017) Multivariate shift testing for
18
19 592 hydrological variables, review, comparison and application. Journal of Hydrology 548; 88–
20
21 593 103. <http://dx.doi.org/10.1016/j.jhydrol.2017.02.033>
22
23
24 594 - Chow, V.T. Maidment, D.R. and Mays, L.W. (1988) Applied Hydrology, McGraw-Hill, New
25
26 595 York, NY.
27
28 596 - Chowdhary, Hemant. (2009) Copula-Based Multivariate Hydrologic Frequency Analysis.
29
30 *LSU Doctoral Dissertations*. 1211.
31 597
32
33 598 - Clarke, RT. (1994) Fitting Distributions. Chapter 4 Statistical modeling in hydrology. John
34
35 599 Wiley & Sons Ltd: Chichester; 39–85.
36
37
38 600 - Corbella, S & Stretch, D.D. (2012) Multivariate return periods of sea storms for erosion risk
39
40 601 assessment. Nat. Hazards Earth Syst. Sci., 12, 2699–2708, 2012 [www.nat-hazards-earth-syst-](http://www.nat-hazards-earth-syst-sci.net/12/2699/2012/)
41
42 602 [sci.net/12/2699/2012/](http://www.nat-hazards-earth-syst-sci.net/12/2699/2012/). DOI:10.5194/nhess-12-2699-2012.
43
44
45 603 - Cunnane, C. (1987) Review of statistical models for flood frequency estimation. In
46
47 604 Hydrologic Frequency Modeling, Singh V.P. (ed.). Reidel, Dordrecht: The Netherlands; 49-
48
49 605 95.
50
51 606 - De Michele, C., & Salvadori, G. (2003) A Generalized Pareto intensity-duration model of
52
53 storm rainfall exploiting 2-copulas. Journal of Geophysical Research, 108(D2), 4067,
54 607
55 DOI:10.1029/2002JD002534.
56 608
57
58
59
60

- 1
2
3 609 - El Adlouni, S., & T. B.M.J. Ouarda. (2008) Étude de la loi conjointe Débit-Niveau par les
4
5 610 copules: Cas de la rivière Châteauguay. *Canadian Journal of Civil Engineering*, 35: 1128-
6
7 611 1137, DOI : 10.1139/L08-054.
8
9
10 612 - Gilbert, Richard O. (1987) *Statistical Methods for Environmental Pollution Monitoring*. VAN
11
12 613 NOSTRAND REINHOLD COMPANY, *New York*, pp. 186-188.
13
14
15 614 - Gräler, B., van den Berg, M., Vandenberghe, S., Petroselli, A., Grimaldi, S., De Baets, B., and
16
17 615 Verhoest, N. (2013) Multivariate return periods in hydrology: a critical and practical review
18
19 616 focusing on synthetic design hydrograph estimation, *Hydrology and Earth System Sciences*,
20
21 617 17, 1281–1296, DOI:10.5194/hess-17-1281-2013.
22
23
24 618 - Grimaldi, S. & Serinaldi, F. (2006a) Asymmetric copula in multivariate flood frequency
25
26 619 analysis. *Advances in Water Resources* 29: 1155–1167.
27
28
29 620 - Grimaldi, S. & Serinaldi, F. (2006b) Design Hyetograph analysis with 3-copula function.
30
31 621 *Hydrological sciences journal*, 51:2, 223-238, DOI:10.1623/hysj.51.2.223.
32
33 622 - Grimaldi, S., Petroselli, A., Salvadori, G., & De Michele, C. (2016). Catchment compatibility
34
35 623 via copulas: A non-parametric study of the dependence structures of hydrological responses.
36
37 624 *Advances in Water Resources*, 90, 116–133.doi:10.1016/j.advwatres.2016.02.003
38
39
40 625 - Hallouz, F., Meddi, M., Mahe, G. (2013) Modification du régime hydroclimatique dans le
41
42 626 bassin de l'Oued Mina (nord-ouest d'Algérie) *Revue des sciences de l'eau / Journal of Water*
43
44 627 *Science*, vol. 26, n° 1, 2013, p. 33-38, DOI: 10.7202/1014917ar.
45
46
47 628 - Hamdi, Y., Chebana, F., & Ouarda, T. B. M. J. (2016) Bivariate drought frequency analysis in
48
49 629 the Medjerda River Basin, Tunisia. *J. Civil Environ. Eng.*, 6 (3): Art. 1000227.
50
51
52 630 - Hosking, JRM. (1990) L-moments: analysis and estimation of distributions using linear
53
54 631 combinations of order statistics. *Journal of the Royal Statistical Society* 52(2): 105–124.
55
56
57
58
59
60

- 1
2
3 632 - Janga Reddy, M. & Poulomi Ganguli. (2012) Bivariate Flood Frequency Analysis of Upper
4
5 633 Godavari River Flows Using Archimedean Copulas. *Water Resour Manage* (2012) 26:3995–
6
7 634 4018. DOI 10.1007/s11269-012-0124-z.
- 9
10 635 - Jiang, X., Yang, L., & Tatano, H. (2019) Assessing Spatial Flood Risk from Multiple Flood
11
12 636 Sources in a Small River Basin: A Method Based on Multivariate Design Rainfall. *Water*,
13
14 637 *11(5)*, 1031. DOI:10.3390/w11051031.
- 16
17 638 - Kao, S.C., & Govindaraju, R. S. (2007) A bivariate frequency analysis of extreme rainfall
18
19 639 with implications for design. *J.Geophys. Res.*, 112, D13119, DOI: 10.1029/2007JD008522.
- 21
22 640 - Karmakar, S., Simonovic, S.P. (2007) Flood frequency analysis using copula with the mixed
23
24 641 marginal distribution. *Water Resources Research Report n°055*. University of Western
25
26 642 Ontario. Department of Civil and Environmental Engineering.
- 28
29 643 - Kendall M. G. (1975) *Rank Correlation Methods*, 4th ed. Charles Griffin, London.
- 30
31 644 - Khaliq, M.N., Ouarda, T.B.M.J., Gachon, P., Sushama, L. and St-Hilaire, A. (2009)
32
33 645 Identification of hydrological trends in the presence of serial and cross correlations: A review
34
35 646 of selected methods and their application to annual flow regimes of Canadian rivers. *Journal*
36
37 647 *of Hydrology*, 368, 117-130.
- 39
40 648 - Kim, G., Silvapulle, M.J., & Silvapulle, P. (2007) Comparison of semiparametric and
41
42 649 parametric methods for estimating copulas, *Computational Statistics & Data Analysis* 51(6),
43
44 650 2836–2850.
- 46
47 651 - Kojadinovic, I. & Yan, J. (2010) Comparison of three semi parametric methods for estimating
48
49 652 dependence parameters in copula models, *Insurance: Mathematics and Economics* 47(1), 52–
50
51 653 63.
- 53
54 654 - Kundzewicz, Z.W., Graczyk, D., Maurer, T., Pinskiwar, I., Radziejewsky, M., Svensson, C.,
55
56 655 and Szwed, M. (2005) Trend detection in river flow series: 1. Annual maximum flow.
57
58 656 *Hydrol.Sci. J.* 50, 797-810.
- 59
60

- 1
2
3 657 - Leonard, M., Metcalfe, A., Lambert M. (2008) Frequency analysis of rainfall and streamflow
4
5 658 extremes accounting for seasonal and climatic partitions. *Journal of Hydrology*, Volume 348,
6
7
8 659 Issues 1-2, 1 January 2008, Pages 135-147.
9
10 660 - Leite Rosa., F (2011) Statistical modeling of flood discharges and volumes in Continental
11
12 661 Portugal: convencional and bivariate analyses. Extended Abstract. Dissertation for obtaining
13
14 662 the degree of master in Civil Engineering. Universidade Técnica de Lisboa.
15
16
17 663 - Liu, Y.R., Li, Y.P., Ma, Y., Jia, Q.M., Su, Y.Y. (2019) Development of a Bayesian-copula-
18
19 664 based frequency analysis method for hydrological risk assessment - the Naryn River in
20
21 665 Central Asia *Journal of Hydrology*, DOI: <https://doi.org/10.1016/j.jhydrol.2019.124349>
22
23
24 666 - Mann, H. B. (1945) Non-Parametric tests against trend, *Econometrica* 13:245-259.
25
26 667 - Mebarki, A. (1999) Approche hydrologique des bassins du Nord-Est Algérien, Actes des
27
28 668 journées d'information et d'étude (Constantine, 20-21 Octobre), la nouvelle politique de
29
30 669 l'Eau, les Agences de bassins hydrographiques et le bassin Constantinois Seybousse-
31
32 670 Méllègue, ABH-CSM., Constantine, pp. 22-30.
33
34
35 671 - Meddi, M., Toumi, S. & Assani, A.A. (2017) Application of the L-moments approach to the
36
37 672 analysis of regional flood frequency in Northern Algeria', *Int. J. Hydrology Science and*
38
39 673 *Technology*, Vol. 7, No. 1, pp.77–102.
40
41
42 674 - NERC. (1975) Flood Studies Report (in five volumes). Natural Environment Research
43
44 675 Council: London.
45
46
47 676 - Nayak, M. A., & Villarini, G. (2016). Evaluation of the capability of the Lombard test in
48
49 677 detecting abrupt changes in variance. *Journal of Hydrology*, 534, 451–
50
51 678 465.doi:10.1016/j.jhydrol.2016.01.016
52
53
54 679 - Rao, A.R., Hamed, K.H. (2000) Flood Frequency Analysis. CRC Press, Boca Raton.
55
56 680 - Renard, B. and Lang, A. (2006) Use of a Gaussian copula for multivariate extreme value
57
58 681 analysis: Some case studies in hydrology. *Advances in Water Resources* 30: 897–912.
59
60

- 1
2
3 682 - Requena, A.I., Mediero, L., & Garrote, L. (2013) A bivariate return period based on copulas
4
5 683 for hydrologic dam design: accounting for reservoir routing in risk estimation. *Hydrology*
6
7
8 684 *And Earth System Sciences*.**17**:p.3023-3038.
- 9
10 685 - Salarpour, M., Yusop, Z., Yusof, F., Shahid, S., Jajarmizadeh, M. (2013) Flood frequency
11
12 686 analysis based on t-copula for Johar River, Malaysia. *Journal of Applied Sciences* 13 (7):
13
14
15 687 1021-1028.
- 16
17 688 - Salvadori, G., & De Michele, C. (2004) Frequency analysis via copulas: theoretical aspects
18
19 689 and applications to hydrological events. *Water Resources Research*, 40:W12511,
20
21 DOI:10.1029/2004wr003133.
22 690
- 23
24 691 - Salvadori, G., De Michele, C. & Durante, F. (2011) On the return period and design in a
25
26 692 multivariate framework, *Hydrol. Earth Syst. Sci.*, 15, 3293–3305, DOI:10.5194/hess-15–
27
28 693 3293-2011.
- 29
30
31 694 - Salvadori, G., Durante, F. & De Michele, C. (2013) Multivariate return period calculation via
32
33 695 survival functions, *Water Resour. Res.*, 49, 2308–2311, doi:10.1002/wrcr.20204.
- 34
35 696 - Serinaldi, F. & C. G. Kilsby. (2013) The intrinsic dependence structure of peak, volume,
36
37 697 duration, and average intensity of hyetographs and hydrographs, *Water Resour. Res.*, 49,
38
39 698 3423–3442, DOI:10.1002/wrcr.20221.
- 40
41
42 699 - Singh, J., Vittal, H., Singh, T., Karmakar, S., Ghosh, S. (2015) A Framework for Investigating
43
44 700 the Diagnostic Trend in Stationary and Nonstationary Flood Frequency Analyses Under
45
46 701 Changing Climate. *Journal of Climate Change*, Vol. 1, No. 1-2, pp. 47–65.
- 47
48
49 702 - Touazi, M., Bhiry, N., Laborde, J.P. & Achour, F. (2011) Régionalisation des débits moyens
50
51 703 mensuels en Algérie du nord. *Revue des sciences de l'eau / Journal of Water Science*, 24 (2),
52
53 704 177–191. <https://doi.org/10.7202/1006110ar>
- 54
55
56 705 - Vittal, H., Singh, J., Kumar, P., Karmakar, S. (2015) A framework for multivariate data-based
57
58 706 at-site flood frequency analysis: Essentiality of the conjugal application of parametric and
59
60

1
2
3
4
5
6
7
8
9
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43
44
45
46
47
48
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51
52
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55
56
57
58
59
60

nonparametric approaches, *Journal of Hydrology* 525; 658–675.
<http://dx.doi.org/10.1016/j.jhydrol.2015.04.024>.

- Wald, A. J. Wolfowitz. (1943) An Exact Test for Randomness in the Nonparametric Case based on Serial Correlation, *Ann. Math. Statist.*, Vol. 14, pp. 378-388.

- Xie, H., D. Li, & L. Xiong (2014) Exploring the ability of the Pettitt method for detecting change point by Monte Carlo simulation, *Stochastic Environ. Res. Risk Assess.*, 28(7), 1643–1655, DOI:10.1007/s00477-013-0814-y.

- Xiong, L., C. Jiang, C.-Y. Xu, K.-x. Yu, & S. Guo (2015). A framework of change point detection for multivariate hydrological series, *Water Resour. Res.*, 51, 8198–8217, DOI:10.1002/2015WR017677.

- Yamani, K., Hazzab, A., Sekkoum, M., Slimane, T. (2016) Mapping of vulnerability of flooded area in arid region. Case study: area of Ghardaia-Algeria. *Model. Earth Syst. Environ*, 2:147.

- Yang, W., Xiao, C., Liang, X., & Zhang, Z. (2019) Two baseflow separation methods based on daily average gage height and discharge. *Water Supply*. DOI:10.2166/ws.2019.074

- Yue, S., Ouarda, T. B. M. J., Bobée, B., Legendre, P. and Bruneau, P. (2002) Approach for describing statistical properties of flood hydrograph, *Journal of Hydrologic Engineering*, 7, 147.

- Zhang, L. (2005) Multivariate hydrological frequency analysis and risk mapping. *LSU Doctoral Dissertations*. 1351.

- Zhang, L., & Singh, V.P. (2006) Bivariate flood frequency analysis using copula method. *J Hydrol Eng* 11(2):150–164.

- Zhou, T., Liu, Z., Jin, J., & Hu, H. (2019). Assessing the Impacts of Univariate and Bivariate Flood Frequency Approaches to Flood Risk Accounting for Reservoir Operation. *Water*, 11(3): 475. DOI:10.3390/w11030475

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2
3 732
4
5
6 733
7
8
9 734
10
11
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13 735
14
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For Peer Review

746 **Table 1.** Historical review of riverine flood in Algeria

Event date		Locations	Provinces	Wadis	Impacts
Start date	End date				
06/10/1966	06/10/1966	West	Chlef	Cheliff	57 deaths
November 1967	November 1967	Northeast and Aurès	Batna, Timgad, Constantine	El Grine, Berghout, Soultès, Hammam Grouz, Rhummel	20 deaths
September 1969	October 1969	North centre	Algiers, Medea, Blida, Boumerdès	Djer, Mazafran, El Harrach, Sébaou, Isser, Corso	76 deaths
March 1973	March 1973	Northeast and Northwest	Tlemcen, Annaba, Taref	Seybousse, Saf-Saf, Kebir, Ressoul, Ksob, Sikkak, Isser	21 deaths
28/03/1974	31/03/1974	Central and Western	Algiers, Blida, Boumerdès, Tizi Ouzou, Tlemcen	El Harrach, Djer, Mazafran, M'leta, Sébaou, Sikkak, Isser	52 deaths and 4,570 destroyed houses
November 1984	November 1984	Western	Chlef, Relizane, Sidi Bel Abbès, Tlemcen	Cheliff, R'hiou, Mekkerà, Tafna	54 deaths
09/01/1985	09/01/1985	North-East	Constantine, Annaba, Sétif, Batna, Khenchela	Rhummel, Seybousse, Ksob, Kissir, Soultès, Yabous, Lahtiba	26 deaths
21/01/1992	29/01/1992	Central and Western	Algiers, Tipaza, Blida, Ain Defla, Chlef	Djer, Mazafran, El Harrach, Hamiz, Tikezal, Cheliff	21 deaths
19/10/1993	20/10/1993	Western	Relizane, Chlef, Tiaret	R'hiou, Cheliff, Menni	22 deaths
21/09/1994	26/09/1994	Eastern and Western	Borj-Bou-Arreridj, Tiaret, Tissemsilt, Bouira, Naama	Ksob, Aïn Zada, Touil, Mina, El Abed	32 deaths
23/04/1996	30/04/1996	Eastern	El Taref	Kebir, Kebir East, Bougous	4 deaths
06/08/2000	06/08/2000	Eastern	Borj-Bou-Arreridj, Oum El Bouaghi	Ksob, Aïn Zada, Boulfreis, El Hammam	7 deaths
13/10/2000	13/10/2000	Western	Ain Temouchent	Maleh, El Kihel	9 deaths
22/10/2000	25/10/2000	Western	Sidi Bel Abbès, Relizane, Tlemcen, Naama	Mekkerà, Taria, Mina, R'hiou, Tlata, Mouillah, Tafna	28 deaths
10/11/2001	14/11/2001	Central and Western	Algiers, Blida, Boumerdès, Tizi Ouzou, Tipaza, Chlef, Ain Defla, Relizane, Mascara, Oran, Mostaganem, Saida and Tlemcen	Bab el Oued, Korich, Djer, Mazafran, Kniss, Béni Messous, Tléta, El Kerma, Sidi Arrack, El Harrach, Annassers, Ardjem, Djemaïa, Isser, Djemâa, Ménaïel, Boumzar, Barbara, Sébaou, Guettar, El Hammam, Merdja, Boumerdès, Corso, Hamiz, Arbatache, Taksebt, M'leta, Cheliff, Tikezal, Zeddine, El Abtal, Goussine, Kramis, Mina, R'hiou, El Melah, Tlélat, Tafna, Mekkerà, Tlata, Taria	921 deaths and 300 million dollars
17/08/2002	28/08/2002	Eastern and Western	Tébessa, Batna, Mila, Guelma, Oum El Bouaghi, Khenchela, Souk Ahras, Sidi Bel Abbès, Annaba	Melegue, Mejerdah, Soultès, Seybousse, Kebir, Rhummel, Mekkerà, Taria	29 deaths and 1.5 million dollars of material damages
09/10/2002	12/10/2002	Eastern	Sétif, Batna, Biskra	Djemourah, Biskra, Abiod, Aïn Zada	13 deaths

1	07/12/2002	07/12/2002	Central	Béjaia, Tizi Ouzou	Sébaou, M'leta, Soummam	6 deaths
2	02/04/2003	05/04/2003	Central and Eastern	Tizi Ouzou, Béjaia, Annaba, Blida, Constantine	Sébaou, M'leta, Soummam, Djer, Mazafran, Rhummel	15 deaths
3	13/11/2004	14/11/2004	Central	Algiers	Bab el Oued, Korich, Djer, Mazafran, Kniss, Béni Messous, Tléta, El Kerma, Sidi Arrack, El Harrach, Annassers	19 deaths
4	21/03/2007	23/03/2007	Western	Relizane, Saida, Oran, Sidi Bel Abbès	Mina, R'hiou, Melah, Tlélat, Tafna, Mekker, Tlata, Taria	
5	24/11/2007	02/12/2007	Central and Western	Algiers, Tipaza, Boumerdès, Oran, Tizi Ouzou, Blida	Bab el Oued, Korich, Djer, Mazafran, Kniss, Béni Messous, Tléta, El Harrach, Annassers, Djemaïa, Isser, Djemâa, Larbaa, Sébaou, El Hammam, Boumerdès, Corso, Hamiz, Chiffa, Arbatache, Taksebt, M'leta, El Melah, Tlélat	14 deaths
6	03/02/2012	14/02/2012	Central and Eastern	Skikda, Boumerdès, Tizi Ouzou	Saf-Saf, Ameer, Boumerdès, Corso, Hamiz, Sébaou, M'leta	49 deaths

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Table 2. Characteristics of the selected stations

Watershed	Station	Wadi	Surface (km ²)	Length of the main thalweg (km)	Latitude	Longitude	Period	Annual average yield (Hm ³)
Cheliff	Ammi Moussa	R'hiou	1890	114.5	35.86	1.12	1961-2012	84
Algiers Coastal	Azzefoun	M'leta	36	13	36.89	4.35	1988-2012	11.8
Algiers Coastal	El Affroun	Djer	395	50	36.45	2.59	1970-1993	37.1
Algiers Coastal	Baraki	El Harrach	970	59	36.68	3.09	1972-2008	142.2
Eastern Constantine Coastal	Khemakhem	Saf-Saf	322	27	36.55	6.93	1975-2005	30.7
Eastern Constantine Coastal	Ain El Assel	Kebir East	680	41	36.58	8.37	1948-2003	207.8
Western Oran coastal	Ghazaoute	Tlata	100	8	35.1	-1.87	1971-2012	2.9
Highlands	Timgad	Soultès	194	29.8	35.49	6.46	1970-2005	5.3
Isser	Lakhdaria	Isser	3620	114	36.61	3.58	1967-2011	304.1
Macta	S.A. Ben Y	Mekkerà	1890	91.1	34.97	-0.74	1950-2009	20.4
Macta	Taria	Taria	1360	98	35.12	0.09	1974-2012	20.5

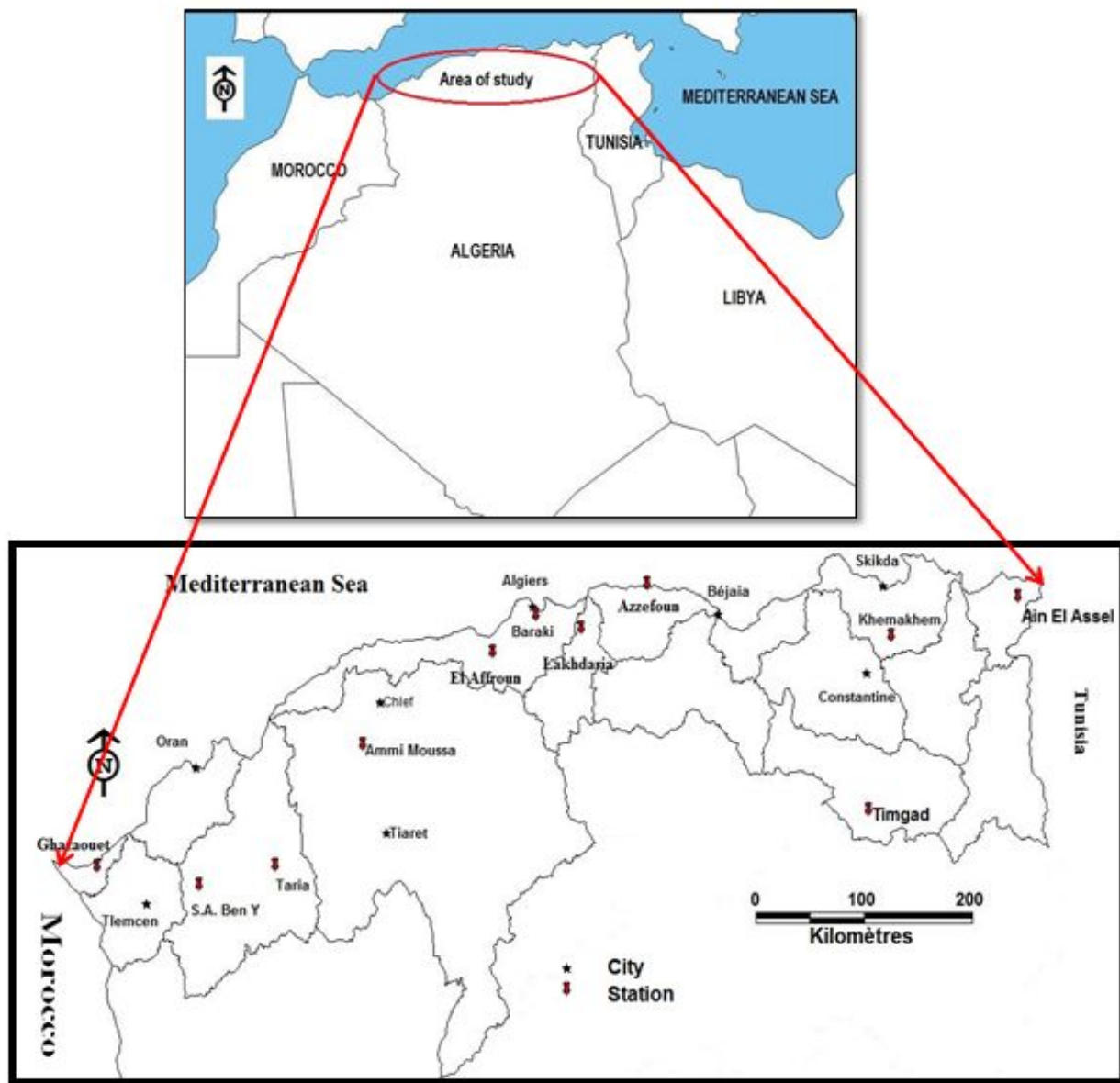


Figure 1. Localization of studied stations

Table 3. Descriptive analysis of univariate variables series

Stations	Ammi Moussa	Azzefoun	El Affroun	Baraki	Khemakhem	Ain El Assel	Ghazaoute	Timgad	Lakhdaria	S.A. Ben Y	Taria	
D	Minimum	3	3	6	7	2	4	2	1	3	1	3
	Maximum	16	11	22	20	17	15	18	12	17	13	19
	Mean	8.7	7.9	14.8	13.3	8.7	8.5	7.9	3.7	8.7	6.5	8.3
	Standard deviation	3.2	1.9	4.01	3.24	4.46	2.48	3.62	2.21	3.6	3.13	3.86
	Median	9	8	15	13	9	8	7	4	8	6	7
	Cv (Variation)	0.4	0.2	0.3	0.2	0.5	0.3	0.5	0.6	0.4	0.5	0.5
	Cs (Skewness)	0.2	-0.9	-0.3	0.1	0.3	0.2	1.1	1.4	0.6	0.4	0.9
	Ck (Kurtosis)	2.3	3.5	2.9	2.5	2	2.5	4.2	6.5	2.5	2.1	3.5
V	Minimum	1.8	0.05	0.2	3.8	0.2	3.1	0.03	0.1	5.1	0.1	0.3
	Maximum	74.3	9.3	57.8	140.4	51.4	143.6	5.5	9.5	340.9	21.2	20.9
	Mean	21.8	3.2	13.9	39.4	9.6	47.4	1.2	1.7	68.1	4.9	4.9
	Standard deviation	17.7	2.7	15.6	36.6	10.5	28.6	1.4	2.3	75.8	5	4.9
	Median	15.9	2.2	6.5	29.2	6.8	43.7	0.6	0.9	43.5	3.1	3.4
	Cv (Variation)	0.8	0.8	1.1	0.9	1.1	0.6	1.1	1.4	1.1	1	1
	Cs (Skewness)	1.4	1	1.4	1.5	2.7	1.3	1.5	2.4	2.2	1.7	1.7
	Ck (Kurtosis)	4.4	2.7	4.2	4.5	10.9	5.3	4.8	7.8	7.4	5.4	5.1
Qp	Minimum	17.3	0.5	0.7	19	1.9	6.4	0.14	0.7	28.7	0.8	2
	Maximum	428.6	71.2	268.1	876.4	132.9	467.9	33.8	77.6	1489.5	177.6	130
	Mean	123.9	21.3	69.7	191.6	42.4	199.7	8.3	15.5	303.5	38.6	34.7
	Standard deviation	86	23	72	223	32.2	98.9	9.3	18.8	302.6	43.5	32.7
	Median	100.2	11.8	42.3	120.1	37.1	200.3	3.8	7.7	213.7	20.7	21.8
	Cv (Variation)	0.7	1.1	1	1.2	0.7	0.5	1.1	1.2	1	1.1	0.9
	Cs (Skewness)	1.3	1.3	1.3	1.7	1.1	0.5	1.3	2.1	2.5	1.7	1.2
	Ck (Kurtosis)	5.2	3.2	3.9	5.2	4.1	3.2	3.6	6.6	8.9	5.3	3.6
Size	47	25	24	28	26	51	42	36	41	51	37	
Period	1961-2012	1988-2012	1970-1993	1972-2008	1975-2005	1948-2003	1971-2012	1970-2005	1967-2011	1950-2009	1974-2012	
Number of missing years	5	0	0	8	0	6	0	0	4	9	2	

Table 4. Correlation matrix of the analysed pairs

	Ammi Moussa	Azzefoun	El Affroun	Baraki	Khemakhem	Ain El Assel	Ghazaoute	Timgad	Lakhdaria	S.A. Ben Y	Taria
(D. V)	0.45	-0.15	0.28	0.03	0.62	0.37	0.55	0.23	0.42	0.26	0.3
(D. Qp)	0.25	-0.23	0.25	-0.2	0.46	0.23	0.38	0.17	0.25	0.15	0.1
(Qp. V)	0.86	0.95	0.92	0.92	0.85	0.78	0.95	0.96	0.88	0.94	0.8

Bold character: significant correlation

Table 5. Checking of the basic assumptions for the margins

	Station	Stationarity				Independence		Homogeneity	
		Mann-Kendall		Spearman		Wald-Wolfowitz		Wilcoxon	
		K	p-value	S	p-value	U	p-value	W	p-value
D	Ammi Moussa	-1.92	0.054	-1.79	0.08	1.86	0.063	1.43	0.15
	Azzefoun	-1.7	0.09	-0.94	0.36	0.94	0.35	0.8	0.43
	El Affroun	-1.33	0.18	-1.41	0.17	1.57	0.12	1.66	0.1
	Baraki	0.4	0.69	0.81	0.42	0.16	0.87	0.67	0.5
	Khemakhem	-0.89	0.38	-0.72	0.48	0.95	0.34	0.49	0.62
	Ain El Assel	-0.72	0.47	-0.4	0.69	1.86	0.06	0.73	0.46
	Ghazaoute	-1.15	0.25	-0.76	0.45	0.063	0.95	1.03	0.3
	Timgad	-0.8	0.42	0.086	0.93	0.8	0.42	1.25	0.21
	Lakhdaria	0.27	0.79	0.46	0.65	0.89	0.37	0.7	0.49
	S.A. Ben Y	1.64	0.1	1.98	0.05	0.76	0.45	2.56	0.01*
Taria	-1.82	0.07	-1.7	0.1	0.16	0.88	1.73	0.08	
V	Ammi Moussa	-1.55	0.12	-1.43	0.16	1.18	0.24	1.15	0.25
	Azzefoun	2.02	0.04*	2.29	0.03*	0.094	0.92	1.9	0.06
	El Affroun	-1.8	0.07	-1.88	0.08	0.48	0.63	0.7	0.48
	Baraki	-1.84	0.07	-2	0.06	0.51	0.61	0.53	0.6
	Khemakhem	0.26	0.79	0.21	0.84	0.3	0.76	1.41	0.16
	Ain El Assel	-0.49	0.63	-0.47	0.64	0.58	0.56	0.58	0.56
	Ghazaoute	-1.68	0.09	-1.66	0.1	0.57	0.57	1.32	0.19
	Timgad	0.041	0.97	-0.017	0.99	1.95	0.05	0.49	0.62
	Lakhdaria	-1.65	0.1	-1.53	0.13	1.24	0.21	1.4	0.16
	S.A. Ben Y	0.96	0.34	1.16	0.25	1.14	0.25	1.6	0.11
Taria	-0.052	0.96	0.17	0.87	1.74	0.08	0.81	0.42	
Qp	Ammi Moussa	-0.42	0.67	-0.25	0.81	0.42	0.67	0.064	0.95
	Azzefoun	1.9	0.06	2.7	0.01*	0.013	0.99	2.38	0.02*
	El Affroun	-1.58	0.11	-1.78	0.09	0.59	0.55	1.09	0.28
	Baraki	-1.8	0.07	-1.96	0.06	0.025	0.98	0.44	0.66
	Khemakhem	1.41	0.16	1.26	0.22	0.38	0.71	2.19	0.03*
	Ain El Assel	-1.66	0.1	-1.62	0.11	0.76	0.45	1.6	0.11
	Ghazaoute	-1.8	0.07	-1.81	0.08	0.11	0.92	1.37	0.17
	Timgad	0.52	0.6	0.42	0.67	1.71	0.09	0.81	0.42
	Lakhdaria	-1.91	0.06	-1.93	0.06	0.77	0.44	1.12	0.26
	S.A. Ben Y	1.22	0.22	1.31	0.2	1.12	0.26	1.66	0.1
Taria	1.53	0.13	1.52	0.14	2.19	0.03*	1.88	0.06	

*: accepted at level of 1% but rejected at 5%

Table 6. Selected marginal distribution

Variable	Station	Marginal laws	Estimation method	Parameter		
				1st	2nd	3
Duration	Ammi Moussa	Weibull	Maximum Likelihood	9.64	3.14	-
	Azzefoun	Weibull	Method of Moments	8.55	4.81	-
	El Affroun	Weibull	Method of Moments	16.33	4.17	-
	Baraki	Weibull	Method of Moments	14.57	4.68	-
	Khemakhem	Weibull	Maximum Likelihood	9.93	2.14	-
	Ain El Assel	Gamma	Maximum Likelihood	1.47	12.72	-
	Ghazaoute	Gumbel	Method of Moments	6.32	2.82	-
	Timgad	Weibull	Method of Moments	4.54	1.87	-
	Lakhdaria	Gumbel	Method of Moments	7.09	2.81	-
	S.A. Ben Y	Weibull	Method of Moments	7.35	2.19	-
Taria	Gumbel	Method of Moments	6.56	3.01	-	
Volume	Ammi Moussa	Log-Normal type 2	Maximum Likelihood	2.77	0.84	-
	Azzefoun	Exponential	Maximum Likelihood	0.85	2.53	-
	El Affroun	Weibull	Maximum Likelihood	15.16	0.99	-
	Baraki	Weibull	Method of Moments	40.5	1.08	-
	Khemakhem	Weibull	Method of Moments	9.22	0.91	-
	Ain El Assel	Weibull	Method of Moments	53.11	1.71	-
	Ghazaoute	Weibull	Maximum Likelihood	1.12	0.88	-
	Timgad	Log-Normal type 2	Maximum Likelihood	-0.15	1.2	-
	Lakhdaria	Log-Normal type 2	Maximum Likelihood	3.78	0.93	-
	S.A. Ben Y	Weibull	Maximum Likelihood	4.96	1.03	-
Taria	Log-Normal type 2	Maximum Likelihood	1.17	0.97	-	
Peak flow	Ammi Moussa	Gamma	Maximum Likelihood	0.017	2.15	-
	Azzefoun	Log-Normal type 2	Maximum Likelihood	2.84	0.87	-
	El Affroun	Weibull	Maximum Likelihood	72.27	0.98	-
	Baraki	Weibull	Maximum Likelihood	186.72	0.95	-
	Khemakhem	Weibull	Method of Moments	46.03	1.33	-
	Ain El Assel	Weibull	Method of Moments	225.49	2.12	-
	Ghazaoute	Weibull	Maximum Likelihood	7.61	0.85	-
	Timgad	Log-Normal type 2	Maximum Likelihood	2.13	1.17	-
	Lakhdaria	GEV	L-Moments	167.29	101.12	-0.39
	S.A. Ben Y	Weibull	Maximum Likelihood	37.86	0.96	-
Taria	Exponential	Maximum Likelihood	3.86	31.86	-	

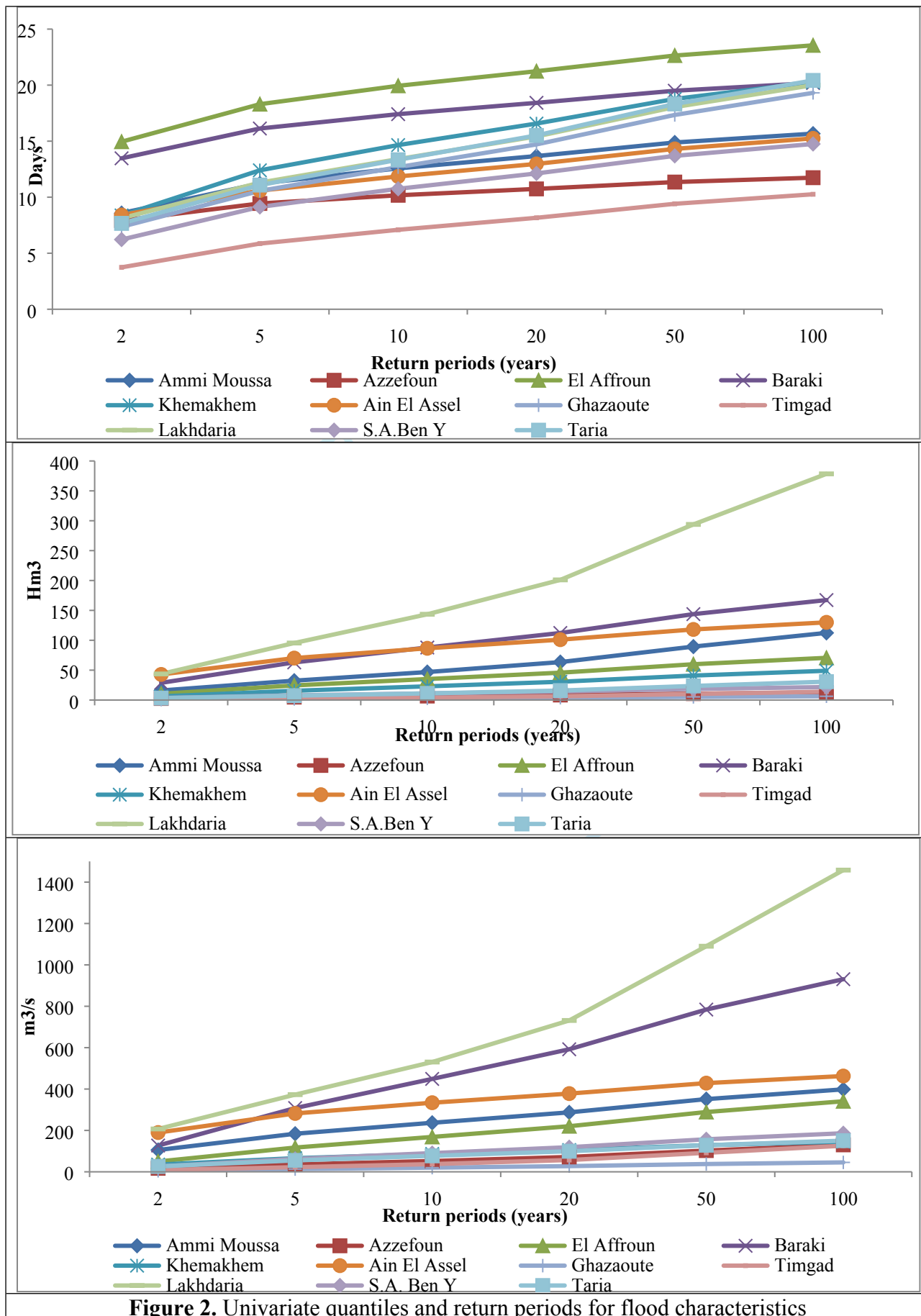


Figure 2. Univariate quantiles and return periods for flood characteristics

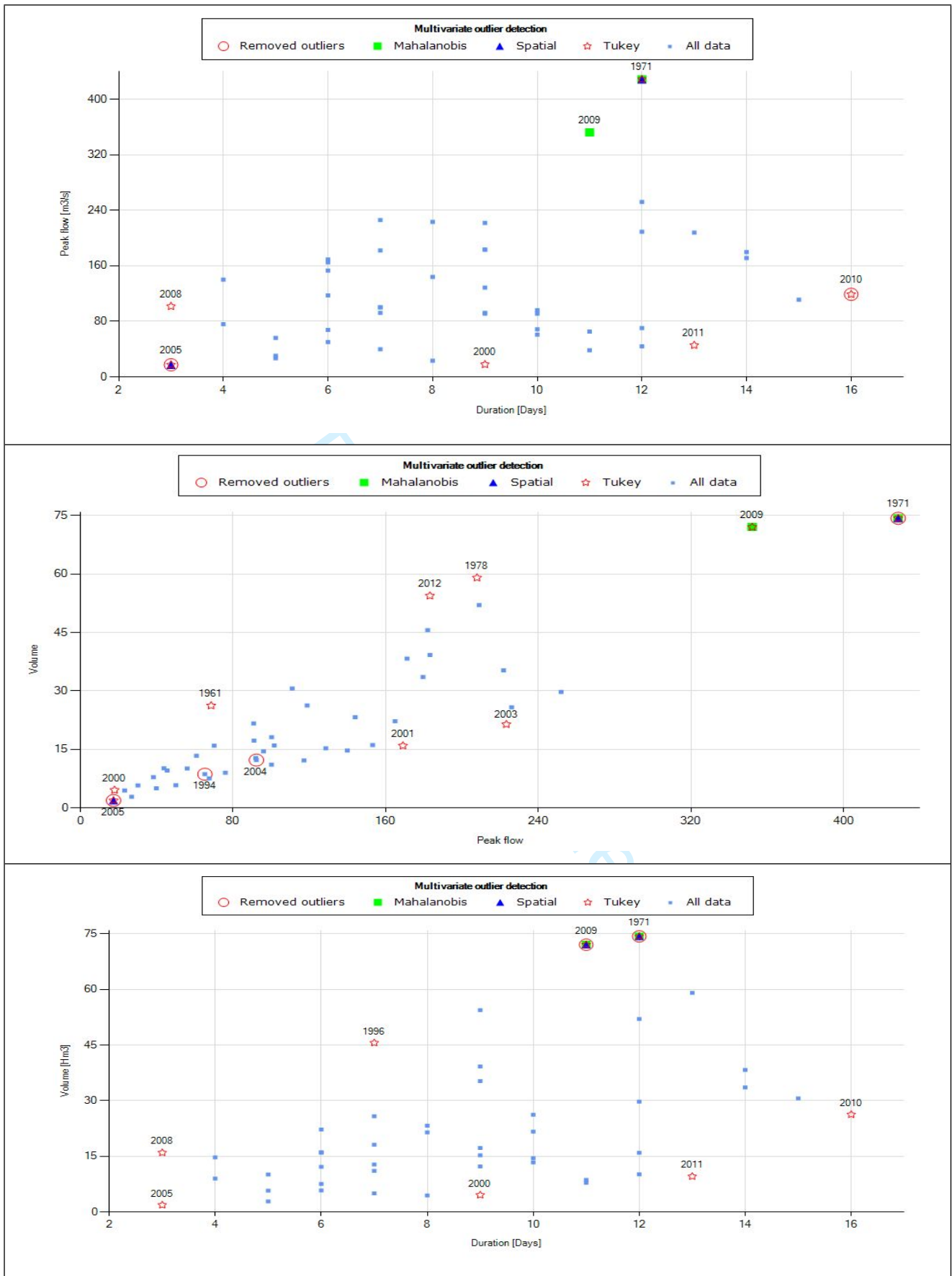


Figure 3. Multivariate outliers detection for the flood characteristics at Ammi Moussa

Table 7. Multivariate trend testing results with Mann–Kendall and Spearman (5% is the considered significance level)

Pair	Station	MK tests						Spearman test					
		CST		CET		CIT		CST		CET		CIT	
		Value	p-value	Value	p-value	Value	p-value	Value	p-value	Value	p-value	Value	p-value
(D, V)	Ammi Moussa	1.18	0.24	56869	0.07	3.77	0.15	0.97	0.33	4929433	0.15	2.73	0.26
	Azzefoun	0.14	0.89	8500	0.01*	7.8	0.02*	0.43	0.67	142885	0.05	5.82	0.05
	El Affroun	1.15	0.25	6074	0.08	3.97	0.14	1.13	0.26	172296.5	0.1	3.75	0.15
	Baraki	0.58	0.56	9277	0.16	3.75	0.15	0.43	0.67	525925	0.12	4.47	0.11
	Khemakhem	0.21	0.83	1850	0.56	2.23	0.33	0.17	0.86	48938.5	0.67	1.46	0.48
	Ain El Assel	0.38	0.7	9077	0.7	0.53	0.77	0.097	0.92	251914.5	1.00	0.16	0.92
	Ghazaoute	1.02	0.31	35572	0.13	3.08	0.21	0.8	0.43	2732876.5	0.22	2.61	0.27
	Timgad	0.12	0.9	2474	0.66	1.02	0.6	0.22	0.83	39218.5	0.9	0.21	0.9
	Lakhdaria	0.49	0.63	22529	0.23	4.61	0.1	0.22	0.83	2517205	0.21	5.89	0.05
	S.A. Ben Y Taria	0.8	0.42	40322	0.23	2.28	0.32	1.12	0.26	12675744.5	0.07	4.78	0.09
(D, Qp)	Ammi Moussa	0.72	0.47	29045	0.24	2.79	0.25	0.49	0.62	2108665	0.44	1.67	0.43
	Azzefoun	0.33	0.74	9316	0.01*	7.94	0.02*	0.71	0.48	242521	0.02*	8.17	0.02
	El Affroun	1.22	0.22	8692	0.07	4.33	0.11	1.27	0.2	331124	0.06	4.76	0.09
	Baraki	0.61	0.54	8905	0.17	3.31	0.19	0.46	0.65	511394	0.13	3.69	0.16
	Khemakhem	0.29	0.77	3145	0.36	2.52	0.28	0.27	0.79	655001	0.54	1.46	0.48
	Ain El Assel	0.87	0.38	42365	0.22	2.58	0.28	0.55	0.58	4903528.5	0.32	2.35	0.31
	Ghazaoute	1.1	0.27	39125	0.1	3.68	0.16	0.88	0.38	3197912.5	0.18	3.12	0.21
	Timgad	0.66	0.51	7081	0.28	2.61	0.27	0.34	0.73	75245	0.77	0.49	0.78
	Lakhdaria	0.37	0.71	20714	0.19	4.16	0.12	0.13	0.9	2492786.5	0.11	5.64	0.06
	S.A. Ben Y Taria	0.79	0.43	29125	0.3	1.91	0.39	1.09	0.27	8789645	0.09	4.15	0.13
(Qp, V)	Ammi Moussa	0.26	0.8	9197	0.5	4.99	0.08	0.13	0.9	874602	0.54	5.56	0.06
	Azzefoun	1.23	0.22	6760	0.05	4.05	0.13	1.32	0.19	191801	0.04*	4.48	0.11
	El Affroun	0.94	0.35	5570	0.13	3.93	0.14	0.95	0.34	167858.5	0.13	3.52	0.17
	Baraki	1.18	0.24	17300	0.06	3.53	0.17	1.21	0.23	872581	0.06	3.67	0.16
	Khemakhem	0.4	0.69	1828	0.47	2.9	0.23	0.26	0.79	38034	0.59	2.7	0.26
	Ain El Assel	0.61	0.54	33410	0.28	3.35	0.19	0.58	0.56	4695522.5	0.31	3.03	0.22
	Ghazaoute	1.12	0.26	52225	0.08	3.32	0.19	1.07	0.28	5297976.5	0.09	3.22	0.2
	Timgad	0.42	0.68	4505	0.5	3.00	0.22	0.37	0.71	270442	0.55	2.57	0.28
	Lakhdaria	1.03	0.31	35482	0.1	4.35	0.11	0.96	0.34	2988449	0.12	4.53	0.1
	S.A. Ben Y Taria	0.45	0.65	13700	0.48	1.05	0.59	0.55	0.58	2987345	0.39	0.89	0.64
		0.14	0.89	1933	0.64	7.33	0.03*	0.066	0.95	82100	0.72	6.04	0.05

CST: Covariance Sum Test, CET: Covariance Eigenvalue Test, CIT: Covariance Inversion Test.

*: accepted at level of 1%

Table 8. Tests of serial independence and homogeneity for multivariate series

Pair	Station	Multivariate independence test		Homogeneity multivariate tests							
				C-test		M-test		W-test		Z-test	
		Value	p-value	Value	p-value	Value	p-value	Value	p-value	Value	p-value
(D, V)	Ammi Moussa	0.0035	0.08	9.02	0.27	0.58	0.14	0.9	1	0.0011	0.97
	Azzefoun	0.000093	0.79	2.19	0.25	0.26	0.13	0.7	1	0.64	0.43
	El Affroun	0.00052	0.24	16.12	0.13	0.22	0.04*	0.61	1	0.31	0.58
	Baraki	0.0003	0.73	15.6	0.43	0.94	0.77	0.97	1	0.099	0.75
	Khemakhem	0.0018	0.19	8.42	0.16	0.4	0.16	0.75	1	0.66	0.41
	Ain El Assel	0.00092	0.51	18.03	0.3	0.85	0.49	0.93	1	1.02	0.31
	Ghazaoute	0.0012	0.28	3.11	0.17	0.81	0.49	0.86	1	0.98	0.32
	Timgad	0.0024	0.07	0.71	0.83	0.88	0.6	0.92	1	0.014	0.91
	Lakhdaria	0.0012	0.34	68.38	0.1	0.63	0.26	0.79	1	0.36	0.55
	S.A. Ben Y Taria	0.0012	0.33	9.17	0.02*	0.56	0.16	0.77	1	0.74	0.39
(D, Qp)	Ammi Moussa	0.0021	0.07	10.76	0.96	0.77	0.35	0.89	1	0.0066	0.93
	Azzefoun	0.00037	0.17	16.89	0.15	0.67	0.51	0.76	1	0.13	0.72
	El Affroun	0.00055	0.31	70.66	0.11	0.31	0.07	0.57	1	0.053	0.82
	Baraki	0.0002	0.56	58.71	0.61	0.97	0.96	0.99	1	0.26	0.61
	Khemakhem	0.00077	0.26	32.32	0.09	0.37	0.11	0.77	1	1.23	0.27
	Ain El Assel	0.0015	0.12	99.41	0.13	0.72	0.15	0.84	1	0.085	0.77
	Ghazaoute	0.00083	0.39	7.57	0.19	0.61	0.14	0.88	1	0.88	0.35
	Timgad	0.0021	0.06	1.67	0.94	0.95	0.92	1.00	1	0.001	0.97
	Lakhdaria	0.0004	0.95	121.94	0.33	0.55	0.2	0.84	1	0.14	0.71
	S.A. Ben Y Taria	0.00082	0.38	52.74	0.06	0.28	0.03*	0.8	1	0.31	0.58
(Qp, V)	Ammi Moussa	0.0084	0.06	22.69	0.68	0.61	0.25	0.89	1	0.59	0.44
	Azzefoun	0.0047	0.06	17.06	0.23	0.69	0.59	0.84	1	0.25	0.61
	El Affroun	0.001	0.56	49.15	0.26	0.28	0.33	0.83	1	1.76	0.18
	Baraki	0.0011	0.74	61.66	0.6	0.76	0.54	0.92	1	0.013	0.91
	Khemakhem	0.0011	0.58	29.72	0.1	0.44	0.42	0.77	1	1.83	0.18
	Ain El Assel	0.00098	0.93	97.26	0.14	0.57	0.07	0.81	1	0.41	0.52
	Ghazaoute	0.0022	0.31	6.76	0.2	0.66	0.12	0.83	1	0.0013	0.97
	Timgad	0.005	0.07	4.35	0.62	0.68	0.28	0.78	1	0.32	0.57
	Lakhdaria	0.001	0.95	148.3	0.28	0.84	0.59	0.89	1	0.18	0.67
	S.A. Ben Y Taria	0.0028	0.25	44.46	0.08	0.53	0.11	0.87	1	1.21	0.27
		0.0029	0.14	6.6	0.8	0.4	0.1	0.68	1	0.024	0.88

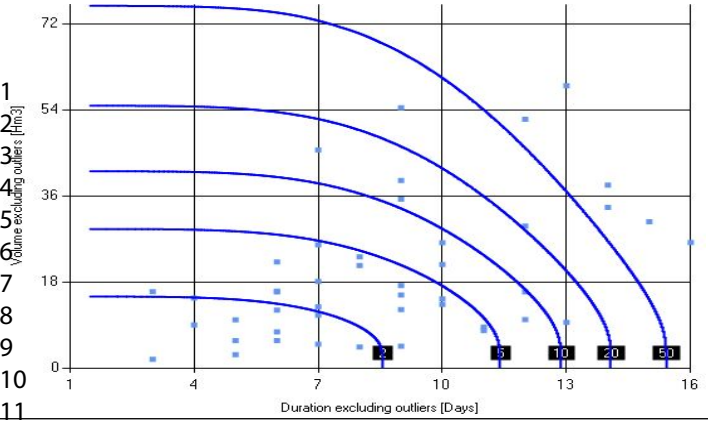
W-test: Wilcox test; Z-test: Zhang test; C-test: Cramer test

Table 9. Retained copula for flood characteristics

	Ammi Moussa	Azzefoun	El Affroun	Baraki	Khemakhem	Ain El Assel	Ghazaoute	Timgad	Lakhdaria	S.A. Ben Y	Taria
(D, V)	Clayton	Frank	Clayton	Clayton	Gumbel	Frank	Joe	Gumbel	Gumbel	Clayton	Gumbel
Parameter	0.85	-1.36	1.15	0.39	2.02	2.72	1.88	1.2	1.61	1.16	1.25
Sn p-value	0.0005	0.0005	0.09	0.05	0.09	0.0005	0.0035	0.0005	0.0005	0.0015	0.0005
Tn p-value	0.17	0.06	0.005	0.03	57	0.0045	0.21	0.05	0.28	0.08	0.19
AIC	-8.56	1.25	-4.76	0.29	-13.84	-7.14	-11.61	0.19	-11.96	-16.02	-1.39
BIC	-6.75	2.25	-3.67	1.63	-12.58	-5.23	-9.87	1.59	-10.25	-14.11	0.22
(D, Qp)	Frank	Ali-Mikhail-Haq	Clayton	Frank	Joe	Frank	Joe	Clayton	Frank	Clayton	Frank
Parameter	1.24	-0.51	0.52	-1.17	1.61	1.53	1.68	0.32	1.67	0.54	-0.62
Sn p-value	0.0015	0.0005	0.004	0.007	0.03	0.0005	0.0015	0.0005	0.0005	0.0005	0.0005
Tn p-value	0.009	0.26	0.45	0.06	0.58	0.017	0.07	0.05	0.07	0.08	0.1
AIC	0.16	1.73	-0.19	1.08	-2.43	-0.86	-7.36	0.97	-0.57	-2.88	1.64
BIC	1.97	2.77	0.98	2.41	-1.25	1.05	-5.62	2.34	1.06	-1.01	3.16
(Qp, V)	Clayton	Joe	Frank	Frank	Gumbel	Clayton	Gumbel	Gumbel	Frank	Gumbel	Frank
Parameter	3.27	6.52	17.65	14.16	3.24	2.62	6.55	10.67	10.44	4.22	19.33
Sn p-value	0.14	0.7	0.51	0.701	0.38	0.17	0.39	0.29	0.83	0.33	0.11
Tn p-value	0.84	0.17	0.13	0.704	0.83	0.13	0.72	0.36	0.15	0.36	0.73
AIC	-54.73	-36.47	-41.59	-46.35	-31.37	-49.42	-115.73	-128.57	-47.61	-95.73	-71.48
BIC	-52.97	-35.47	-40.5	-45.02	-30.19	-47.51	-113.99	-127.03	-45.94	-93.84	-69.98

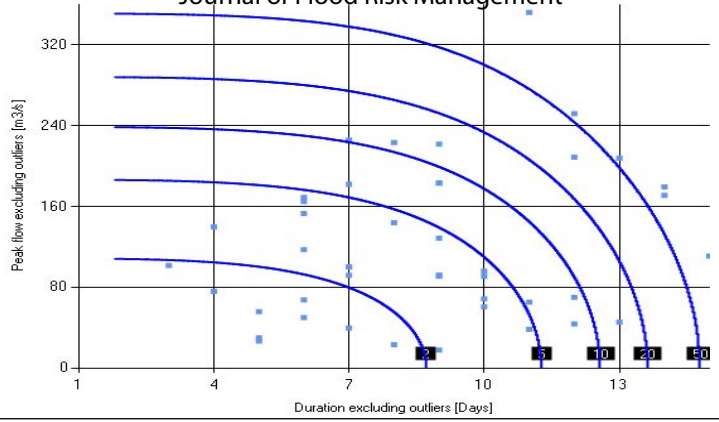
Sn test: based on empirical copula valid for any copula; Tn test: based on K function, valid for Archimedean copulas.

a1



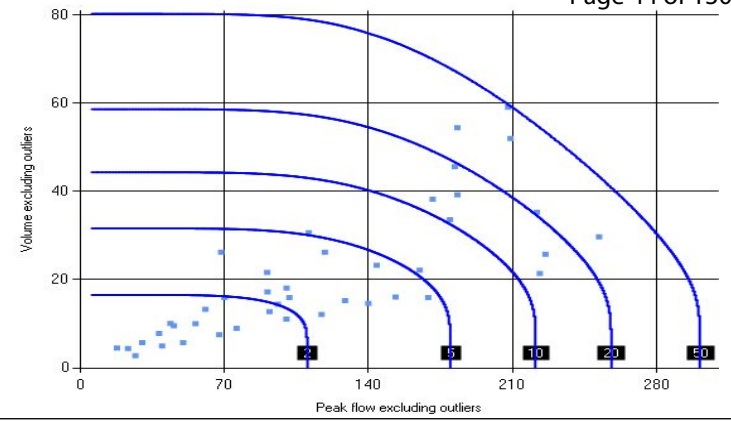
(D, V) Ammi Moussa (Clayton, 0.45)

a2



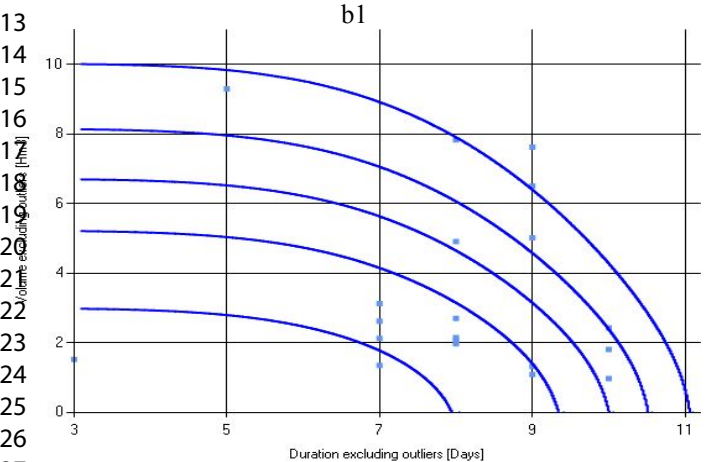
(D, Qp) Ammi Moussa (Frank, 0.25)

a3



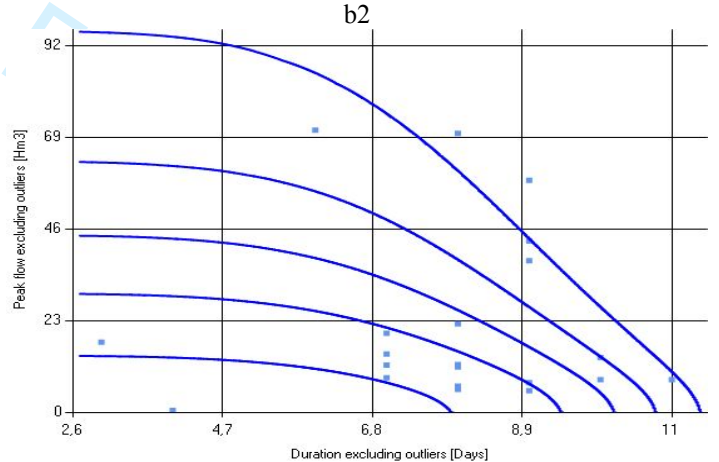
(Qp, V) Ammi Moussa (Clayton, 0.86)

b1



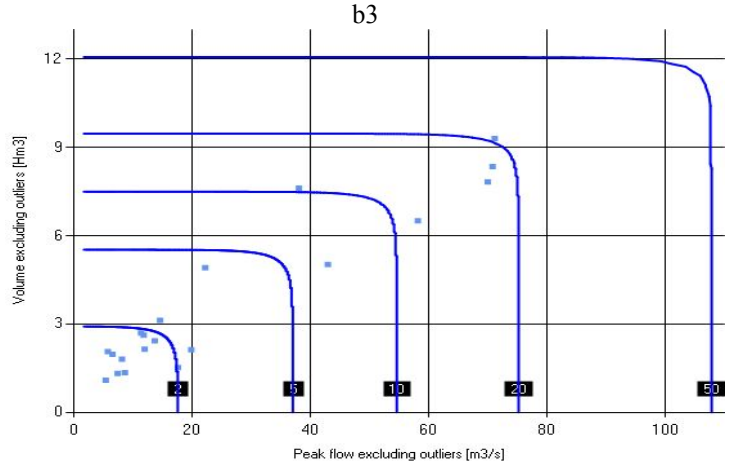
(D, V) Azzefoun (Frank, -0.15)

b2



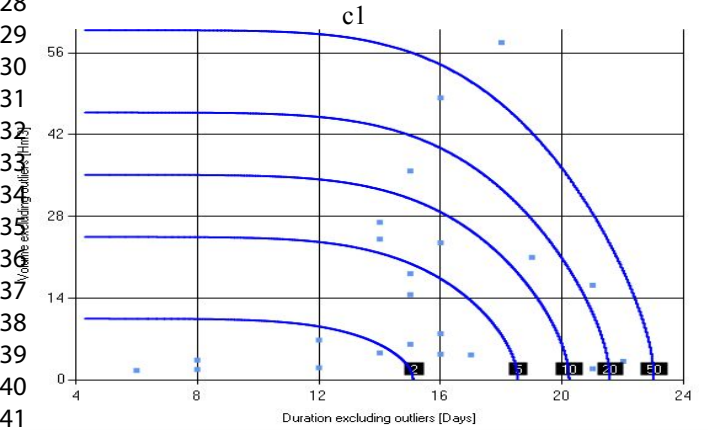
(D, Qp) Azzefoun (Ali-Mikhail-Haq, -0.23)

b3



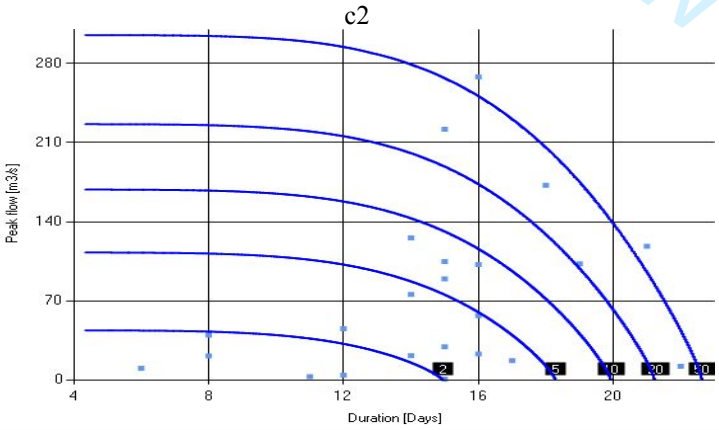
(Qp, V) Azzefoun (Joe, 0.95)

c1



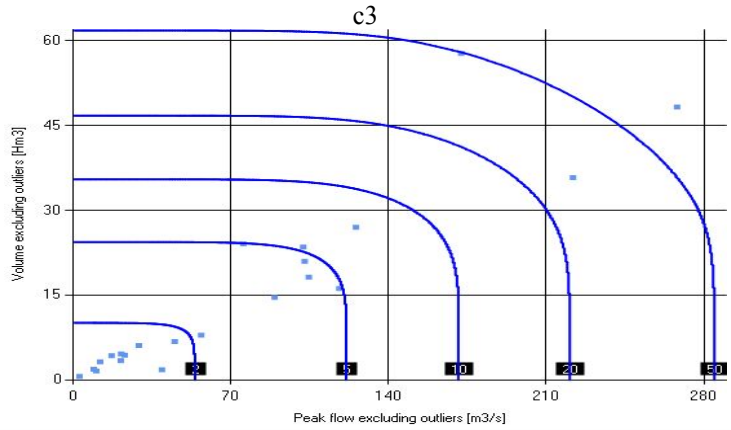
(D, V) El Affroun (Clayton, 0.28)

c2



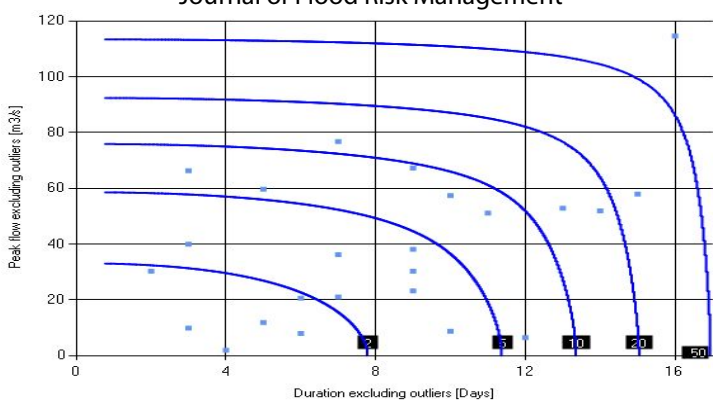
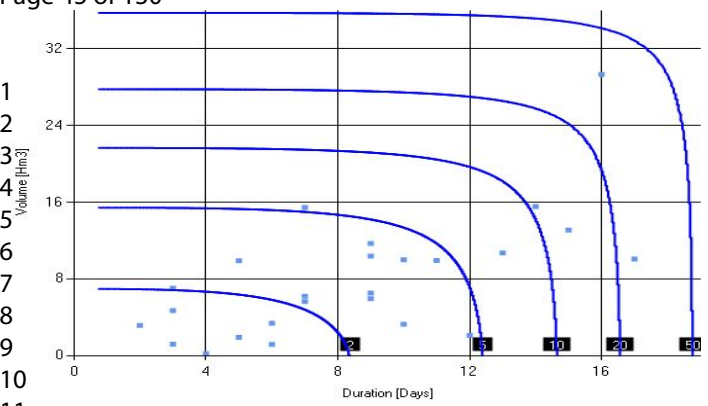
(D, Qp) El Affroun (Clayton, 0.25)

c3

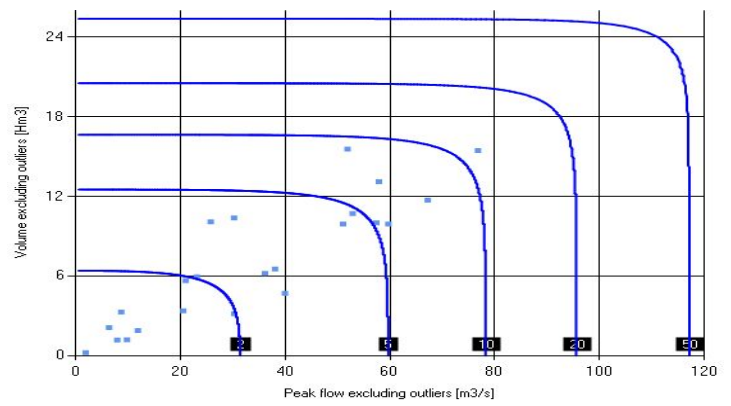


(Qp, V) El Affroun (Frank, 0.92)

d1



d3

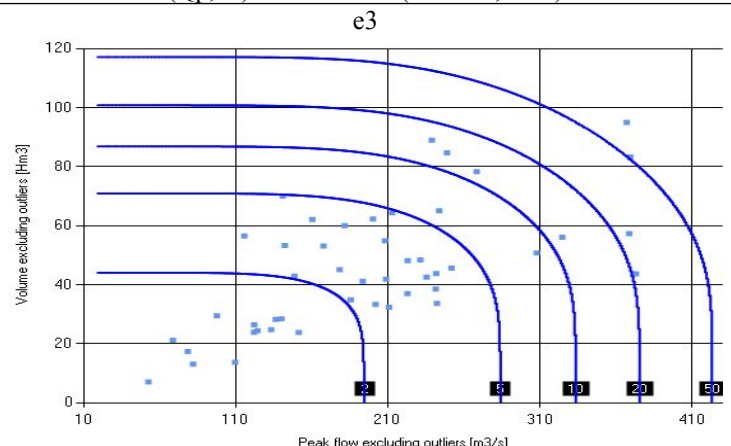
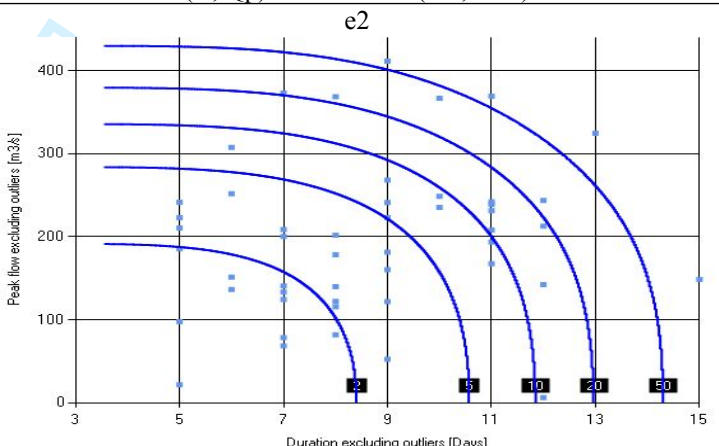
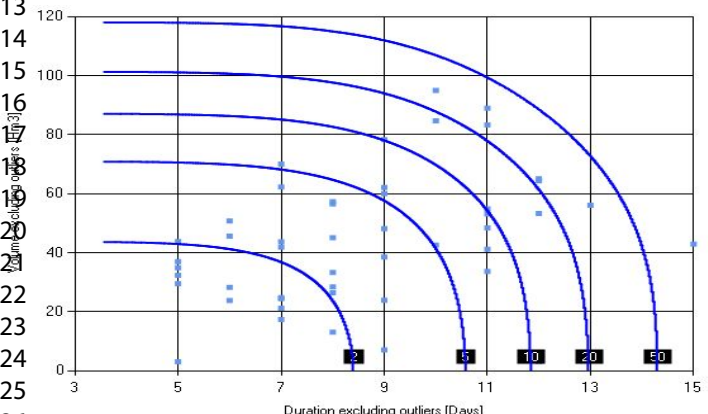


(D, V) Khemakhem (Gumbel, 0.62)

(D, Qp) Khemakhem (Joe, 0.46)

(Qp, V) Khemakhem (Gumbel, 0.85)

e1

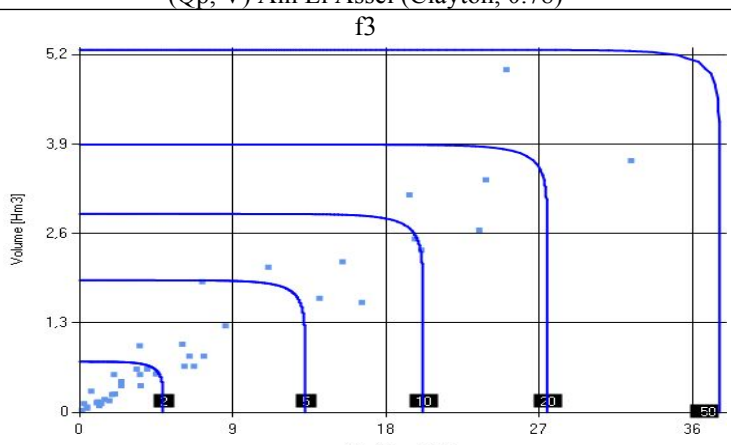
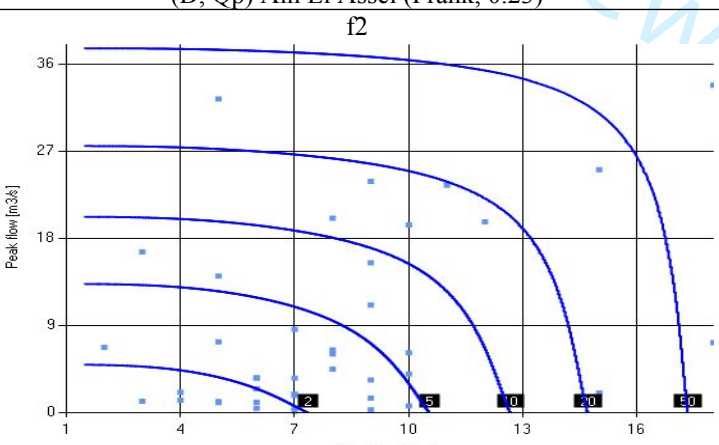
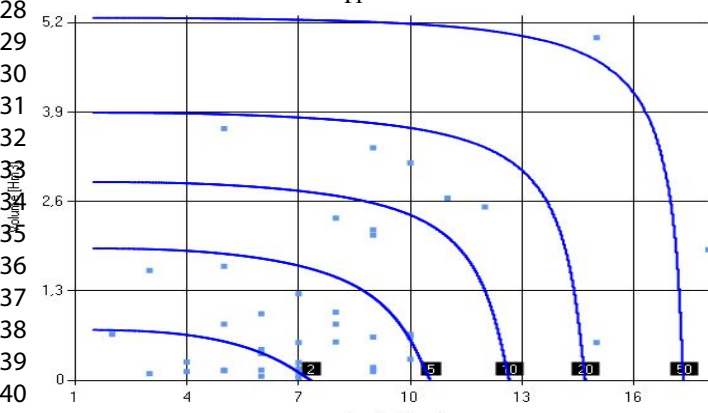


(D, V) Ain El Assel (Frank, 0.37)

(D, Qp) Ain El Assel (Frank, 0.23)

(Qp, V) Ain El Assel (Clayton, 0.78)

f1

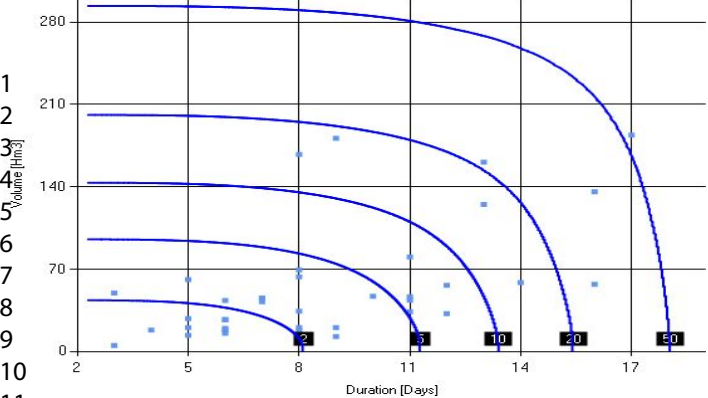


(D, V) Ghazaouet (Joe, 0.55)

(D, Qp) Ghazaouet (Joe, 0.38)

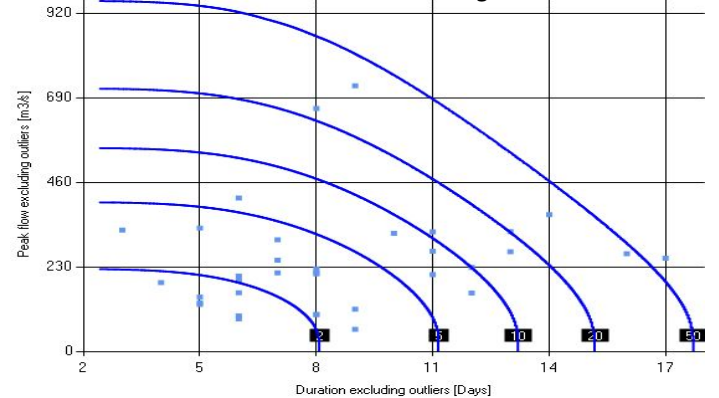
(Qp, V) Ghazaouet (Gumbel, 0.95)

g1



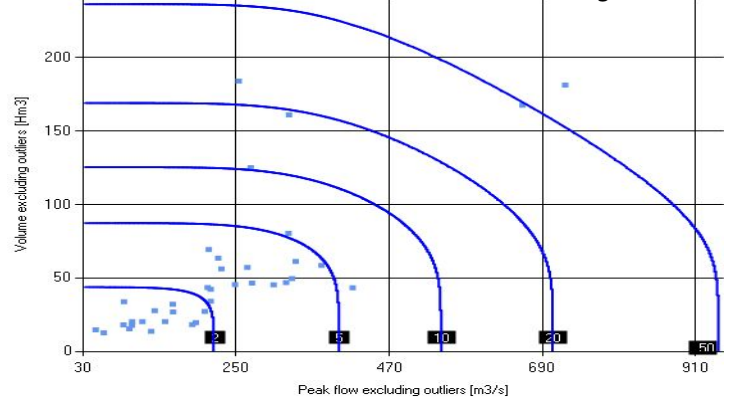
(D, V) Lakhdaria (Gumbel, 0.42)

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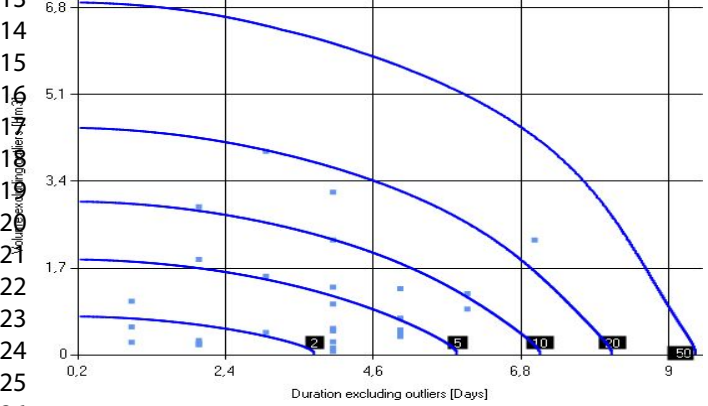
(D, Qp) Lakhdaria (Frank, 0.25)

g3



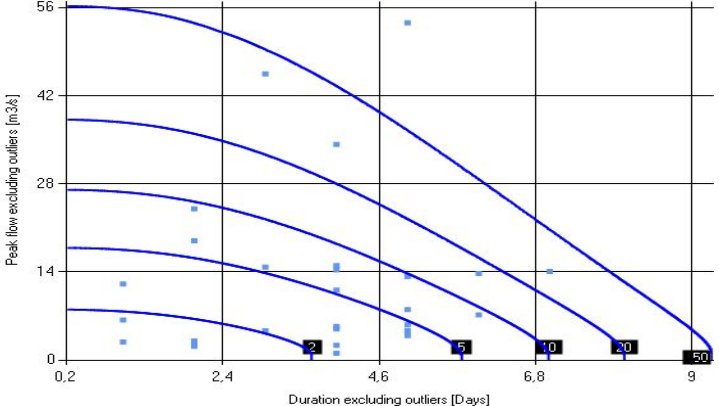
(Qp, V) Lakhdaria (Frank, 0.88)

h1



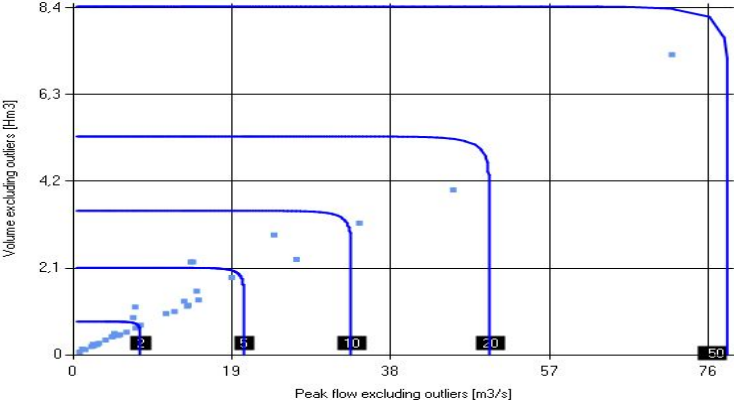
(D, V) Timgad (Gumbel, 0.23)

h2



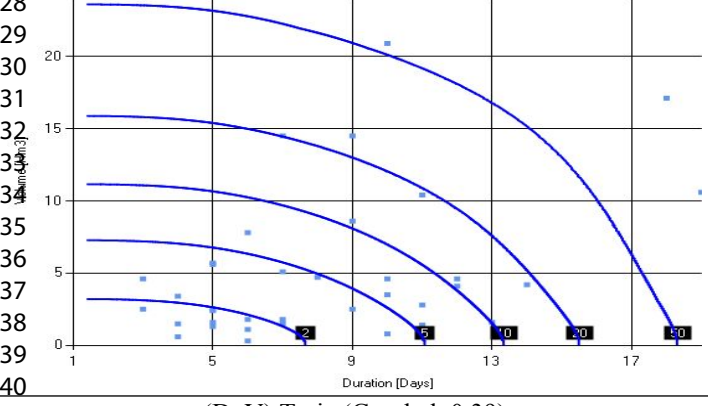
(D, Qp) Timgad (Clayton, 0.17)

h3



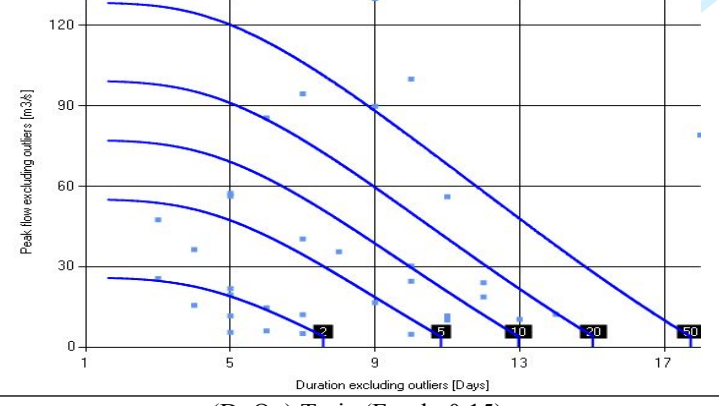
(Qp, V) Timgad (Gumbel, 0.96)

i1



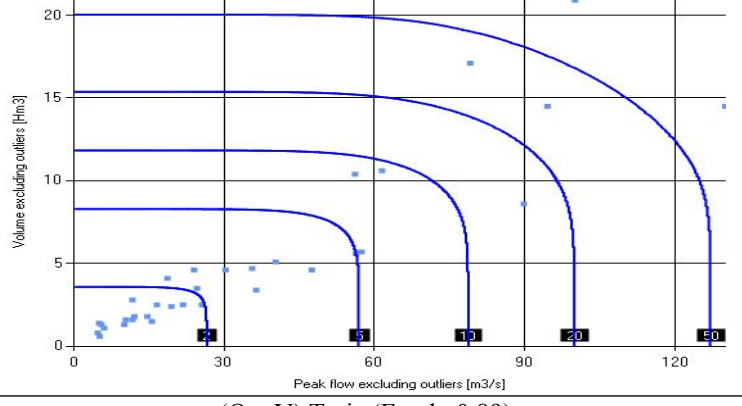
(D, V) Taria (Gumbel, 0.39)

i2



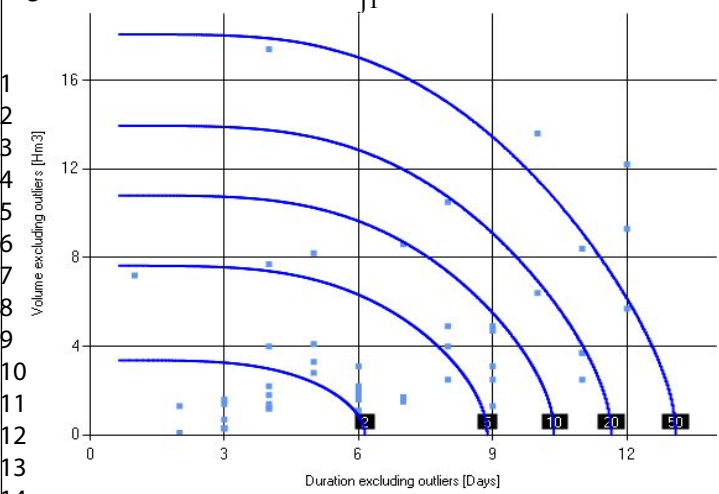
(D, Qp) Taria (Frank, 0.15)

i3

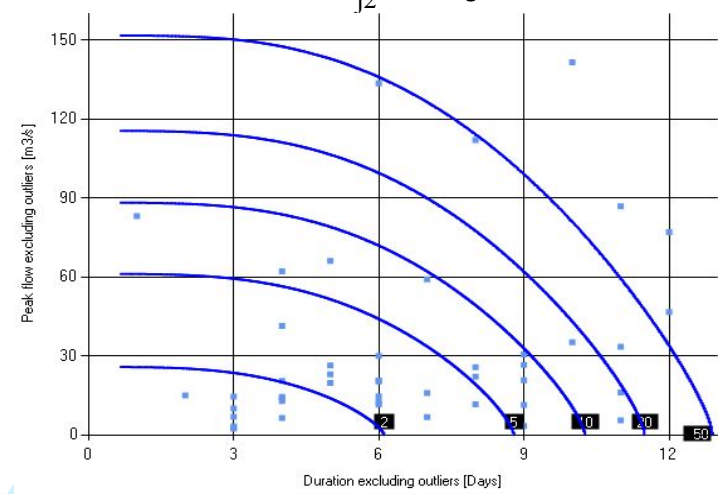


(Qp, V) Taria (Frank, 0.89)

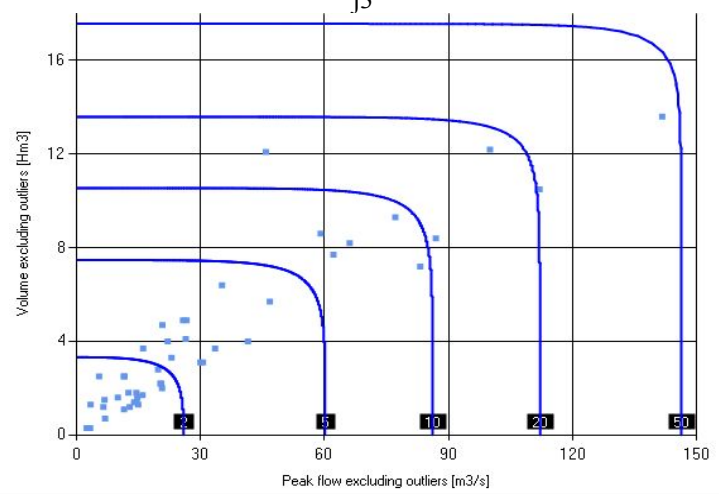
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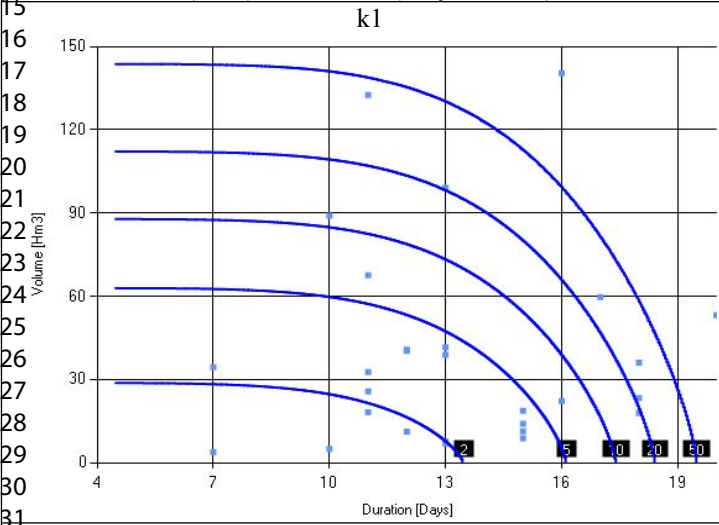
(D, V) S.A. Ben Y (Clayton, 0.26)



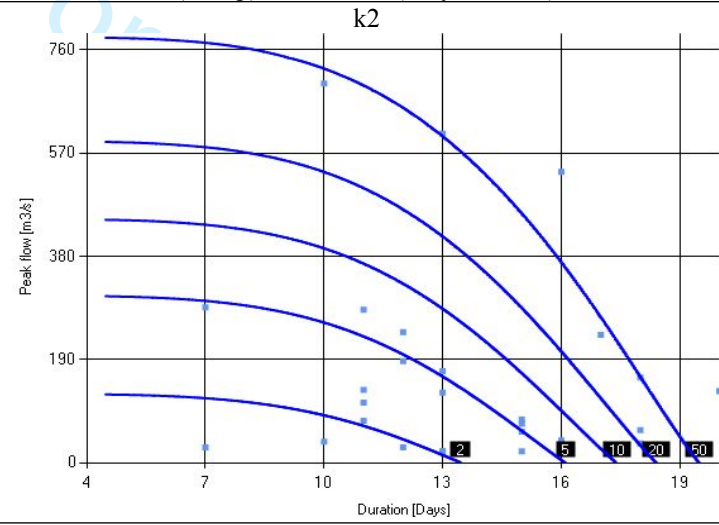
(D, Qp) S.A. Ben Y (Clayton, 0.15)



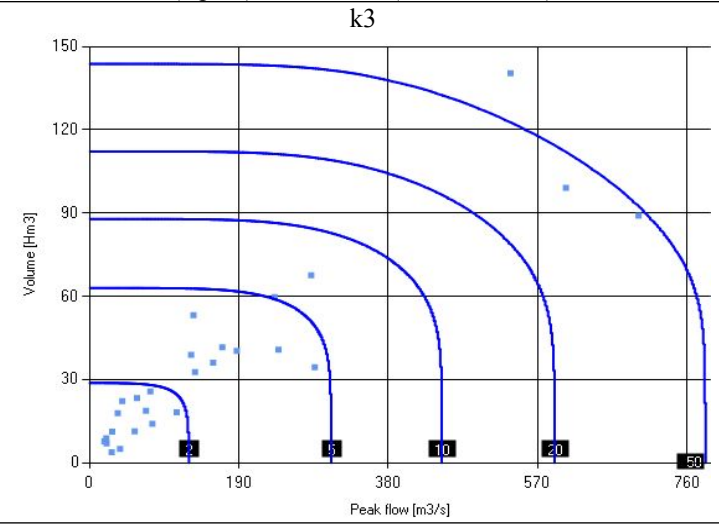
(Qp, V) S.A. Ben Y (Gumbel, 0.94)



(D, V) Baraki (Clayton, 0.03)



(D, Qp) Baraki (Frank, -0.2)



(Qp, V) Baraki (Frank, 0.92)

Figure 4. Contour line of pairs according to the selected copula for return period event of $X \geq x \cap Y \geq y$

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